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## Introduction

- Saturated hydraulic conductivity ( $K_{a}$ )
- is one of the most important parameters
- determining groundwater flow and contaminant transport
- in both unsaturated and saturated porous media
- Small-scale variability of  $K_{s}$  is key to obtain effective transport parameters and
- to explain K measurements or inverse estimates at the larger scale
- ~350 m of borehole core is available at Mol/Dessel, Belgium (Fig. 1; Beerten et al. 2010)
- sediments of Miocene to Pleistocene age, marine to continental origin
- sand to clayey sand with distinct clay lenses, with varying glauconite content • 2 samples each 2 meters with  $K_{c}$  from constant head permeameter tests in the lab
- K range of 7 orders of magnitude
- Thin slabs separated from the cores are analysed in this study (Fig. 1)

## Methods

- Measurements
- Use of the TinyPerm II hand-held air permeameter device (Fig. 2; New England Research & Vindum Engineering 2011)
- > 5000 measurements on the dry borehole core slabs at 5 cm resolution, performed within 5 days (Figs. 3 & 6)
- Equation of Loll et al. (1999) to convert air permeability to a K estimate, since perfect agreement between intrinsic permeability estimated from measurement of air and water flow cannot be expected
- Additional measurements to quantify measurement error and operator influence (Fig. 4) Calibration with the lab K measurements with a linear mixed-effects model, with random effects for
- both the stratigraphy and borehole factors (Fig. 5)
- Spatial analysis
  - Variography for the lab measurements and air permeameter estimates after standardisation
  - Fitting an intrinsic model of co-regionalisation (Goovaerts 1997)
  - Interpolation of lab  $K_{s}$  data with air permeameter estimates as secondary variable • Leave-two-out cross-validation to quantify the predictive uncertainty on  $K_{s}$ , and the accuracy gain with using the secondary data (samples at a distance of about 10 cm are left out together)

## Results

- The relative differences between the stratigraphical units corresponds to the lab analyses observations (Fig. 3)
- A systematic bias and smaller range of  $K_{s}$  values is predicted using the equation from Loll et al. (1999) Measurement error as well as the systematic bias introduced by the operator are small compared to
- the intrinsic  $K_{s}$  variability (compare Figs. 3 & 4) Correlation between lab measurements and air permeability estimates is 0.74 and increases after
- calibration to 0.84 (see Fig. 5)
- After standardisation of the data, an intrinsic model of coregionalisation was fitted to the experimental variograms with two nested spherical models (Fig. 6). One for the short range (0.4 m) and one for the long range (12 m).
- Predictions are presented on Fig. 7, and show a lot of small-scale heterogeneity, as well as clear zones of lower  $K_{a}$  values. In the zones for which core slabs were missing, important uncertainty remains, as indicated by the larger confidence intervals.
- Cross-validation results (see Fig. 8)
- Performance kriging: MSE: 1.13; ME: 0.04; R<sup>2</sup>: 0.31
- Performance co-kriging: MSE: 0.79; ME: -0.02; R<sup>2</sup>: 0.71
- Especially the low  $K_{a}$  range predictions are improved

## Conclusions

- Hand-held air permeameter measurements on undisturbed borehole cores provide a very cost-effective way to obtain high-resolution  $K_{s}$  data
- Even core slabs that have been lying open to air, and have been subject of several investigations during a few years, provide useful information
- Without calibration, reliable relative  $K_{s}$  estimates can be obtained, and equations from literature provide absolute  $K_{s}$  estimates (e.g. Loll et al. 1999)
- Calibration with laboratory measurements improves the accuracy, and is recommended for core slabs of this state



# **Centimeter-scale secondary information on hydraulic conductivity** using a hand-held air permeameter on borehole cores

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