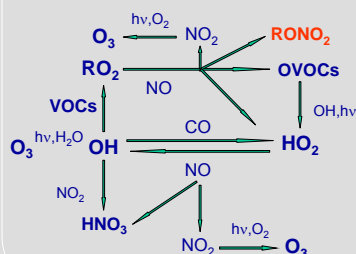


Introduction

Reactions of alkyl peroxy radicals with NO are recognized to potentially reduce tropospheric ozone formation through their minor channel (b) producing organic nitrates, which are both ROx and NOx reservoir or sink species:

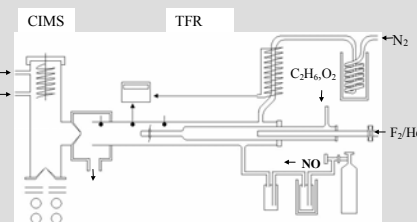


The available nitrate yields measured in these reactions range from $\approx 4\%$ for propyl nitrate to $\approx 30\%$ for long-chain alkyl nitrates ($C \geq 7$) at $T = 298 \text{ K}$ and $P = 760 \text{ Torr}$ [eg. 1]. The yields also increase with increasing pressure and with decreasing temperature. For the short-chain alkyl nitrates ($C \leq 4$), very few data are available due to the difficulties to precisely measure low nitrate yields. Yield data are reported here for ethyl nitrate, as a function of pressure and temperature, and for iso-propyl nitrate and methyl nitrate as a function of pressure at ambient temperature.



Turbulent Flow Reactor (TFR) with Chemical Ionization Mass Spectrometer (CIMS)

Reactor conditions:
P = 50-600 Torr; T = 220-298 K; Re = 4700-12000.



• CIMS detection of the species in positive (PTR, DTR) and/or negative (NICI) mode:

- PTR (H_3O^+): CH_2O , CH_3OH , CH_3ONO , CH_3ONO_2 , CH_3CHO , $\text{C}_2\text{H}_5\text{OH}$, $\text{C}_2\text{H}_5\text{ONO}$, $\text{C}_2\text{H}_5\text{ONO}_2$
 - DTR (D_3O^+): CH_2O , CH_3OH , CH_3ONO , CH_3ONO_2
 - NICI (SF_6^-): NO_2
 - NICI (F^- from NF_3): $\text{C}_2\text{H}_5\text{OH}$, CH_3CHO , $\text{CH}_3\text{C}(\text{O})\text{CH}_3$, $\text{C}_2\text{H}_5\text{ONO}_2$, $i\text{-C}_3\text{H}_7\text{ONO}_2$ (a new sensitive method)
- Calibration of the species using commercial compounds or producing them *in situ*.

Formation of ethyl nitrate (reaction 1b)

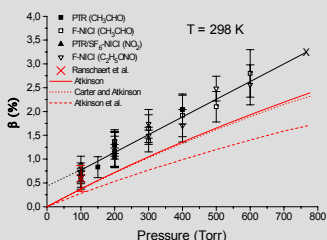
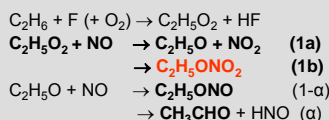


Figure 1. Pressure dependence of β for ethyl nitrate formation. Red lines are calculated values from empirical functions [4, 5, 6].

Main reactions in the TFR:



Determination of the branching ratio

$$\beta = k_{1b}/k_{1a}$$

$$\beta = \Delta[\text{C}_2\text{H}_5\text{ONO}_2] / \Delta[\text{NO}_2]$$

$$\beta = \alpha \cdot \Delta[\text{C}_2\text{H}_5\text{ONO}_2] / \Delta[\text{CH}_3\text{CHO}]$$

$$\beta = (1-\alpha) \cdot \Delta[\text{C}_2\text{H}_5\text{ONO}_2] / \Delta[\text{C}_2\text{H}_5\text{ONO}]$$

The overall P and T dependence of β is expressed as [2]:

$$\beta(P, T)(\%) = k_{1b}/k_{1a} = (3.88 \times 10^{-3} \cdot P(\text{Torr}) + 0.365) \cdot (1 + 1500/(T(\text{K}) - 1/298))$$

($100 \leq P \leq 760 \text{ Torr}$ and $215 \leq T \leq 300 \text{ K}$)

→ Only small changes of β and $k_{1b} \approx 0.01 \cdot k_1(T) \cdot \beta(P, T)$ throughout the whole troposphere:

$$\beta \approx 3.3\% \quad k_{1b} \approx 2.8 \times 10^{-13} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1} \text{ at the Earth surface}$$

$$\beta \approx 2.8\% \quad k_{1b} \approx 3.9 \times 10^{-13} \text{ " " " in the upper troposphere}$$

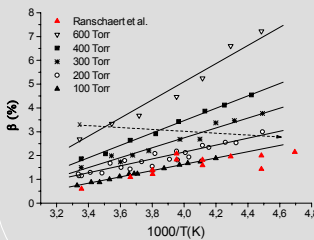


Figure 2. Temperature dependence of β for ethyl nitrate formation at different pressures.

Atmospheric application

The present work provides values of k_{1b} , k_{2b} and k_{3b} for use in chemistry-transport models to calculate the impact of chain termination reactions 1b, 2b and 3b on the tropospheric composition, in particular for species involved in the VOC/NOx/ozone chain mechanism. This impact is likely to be not negligible considering both the efficiency of radical chain termination reactions in this mechanism and the rather long lifetimes of methyl-, ethyl- and i-propyl nitrate as NO_x reservoirs (typical lifetimes of 1 month, 2 weeks and 1 week, respectively, controlled mainly by photolysis [9]). Besides, the apparent inconsistency between the k_{3b}/k_{3a} ratio obtained in this work and atmospheric measurements of methyl nitrate requires further investigation.

Formation of iso-propyl nitrate (reaction 2b)

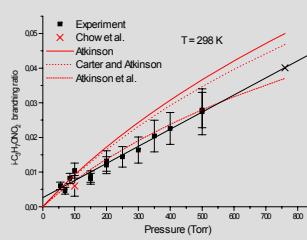
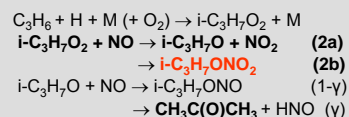


Figure 3. Pressure dependence of β for iso-propyl nitrate formation. Red lines are calculated values from empirical functions [4, 5, 6].

Main reactions in the TFR:



Determination of the branching ratio

$$\beta = k_{2b}/k_{2a}$$

$$\beta = \gamma \cdot \Delta[i\text{-C}_3\text{H}_7\text{ONO}_2] / \Delta[\text{CH}_3\text{COCH}_3]$$

A linear pressure dependence of β is observed providing the following expression [7]:

$$\beta(P)(\%) = k_{2b}/k_{2a} = (4.93 \pm 0.24) \times 10^{-3} \cdot P(\text{Torr}) + (0.265 \pm 0.067)$$

($55 \leq P \leq 500 \text{ Torr}$, $T = 298 \text{ K}$)

→ Only small changes of $k_{2b} \approx 0.01 \cdot k_2(T) \cdot \beta(P, T)$ are expected throughout the whole troposphere by combining our value of $\beta \approx 4\%$ at 760 Torr and 298 K and that of Chow et al. [8], $\beta \approx 2\%$ at 100 Torr and 210 K:

$$\beta \approx 4\% \quad k_{2b} \approx 3.6 \times 10^{-13} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1} \text{ at the Earth surface}$$

$$\beta \approx 2\% \quad k_{2b} \approx 3.0 \times 10^{-13} \text{ " " " in the upper troposphere}$$

Formation of methyl nitrate (reaction 3b) (preliminary results)

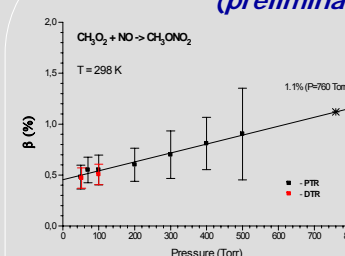
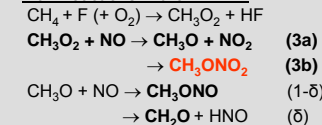


Figure 4. Pressure dependence of β for methyl nitrate formation.

Main reactions in the TFR:



Determination of the branching ratio

$$\beta = k_{3b}/k_{3a}$$

$$\beta = \delta \cdot \Delta[\text{CH}_3\text{ONO}_2] / \Delta[\text{CH}_2\text{O}]$$

$$\beta = (1-\delta) \cdot \Delta[\text{CH}_3\text{ONO}_2] / \Delta[\text{CH}_3\text{ONO}]$$

PTR detection of methyl nitrate ($\text{CH}_3\text{ONO}_2\text{-H}^+$ ion, $m/z 78$) was complicated by formation of $\text{CH}_2\text{O} \cdot \text{NO} \cdot \text{H}_2\text{O}$ complex with the same mass. Its contribution was determined separately and subtracted from the reaction signal. Deuteron Transfer Reaction (DTR) was used along with PTR detection in order to avoid this interference by detection of methyl nitrate at $m/z 79$ ($\text{CH}_3\text{ONO}_2\text{-D}^+$ ion), while the complex is moved to mass $m/z 80$ ($\text{CH}_2\text{O} \cdot \text{NO} \cdot \text{D}_2\text{O}$ ion).

A linear pressure dependence of β is observed providing the following expression:

$$\beta(P)(\%) = k_{3b}/k_{3a} = (8.8 \pm 0.7) \times 10^{-4} \cdot P(\text{Torr}) + (0.45 \pm 0.03)$$

($50 \leq P \leq 500 \text{ Torr}$, $T = 298 \text{ K}$)

→ $\beta \approx 1\%$ $k_{3b} \approx 7 \times 10^{-14} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ at the Earth surface

References:

- [1] Atkinson R. et al., *J. Phys. Chem.* 1982, 86, 4563; [2] Butkovskaya N. et al., *J. Phys. Chem. A* 2010, 114, 956; [3] Ranschaert D.L. et al., *J. Phys. Chem. A* 2000, 104, 5758; [4] Atkinson R. et al., *J. Phys. Chem.* 1983, 87, 2012; [5] Carter W.P.L. and Atkinson R., *J. Atmos. Chem.* 1989, 8, 165; [6] Atkinson R., *Phys. Chem. Ref. Data* 1994, 1, Monograph 2; [7] Butkovskaya et al. *Z. Phys. Chem.* 2010, 224, 1025; [8] Chow J.M. et al., *J. Phys. Chem. A* 2003, 107, 3040; [9] Talukdar R. et al., *J. Chem. Soc. Faraday Trans.* 1997, 93, 2797

Acknowledgement

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