





Air permeametry on outcrop analogues: a composite image of the Neogene aquifer, Belgium.

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Introduction	Methods	Results	Conclusions

- Introduction
- Methods
- Results
- Conclusions







Introduction	Methods	Results	Conclusions

- Introduction
 - Research context
 - Neogene aquifer
- Methods
- Results
- Conclusions







Introduction	Methods	Results	Conclusions

Research context

- Saturated hydraulic conductivity (K_s)
 - one of the most important parameters
 - determining groundwater flow and contaminant transport
 - in both unsaturated and saturated porous media
- In-situ K_s measurements
 - remain very complex and scale-dependent
 - air permeameters have being used effectively in the field as an indirect method to determine $K_{\rm s}$
- Case study: Neogene aquifer, Belgium
 - important groundwater resource
 - subject to hydrogeological assessments in the framework of potential future disposal of radioactive waste







Introduction	Methods	Results	Conclusions

STUDIECENTRUM VOOR KERNENERGIE CENTRE D'ETUDE DE L'ENERGIE NUCLEAIRE

Neogene aquifer - Location











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Introduction

- Methods
 - Air permeameter measurements
 - Validation in the lab
 - Numerical upscaling
- Results
- Conclusions







Introduction	Methods	Results	Conclusions

Air permeability measurements

Tinyperm II air permeameter

New England Research & Vindum Engineering 2011

• Equation from literature to convert k_a to K_s Iversen et al 2003











Introduction	Methods	Results	Conclusions

Validation in the lab

- Kopecky ring samples of different lithologies
- Constant head permeameter tests











Introduction	Methods	Results	Conclusions

Numerical upscaling

- Derive one single K tensor with flow conservation
- Approaches
 - Permeameter-type setup
 - Prescribed head at in- and outflow boundaries
 - No-flow at parallel boundaries
 - Combination of 4 boundary condition setups
 - Prescribed head everywhere
 - 4 flow directions

$$-\begin{bmatrix} \nabla \bar{h}_{x1} & 0 & \nabla \bar{h}_{z1} \\ 0 & \nabla \bar{h}_{z1} & \nabla \bar{h}_{x1} \\ \nabla \bar{h}_{x2} & 0 & \nabla \bar{h}_{z2} \\ 0 & \nabla \bar{h}_{z2} & \nabla \bar{h}_{x2} \end{bmatrix} \begin{bmatrix} K_{xx} \\ K_{zz} \\ K_{xz} \end{bmatrix} = \begin{bmatrix} \bar{q}_{x1} \\ \bar{q}_{z1} \\ \bar{q}_{z2} \\ \bar{q}_{z2} \\ \dots \end{bmatrix}$$

e.g. Zhou et al 2010; Li et al 2011

$$-\begin{bmatrix} \nabla \bar{h}_x & 0\\ 0 & \nabla \bar{h}_z \end{bmatrix} \begin{bmatrix} K_{xx}\\ K_{zz} \end{bmatrix} = \begin{bmatrix} \bar{q}_x\\ \bar{q}_z \end{bmatrix}$$





Introduction	Methods	Results	Conclusions

- Introduction
- Methods
- Results
 - Air permeameter measurements
 - Validation in the lab
 - Numerical upscaling
- Conclusions







Introduction	Methods	Results	Conclusions

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Quaternary

Pliocene

Miocene

Air permeameter measurements









Introduction	Methods	Results	Conclusions

STUDIECENTRUM VOOR KERNENERGIE CENTRE D'ETUDE DE L'ENERGIE NUCLEAIRE

Quaternary

Pliocene

Miocene

Air permeameter measurements



14





Introduction	Methods	Results	Conclusions

CENTRE D'ETUDE DE L'ENERGIE NUCLEAIRE

Air permeameter measurements – 2 examples



 $\log_{10}(K)$ -6.29 -4.80 -4.43 -3.58 -3.37 -3.21 -1.63 (*m*/s) 10 20 30 cm 0

10 20 30 cm 0

Poederlee Sands







Introduction	Methods	Results	Conclusions

Air permeameter measurements

חו	Sediment	log ₁₀ (K _s) (m/s)			
		Min	Max	Mean	Range
2	Kleine Nete point bar sands	-4.3	-3.4	-3.7	0.8
d	Kleine Nete channel fill	-9.2	-8.3	-8.9	0.9
b	Campine clay-sand complex profile	-8.1	-3.6	-5.1	4.6
С	Ice wedge & cryoturbation structures	-7.8	-2.5	-4.9	5.3
d	Lommel Sands	-6.4	-2.3	-4.2	4.1
е	Mol Sands	-4.0	-3.1	-3.7	0.9
f	Top Poederlee Formation	-6.5	-3.7	-4.5	2.8
g	Bottom Poederlee Formation 1	-7.8	-4.4	-5.9	3.4
h	Bottom Poederlee Formation 2	-5.4	-1.6	-3.9	3.9
i	Hukkelberg stratigraphical boundary	-8.2	-3.7	-5.3	4.5
j	Kasterlee Sands	-6.7	-2.0	-3.9	4.6
k	Kasterlee Clay	-6.6	-2.2	-4.4	4.4
I	Clayey top Diest Formation	-6.1	-1.7	-3.7	4.5
m	Diest Sands	-5.0	-3.6	-4.1	1.3
n	Bolderberg Sands 1	-4.2	-3.5	-3.9	0.8
0	Bolderberg Sands 2	-4.3	-3.5	-3.9	0.8







Introduction	Methods	Results	Conclusions

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Validation in the lab







Introduction	Methods	Results	Conclusions

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Numerical upscaling – 2 examples











IntroductionMethodsResultsConclusionsIntroductionMethodsResultsConclusions

Numerical upscaling

	Carlina ant	Permeameter setup		
IJ	Sediment	log ₁₀ (K _{xx})	log ₁₀ (K _{zz})	VANI
С	Ice wedge & cryoturbation structures	-4.15	-5.57	26.42
d	Lommel Sands	-4.19	-4.19	1.00
е	Mol Sands	-3.64	-3.60	0.92
f	Top Poederlee Formation	-3.88	-5.31	26.71
g	Bottom Poederlee Formation 1	-4.80	-6.23	26.86
h	Bottom Poederlee Formation 2	-3.84	-4.02	1.51
j	Kasterlee Sands	-3.86	-3.85	0.98
k	Kasterlee Clay	-4.13	-5.71	38.31
I	Clayey top Diest Formation	-3.56	-3.88	2.08
m	Diest Sands	-4.06	-4.14	1.21
n	Bolderberg Sands 1	-3.81	-3.84	1.07
0	Bolderberg Sands 2	-3.86	-3.88	1.05







Introduction	Methods	Results	Conclusions

- Introduction
- Methods
- Results
- Conclusions







Introduction	Methods	Results	Conclusions

Conclusions

- A handheld air permeameter is an efficient tool to characterise hydraulic conductivity
- It offers access to spatial variability on scales that are not feasible by classical core-based techniques
- Validation shows that the K_s estimates are within one order of magnitude from the laboratory analyses results
- Complex stratigraphical settings can be effectively characterised if outcrops of the different components are available







Introduction	Methods	Results	Conclusions

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Centimeter-scale secondary information on hydraulic conductivity using a hand-held air permeameter on borehole cores

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> Figure 1: Core sampling at SCK-CEN: a) schematic of sisb removal from main core, b) inclaion is

made in main core to the appropriate depth for also removal, c) 100 cm² steel

conductivity determination is the lab, c) steel core inserted in core material,

el vecuum sealing of mail

core for long-term storage fi display of core slabs for

core for hydrautic

Introduction

- Saturated hydraulic conductivity (K_s)
- Is one of the most important parameters
- determining groundwater flow and contaminant transport
 in both unsaturated and saturated porous media
- Small-scale variability of K, is key to obtain effective transport parameters and
- to explain K, measurements or inverse estimates at the larger scale ~350 m of borehole core is available at Mol/Dessel, Beiglum (Fig. 1; Beerlen et
- al. 2010)
- sediments of Miocene to Pleistocene age, marine to continental origin
- sand to clayey sand with distinct clay lenses, with varying glauconite content
- 2 samples each 2 meters with K, from constant head permeameter tests in
- the lab K, range of 7 orders of magnitude
- Thin slabs separated from the cores are analysed in this study (Fig. 1)

Methods

Measurements

- Use of the TinyPerm II hand-heid air permeameter device (Fig. 2; New England Research & Vindum Engineering 2011)
- > 5000 measurements on the dry borehole core slabs at 5 cm resolution, performed within 5 days (Flos. 3 & 6)
- Equation of Loi et al. (1999) to convert air permeability to a K estimate, since perfect agreement between intrinsic permeability estimated from measurement of air and water flow cannot be expected.
- Additional measurements to quantify measurement error and operator influence (Fig. 4)
 Calibration with the lab K, measurements with a linear mixed-effects model, with random effects for

both the stratigraphy and borehole factors (Fig. 5)

- Spatial analysis
- Variography for the lab measurements and air permeameter estimates after standardisation
 Fitting an intrinsic model of co-regionalisation (Goovaerts 1997)
- Interpolation of lab K, data with air permeameter estimates as secondary variable
- Leave-two-out cross-validation to quantify the predictive uncertainty on K, and the accuracy
 gain with using the secondary data (samples at a distance of about 10 cm are left out together)

Results

The relative differences between the stratigraphical units corresponds to the lab analyses observations (Fig. 3)

- A systematic bias and smaller range of K values is predicted using the equation from Loil et al. (1999)
 Measurement error as well as the systematic bias introduced by the operator are small compared to the intrinsic K variability (compare Fibs) 3.8.4)
- Correlation between lab measurements and air permeability estimates is 0.74 and increases after calibration to 0.84 (see Fig. 5)
- After standardisation of the data, an intrinsic model of coregionalisation was fitted to the experimental variograms with two nested spherical models (Fig. 6). One for the short range (0.4 m) and one for the long range (12 m).
- Predictions are presented on Fig. 7, and show a lot of small-scale heterogeneity, as well as clear zones of lower K values. In the zones for which core slabs were missing, important uncertainty remains, as indicated by the larger confidence intervals.
- Cross-validation results (see Fig. 8)
- Performance kriging: MSE: 1.13; ME: 0.04; R^a: 0.31
- Performance co-kriging: MSE: 0.79; ME: -0.02; Rº: 0.71
- Especially the low K, range predictions are improved

Conclusions

- Hand-heid air permeameter measurements on undisturbed borehole cores provide a very cost-effective way to obtain high-resolution K data
- Even core slabs that have been lying open to air, and have been subject of several investigations during a few years, provide useful information
- Without calibration, reliable relative K estimates can be obtained, and equations from literature provide absolute K estimates (e.g. Loil et al. 1999)
- Calibration with laboratory measurements improves the accuracy, and is recommended for core slabs of this state



Figure 1: Despits of K, values from a) the lab analysis, and b) the air permeanater K, astimates (based on Loif et al., 1996), for each of the stratigraphical units. The whistens astend to the most extreme date point that is no more than 1.5 times the intergrandia range from the box.



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Figure 5: Calibration of the sir permeaneter K, estimates (x; using Loif et al., 1999), with laboratory analyses deb, by means of a linear interd effects model with a random effect for each strategratioal unit and each forehold; (b).



long range (12 m) spherical variogram model



Acknowledgements

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ente on borehole

Figure 2: Performing air pert

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Figure 7: Vertical profiles of K, values from the leb analyses, the air permeaneler K, astimates after calibration, and the coloriging estimates and 85% confidence interval

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Introduction	Methods	Results	Conclusions
Introduction			

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Introduction	Methods	Results	Conclusions

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