ABSTRACT

The thermomechanical evolution of the Earth remains a highly disputed subject. There are, however, some seemingly robust observations. A few are the focus in this study: (a) Tomographic studies, reconstruction of continental motion and rotation, and geodynamic modeling infer that the antipodally located Large Low Shear Velocity Provinces (LLSVPs) at the base of the mantle are stable, long-lived and impose the planform of flow in the mantle and of plate tectonics at the surface (Dziewonski 2010). (b) Studies of post-glacial rebound data, geoid highs and lows and geochemical mixing suggest that the average lower mantle viscosity is between one or two orders of magnitude greater than that of the upper mantle (Dziewonski 2010).

In this study, the geometry of the mantle is approximated by a hollow cylinder. The inner and outer mechanical boundary conditions are free slip and free surface, respectively. Fixed temperatures are prescribed for the thermal boundary conditions. The mantle flow is in the Stokes regime and the Boussinesq approximation applies. Newtonian rheological behavior is employed with a viscosity that depends exponentially on temperature. The large-scale heterogeneities at the base of the mantle are introduced as material of an anomalous viscosity and density.

Both structured and unstructured FEM codes are developed for solving the conservation equations of energy and momentum together with the incompressibility constraint. The advection-diffusion equation is solved in two separate steps. A FEM solver is developed to model diffusion process. A mechanical solver from MILAMIN is utilized to calculate velocities. An ODE-solver is developed to model advection.

One of the objectives in this study is to develop a robust non-diffusive advection solver. Eulerian-Lagrangian advection algorithms are conventionally used to minimize numerical diffusion. In this study, three different ode-solvers are devised and compared based on their numerical cost and the obtained accuracy.

Both structured and unstructured codes are benchmarked by reproducing critical Rayleigh number and Nusselt-Rayleigh scaling for rectangular geometry, provided by the linear stability analysis and the thermal boundary layer theory, respectively. Methods are tested based on convergence of measures, such as Nusselt number, with successive grid and time-step refinement. The convection pattern and Nusselt-Rayleigh relation are compared to results from previous numerical modeling by other authors. Analytical solutions are derived where possible and used for benchmarking.



Figure 1: Reconstructed large igneous provinces, hotspots argued to have deep origin and kimberlites for the past 320 Myr with respect to shearwave anomalies at the base of the mantle.

OBSERVATIONS

- Seismic tomography reveals two antipodally located Large Low Shear Velocity Provinces (LLSVPs) in the lower mantle, with steep margins and extending up to 1000 km from the core-mantle boundary.
- Bulk sound velocity and shear velocity within the LLSVPs are negatively correlated.
- Restored eruption sites of LIPs and hotspot volcanoes, as old as 200 Ma, project radially downwards to the margins of the LLSVPs

- **INTERPRETATIONS**
- The LLSVPs are stable, long-lived and impose the planform of flow in the mantle.
- The seismic velocity reduction is due to thermal as well as chemical and/or phase variations, inferring that LLSVPs are comprised of material with a moderate density contrast (2-5 %).



Figure 2: A cartoon of the equatorial cross-section outlining a proposed circulation in the mantle.

CONCLUSIONS

Using two separate grids for calculating velocities and temperatures allows having a coarser grid for the mechanical solver without significantly affecting the result.

It is advantageous to use higher-order ODE-solver schemes, such as the fourth-order Runge-Kutta, to model advection, since their computational cost is still low compared to the cost of the thermal and mechanical solvers.

Studying convergence of the Nusselt number with subsequent refining of the thermal and mechanical grids has shown significant dependence of the numerically obtained result on the grid resolution. The effect of numerical diffusion is influential on the result in a poorly refined grid. To model convection with Rayleigh number equal to 10^7 in a rectangular domain, a resolution of 400x4000 for the thermal grid is required to obtain a convergent result.

Thermomechanical Modeling of Stability of Large Low Shear Velocity Provinces

ка ≡ ——

 $\mu_{TOP}k$

 $\langle \mu_{BOTTOM} \rangle$

$$Ra_{cr} = 657.5$$



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FEM DIFFUSION SOLVER

The conductive heat transport is modeled by solving the discretized equation of the transient heat diffusion. To do this, a FEM diffusion solver is developed. Performance of the solver is evaluated for different element types and grid resolutions. An analytical solution is found for the problem of a cooling cylinder in two dimensions, and used to estimate the error of the numerically obtained solu-





Figure 4: Scaling of the relative error norm L2 with grid size for triangular (left) and quadratic (*right*) elements with first- and second-order shape functions used for space-discretization.



Figure 5: Analytical and numerical transient solutions to the cooling cylinder problem. First-order elements are used for the space-discretization. Left: Mass matrix is not lumped. *Right*: Mass matrix is lumped.

ODE SOLVERS

Transport of material due to convective dynamics is modeled by solving the discretized advection equation. To do this, three ordinary differential equation (ODE) solvers are developed and compared: Euler's method, fourth-order Runge-Kutta method and the embedded fifthorder Runge-Kutta method. An analytical solution is found for the problem of a passive marker advecting in a shear cell setup. The time it takes for a marker to complete a full rotation cycle and return to its original Figures 10-13: Nusselt number composition is calculated and used to estimate the error of the numerically obtained positions.





Figure 7: Scaling of error with stepsize for Euler and fourth-order Runge-Kutta methods.



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