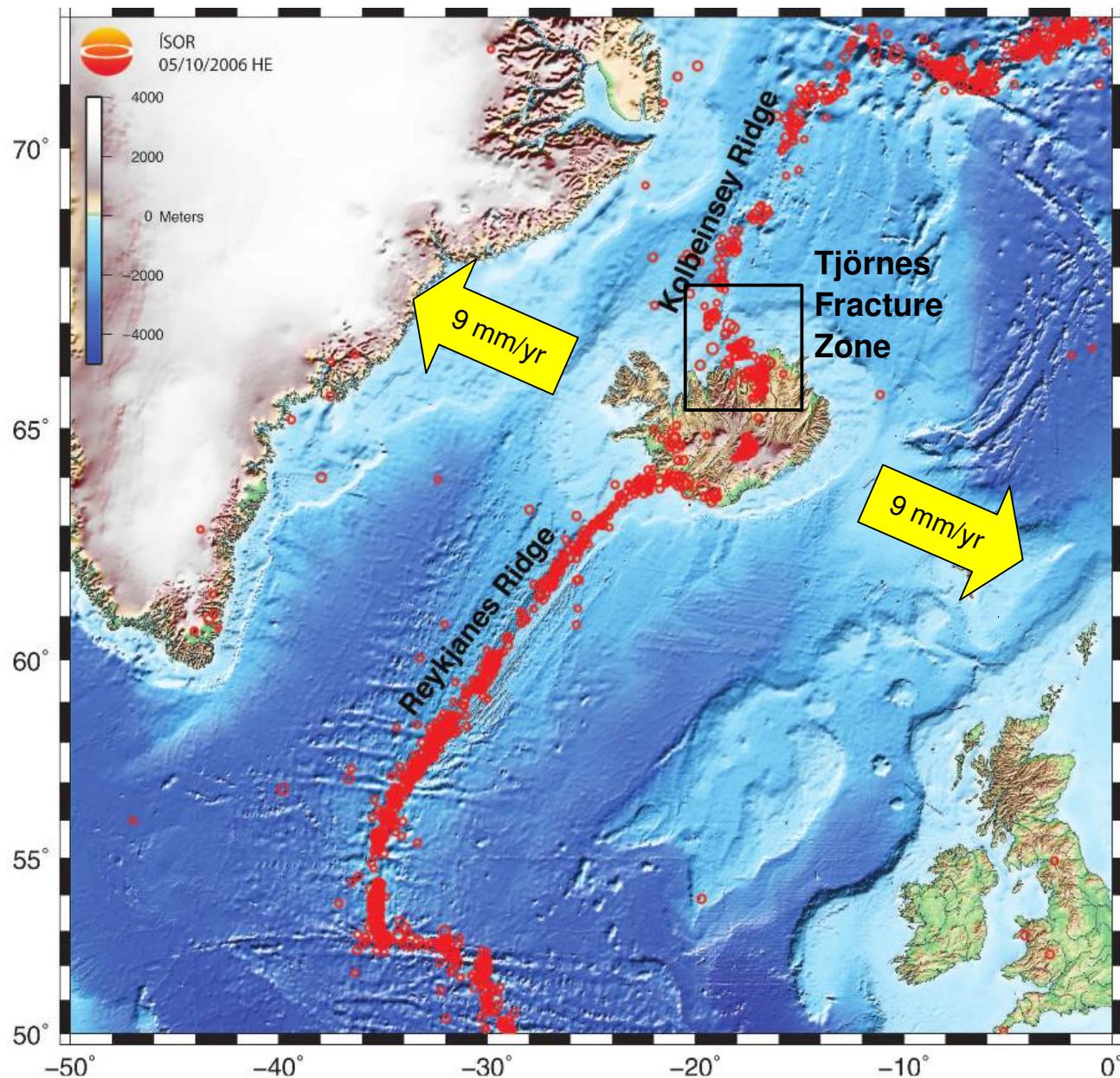


The accumulated seismic moment of the locked Húsavík-Flatey fault, North Iceland, derived from an interseismic model using GPS and InSAR

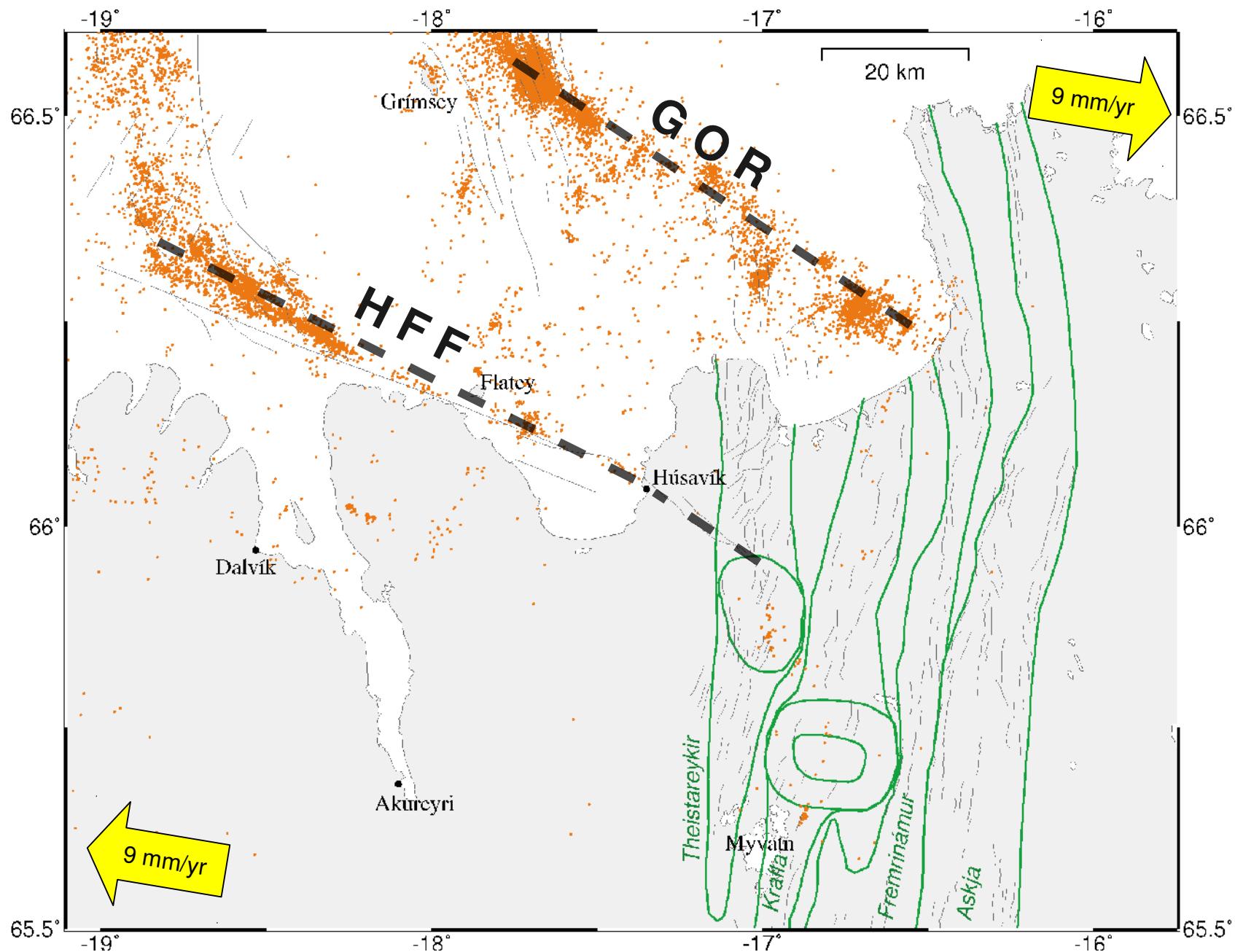
S. Metzger, S. Jónsson, G. Danielsen, S. Hreinsdóttir, F. Jouanne, T. Villemain



Seismicity defines the plate boundary

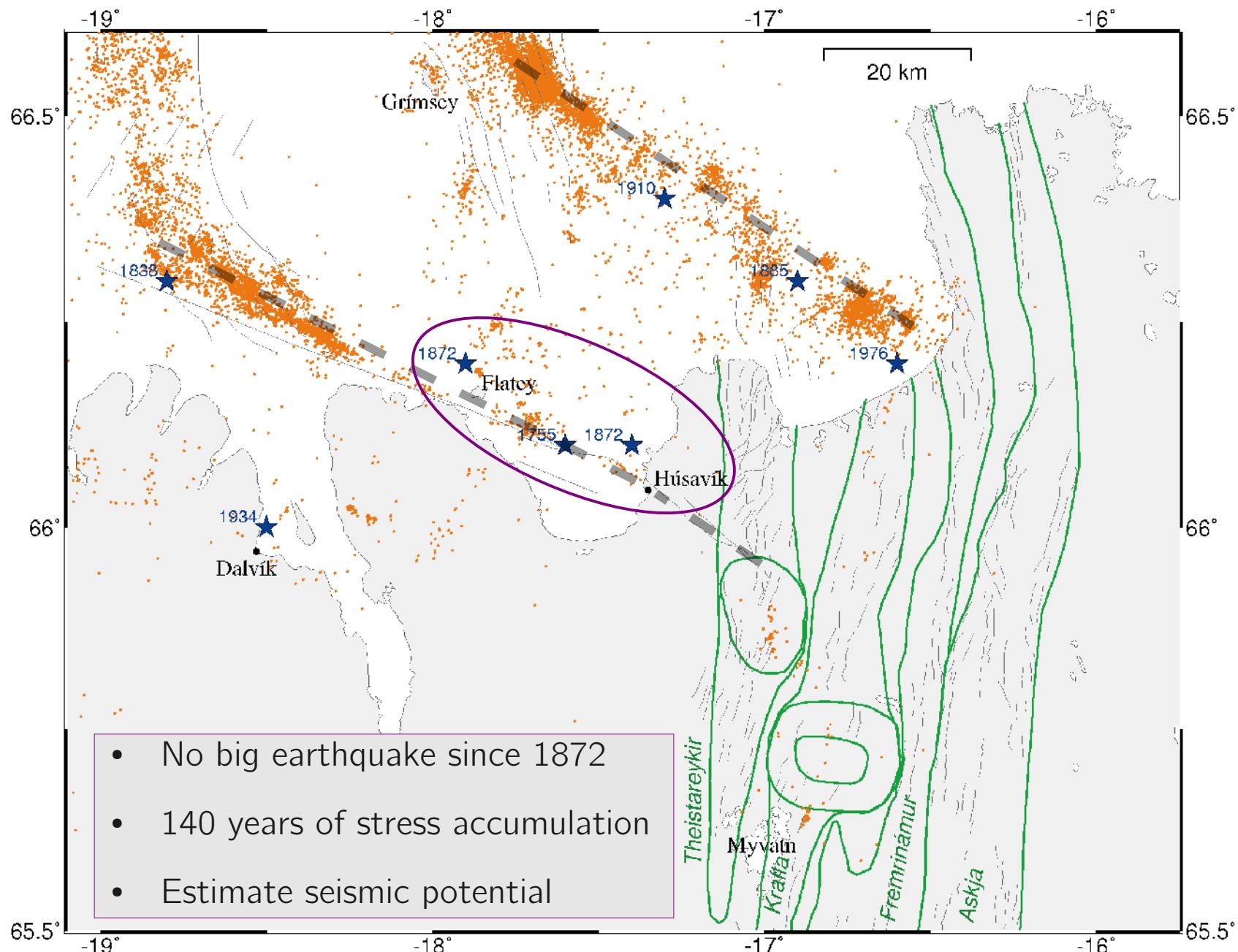


Seismicity defines the plate boundary



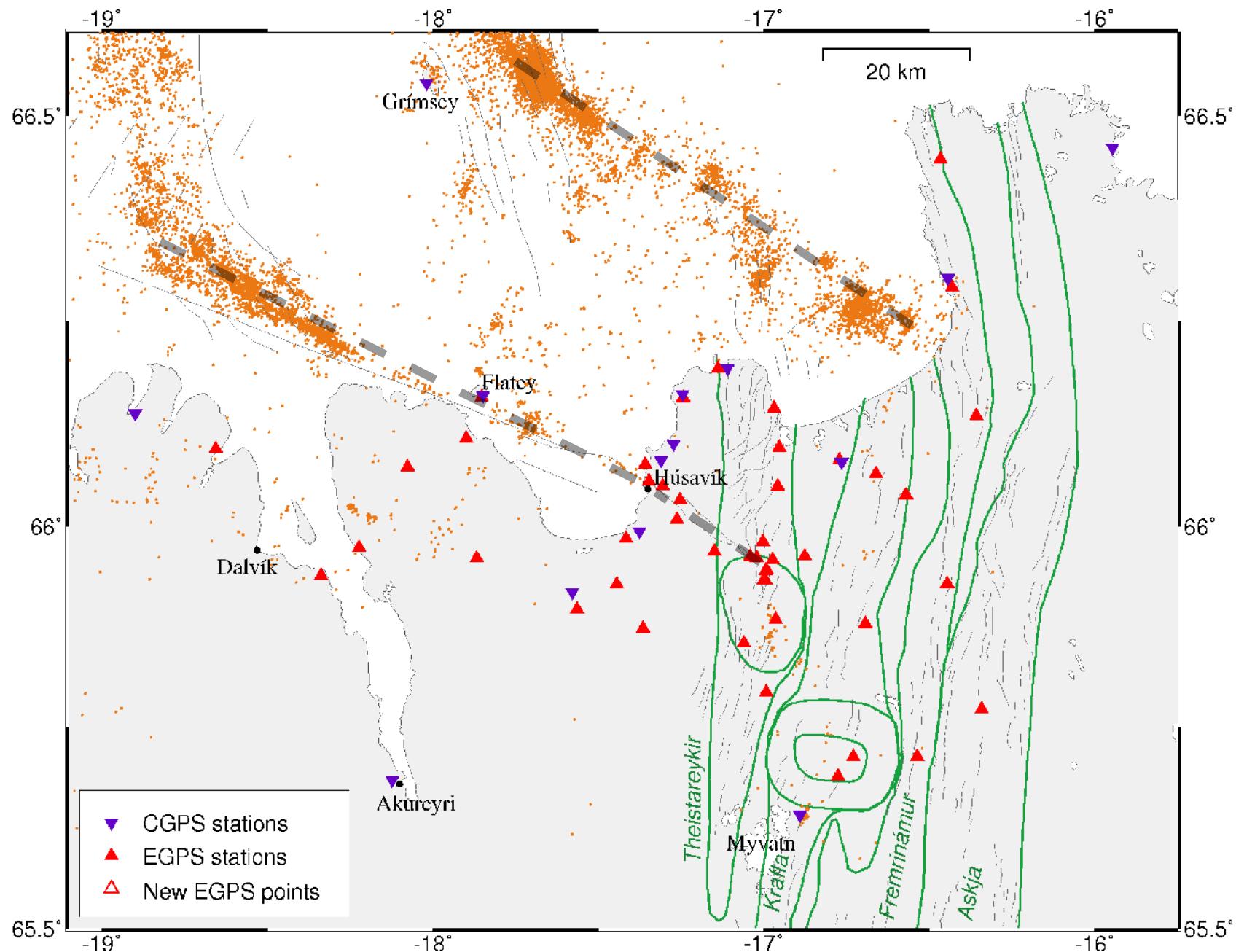


Historical seismicity





GPS network North Iceland



How to estimate steady-state deformation?



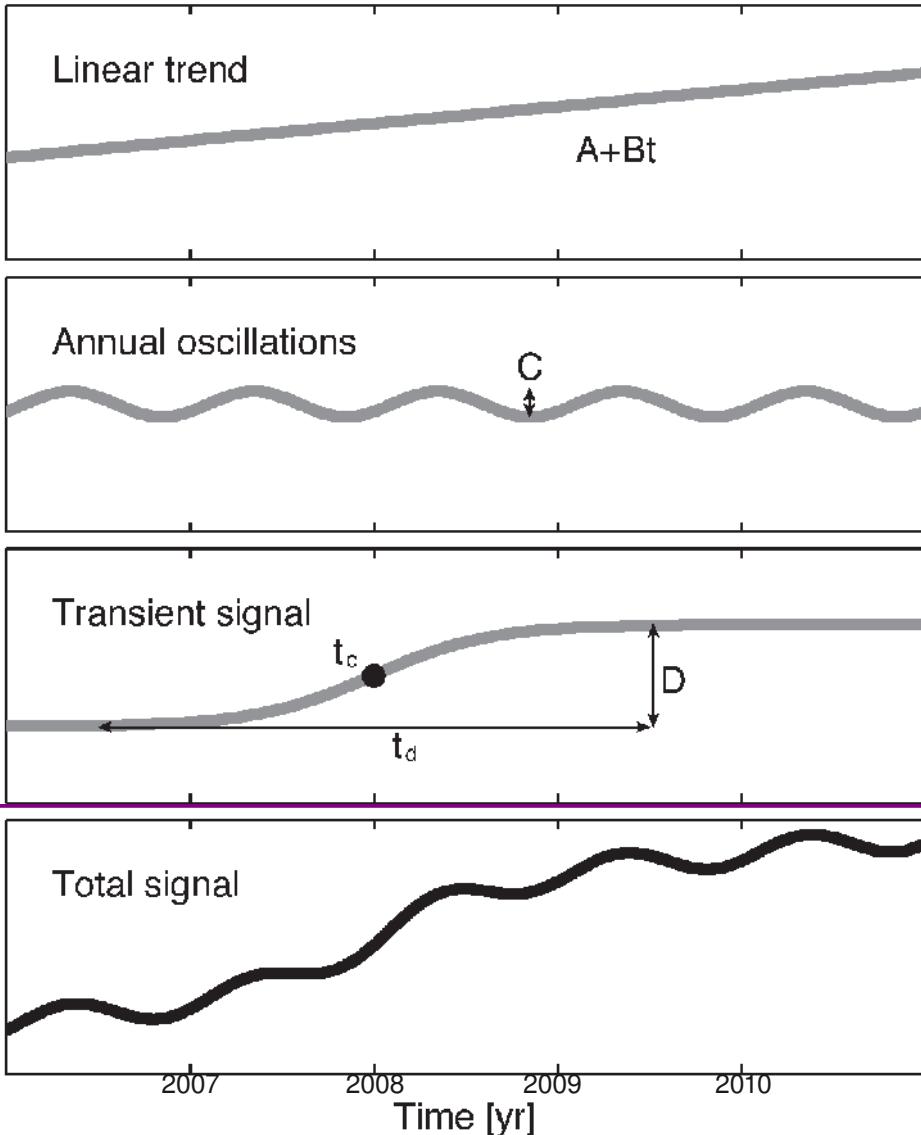
The GPS time-series contain

- 1) steady-state velocity (plate motion)
- 2) annual oscillation (permafrost?)
- 3) transient signal (volcanic uplift@Theistareykir)

⇒ Estimate transient and oscillatory signal

⇒ Extract steady-state velocities

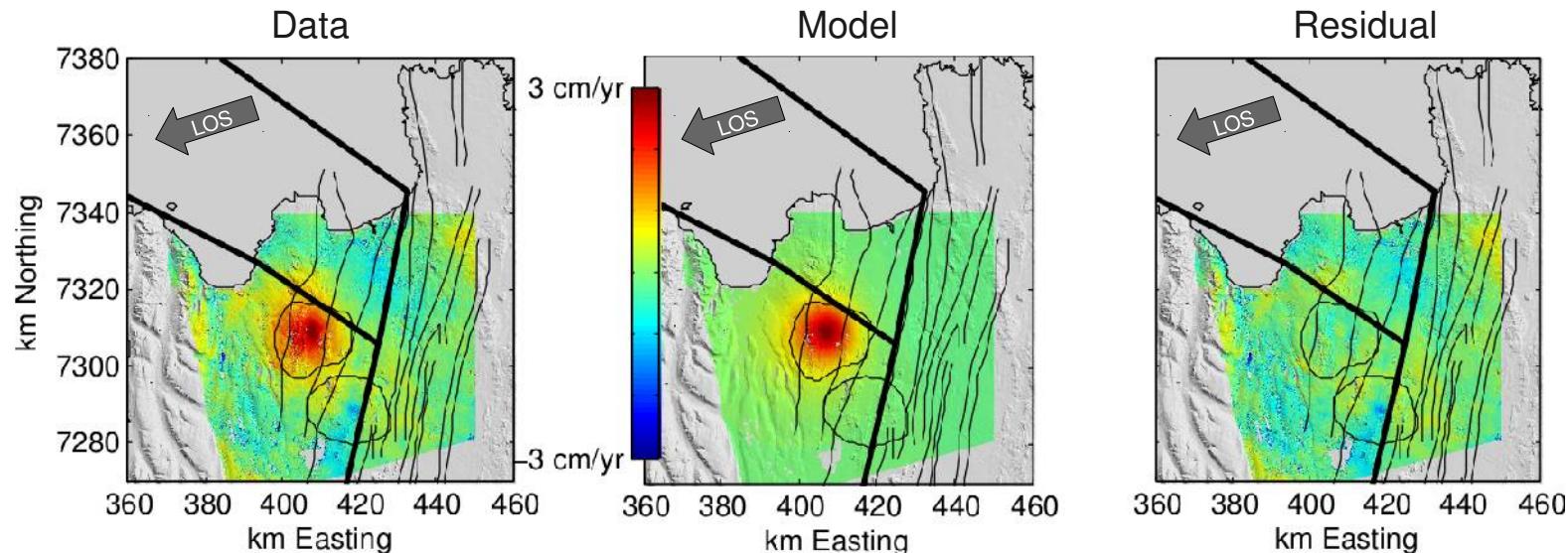
$$A + Bt + C \cos(2\pi t + \Phi) + D \arctan\left(\frac{t - t_c}{t_d}\right)$$



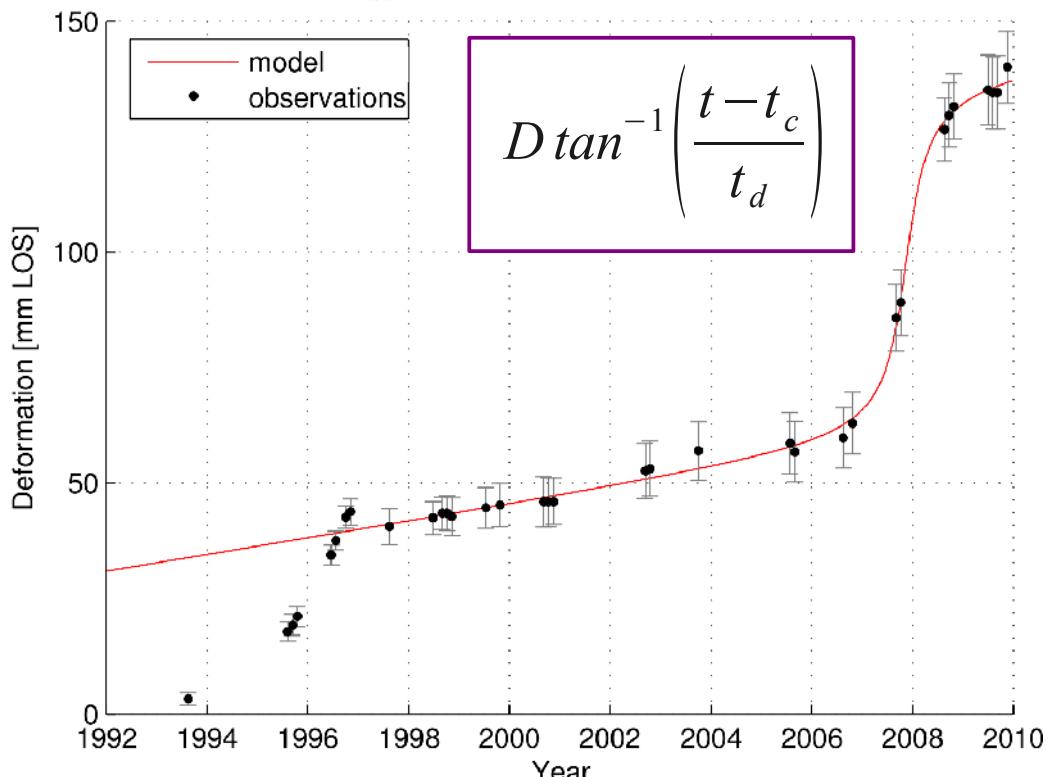
Inflation at Theistareykir volcano



1-yr interferogram
(Metzger et al., 2011)

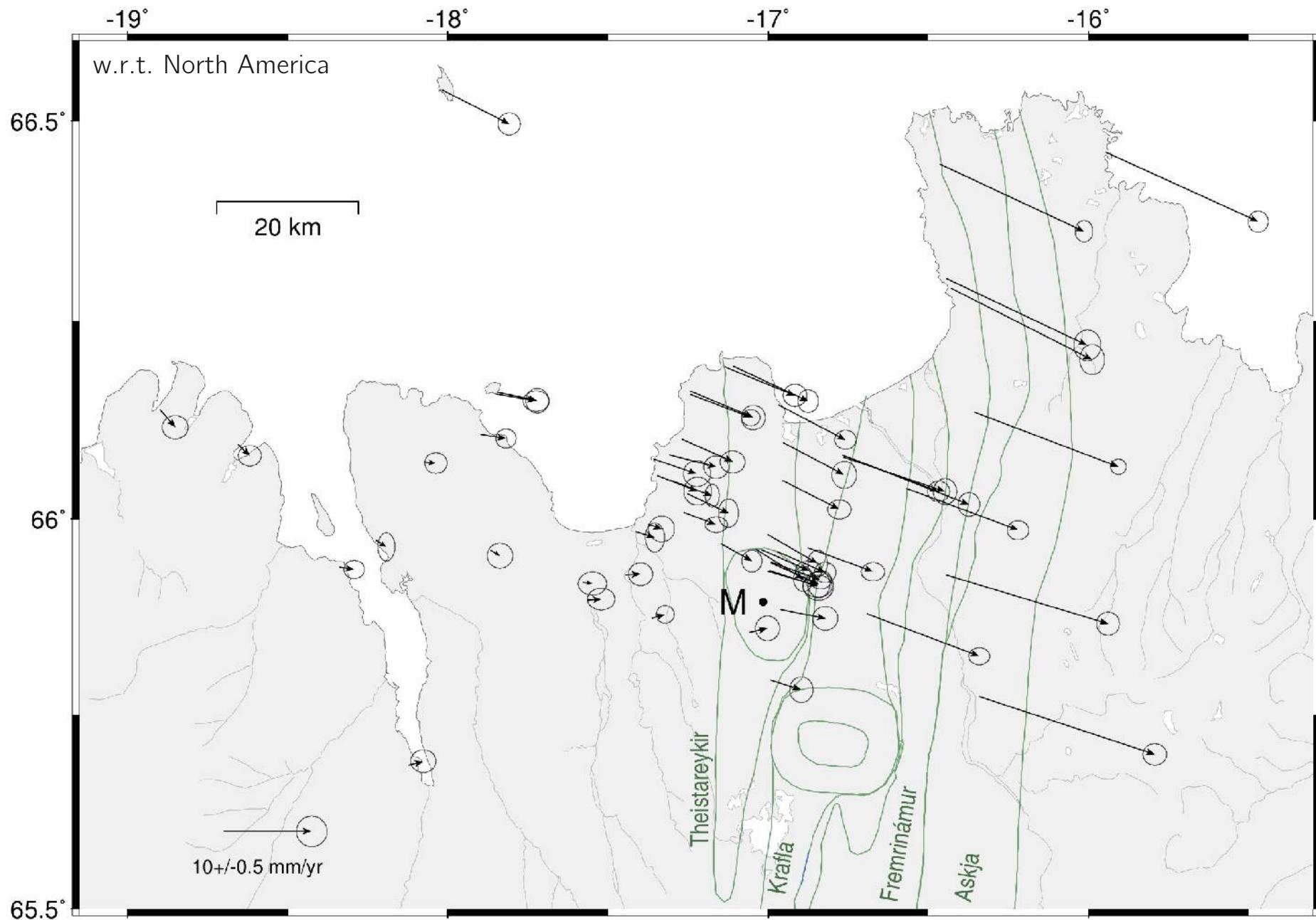


- Mogi model and InSAR data
- Location: Theistareykir volcano
- Duration of inflation: 2007-2009
- Absolute volume change: $25 \times 10^6 \text{ m}^3$
- Depth: 8.5 km
- Absolute uplift: 8 cm





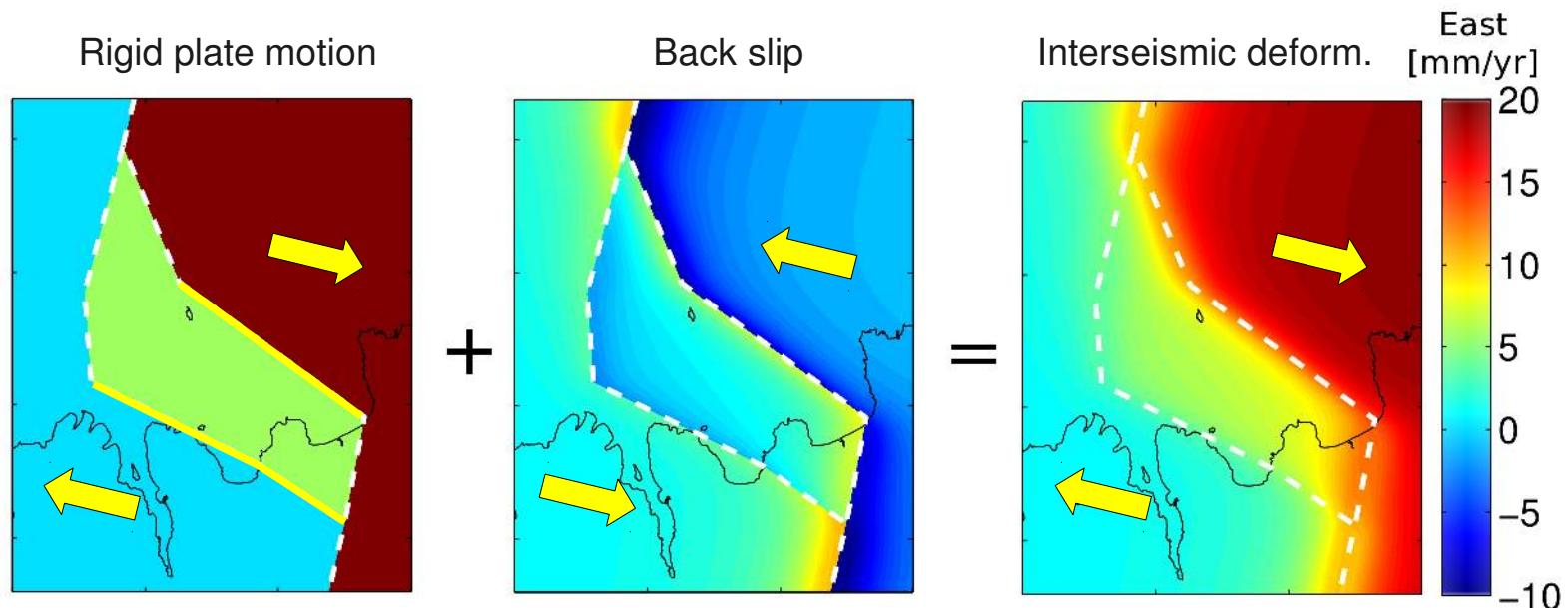
Steady-state velocities





Model concepts

- Nine dislocation segments in an elastic half-space (Okada, 1985)
- Superposition of two signals:

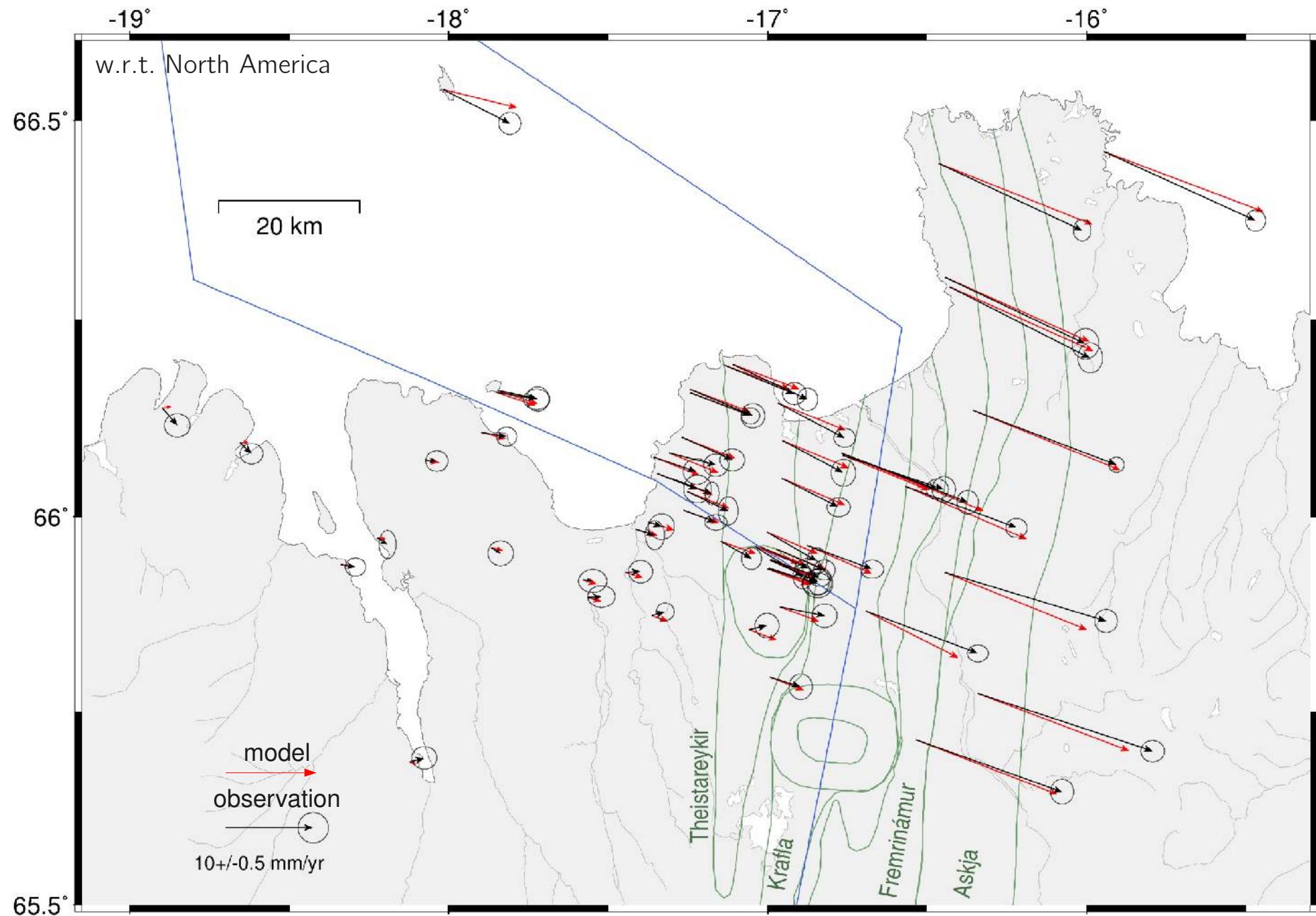


Metzger et al. 2011

- Fixed geometry, no dip-slip etc. → only seven free parameters, e.g.,
 - **Locking depth**
 - Total plate motion + motion partition between HFF/GOR → **Slip-rate**
- Find best-fit parameters and estimate parameter uncertainties



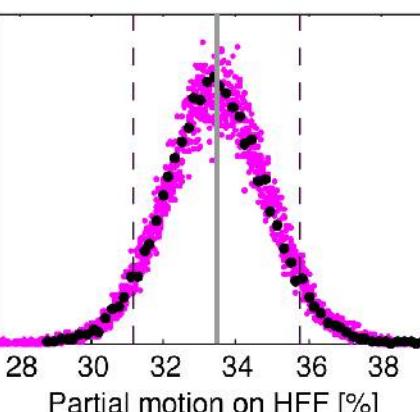
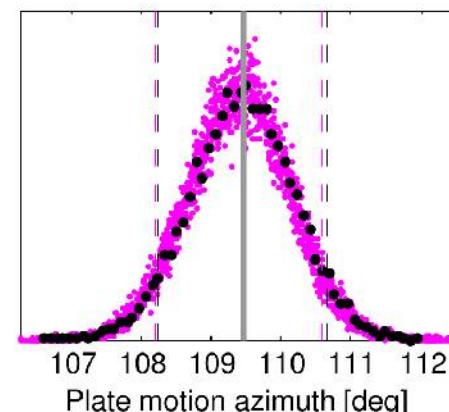
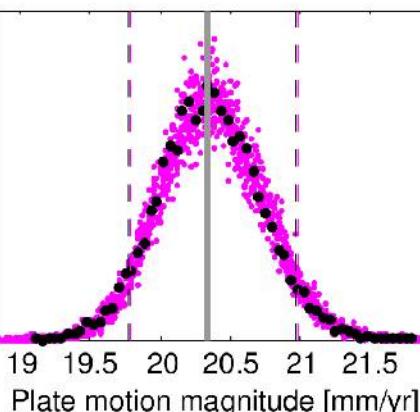
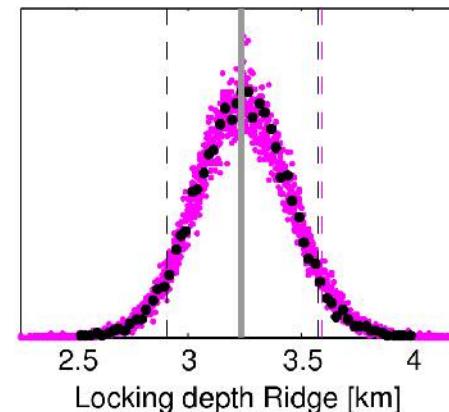
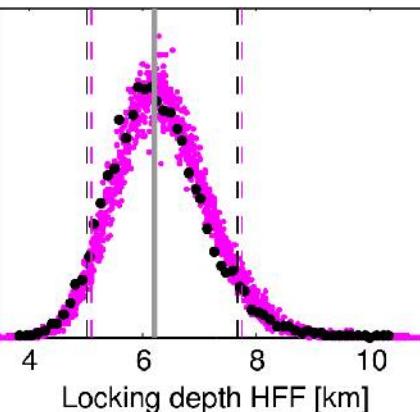
Data and best-fit model



Uncertainty estimation



- How well defined are best-fit model parameters?
- Compare two independent methods
 - 1) Error propagation
Perform statistics with “perturbed” input data
 - 2) Bayesian approach
estimate uncertainties from a posterior probability density function
- Both methods are in good agreement and support each other



- | | |
|-------|------------------------|
| • | Bayesian approach |
| • | Error propagation |
| - - - | 95% conf (Bayesian) |
| - - - | 95% conf (error prop.) |
| — | Best fit (error prop.) |

Model results in review



	This study	Metzger et al. (2011)	Rögnvaldsson et al., 1998
Data points	58 GPS points	14 GPS points	offshore seismicity
Locking depth HFF [km]	6.2 ^{+0.8} _{-0.7}	6.3 ^{+1.7} _{-1.2}	10 – 12
Partial motion HFF [%]	33.4 ± 1.4	33.9 ^{+3.1} _{-3.2}	

→ Locking depth varying along the fault?

- Best parameters confirmed and better constrained, **but:**
- The model uncertainties/assumptions now probably **outweigh** data uncertainties
- The accumulated seismic moment of the Húsavík-Flatey fault corresponds to a **M6.8 ± 0.1** event

Assumptions

- Complete stress-release in 1872
- Steady stress-accumulation and slip-rate
- Complete rupture along the whole fault

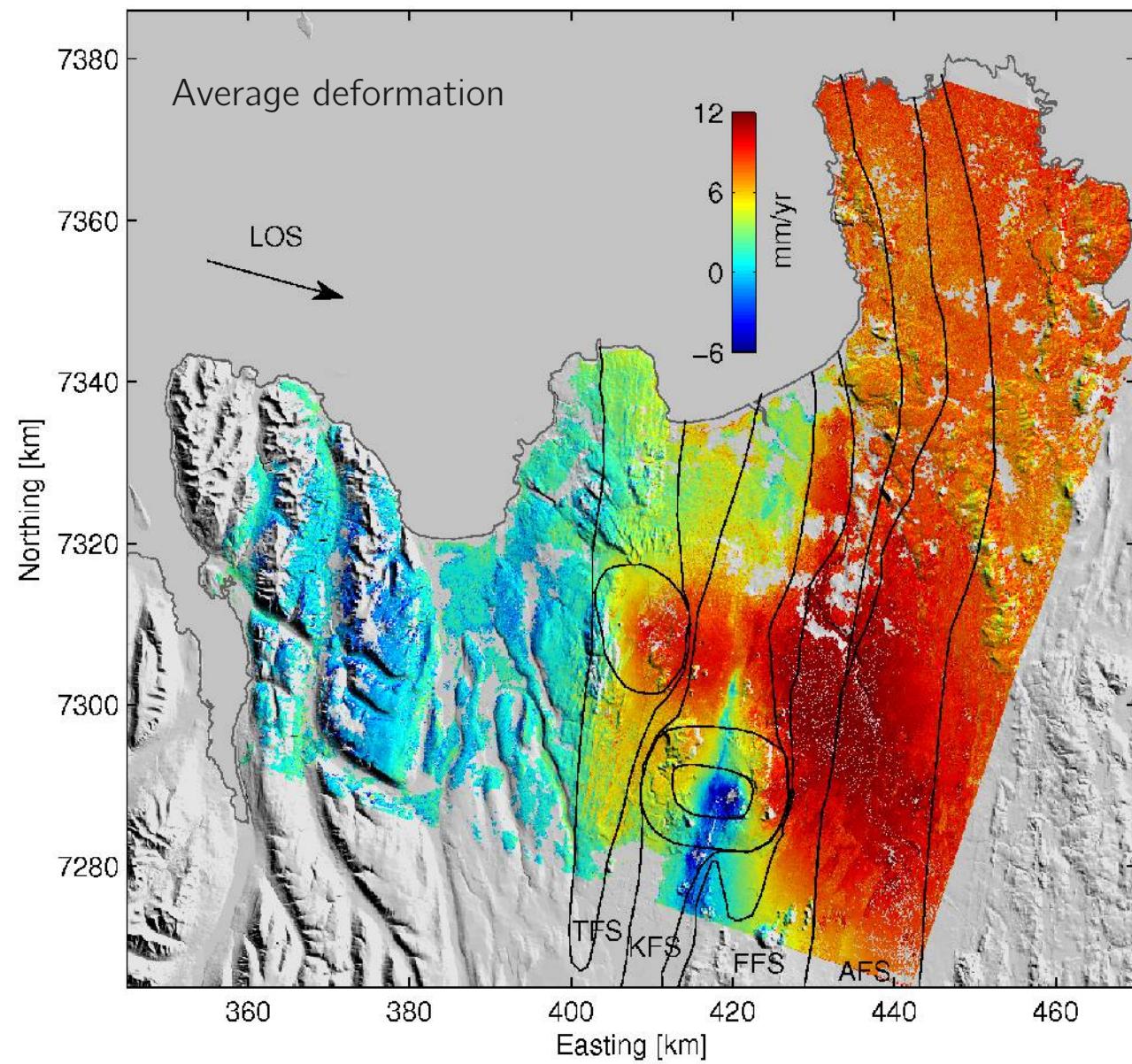
Conclusions



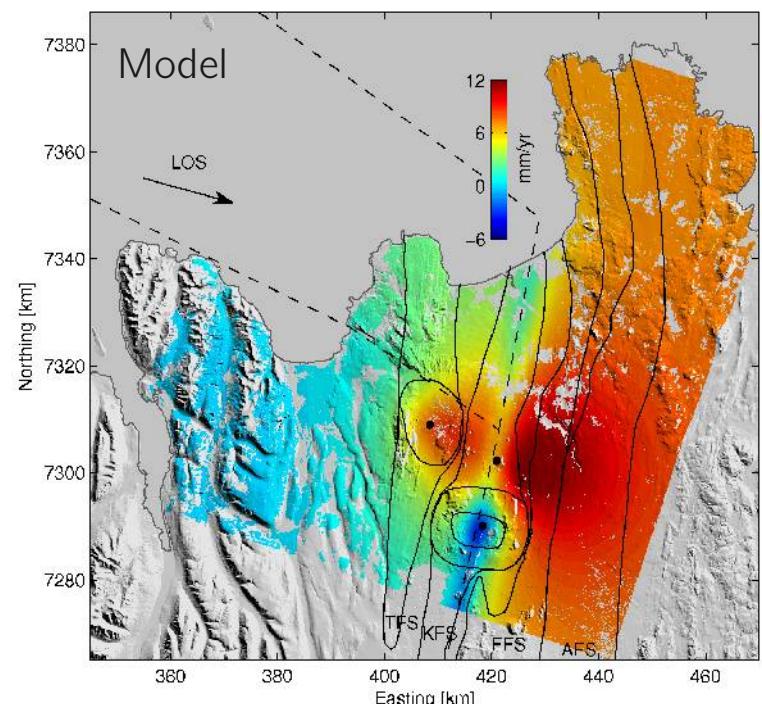
- We use GPS and InSAR time-series to describe the deformation of the North Icelandic plate boundary with a **kinematic model**
- Assuming complete stress-release in 1872 and steady stress accumulation since then, **the seismic potential** of the Húsavík-Flatey fault is equal to a **M 6.8 ± 0.1** event
- Offshore Earthquake relocations suggest a **varying fault locking depth**
- Additional campaign GPS data confirm our earlier results and decreases the error boundaries of the best-fit model parameters significantly → **uncertainties caused by model assumptions probably now outweigh the data uncertainties**



InSAR time-series analysis

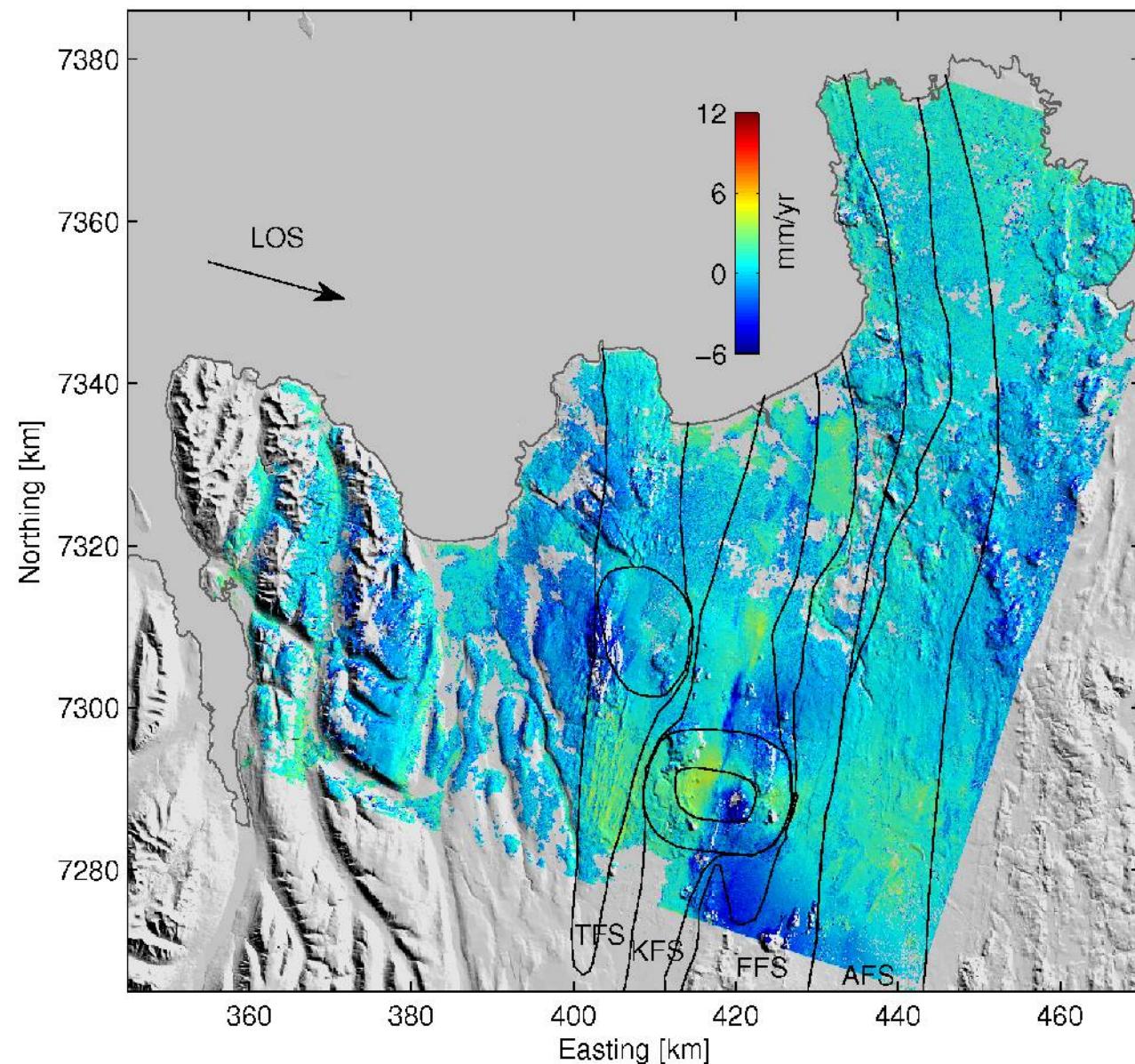


- Steady-state plate extension
- Remains of the Krafla rifting episode 1975-1984 (de Zeeuw-van Dalsen, 2004)
- Theistareykir inflation





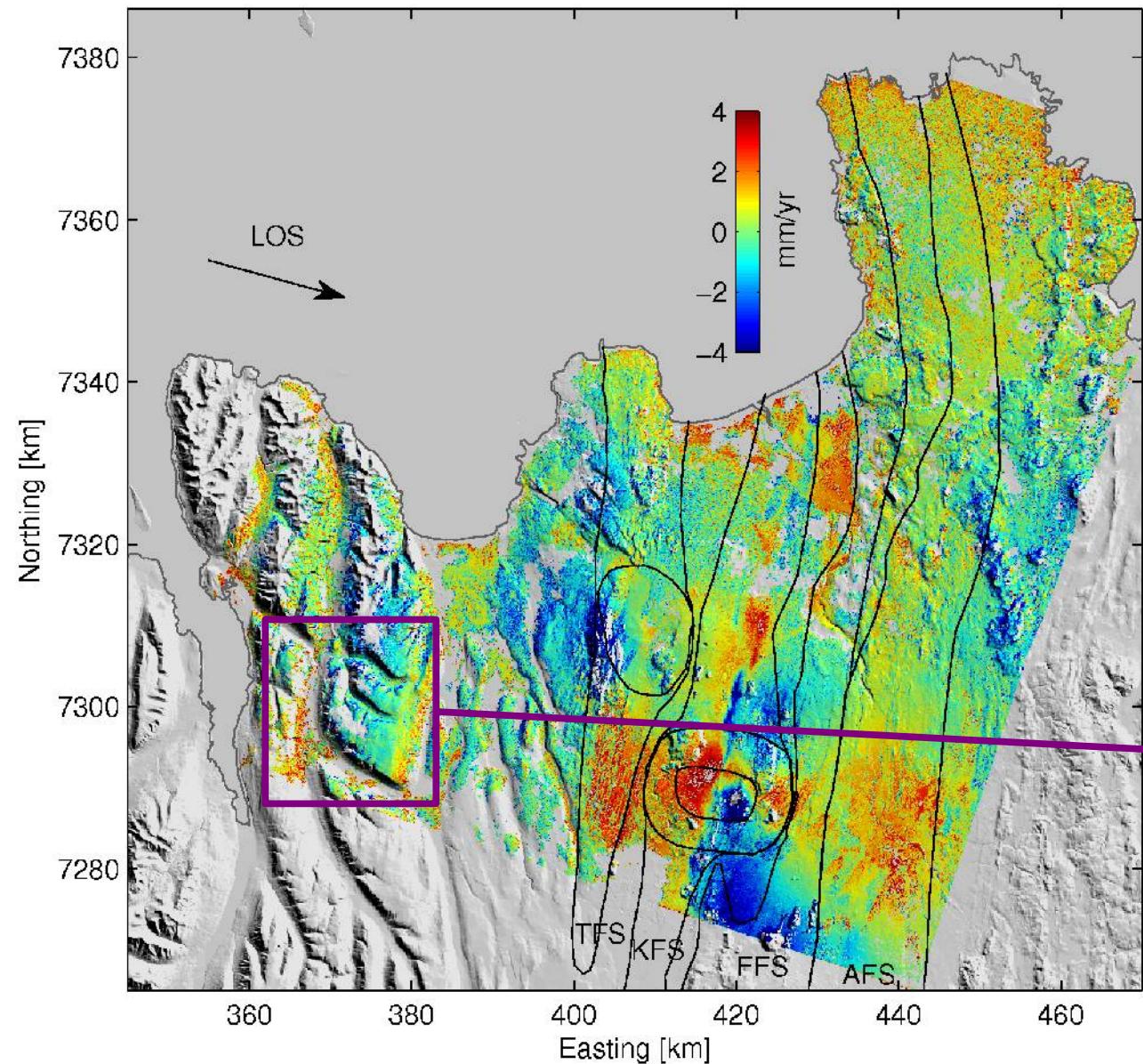
Residuals



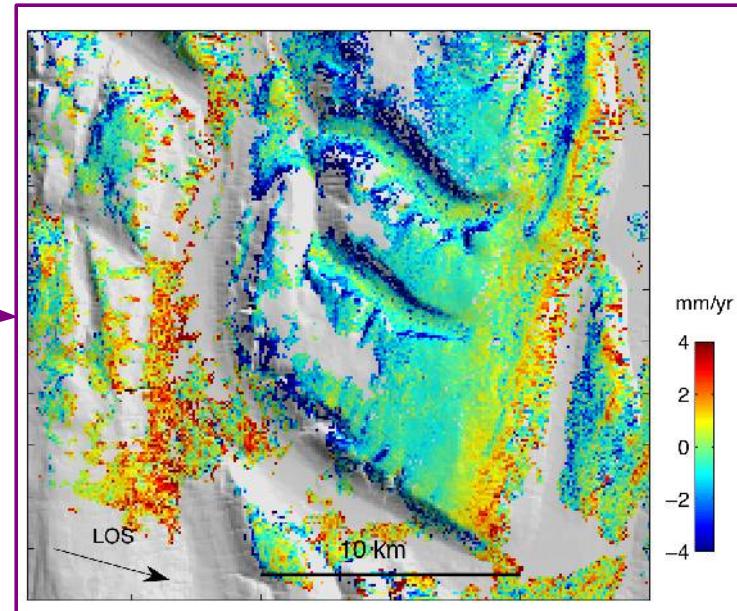
- Average RMS < 1.5 mm/yr



Residuals



- Average RMS < 1.5 mm/yr
- Residual tectonic deformation
- Non-tectonic residuals, e.g. gravity-driven slope motion?



Thank you!

