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EARLY EARTH: Solid Earth Processes

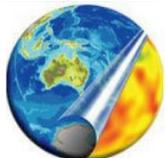
# Subaqueous Archean continental flood basalts were emplaced on hot continental crust

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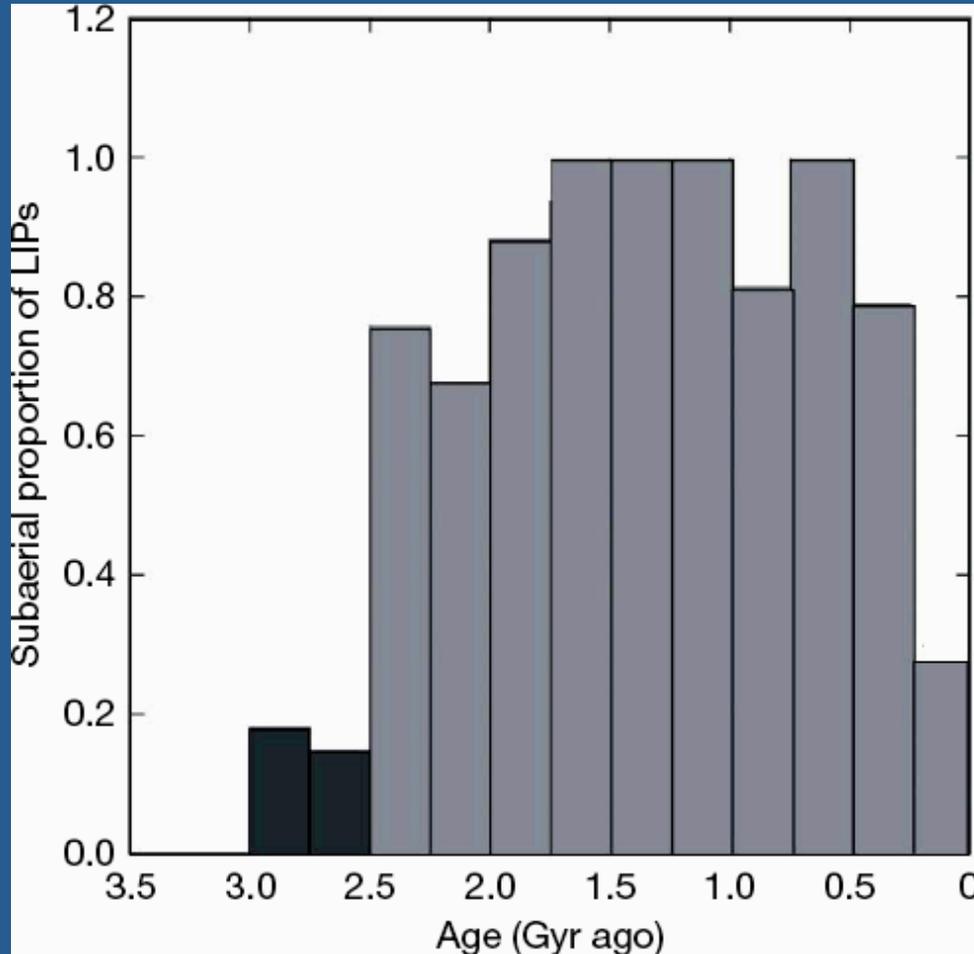


Lyon 1

Archaean pillow basalts in  
flood basalts up to 10 km thick



# Predominance of subaqueous flood basalts in the Archean



*Kump & Barley (2007)*

Subaqueous flood volcanism on continental platforms is:

- common in the Precambrian
- rare to absent in the Phanerozoic

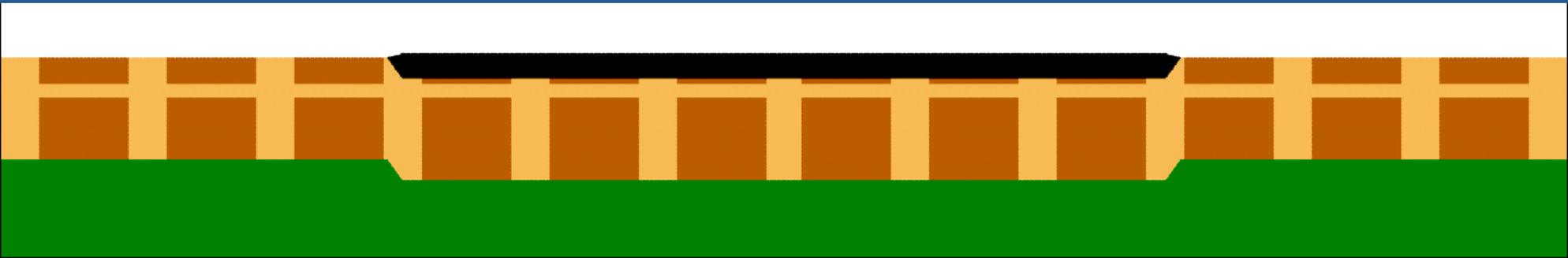
*Arndt (1999)*

# Examples of subaqueous Archean CFBs

Name	Craton	Events*	Thickness (km)	Age (Ma)	Duration (Myr)	References
Lower Maddina Fm., Fortescue Gp.	<b>Pilbara</b>	1	~ 0.6	2718 ± 3 to 2713 ± 3	≤ 11	Blake et al. (2004)
Kylena Fm., Fortescue Gp.	<b>Pilbara</b>	1	≤ 1.4	2749-2735	14	Thorne and Trendall (2001)
Honman Fm., Lake Johnston gr. b.	<b>Yilgarn</b>	1	≤ 1.2	2921 ± 4 to 2903 ± 5	≤ 28	Wang et al. (1996)
Ngezi Gp., Belingwe gr. b.	<b>Zimbabwe</b>	2	6.5	2692 ± 9	18?	Chauvel et al. (1993)
Upper Kam Gp., Yellowknife belt	<b>Slave</b>	3	~ 6	2722 ± 2 to 2701 ± 1	≤ 24	Isachsen and Bowring (1994)
Kambalda gr. b.	<b>Yilgarn</b>	4	~ 4	2726 ± 30 to 2690 ± 5	≤ 71	Tomlinson and Condie (2001)
Lower Warrawoona Gp., Marble Bar gr. b.	<b>Pilbara</b>	sequence	~ 3	3490-3469	> 21	Van Kranendonk et al. (2004)
Balmer assemblage	<b>Superior</b>	sequence	< 10	2992-2964	> 28	Tomlinson and Condie (2001)
Lumby Lake gr. B.	<b>Superior</b>	sequence	≤ 7	< 2963-2898	< 70	Tomlinson and Condie (2001)
Kolar schist belt	<b>Dharwar</b>	sequence	~ 4	~ 2700	?	Krogstad et al. (1989)

Subaqueous flood basalts erupted through continental crust  
*(Arndt, 1999; Tomlinson and Condie, 2001)*

# Topography associated with CFBs



• Elevation

$$e = t_{CFB} \times \left( 1 - \frac{\rho_{CFB}}{\rho_M} \right)$$

•  $e \approx 950$  m for  $t_{CFB} = 6$  km

•  $t_{CFB}$ : thickness

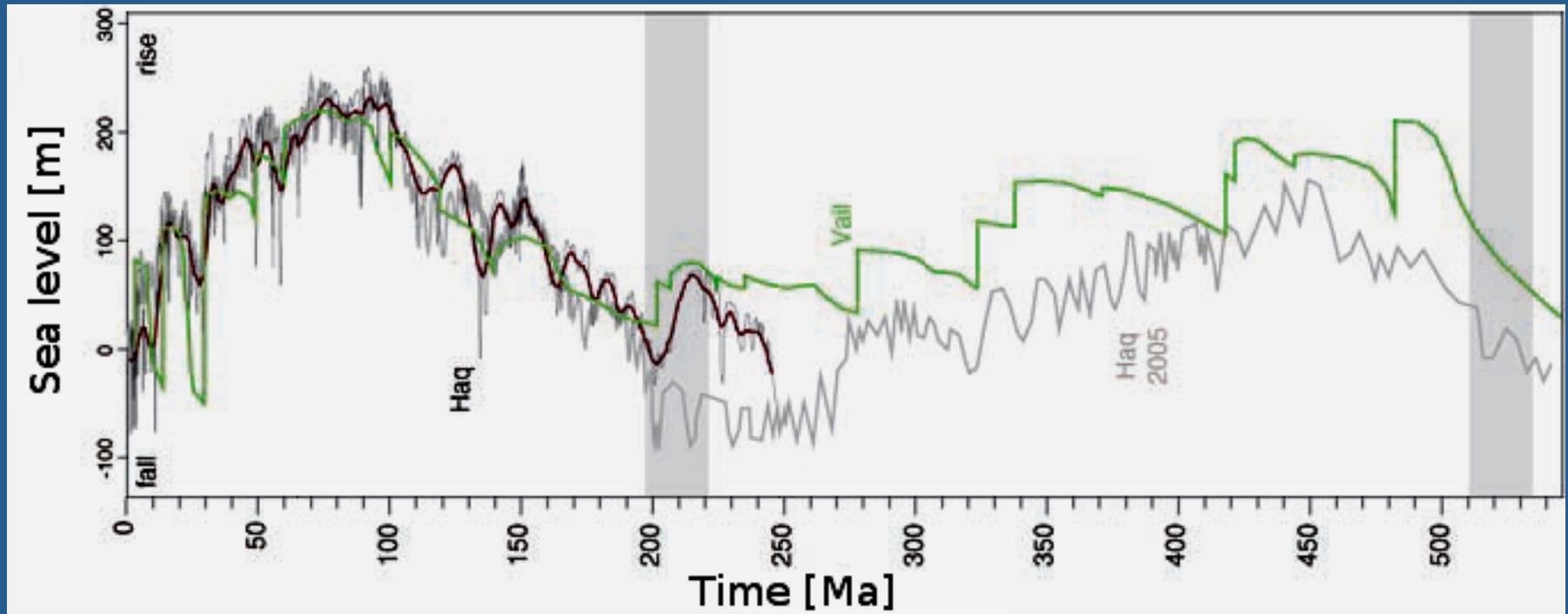
•  $\rho_{CFB} = 2840 \text{ kg m}^{-3}$

•  $\rho_M = 3370 \text{ kg m}^{-3}$

•  $e \approx 1500$  m for  $t_{CFB} = 10$  km

→ Need to explain a change in sea level of up to 1.5 km in  $< 70$  Myr

# Not explained by change in eustatic sea level

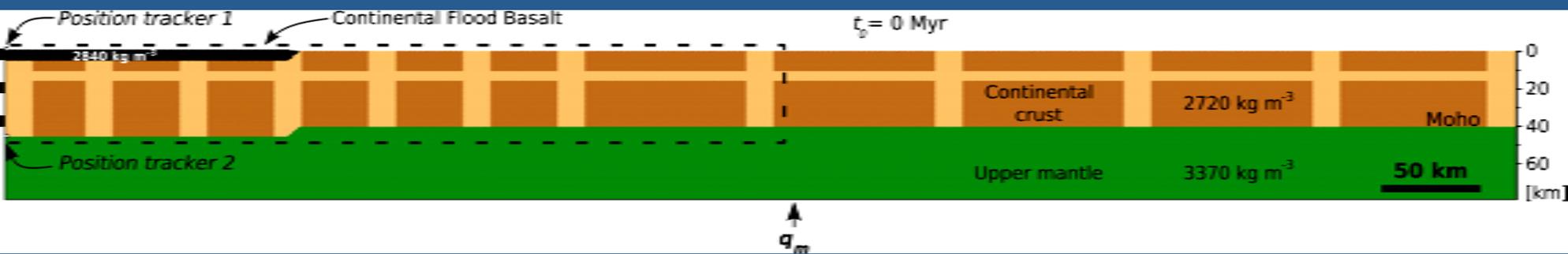


- Phanerozoic rate  $\sim < 2 \text{ m Myr}^{-1}$
- Secular rate  $\sim 0.5 \text{ m Myr}^{-1}$  (Flament et al., 2008)
- Subaqueous Archean CFBs  $15 \text{ m Myr}^{-1}$

Miller et al. (2005)

# Hypothesis and model

Hypothesis: flow of continental crust can remove thickness anomalies associated with a CFB in a few tens of million years



- Ellipsis, particle-in-cell, finite element code

*(Moresi et al., 2003)*

# Model setup – geotherm & rheology

- 1D, steady-state

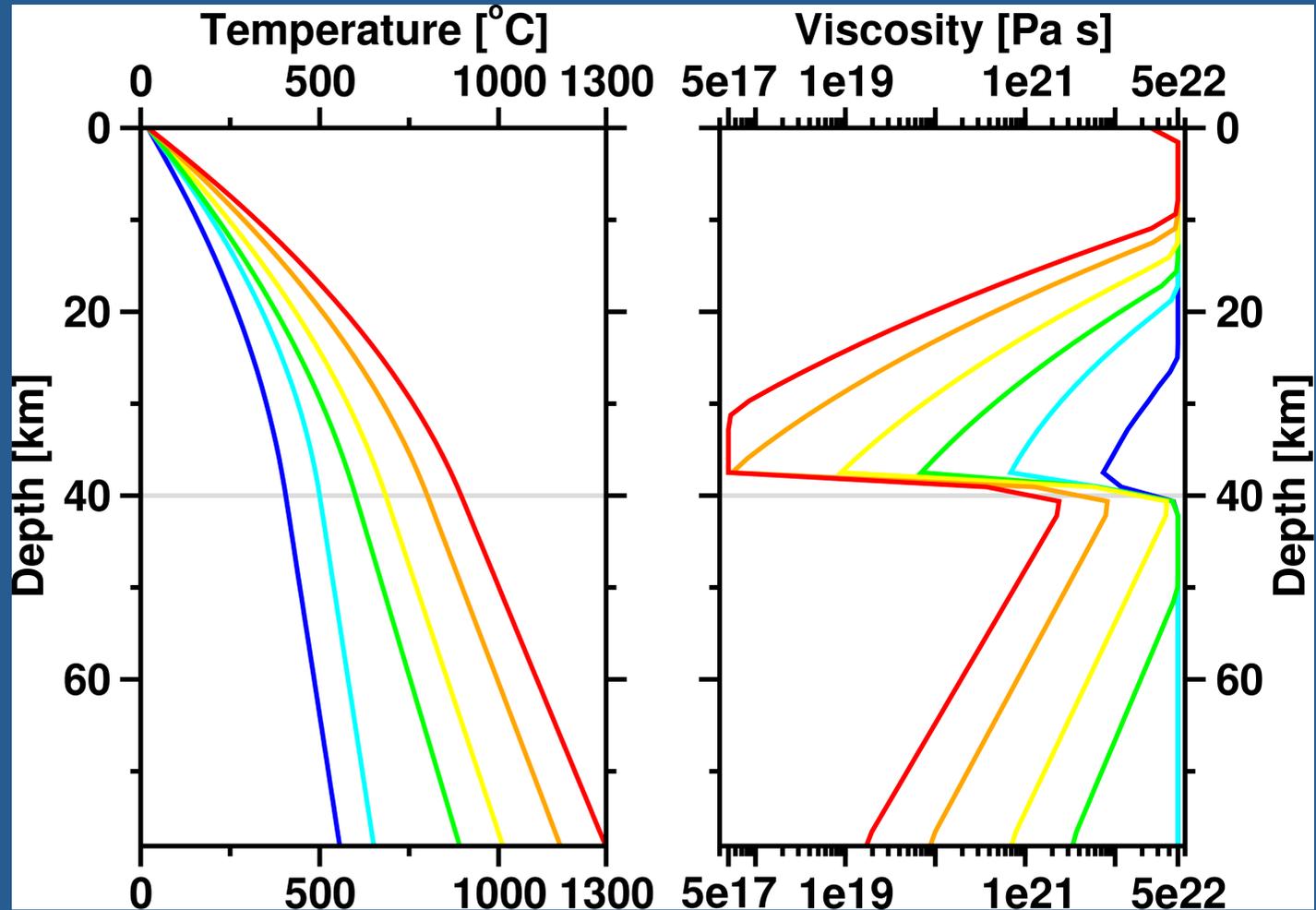
$$k \frac{d^2 T}{dz^2} + \rho_{cc} H(t) = 0$$

- $H(0)$  for present-day cratons

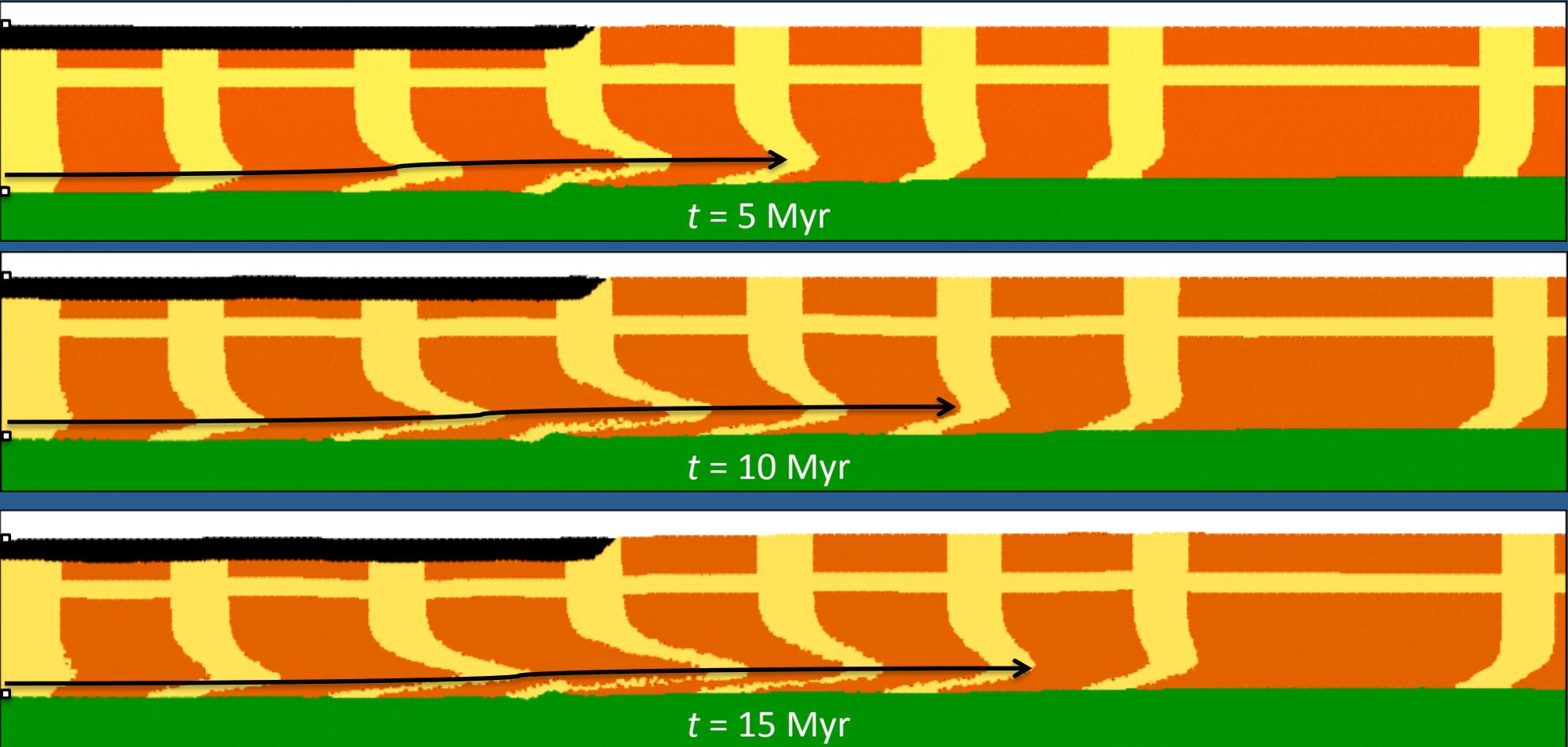
*Taylor & Mclennan (1995)*

- Visco-plastic
- T-dependent viscosity:

$$\eta = \eta_0 \exp(-\gamma T)$$

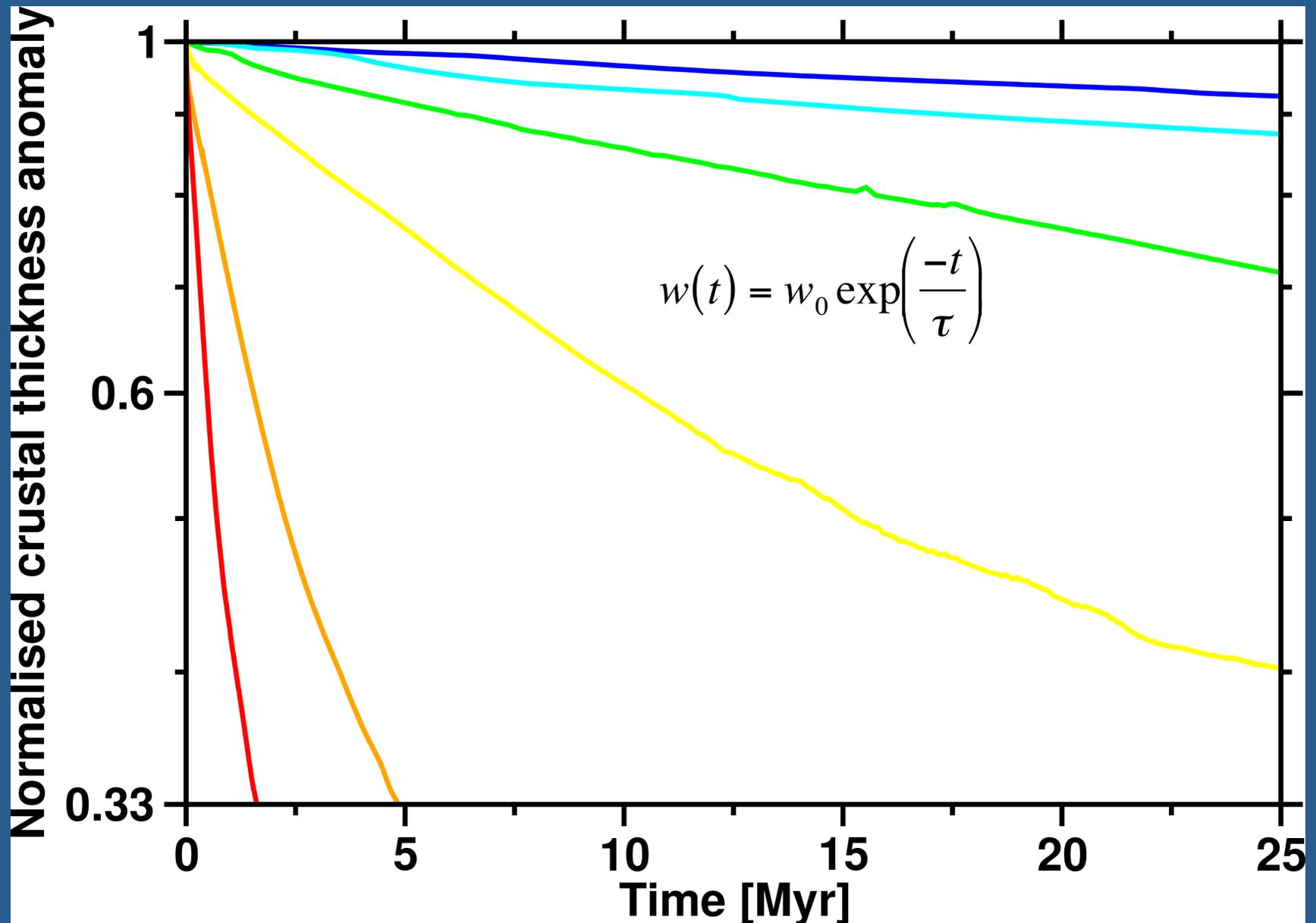


# Relaxation of the thickness anomaly via gravity-driven lower crustal flow

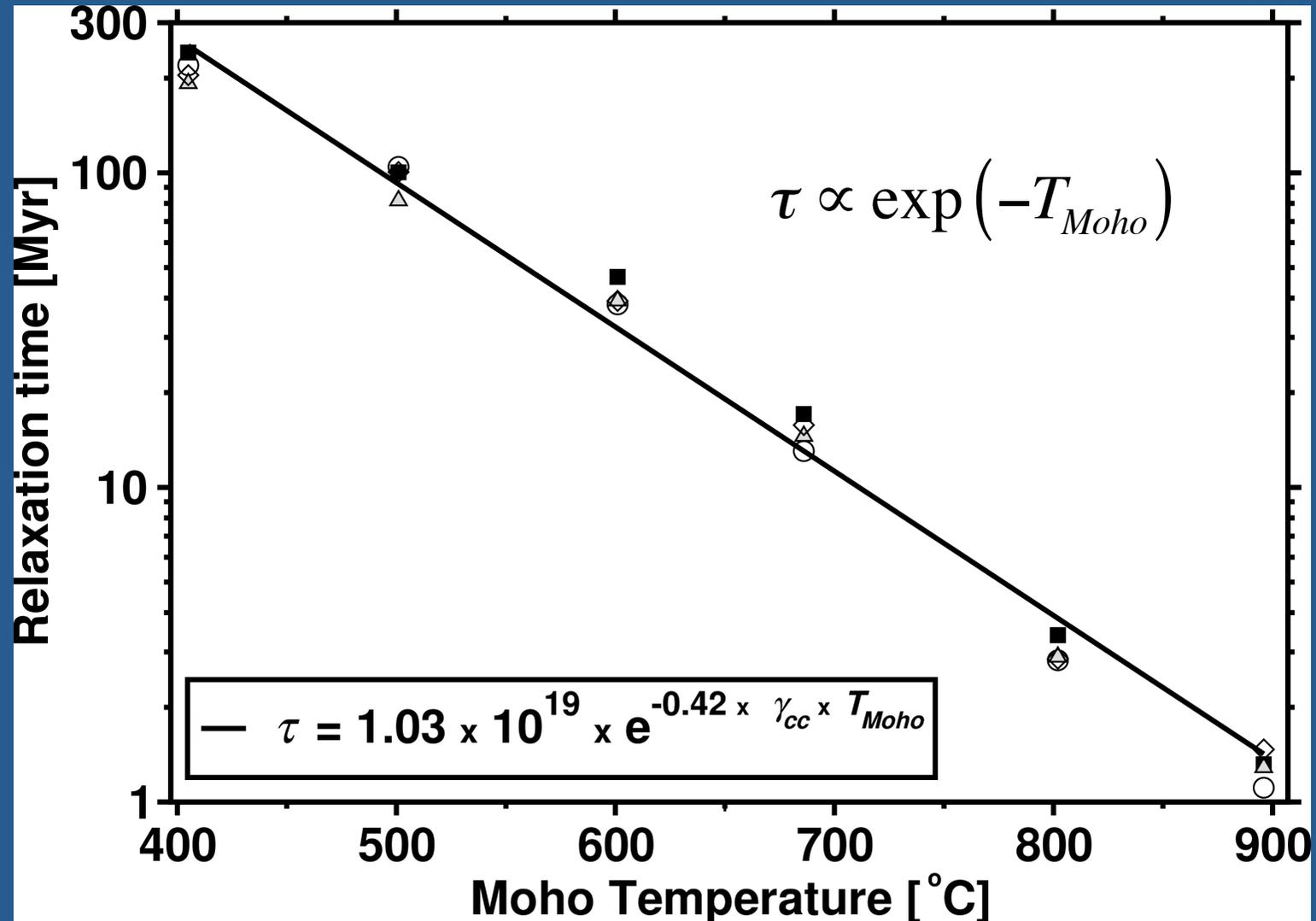


- CFB 6 km thick
- $T_{Moho} \approx 700^{\circ}\text{C}$

# Characteristic relaxation time $\tau$

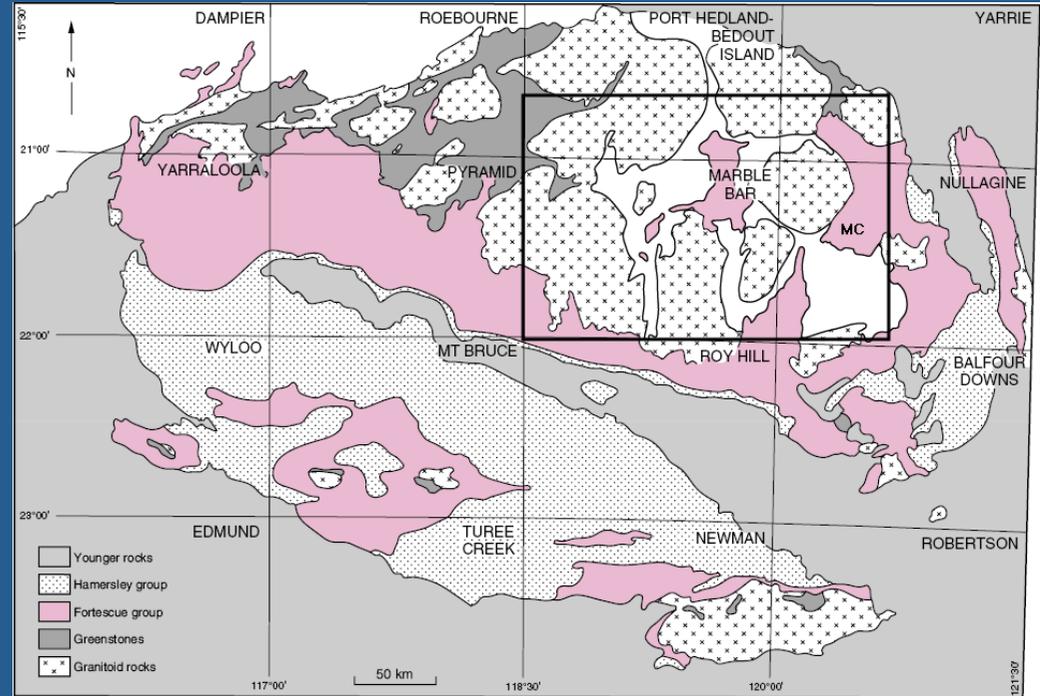
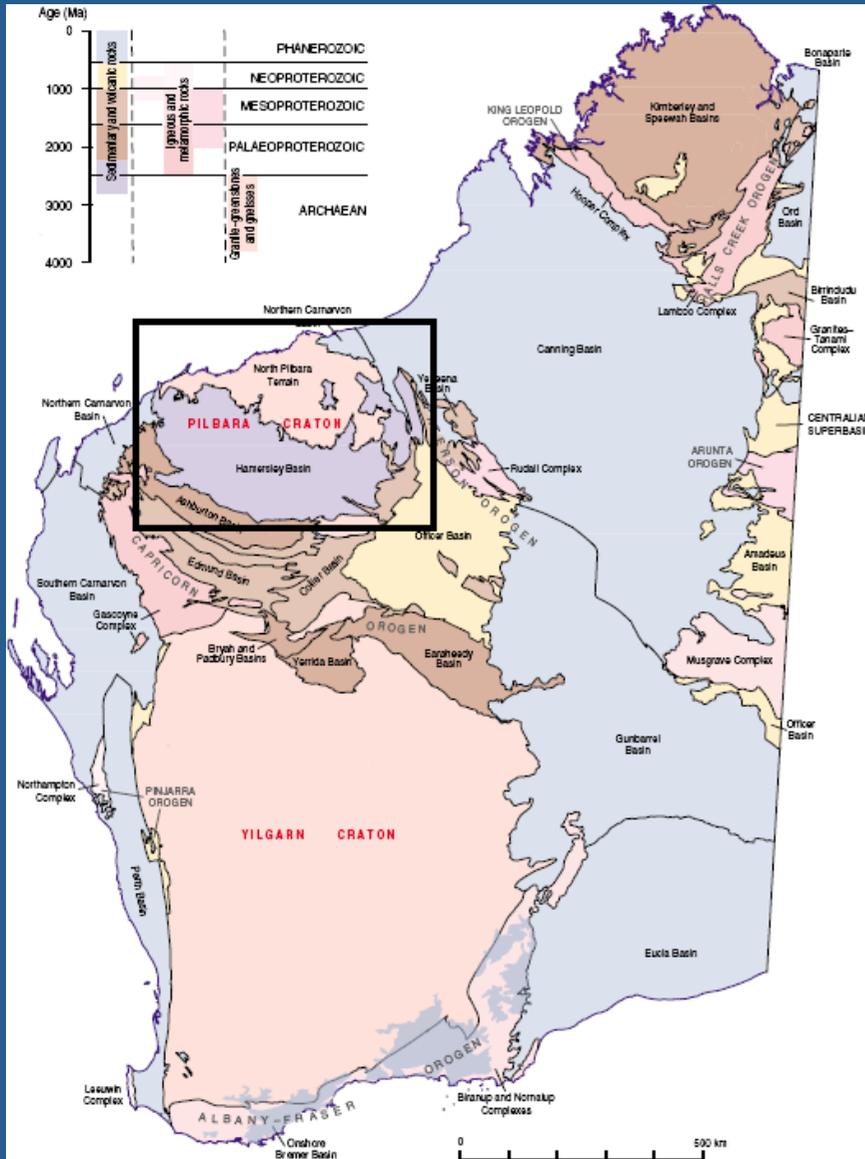


# Scaling laws for $\tau$ as a function of $T_{Moho}$



- Efficiency of lower crustal flow increases with  $T_{Moho}$
- Estimate  $T_{Moho}$  from the subsidence history of a CFB

# Case study: Fortescue Group, Pilbara Craton



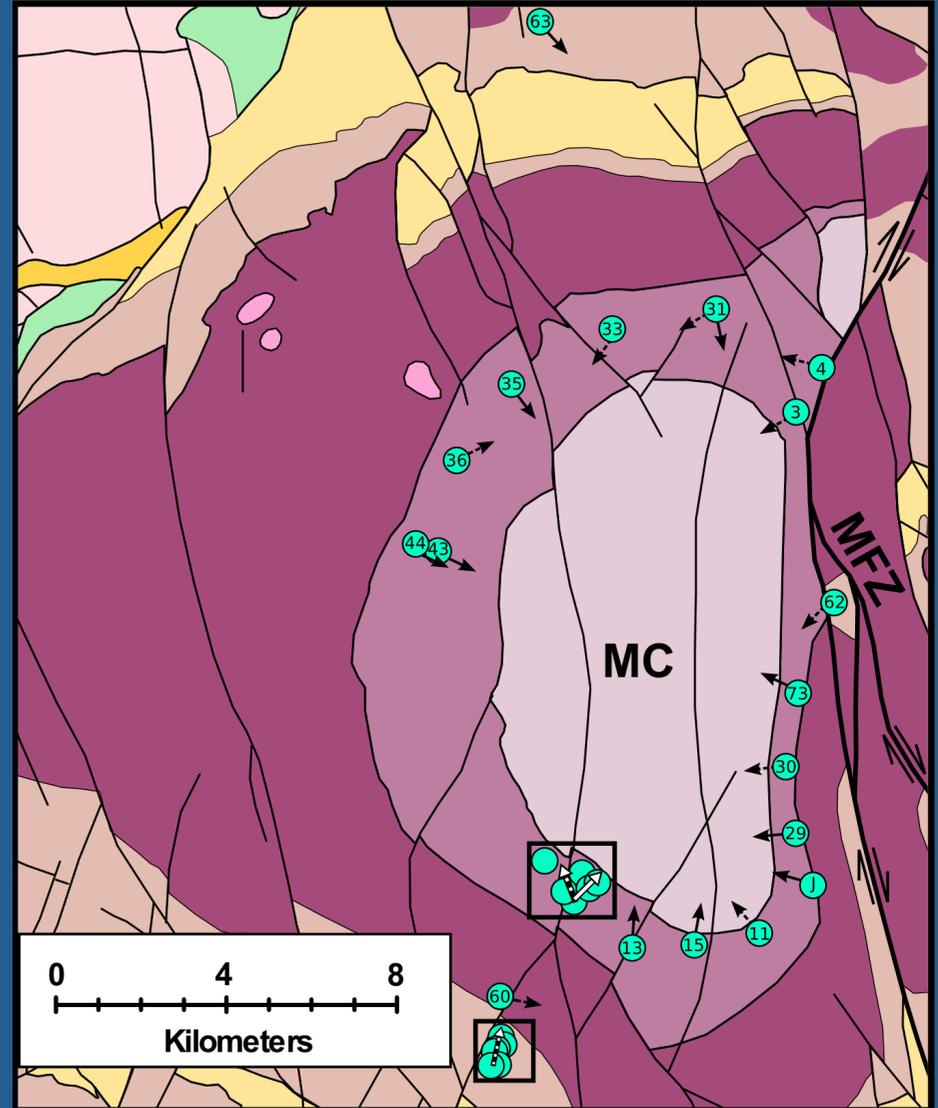
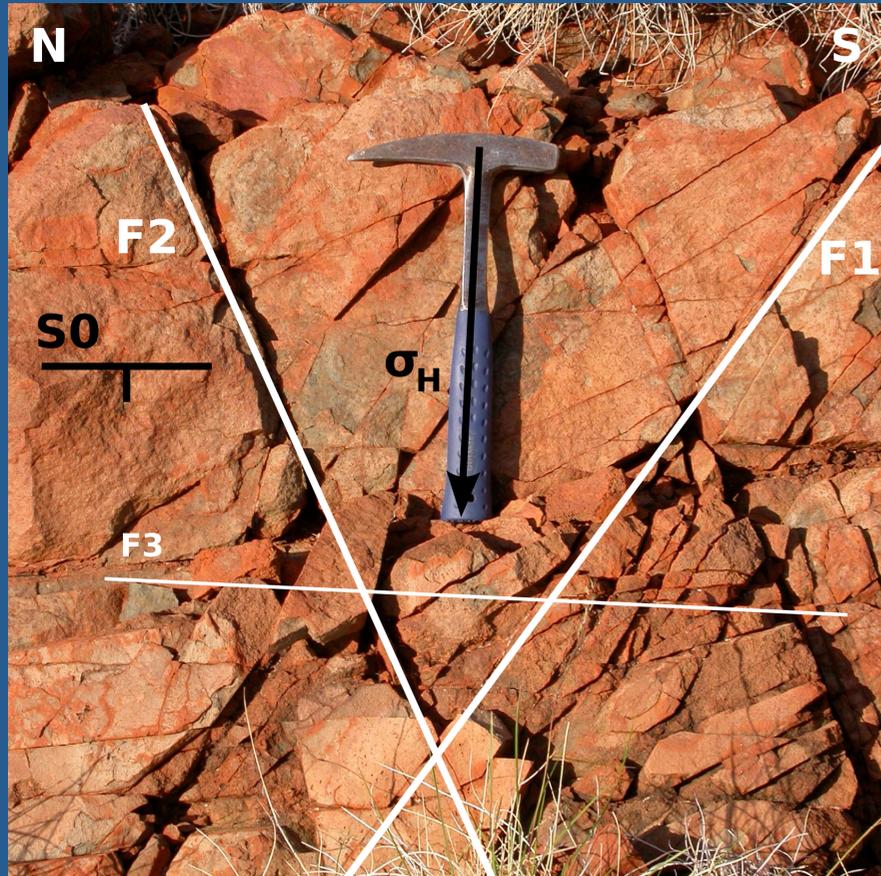
*Thorne & Trendall (2001)*

- 6.5 km thick CFB
- 2775-2630 Ma

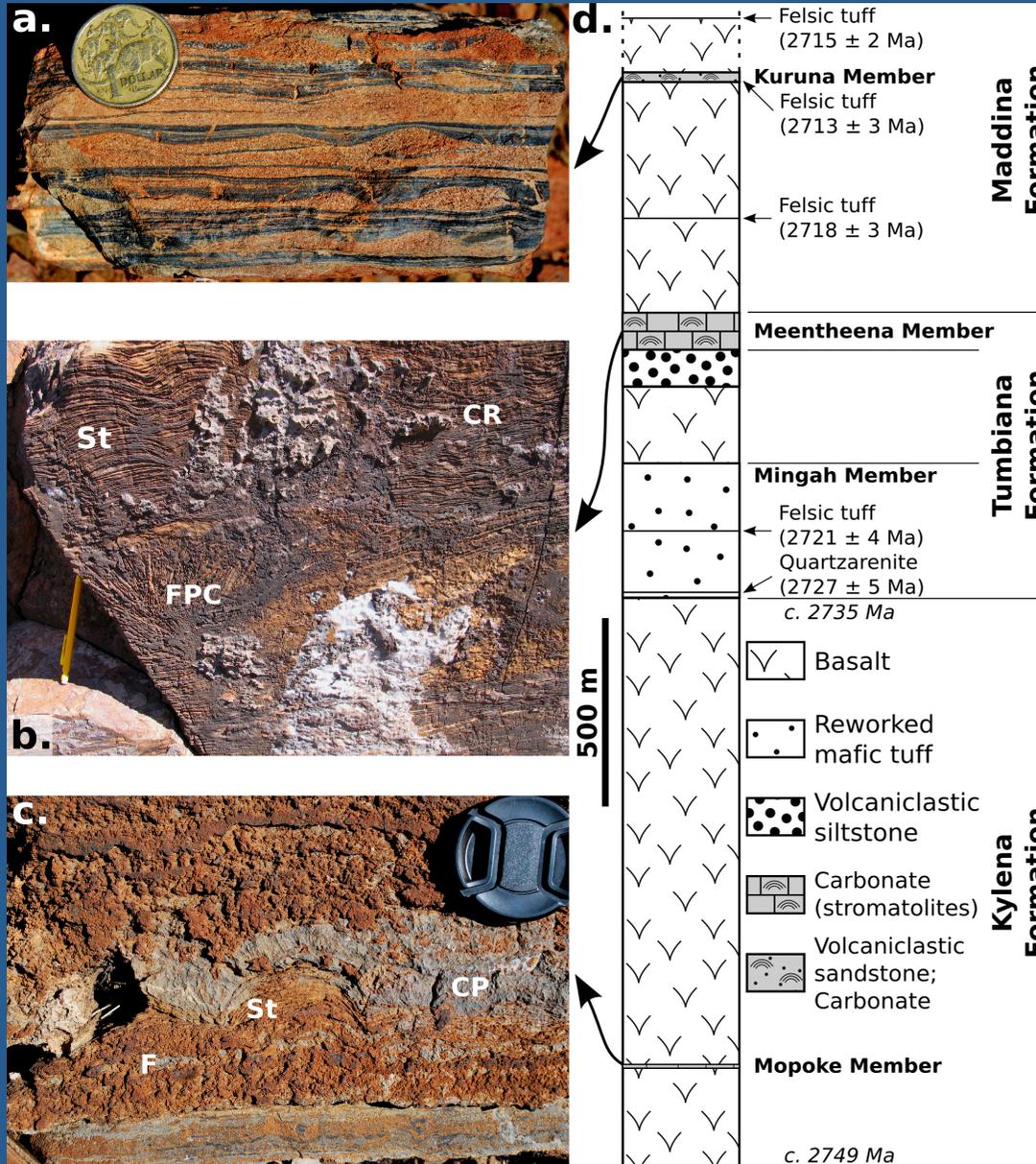
GSWA



# Directions of palaeostress $\sigma_H$ and gravitational subsidence



# An exceptional stratigraphic column

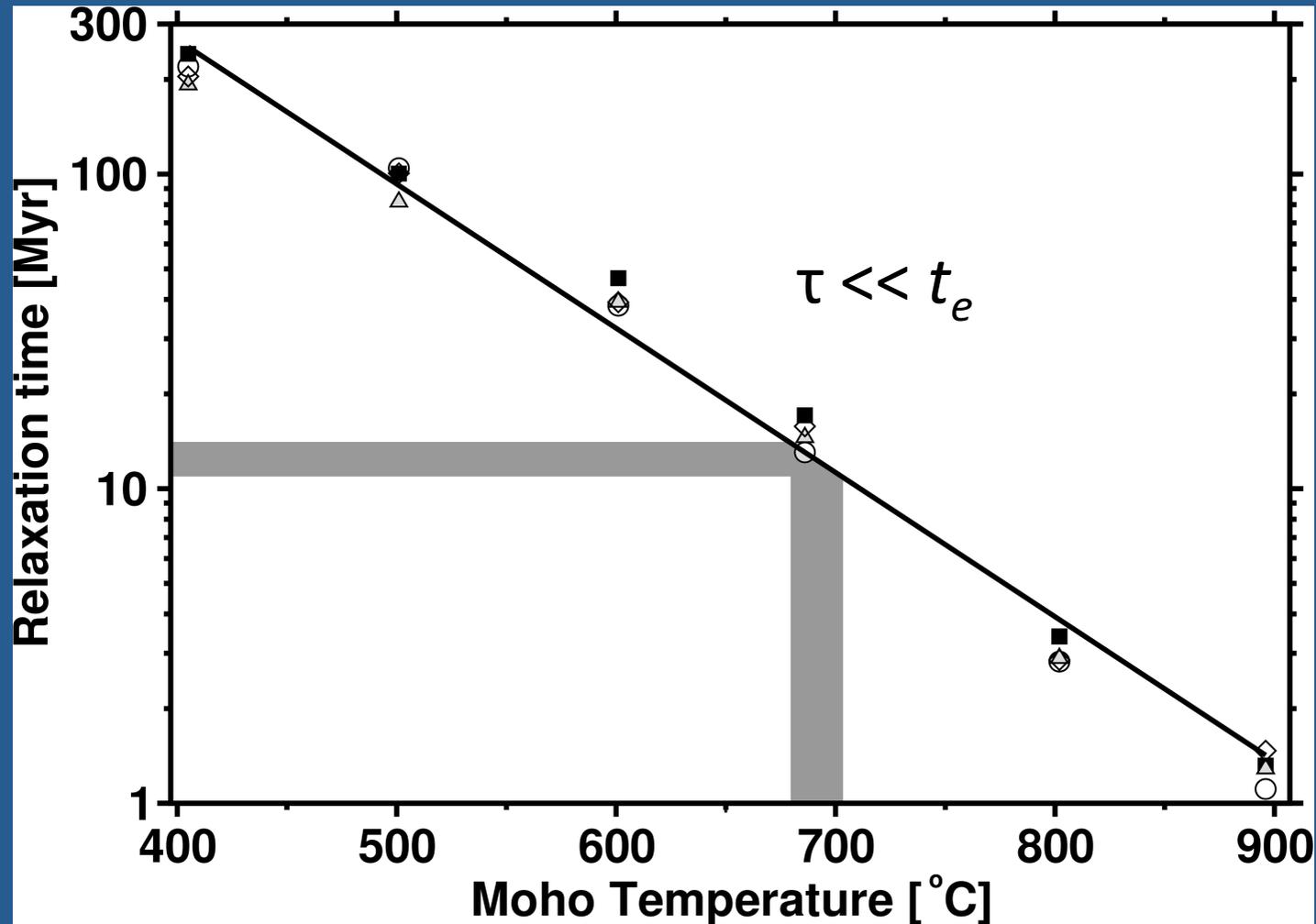


- a. wavy bedding
- b. climbing ripples and flat-pebble conglomerate
- c. Fenestrae and carbonate precipitates

- Maddina Formation  $t_e \leq 11$  Myr
- Kylenea Formation  $t_e < 14$  Myr

Geochronology from *Blake et al. (2004)*

# Cooling of the crust in the East Pilbara



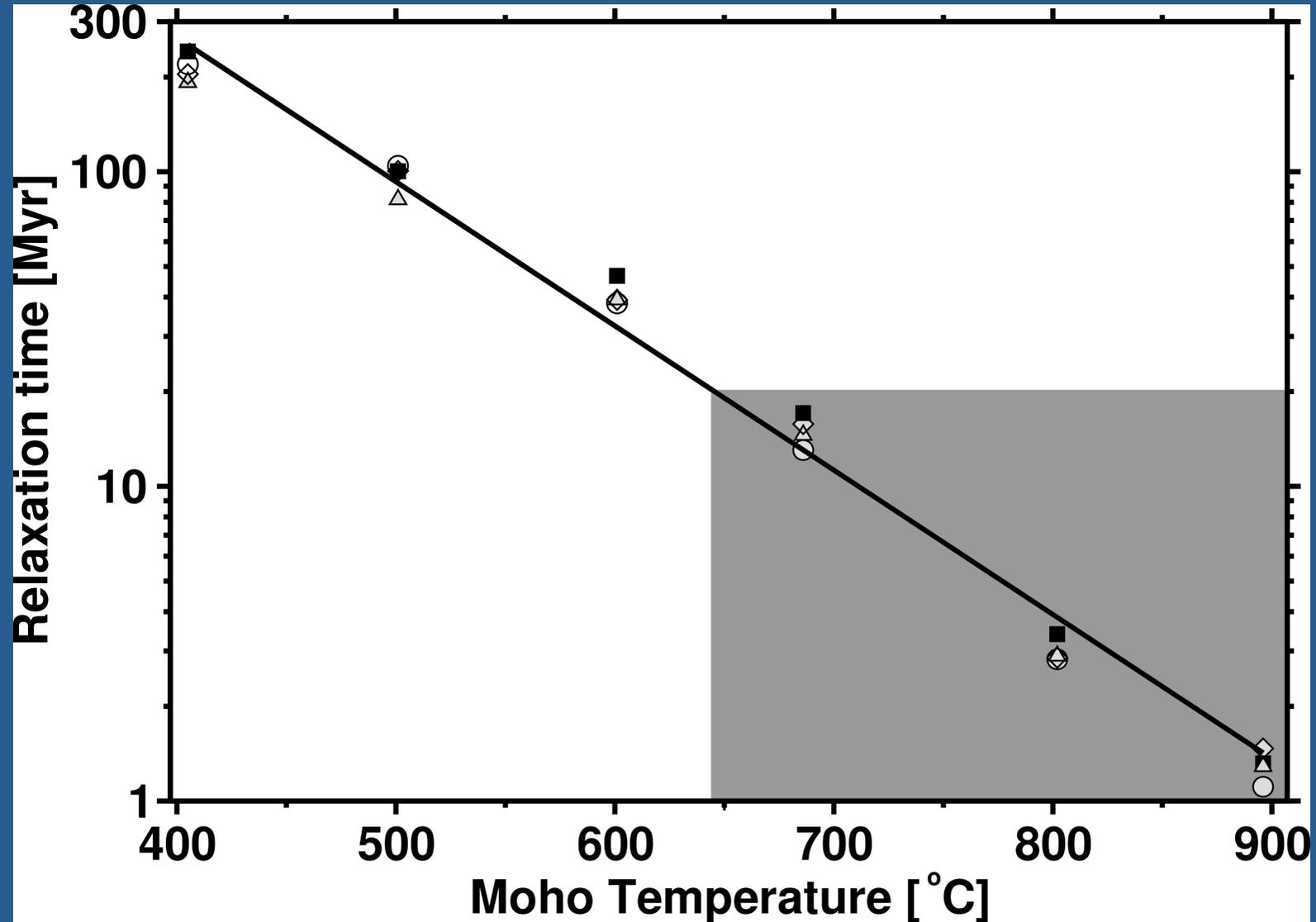
- At 2.7 Ga:  $T_{Moho} \geq 670^\circ\text{C}$
- Present-day:  $T_{Moho} \approx 480^\circ\text{C}$  (surface heat flow)
- Cooling by at least  $\sim 200^\circ\text{C}$

# Relaxation time of subaqueous Archean CFBs

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$$\tau \ll t_e < 20 \text{ Myr}$$

# Archean geotherms were hot



→ Eruption duration of well-constrained events implies  $T_{Moho} \gg 650^{\circ}\text{C}$

# Conclusions

- The flow of hot, ductile lower continental crust was a key process that maintained Archean CFBs below sea level
- Neoproterozoic cooling of the lithosphere explains the contrasted abundance of subaqueous CFBs in Archean vs. post-Archean
- The subsidence history of CFBs can be used to place broad constraints on Archean geotherms
- Well constrained (geochemistry, geology, geochronology) CFBs suggest Archean Moho temperatures  $\gg 650^{\circ}\text{C}$



# Extended results

