

Where no wave has gone before: unconventional elastic wave fields in exotic regimes

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Crew of Enterprise

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Outline

- Introduction
- Unconventional waves
 - Tunnel
 - Borehole
 - Seafloor
- Continuing mission
 - full waveform inversion



Borehole





Full Waveform Inversion



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Goals of seismic technique



1. Imaging: localisation of interfaces (migration)

2. Material parameter (tomography)

- P-wave velocity
- S-wave velocity
- Density
- Attenuation
- Anisotropy

petrophysical relations: pore scale

- porosity
- gas/fluids
- composition



Conventional seismic methods

- Analyse specific wave types only, e.g.
 - P-wave reflection seismics
 - First arrival tomography
 - Surface waves
- Use only a small portion of the information available
 - Travel times
 - Amplitudes of specific waves
- Most information is neglected



Send out new waves and methods

Use more information

- multi-parameter imaging
- higher resolution
- reduction of ambiguities

Strategies

- 1. "unconventional" elastic wavefields
- 2. full waveform inversion

Simulation of elastic wavefields



- Motivation
 - understand complex wave propagation
 - verify new methods
 - kernel of full waveform inversion
- Requirements
 - full elastic wave field
 - arbitrary media including strong contrasts



Parallel Finite-Difference Method



Domain Decomposition



MPI

(Bohlen, 2002; Bohlen & Saenger, 2006)

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Full Waveform Inversion



Tunnel Seismic Prediction







(Bohlen et al., 2007)















(Bohlen et al., 2007)



Prediction with RSSR waves



Field application



TM 5520

TM 5500

(Bohlen et al., 2007, Lüth et al., 2007)

5m

5g

S

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Field application



(Bohlen et al., 2007)



Field application



Tunnel seismic prediction - conclusions

- "New" wave type discovered by wave simulation
- RSSR is becoming practise for tunnel seismic prediction



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Borehole seismic prediction

Prediction ahead of the bit \rightarrow Logging While Drilling (LWD)

- Depth reference for marker horizons
- Fault detection
- Geo-steering



Problem: Vicinity of the bit not suited

for seismic sources and receivers



Borehole seismic prediction

Prediction ahead using TSST-waves ?























(Hellwig & Bohlen, 2008)






































Borehole seismic prediction - conclusions

- TSST is strongest seismic response from reflectors ahead of the bit
- Maximum between 10° and 30° reflector dip



(Hellwig & Bohlen, 2008)



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Motivation

- Inversion of shear wave velocity Vs
- Vs correlates with shear strength
- Applications
 - Offshore constructions, e.g. windparks





Elastic waves in shallow water

DIV (T=0.02s) uvater uvater

ROT (T= 0.02s)



() 45

Scholte wave acquisition



(Bohlen et al., 2004)



Scholte wave observations



(Bohlen et al., 2004; Kugler et al., 2005; Klein et al., 2005; Kugler et al., 2007)



Scholte wave tomography



(Kugler et al., 2007)

Scholte wave tomography



(Kugler et al., 2007)

(Müller et al., 2002)

Additional information on shear properties

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- Scholte waves are most prominent wavefield in shallow water
- Scholte tomography yields 3-D Vs models
- Important for seafloor stability exploration







Strange new wave: Love wave ?!







Strange new wave: Love wave ?!







(Bussat, 2006, PhD thesis)



Surprising observation: Love wave excited by airgun





- Joint inversion of Scholte and Love wave dispersion curves yield consistent 1-D models of Vs and density
- Higher constrains on density model

(Bussat, 2006, PhD thesis)



Sea floor roughness/ripples ?







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Full Waveform Inversion



Seismologist dream

Exploit the full information content of seismic signals !





Numerical implementation of adjoint FWI



Starting point of FWT



(Kurzmann, 2012)

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Starting point of FWT



⁽Kurzmann, 2012)

Inverted **P-wave** velocity model



(Köhn et al., 2012)



Inverted **P-wave** velocity model



Inverted **S-wave** velocity model



V_s [m/s]

(Köhn et al., 2012)



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Inverted **S-wave** velocity model



Inverted **density** model



(Köhn et al., 2012)



Inverted density model



Application to marine streamer data (North Sea)



Data pre-processing: Noise attenuation, spreading transformation, near offsets

(Przebindowska et al., 2011, 2012)



Data fitting



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(Przebindowska et al., 2011, 2012)

Preliminary results





Summary

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Borehole





Full Waveform Inversion



... to boldly go where no man has gone before

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Appendix (additional slides)



Anatomy of elastic wave field around a tunnel



S-wave velocity model



2.5-D Simulation



Born approximation for scattered wave fields

Helmholtz equation:
$$\left[\nabla^2 + \frac{\omega^2}{V^2(x)}\right] u(x, x_s, \omega) = -f(\omega)\delta(x - x_s)$$

I) Linear approximations

Model:
$$\frac{1}{V^2(x)} := m(x) = m_0(x) + \delta m(x)$$

Wave field: $u_{obs}(x,x_s,\omega) = u_0(x,x_s,\omega) + \delta u(x,x_s,\omega)$

II) Born approximation: $\delta m \cdot \delta u := 0$

Scattered wave field:

$$\delta u(x_g, x_s, \omega) = \omega^2 f(\omega) \int_V \delta m(x) g(x, x_s, \omega) g(x_g, x, \omega) d^3 x$$



Interpretation of scattered wave field

Scattered wave field:





Derivation of FWI

Misfit:
$$E = \frac{1}{2} \int d\omega \sum_{s} \sum_{g} \left| \delta u(x_g, x_s, \omega) \right|^2$$

Iterative model update: $m(x)^{n+1} = m(x)^n + \mu^n \gamma(x)^n$

along the gradient: $\gamma(x) = -\nabla_m E = \frac{\partial E}{\partial m}$

using the Born approximation:

$$\delta u(x_g, x_s, \omega) = \omega^2 f(\omega) \int_V \delta m(x) g(x, x_s, \omega) g(x_g, x, \omega) d^3 x$$

we obtain:

$$\gamma(x) = \int d\omega \sum_{s} \sum_{g} \omega^{2} \operatorname{Re} \left[\left(f(w)g(x,x_{s},\omega) \right) \cdot \left(\delta u^{*}(x_{g},x_{s},\omega)g(x_{g},x,\omega) \right) \right]$$
forward wave field
from source
backward wave field
from receiver



Advantage I: maximum resolution



Traveltime tomography resolution: Fresnel Zone

Waveform tomography resolution: wavelength



2.5-D Simulation



Model:

- dipole source (n = 1)
- vol. injection source
- Ricker-wavelet
- 5000 Hz center

frequency

Elastic material parameters			
	v_P in m/s	v_S in m/s	ρ in kg/m^3
formation 1	2520	1000	2300
formation 2	1850	850	2000
orehole fluid	1600	0	1500
drilling tool	5860	3310	7700



Scholte wave tomography



20 m

Scholte wave tomography



local 1-D inversion





Viscoelastic wave equations

stress-strain relation

$$\begin{split} \dot{\sigma}_{ij} &= \frac{\partial v_k}{\partial x_k} \left\{ \pi \left(1 + \tau^p \right) - 2\mu \left(1 + \tau^s \right) \right\} + 2 \frac{\partial v_i}{\partial x_j} \mu \left(1 + \tau^s \right) + \sum_{l=1}^L r_{ijl} \quad \text{if } i = j \\ \dot{\sigma}_{ij} &= \left(\frac{\partial v_i}{\partial x_j} + \frac{\partial v_j}{\partial x_i} \right) \mu \left(1 + \tau^s \right) + \sum_{l=1}^L r_{ijl} \quad \text{if } i \neq j \end{split}$$

memory variables

 $\varrho \frac{\partial v_i}{\partial t} = \frac{\partial \sigma_{ij}}{\partial x_j} + f_i$

$$\begin{aligned} \dot{r}_{ijl} &= -\frac{1}{\tau_{\sigma l}} \left\{ (\pi \tau^p - 2\mu \tau^s) \frac{\partial v_k}{\partial x_k} + 2 \frac{\partial v_i}{\partial x_j} \mu \tau^s + r_{ijl} \right\} & \text{if } i = j \\ \dot{r}_{ijl} &= -\frac{1}{\tau_{\sigma l}} \left\{ \mu \tau^s \left(\frac{\partial v_i}{\partial x_j} + \frac{\partial v_j}{\partial x_i} \right) + r_{ijl} \right\} & \text{if } i \neq j \end{aligned}$$

equation of momentum conservation



Scholte wave acquisition

OBH/OBS



Airgun







Problem: non-linearity of the misfit function

starting model



Bunks et al. (1995)

low frequencies

