

Non-orographic and orographic sources of gravity waves above Antarctica and the Southern Ocean



Riwal PLOUGONVEN

Laboratoire de Météorologie Dynamique, IPSL, École Normale Supérieure, Paris, France

Albert HERTZOG

Laboratoire de Météorologie Dynamique, École Polytechnique, UPMC, Paris, France

Lionel GUEZ

Laboratoire de Météorologie Dynamique, École Normale Supérieure, Paris, France

Introduction

Context and motivation: Importance of gravity waves for the stratosphere:

- **locally:** temperature fluctuations, turbulence when breaking occurs;
- **globally:** contribution to the stratospheric circulation (momentum fluxes, *Alexander et al 2010*).

Vorcore campaign (Sept. 2005 - Feb. 2006): 27 super-pressure balloons in the stratospheric polar vortex (16-19km) → 150,000 independent observations (*Hertzog et al 2007, 2008*).

Ongoing project: Worcore: Complement data analysis from Vorcore by mesoscale simulations, to investigate sources in particular. Implementation:

1. case study of a large-amplitude mountain wave (*Plougonven, Hertzog & Teitelbaum 2008*);
2. sensitivity study for simulations in a large domain (*Plougonven et al 2010*);
3. comparisons between the simulations and the observations: (**present study**);

Numerical setup

The Weather Research and Forecast Model (*Skamarock et al 2005*) is used with:

- a large domain: $10,000 \times 10,000$ km with $dx = 20$ km, 120 vertical levels (about every 300m) up to 5 hPa;
- runs lasting 3 days, with 1 day of spin-up;
- 29 runs, yielding 58 days to be investigated, from 21/10/2005 to 18/12/2005 (output every 6 hours);
- sponge layer in the last 6 km;

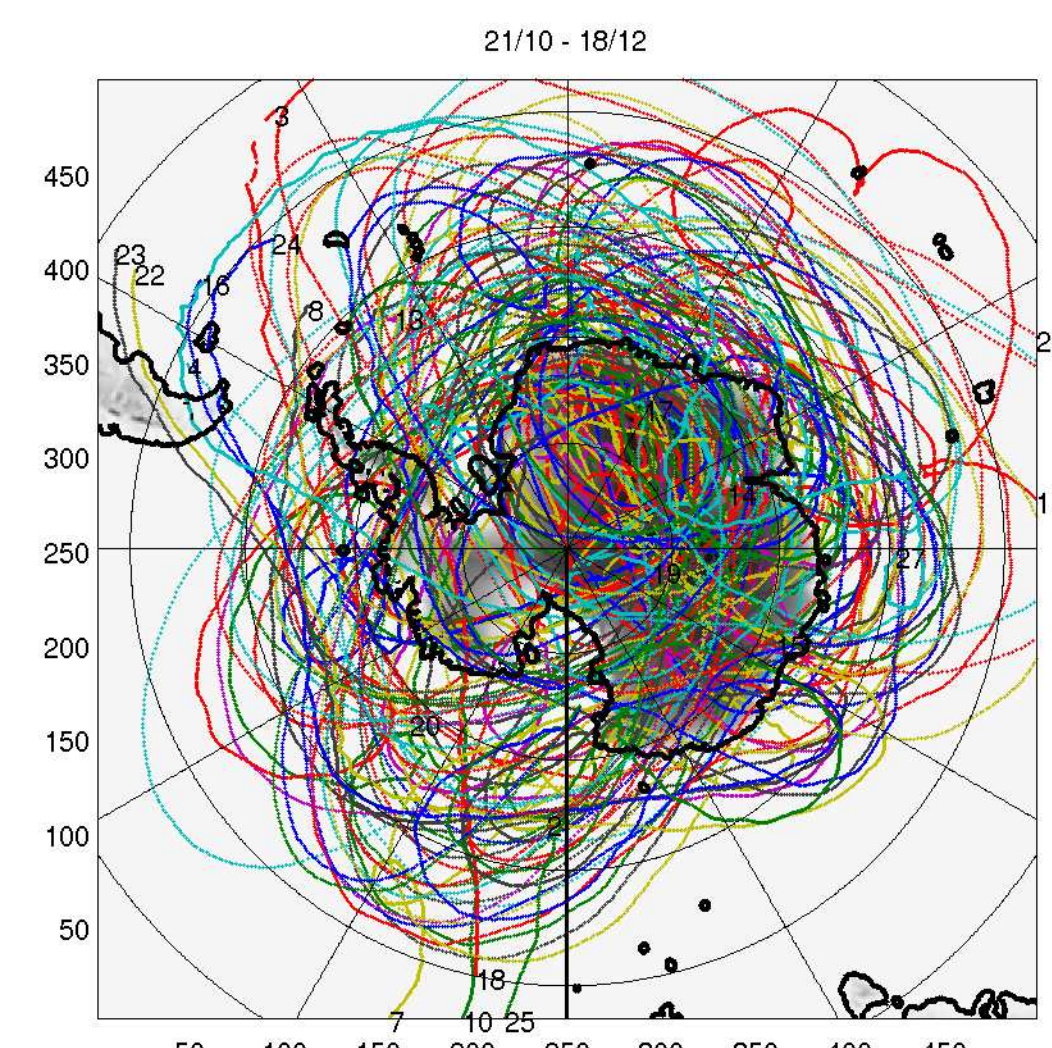


FIG. 1: Domain for the mesoscale simulations, topography of Antarctica and trajectories of the balloons during the simulated period, from 21/10/2005 to 18/12/2005.

Time average and comparison with the balloons

First step to assess the realism of the simulations: comparison of the simulated and observed **momentum fluxes** ($u'w'$, Fig. 2, and $v'w'$, Fig. 3):

- similar orders of magnitude;
- comparable overall structure, with dominance of topographic regions;
- some overestimation above topography, significant underestimation (by about a factor 2) over oceanic regions.

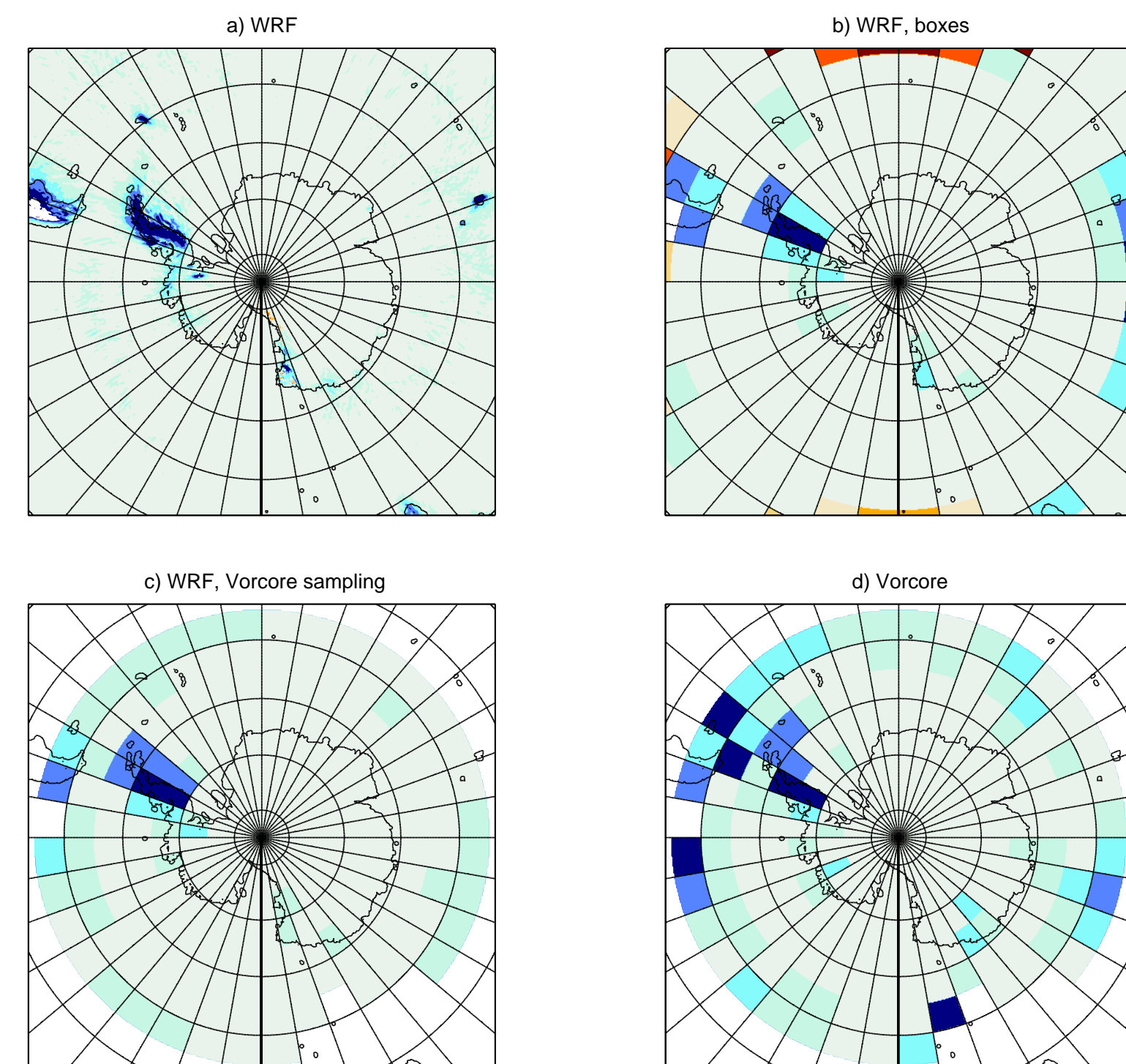


FIG. 2: At $z = 19$ km, zonal momentum fluxes in the simulations (full resolution: a); in boxes $10^\circ \times 5^\circ$ in longitude - latitude b); and sampled only at times when a balloon was present c)) and observations (panel d). Color range: -10 to 10 mPa.

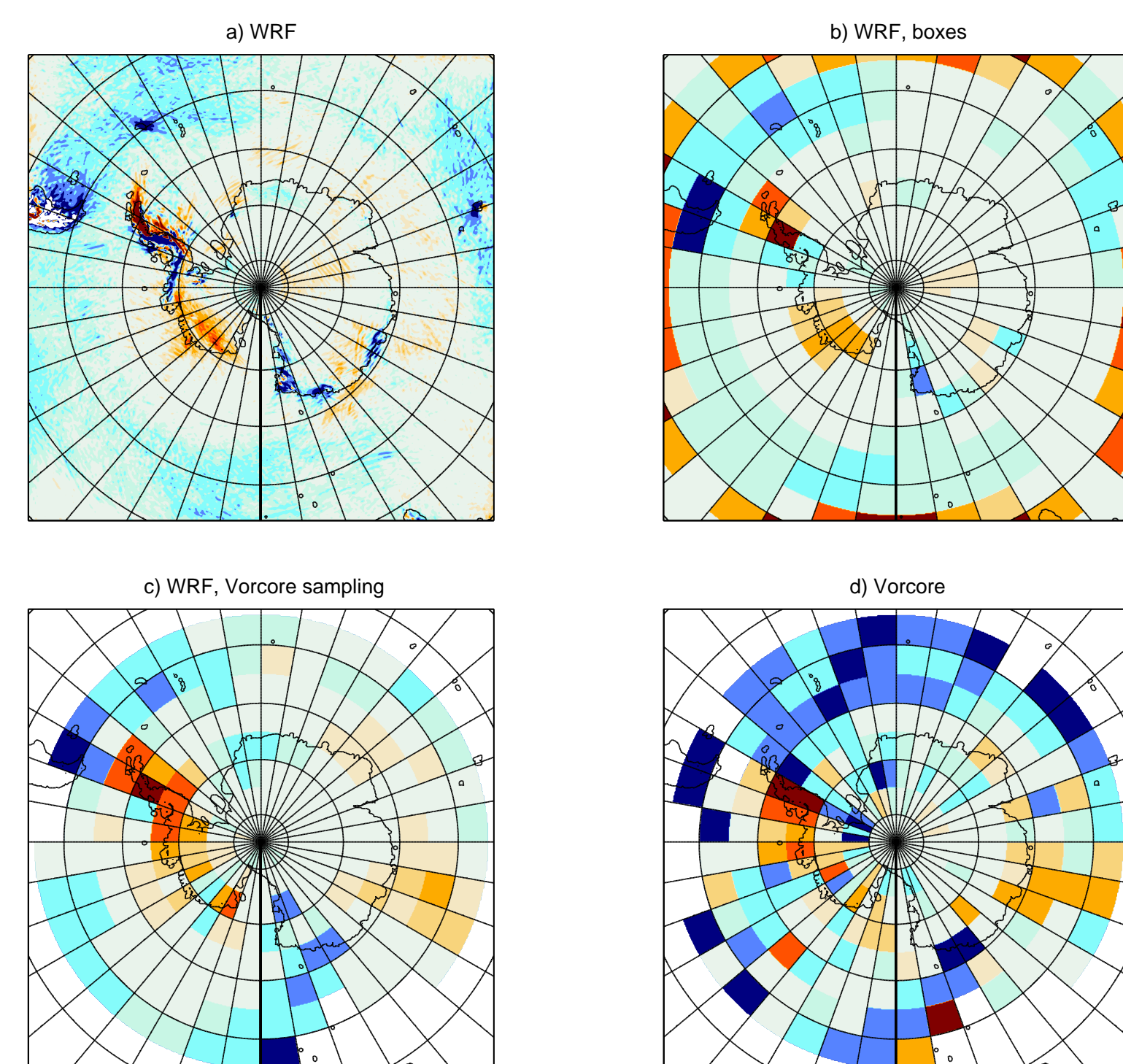


FIG. 3: Same as Fig. 2, but for the meridional momentum fluxes. Color range: from -2 to 2 mPa.

Variations in time

7 regions are defined as shown in **Figure 4**.

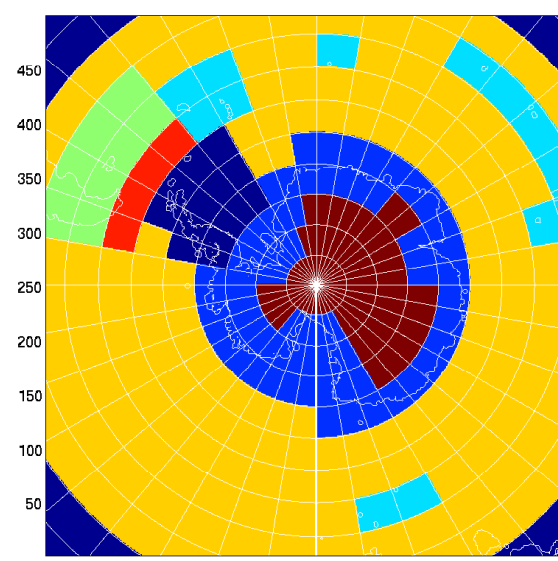


FIG. 4: Regions into which the domain has been subdivided: **1:** (dark blue) Antarctic Peninsula, **2:** (blue) Antarctic coastline, **3:** (Islands), **4:** (tip of the Andes), **5:** (Southern Ocean), **6:** (Drake Passage), **7:** (Plateau).

The time series of the domain-averaged zonal momentum fluxes (**Figure 5**) show that:

- significant intermittency during the first month, much less during the second, and an overall decrease in time as the winds decrease;
- two main contributions: the Antarctic Peninsula (small region 1), and the Southern Ocean (vast region 5);
- intermittency dominated by the strong intermittency over the Antarctic Peninsula.

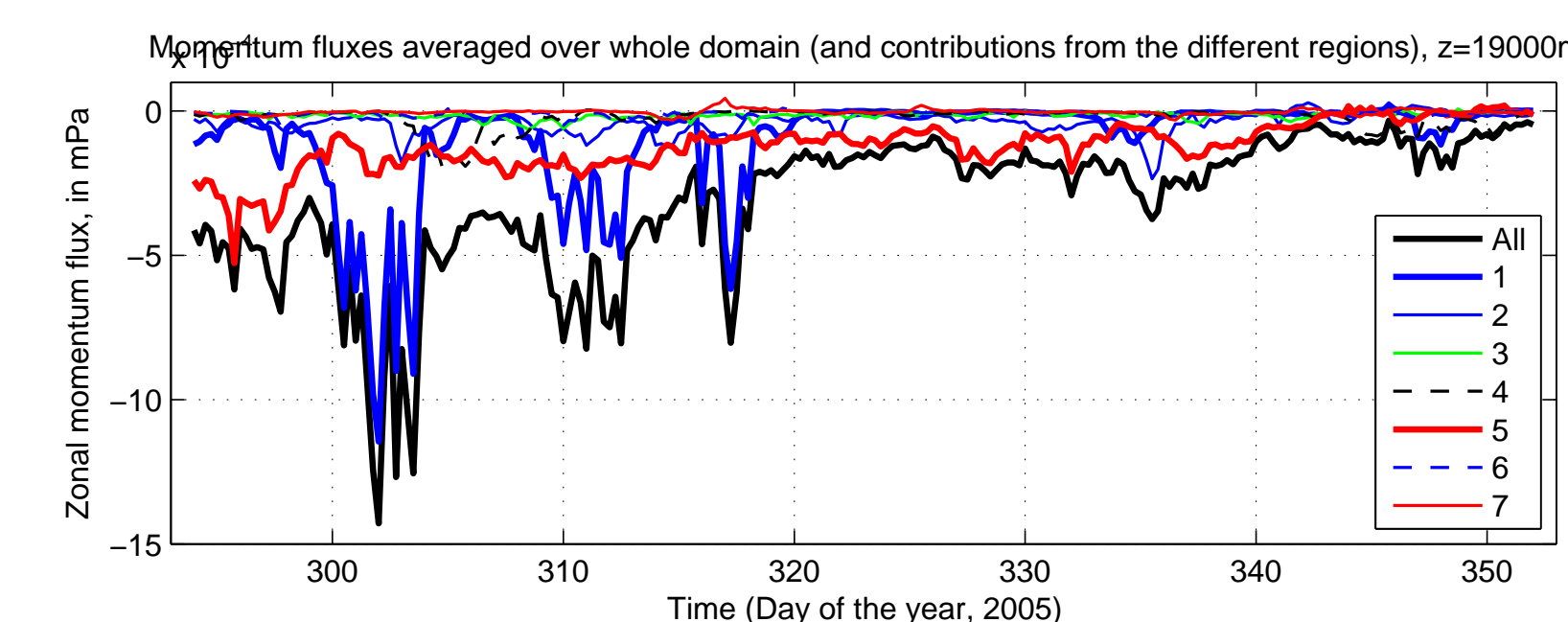


FIG. 5: Time series of the zonal momentum fluxes averaged over the whole domain (black curve), and of the contributions from the seven different regions shown in Fig. 4.

Zonal average

Figure 6 confirms the **comparable importance of orographic and non-orographic** regions found in the observations (*Hertzog et al 2008*).

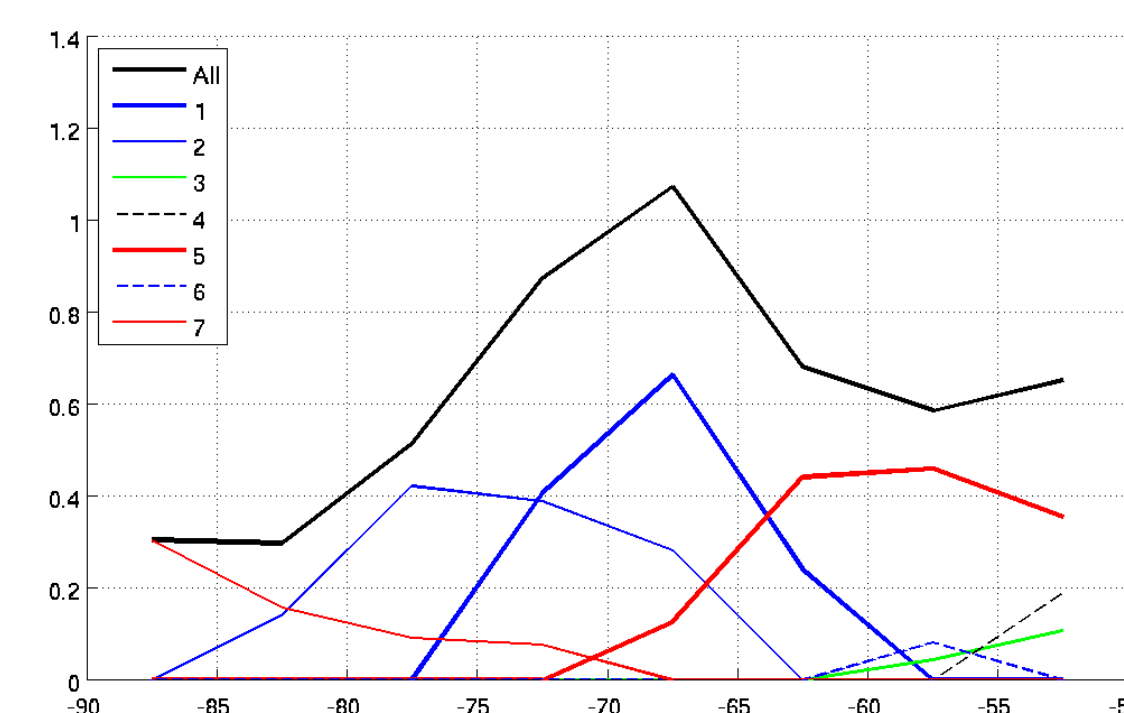


FIG. 6: Zonal average of the absolute momentum fluxes and of the contribution to these of the different regions.

Intermittency

Two definitions are used:

- as in *Hertzog et al 2008*, the ratio of the 50th to the 90th percentile of the timeseries, left panel in **Fig. 7**;
- the percentage of time containing the strongest values, amounting to 50% of the total, right panel in **Fig. 7**.

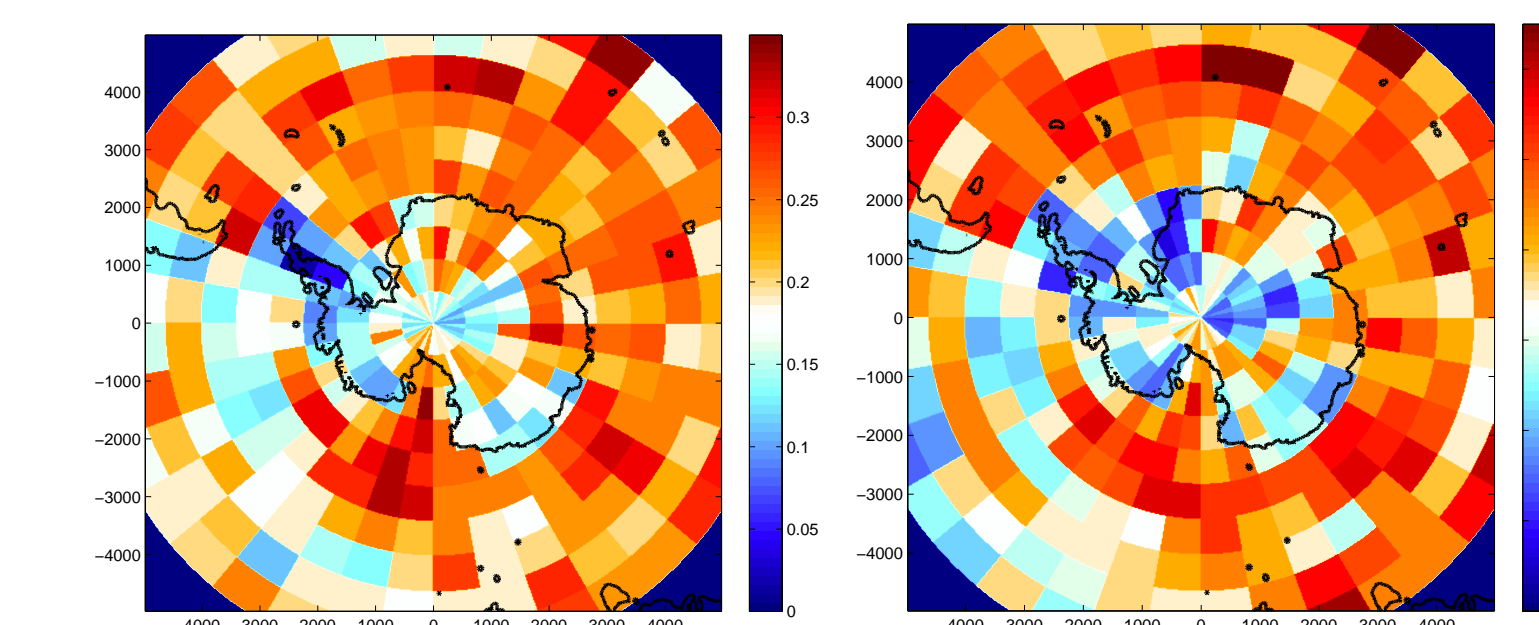


FIG. 7: Intermittency of absolute momentum fluxes estimated in two different ways.

Structure of the average gravity wave field over the oceans

The fine structure of the gravity wave field due to non-orographic waves is revealed in plots of the logarithm of the average absolute momentum fluxes (**Figure 8**). Such plots and movies suggest a strong correlation between the waves and the background winds (polar vortex) in which the waves propagate.

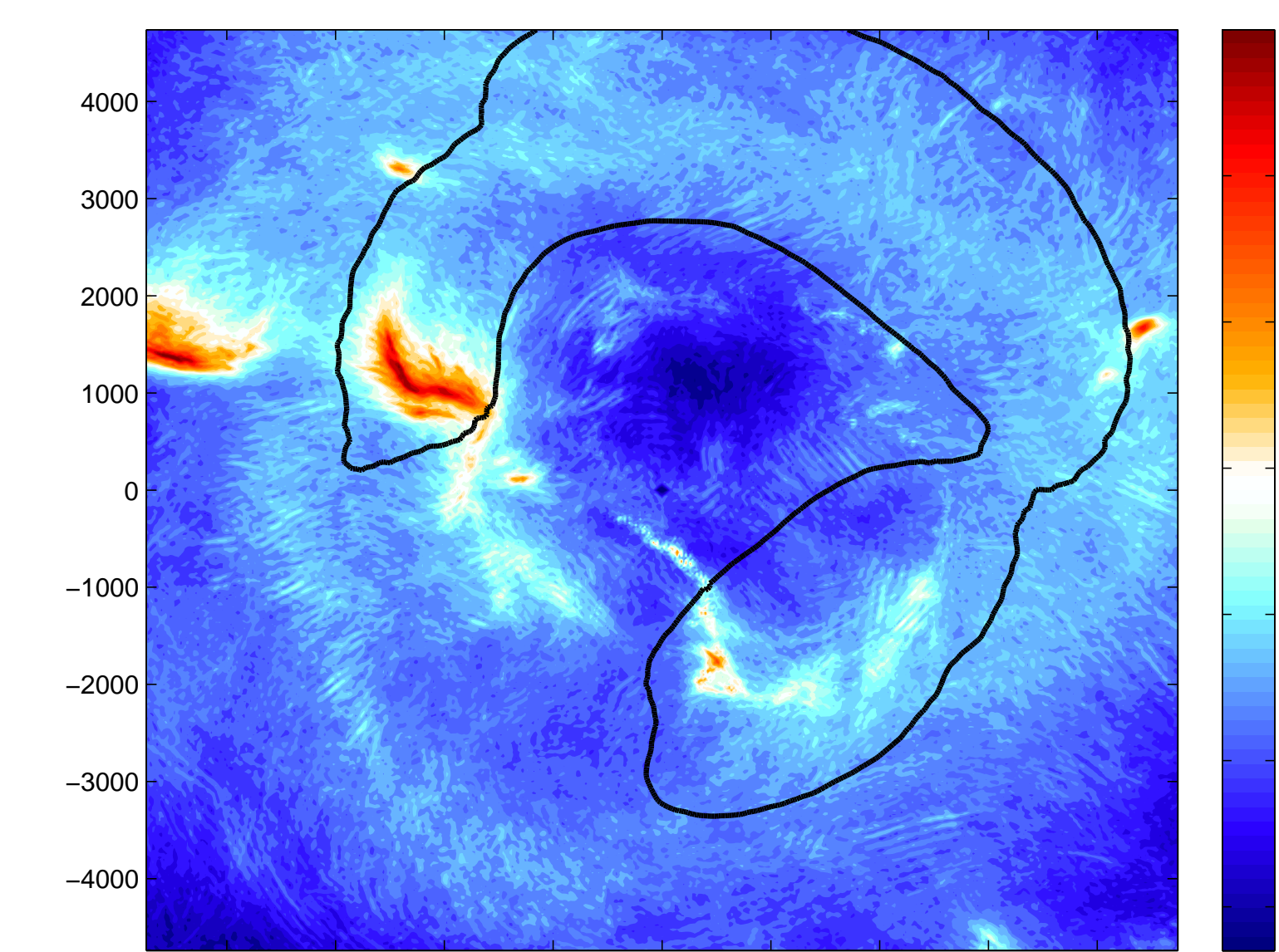


FIG. 8: Logarithm of the mean of the absolute momentum fluxes, averaged over the whole period. Also shown is the 40 m s^{-1} average windspeed isotach.

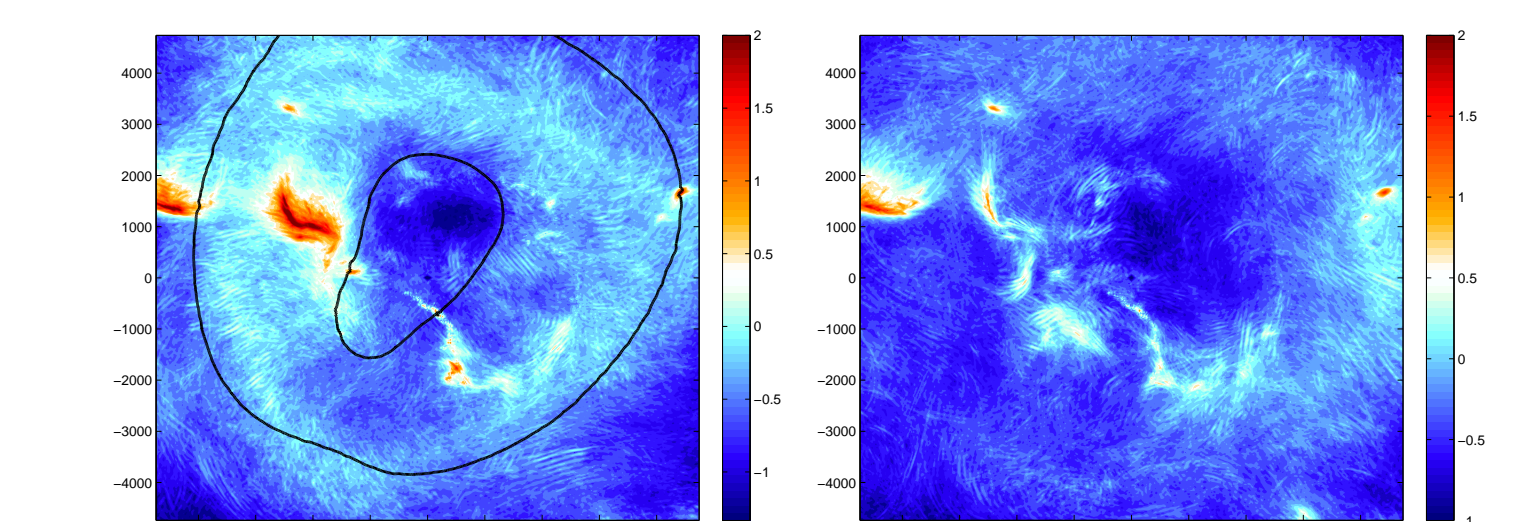


FIG. 8 (continued): During the first month (21/10-19/11/2005, left) and during the second month (19/11 - 18/12/2005, right).

Conclusions

Simulations show fair agreement with observations, though momentum fluxes due to non-orographic waves are underestimated, roughly by a factor 2.

Dominance of very strong, intermittent events above the Antarctic Peninsula, with local momentum fluxes reaching a few Pa (*Plougonven et al 2008*).

Nevertheless, the contribution from non-orographic waves over the Southern Ocean is comparable, corresponding to weaker, less intermittent waves.

Perspectives

The simulations will be further analyzed to investigate:

- vertical variations of the wave field, and implied forcing of the background flow;
- specific sources of the non-orographic waves, through both statistical and case studies;
- relevance of the mechanisms highlighted by idealized studies (e.g. wave capture, *Plougonven and Snyder 2005, 2007*);
- new ways to characterize the gravity wave field and relate it to the background flow (e.g. probability distribution function of momentum fluxes, in model and observations, **figure 9**).

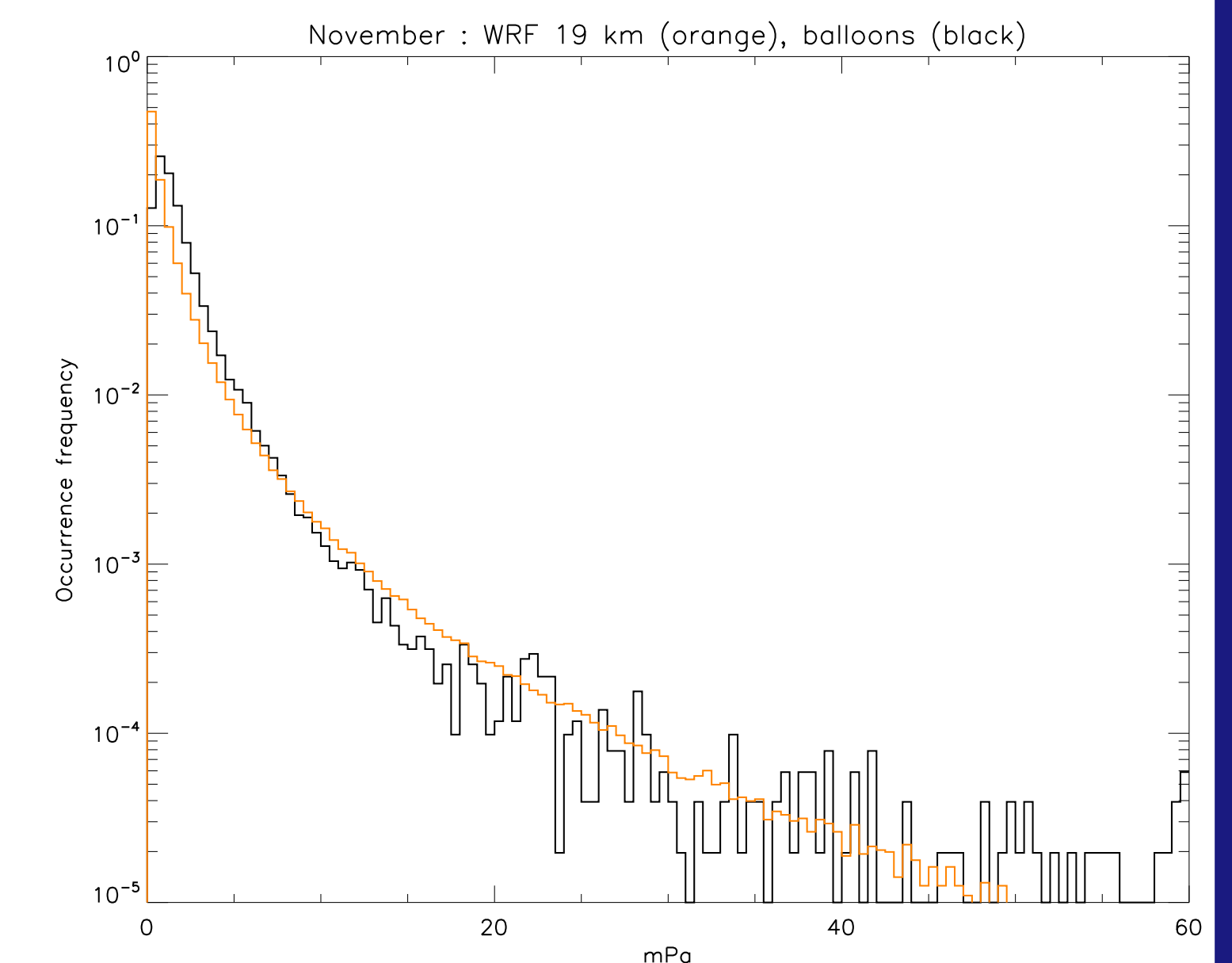


FIG. 9: Probability Distribution Function of absolute momentum fluxes over ocean regions, for November: balloons (black) and simulations (orange).

References

- [1] Alexander et al (2010), *Q.J. Roy. Met. Soc.*.
- [2] Hertzog, A., G. Boccaro, R. Vincent, F. Vial, and P. Coquerez (2008), *J. Atmos. Sci.*.
- [3] Hertzog, A., et al. (2007), *J. Ocean. Atmos. Tech.*, 24, 2048–2061.
- [4] Plougonven, R., A. Hertzog and H. Teitelbaum (2008), *J. Geophys. Res.*, 113, D16113.
- [5] Plougonven, R., A. Arsac, A. Hertzog, L. Guez and F. Vial (2010), *Q.J. Roy. Met. Soc.*.
- [6] Skamarock, W., J. Klemp, J. Dudhia, D. Gill, D. Barker, W. Wang, and J. G. NCAR Technical Note.