



STRUCTURAL, PETROPHYSICAL AND GEOMECHANICAL CHARACTERIZATION OF THE BECANCOUR CO₂ STORAGE PILOT SITE (QUEBEC, CANADA)

SÉQUESTRATION GÉOLOGIQUE DU CO₂

CHAIRE DE RECHERCHE

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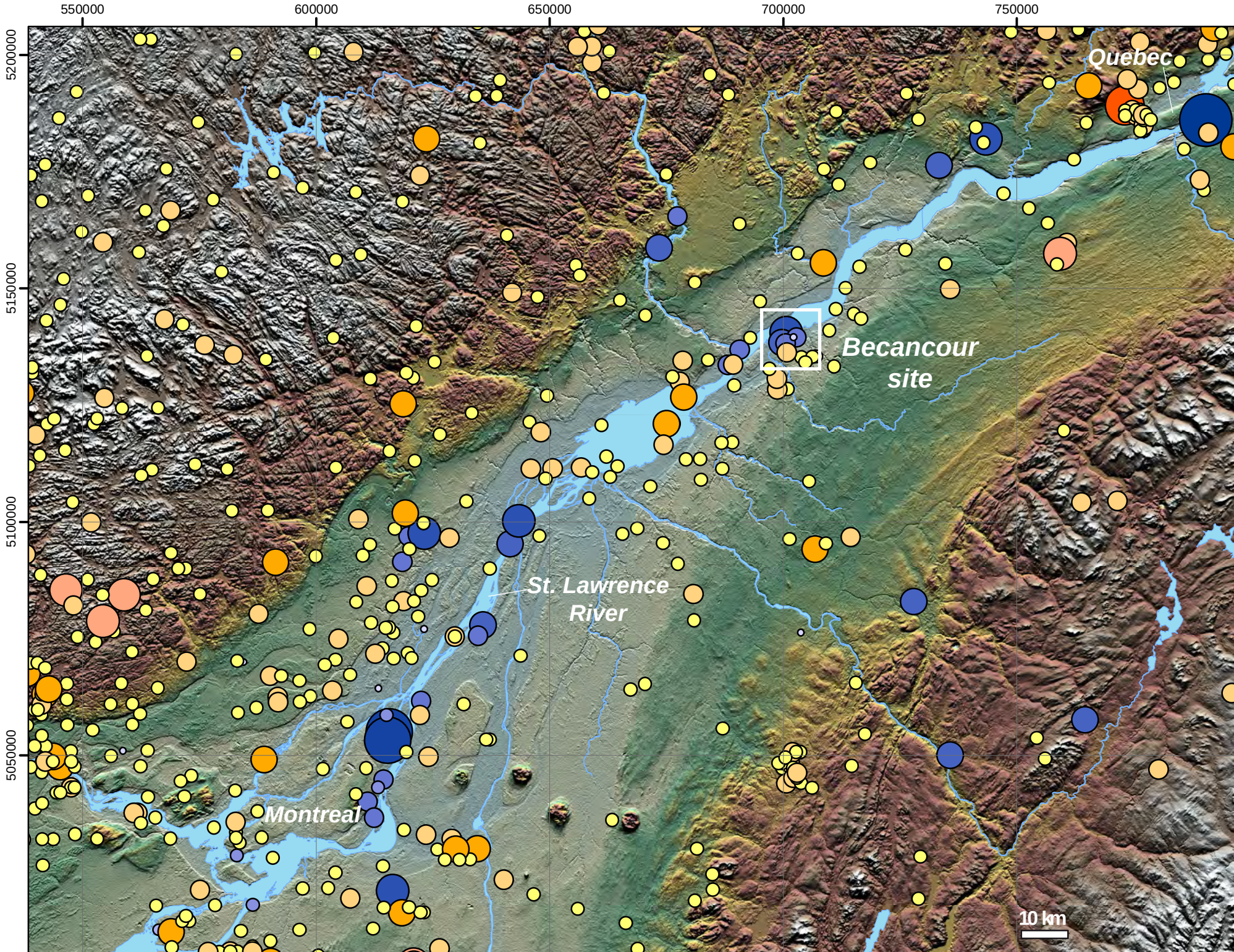


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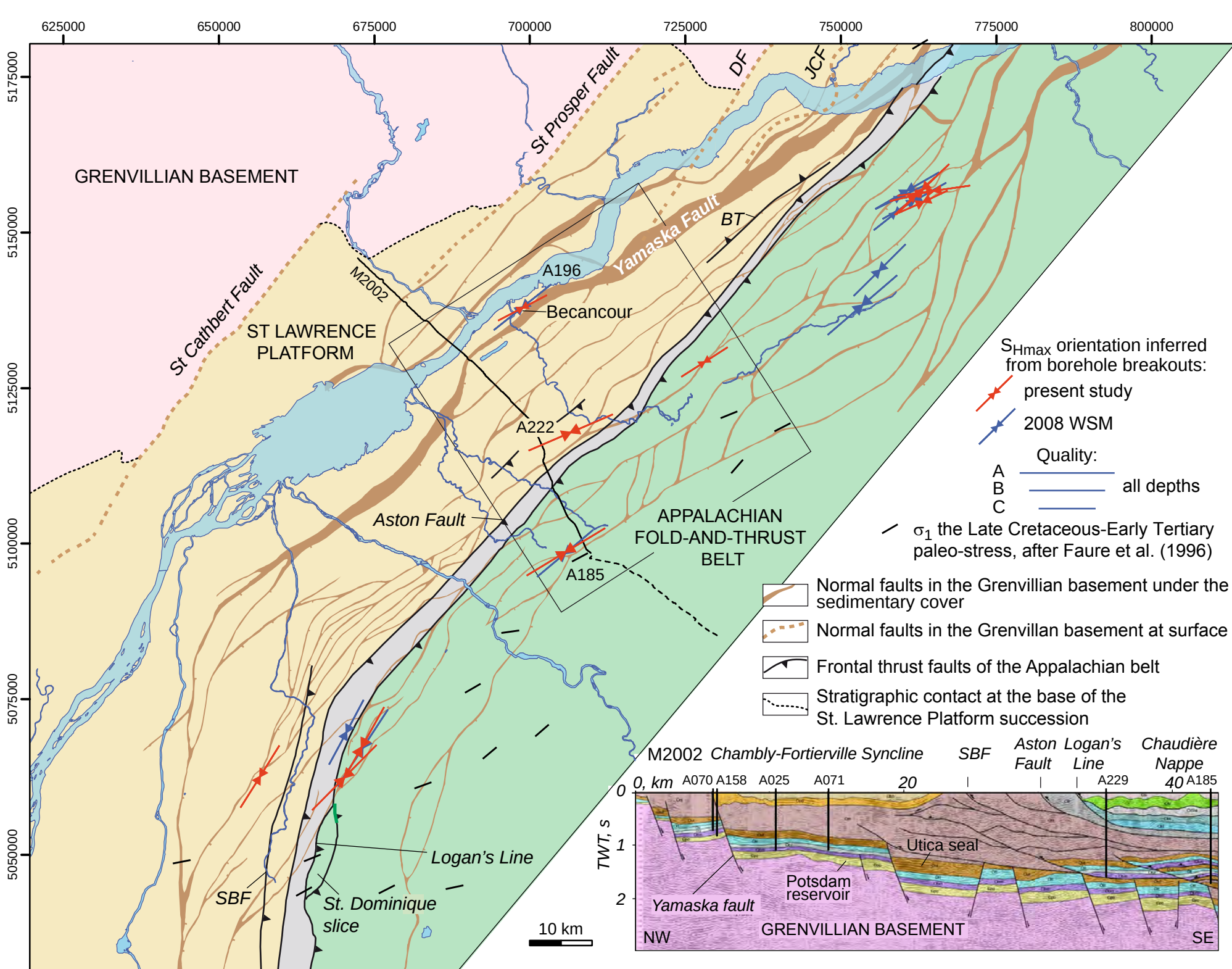
Développement durable,
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Seismicity and CO₂ emissions in the St. Lawrence Lowlands



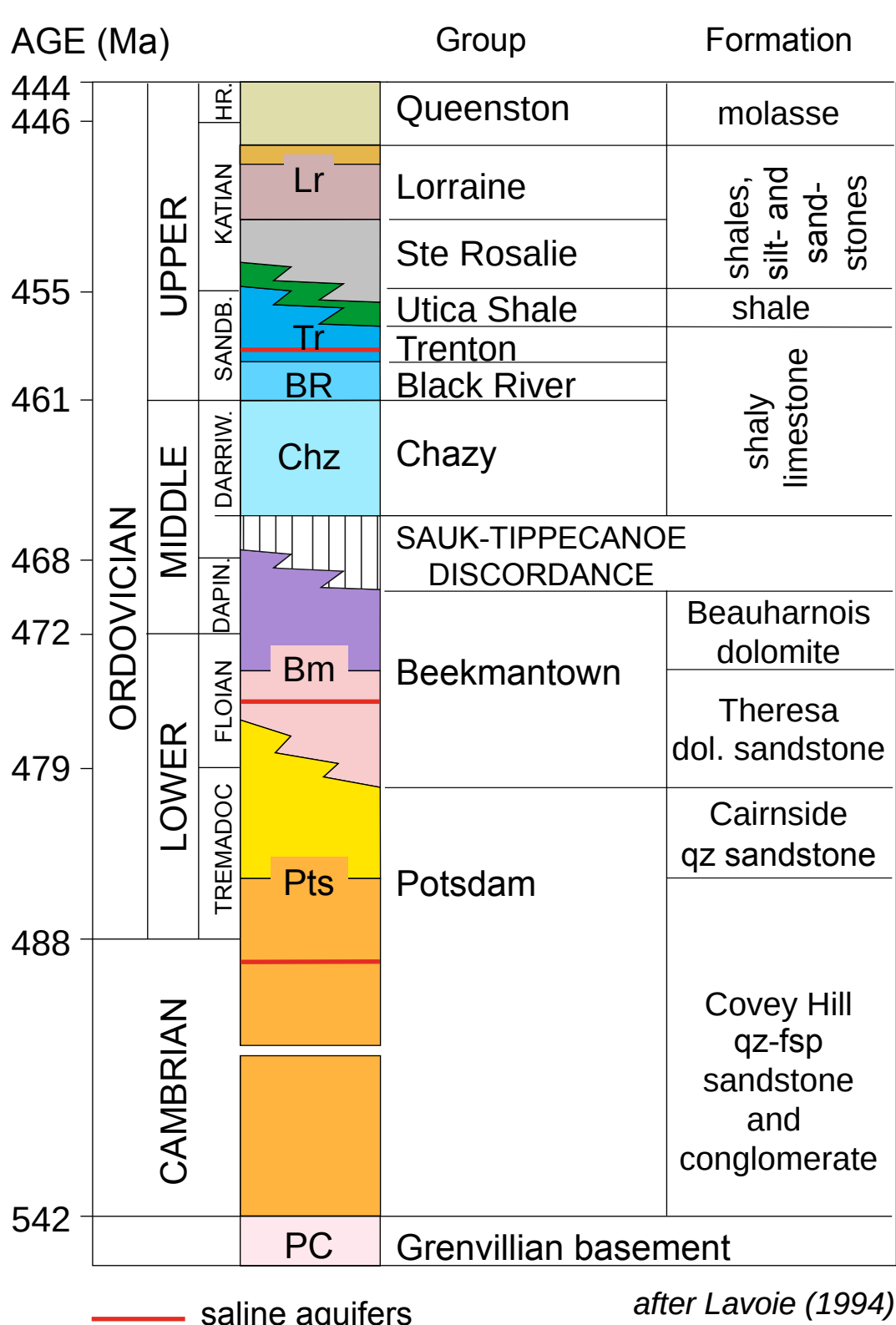
The Becancour site is located in the area of high CO₂ emissions (data after Canada Environment). The southern shore of the St. Lawrence River is characterized by relatively low seismicity (data after Geological Survey of Canada Catalog).

Regional fault patterns and the S_{Hmax} orientations in the St. Lawrence Lowlands



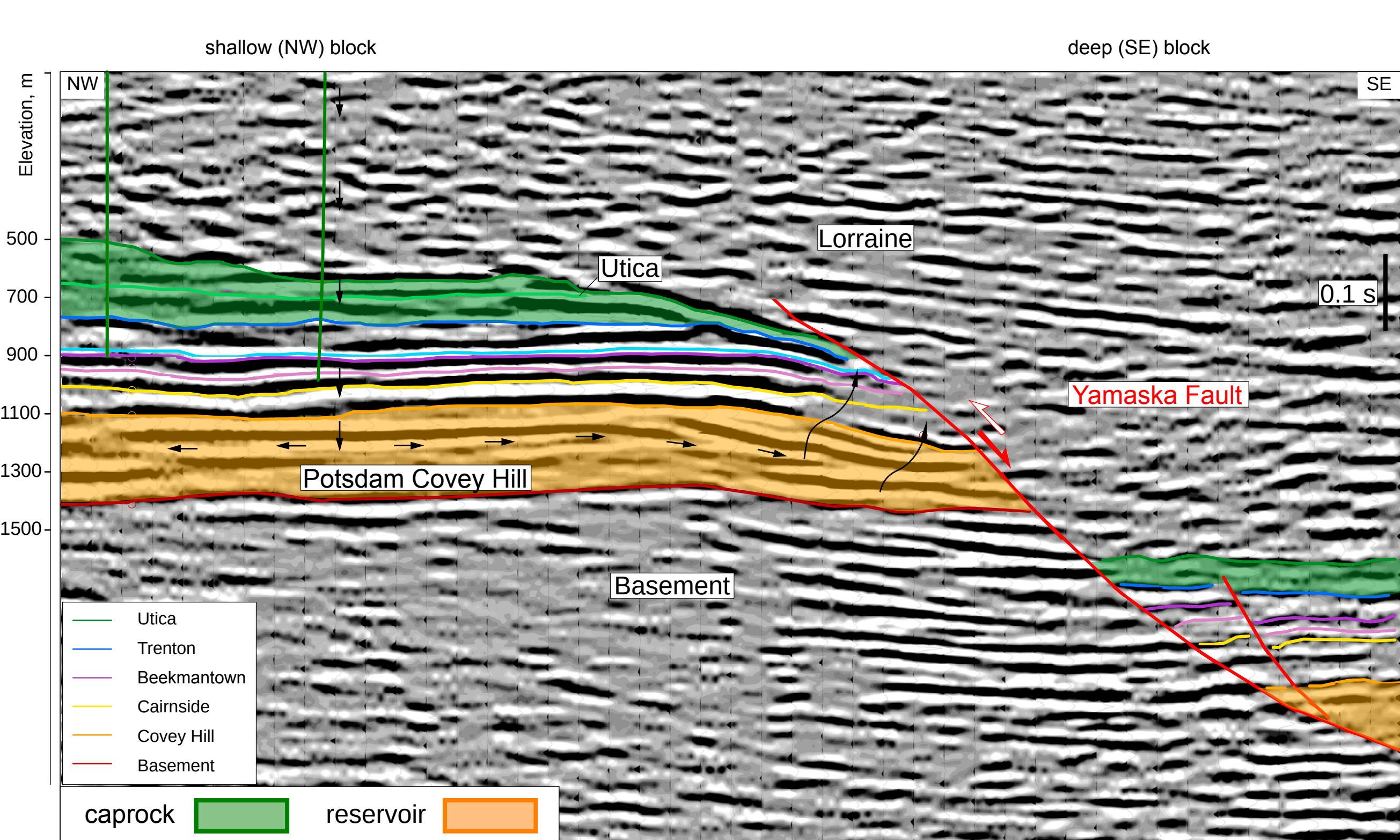
The normal faults in the Grenville basement are projected on the seismic map of Theriault et al. (2005). The thrust faults are shown after Clark (1984a, 1984b), Golevsky (1987), Konstantinovskaya et al. (2009). The interpreted seismic line M2002 is after Castonguay et al. (2006). Abbreviations: DF, Deschambault Fault; JCF, Jacques Cartier Fault; NF, Neville Fault; SBF, Saint-Barnabé Fault; BT, backscarp of the triangle zone. Box indicates Becancour - Notre Dame area.

Sedimentary succession of the St. Lawrence Platform



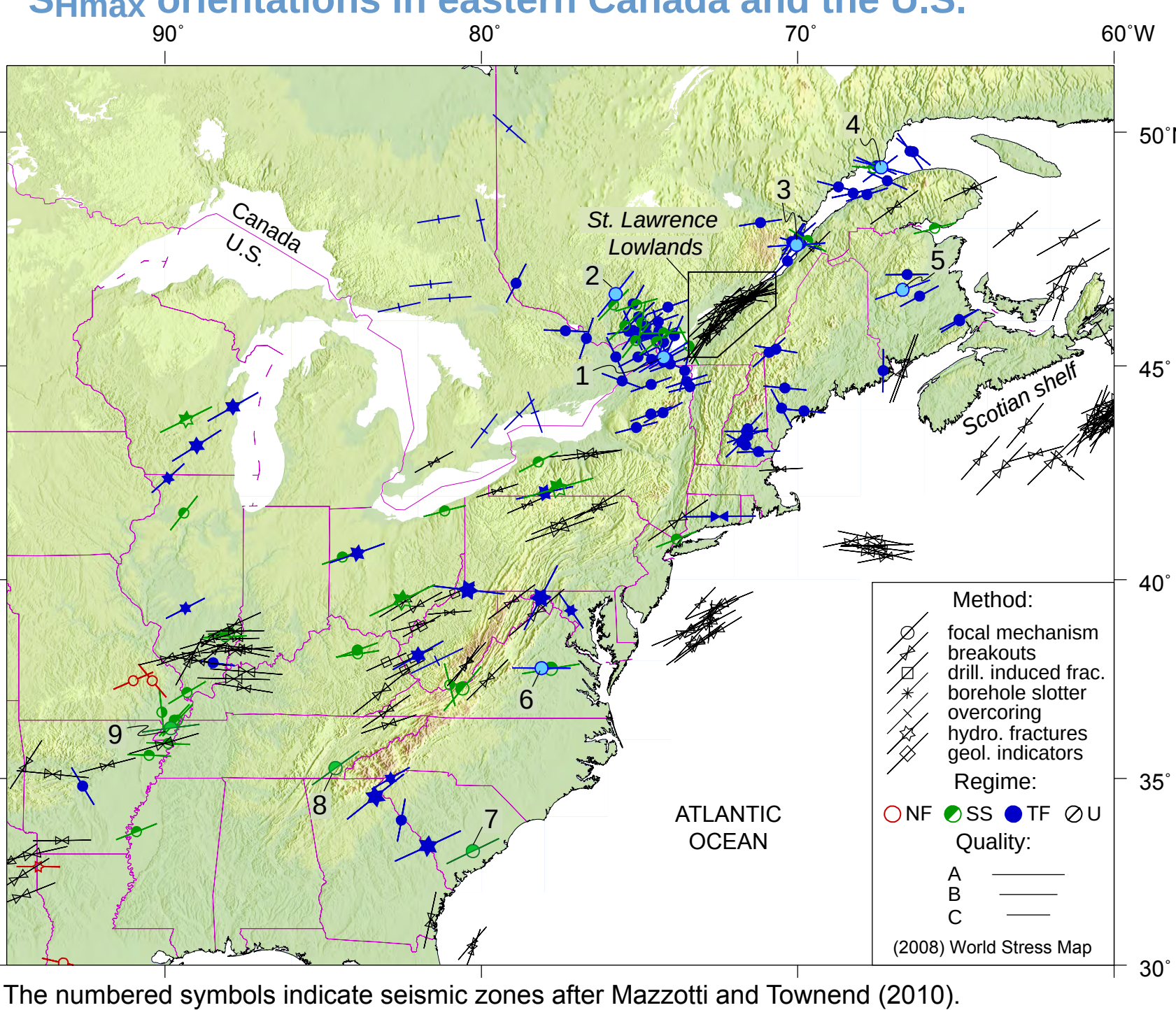
The three potential storage units of Potsdam, Beekmantown and Trenton saline aquifers are known in the St. Lawrence Platform succession of the Becancour area. The Potsdam reservoir rocks are overlain by a multiple caprock system of Utica shales and Lorraine siltstones and sandstones.

The Yamaska normal fault in the Becancour area

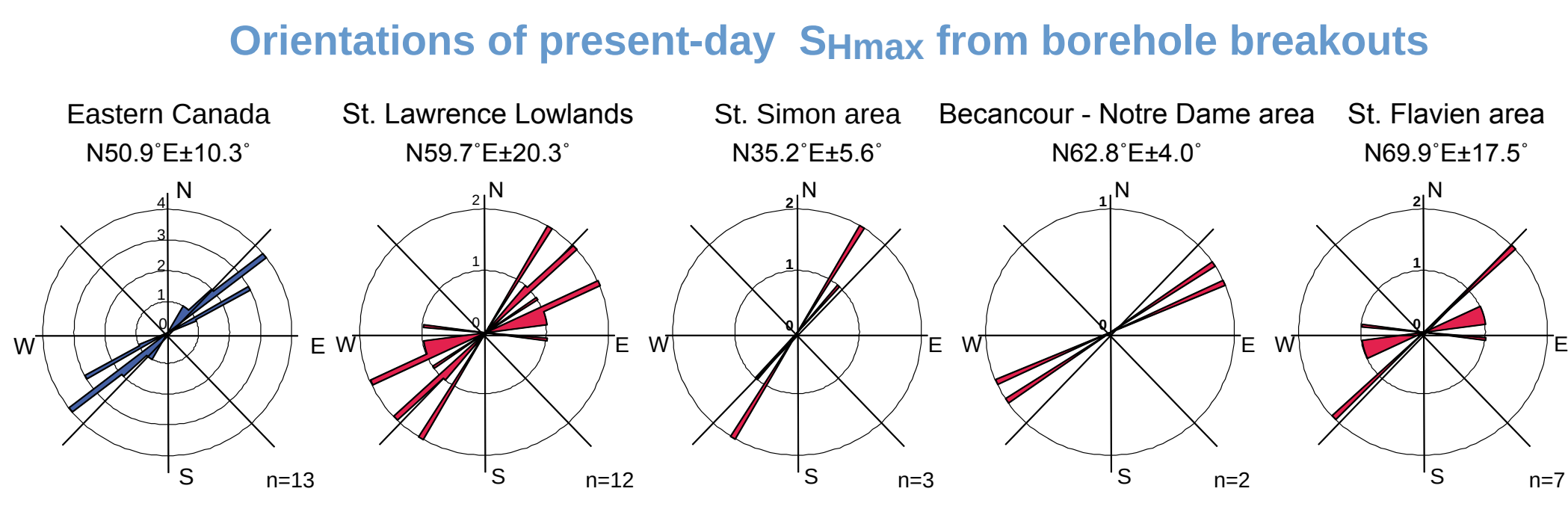


The NE-SW syn-sedimentary rift-related Yamaska normal fault is dipping to SE at about 55°. This fault affects the subhorizontal sedimentary succession of the St. Lawrence Platform and the Grenville basement with 870 m of vertical separation. It was reactivated as reverse fault at the end of Taconian orogeny or during the later compressive phases. The seismic line is about 2 km of length. The ratio of vertical to horizontal scale is 1:2.

Present-day maximum horizontal stress S_{Hmax} orientations in eastern Canada and the U.S.

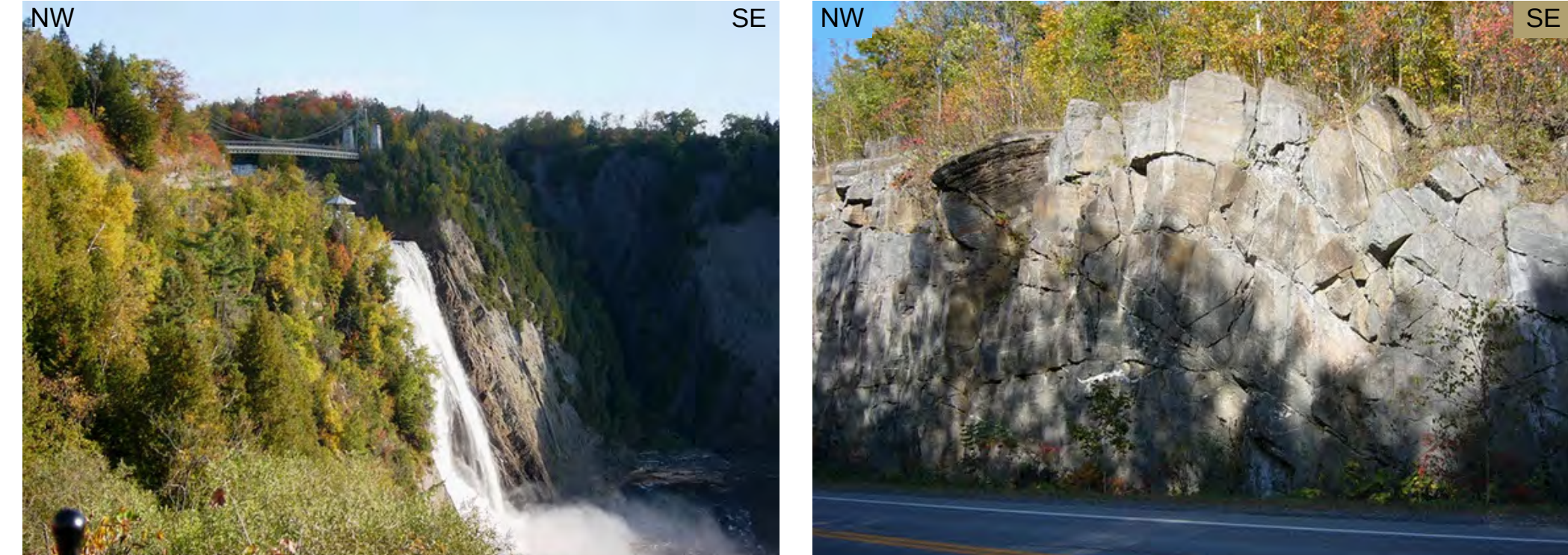


The numbered symbols indicate seismic zones after Mazzotti and Townsend (2010).



after Plumb and Cox (1987) after Konstantinovskaya et al. (2012)
The average S_{Hmax} orientation is NE-SW in Eastern Canada (N51°E). The S_{Hmax} is oriented N60°E in the St. Lawrence Lowlands rotating along the Appalachian front. The S_{Hmax} orientation is N63°E in the Becancour-Notre Dame area.

NE-SW rifting related normal faults and NE-SW reverse faults and folds of the Taconic orogeny



The Montmorency normal fault in Quebec City represents the surface analog of high-angle SE-dipping normal faults affecting the Grenville basement and the St. Lawrence Platform.

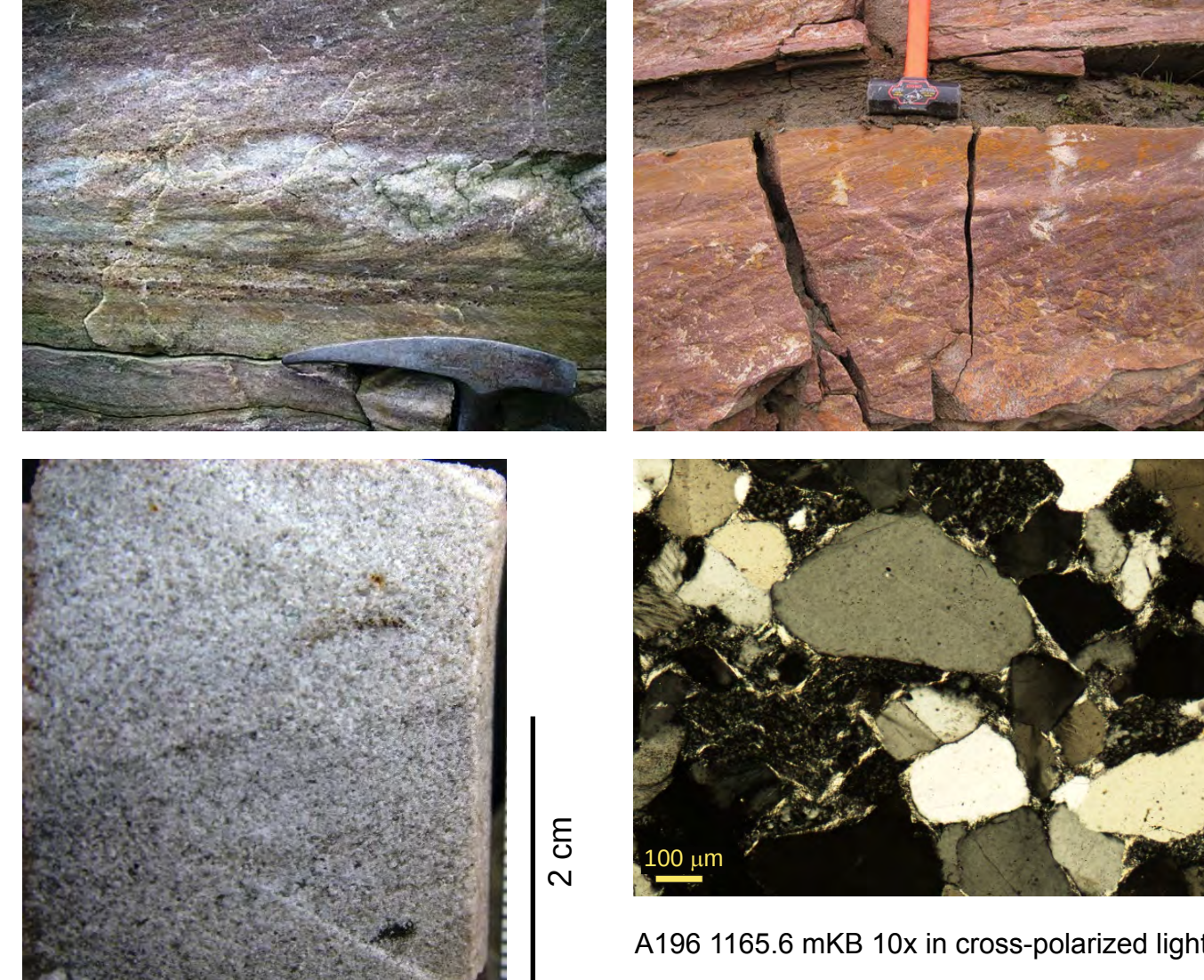
The NW-vergent thrust of the Saint-Dominique tectonic slice represent an example of low-angle normal reverse faults affecting the Appalachian fold-and-thrust belt.

Reservoir and caprock lithology



The Utica shales have low permeability $K < 10^{-3}$ mD.

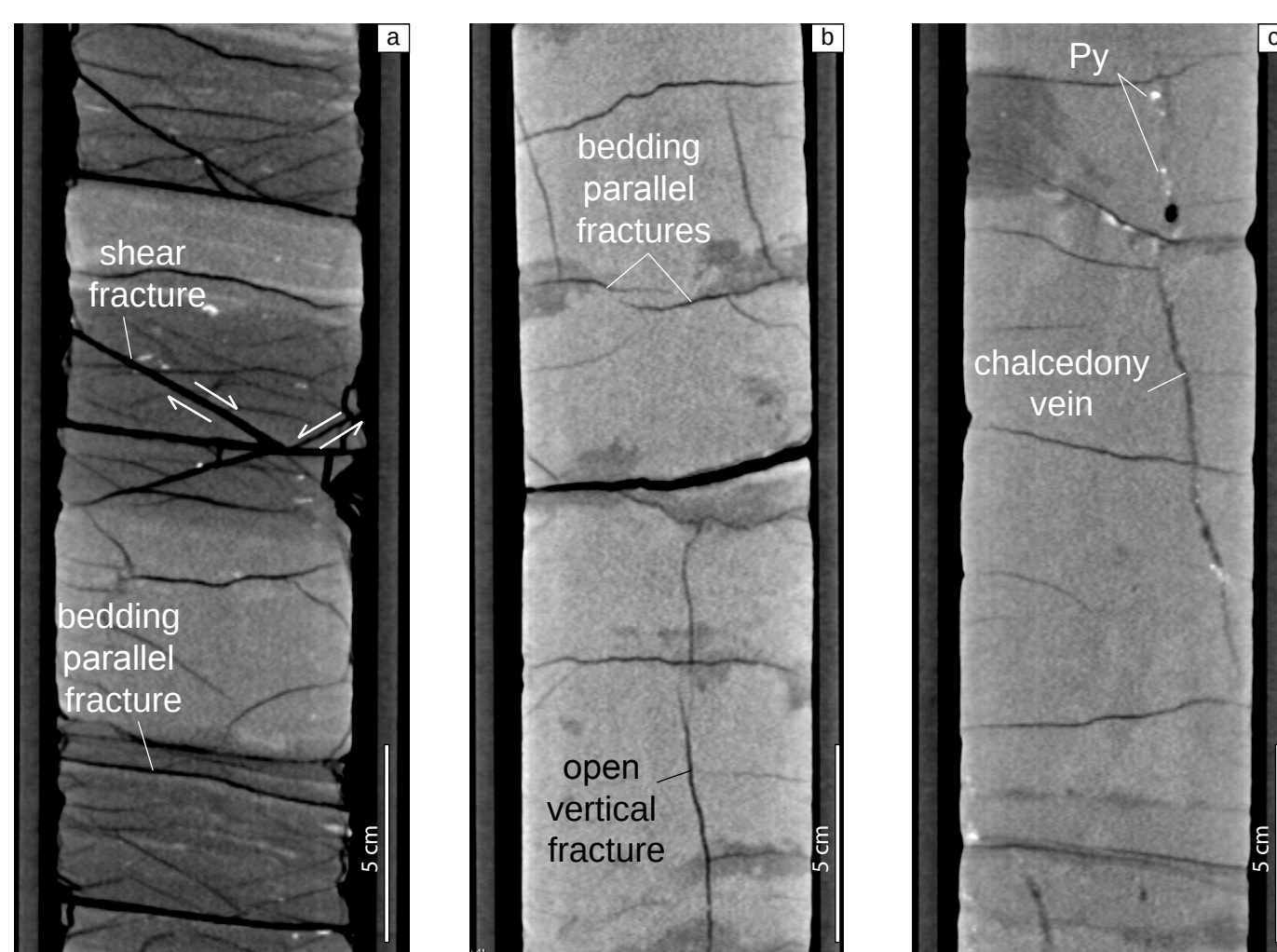
Petrography of the Covey Hill reservoir sandstones



The Covey Hill sandstones have the lowest static Young modulus and compressive strength relative to other sandstone units. The calcareous Utica shale of the regional seal is more brittle and has higher Young modulus and lower Poisson's ratio than the overlying Lorraine siltstones.

The Paleozoic sedimentary succession of the St. Lawrence Platform was characterized to estimate the CO₂ storage capacity, the caprock integrity and the fracture/fault stability at the Becancour pilot site. Results are based on the structural interpretation of 25 seismic lines and analysis of 11 well logs and petrophysical data.

Natural fractures observed in cores of the Utica shales (a) and the Trenton limestones (b, c) in the Becancour area

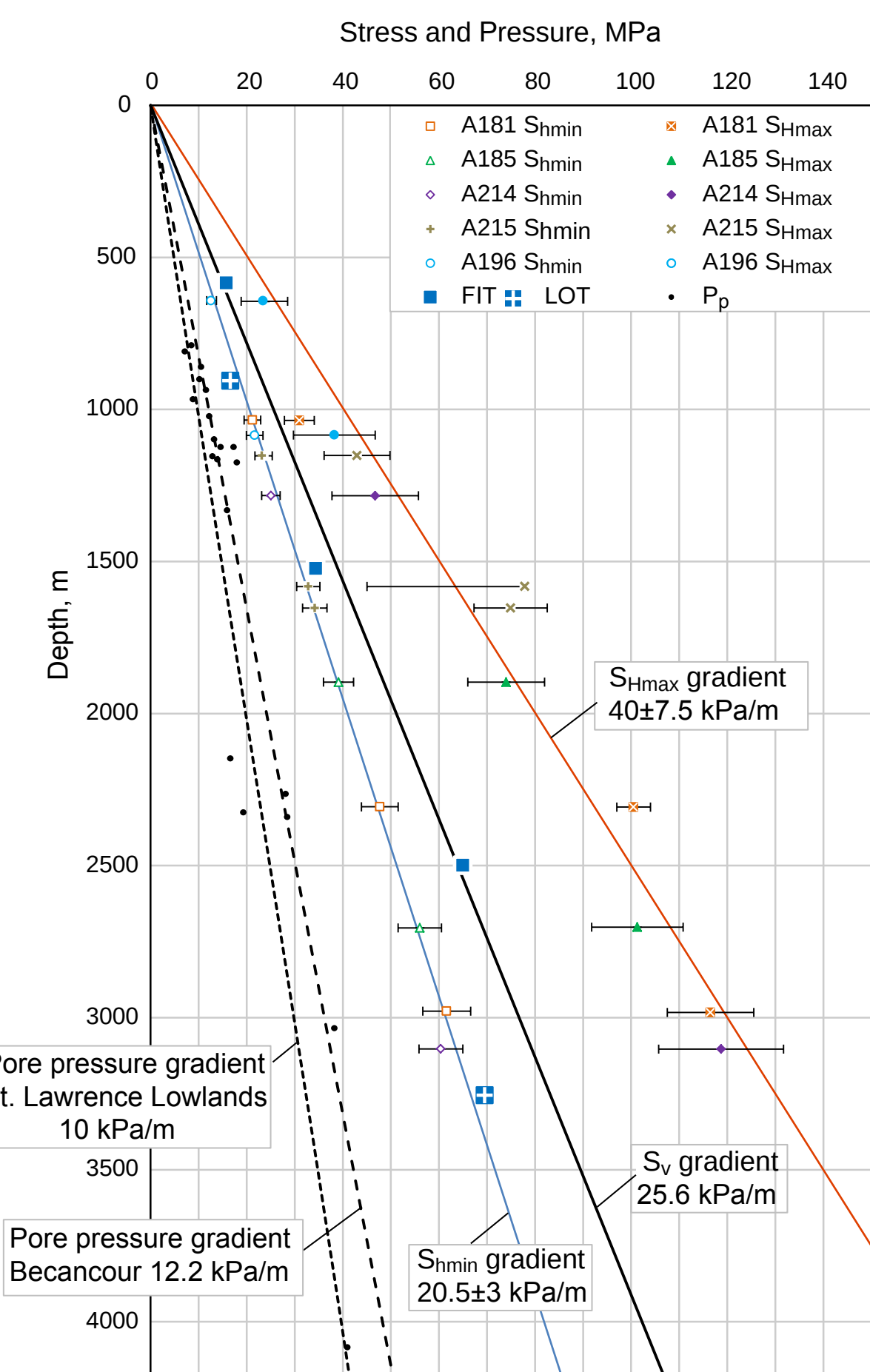


CT (computed tomography) scanner images of the longitudinal (vertical) sections of cores. Stratification is subhorizontal. Shale beds are dark gray, limestones are light gray, shaly limestones are medium gray, open fractures (filled with air) are black, the chalcodony vein is dark gray to black and sulfide mineralization (Py) is bright white spots in the vein. The bedding-parallel and inclined shear fractures are observed in Utica shales.

ACKNOWLEDGEMENTS

The present study is supported by the Ministère du Développement Durable, de l'Environnement et des Parcs du Québec. We are grateful to the Reservoir Development Services group within Baker Hughes Inc. for supporting the stress analysis. Special thanks are due to F. Brunner, L. Dean, R. Mireault and FKEETE Associates Inc. (Calgary) for the pore pressure evaluations from DST analyses, L. Massé and JUNEX for providing regional seismic and well log data, O. Heidbach and N. Pinet for technical advices and discussions. IHS-Seismic Micro-Technology kindly provided the seismic interpretation software used in this study.

Variations of pore pressures and stresses with depth determined from the stress magnitude analysis in the St. Lawrence Lowlands

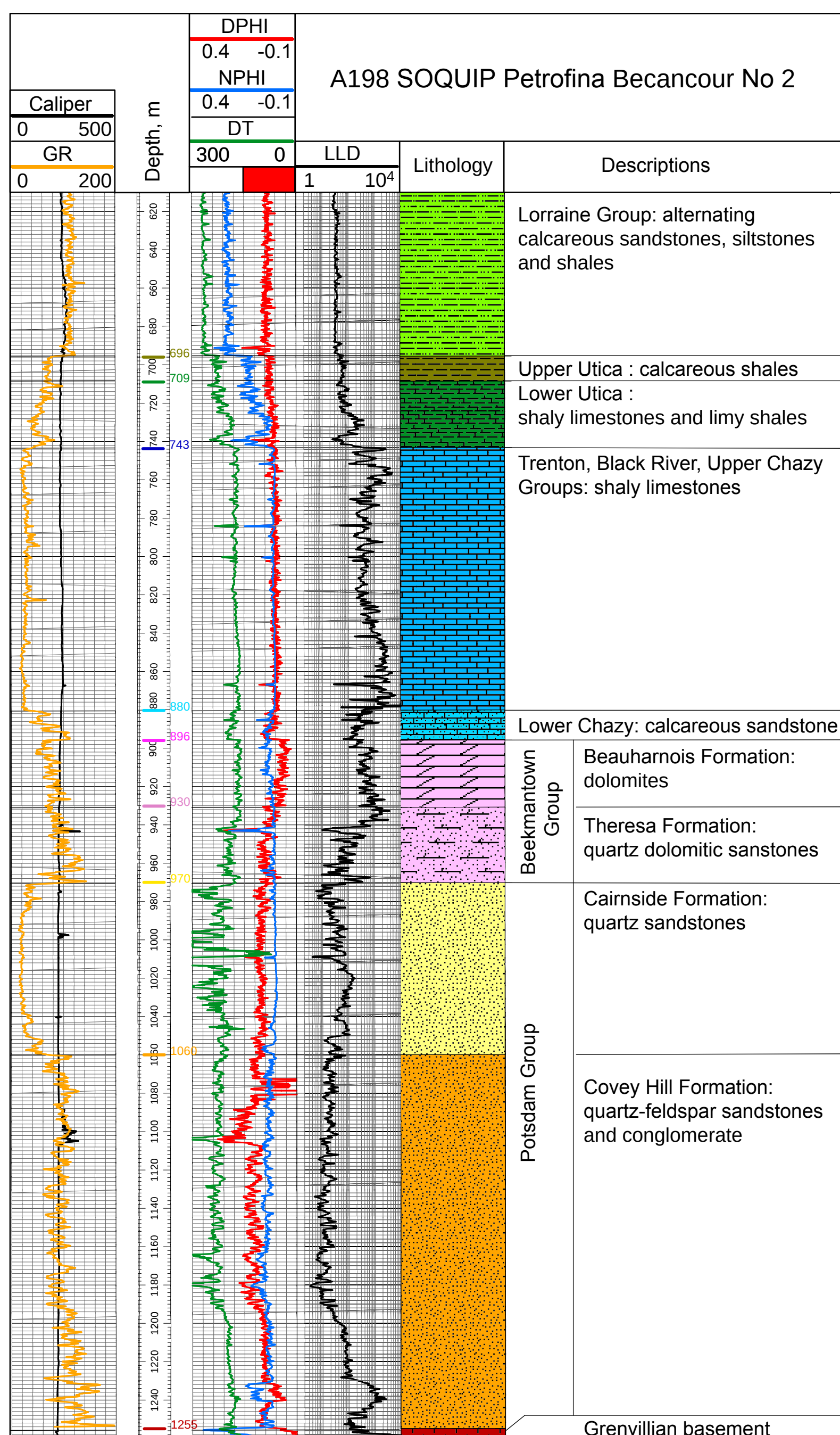


A geomechanical analysis of the St. Lawrence Platform sedimentary basin provides the maximum sustainable fluid pressures for CO₂ injection that will not induce tensile fracturing and shear reactivation along pre-existing fractures and faults in the caprock.

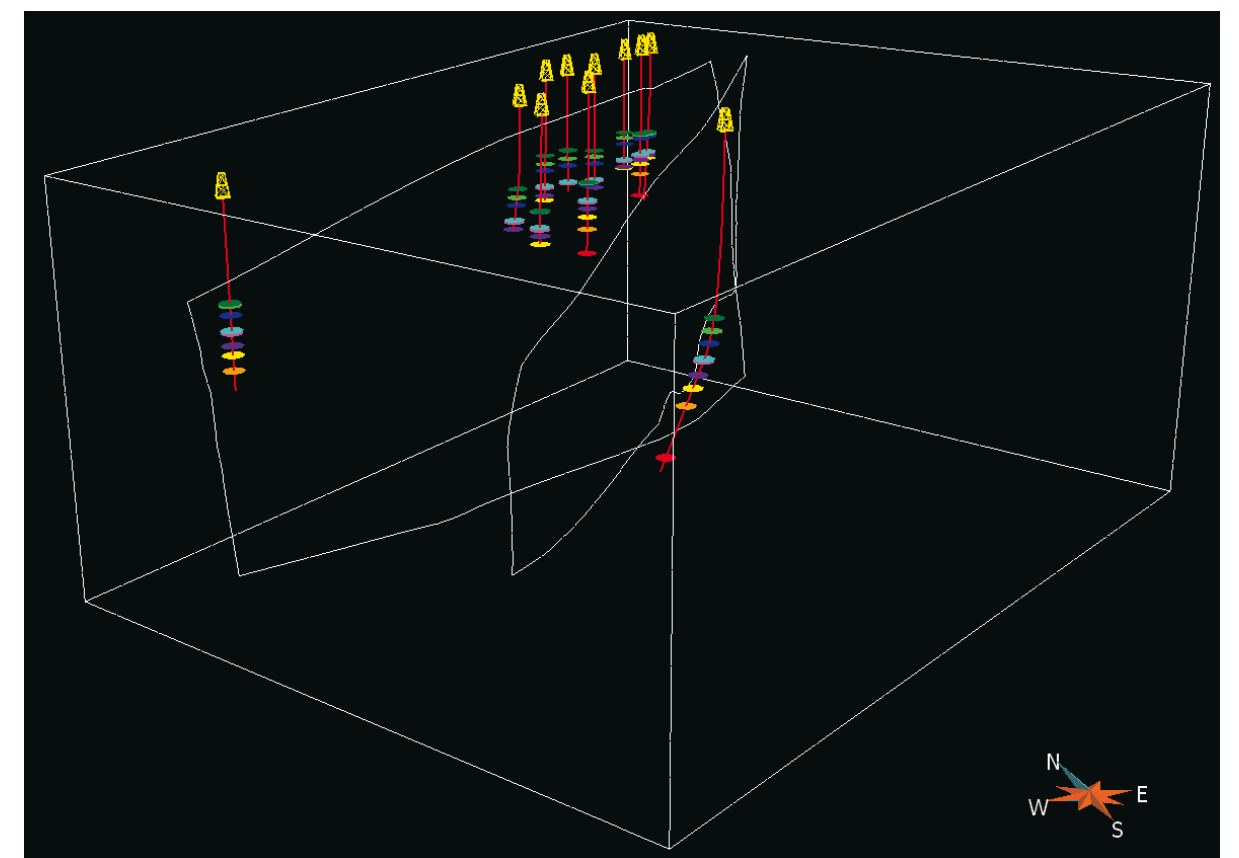
The regional stresses/pressure gradients estimated for the Paleozoic sedimentary basin (depths < 4 km) indicate a strike-slip stress regime.

The Covey Hill sandstones have the lowest static Young modulus and compressive strength relative to other sandstone units. The calcareous Utica shale of the regional seal is more brittle and has higher Young modulus and lower Poisson's ratio than the overlying Lorraine siltstones.

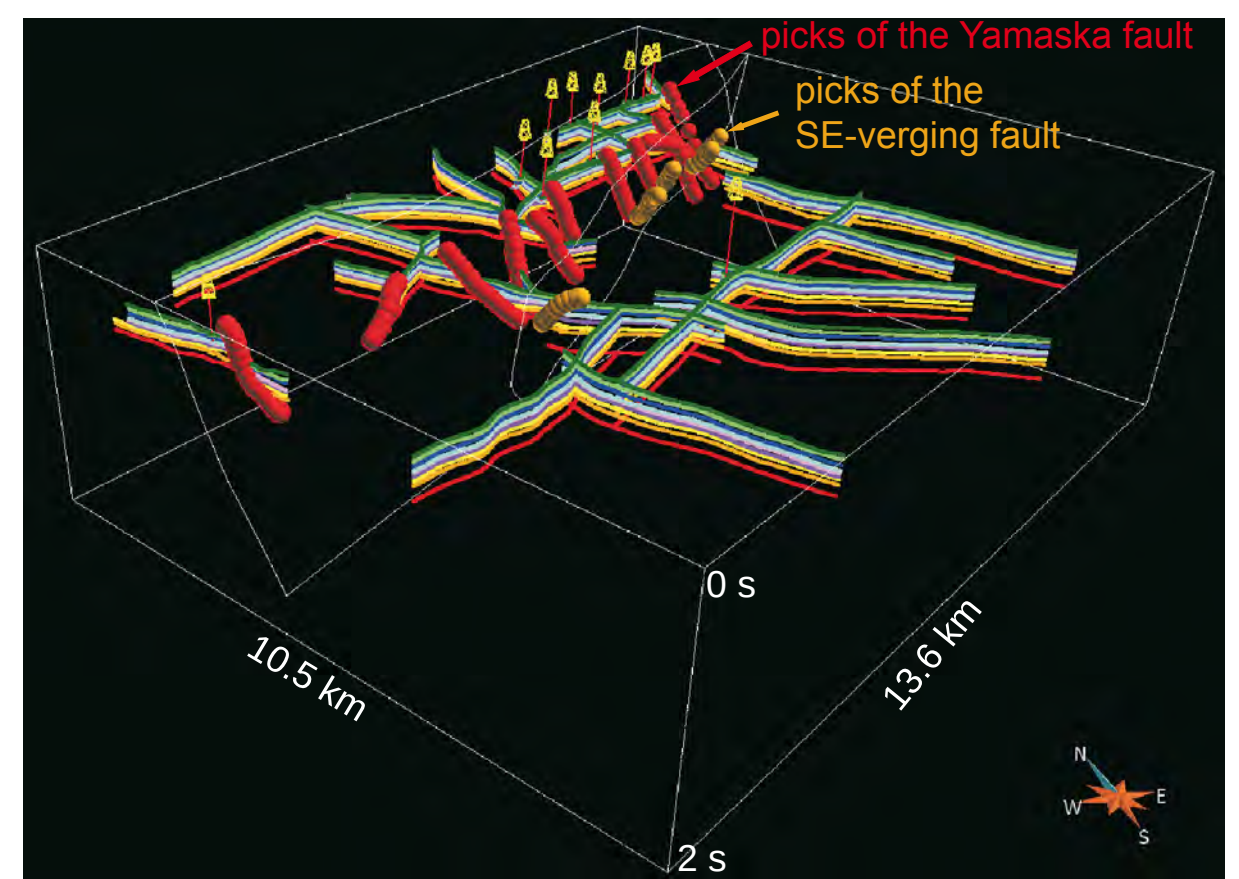
Petrophysical characteristics and formation tops of the stratigraphic units in the St. Lawrence Platform



1- Identification of formation tops in depth (m)



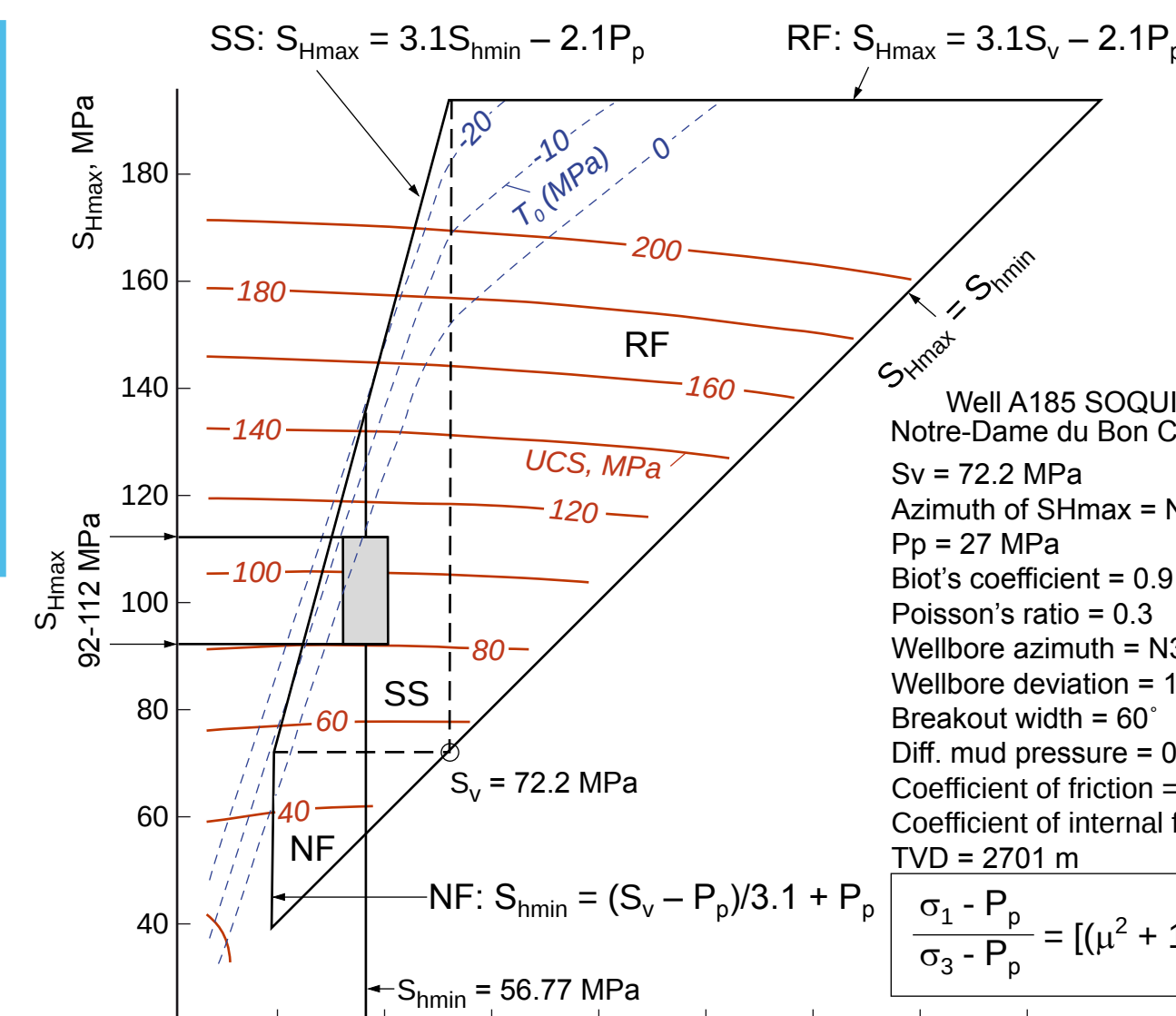
2- Picking of seismic horizons in TWT (s)



CONCLUSIONS

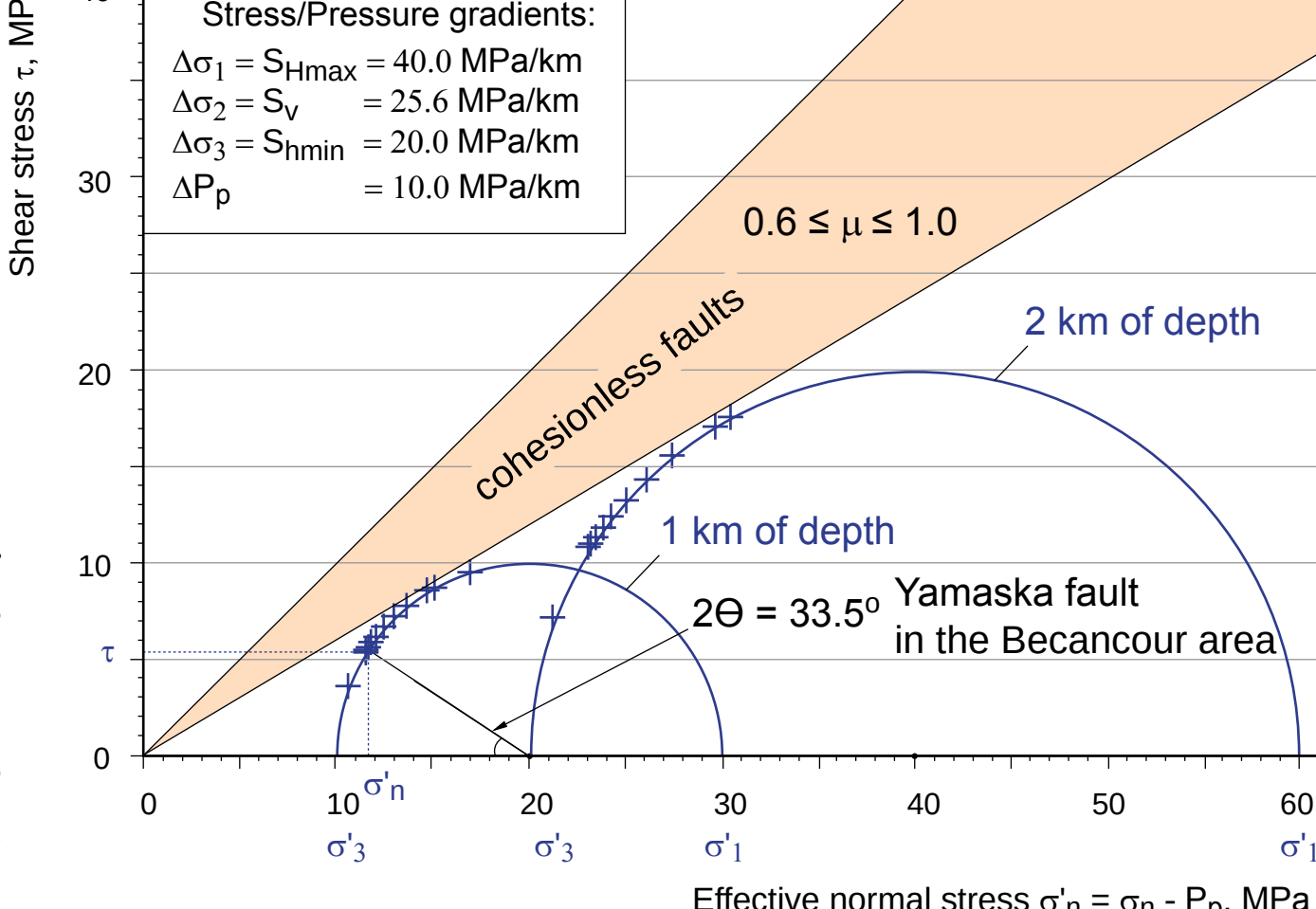
1. The Cambrian-Lower Ordovician Covey Hill reservoir sandstones represent the only unit with significant CO₂ sequestration potential in the Becancour area. These reservoir rocks are characterized by the highest porosity (6%), net pay thickness (188 m) and located at depths (>1026 m) necessary for the supercritical state of CO₂. The relatively low matrix permeability (0.28 mD) may be enhanced by natural or hydraulic fracturing.
2. The regional caprock units of the Upper Ordovician Utica shales and the Lorraine sandstones-siltstones are laterally extended with minimum thickness of 700-800 m. The extremely tight Trenton limestones and Beauharnois dolomites may provide additional aquitard levels above the reservoir rocks.

Polygon of S_{Hmax} estimation constrained by wellbore failure



Variations of pore pressures and stresses with depth determined from the stress magnitude analysis in the St. Lawrence Lowlands are compatible with the strike-slip stress regime $S_{Hmin} < S_v < S_{Hmax}$.

Shear and effective normal stresses acting on the NE-SW high-angle normal faults in the Grenvillian basement of the St. Lawrence Lowlands



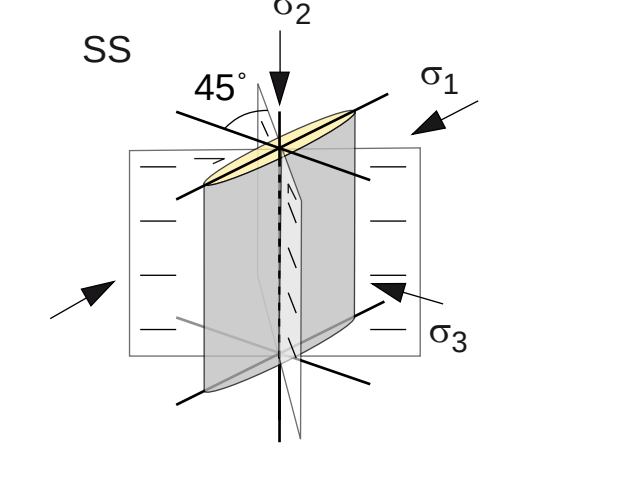
The high-angle NE-SW Yamaska normal fault is oriented at 16.7° to the S_{Hmax} orientation in the Becancour site. The slip tendency along the fault in this area is estimated to be 0.47 based on the stress magnitude and rock strength evaluations for the borehole breakout intervals in local wells. The regional pore pressure-stress coupling ratio under assumed parameters is about 0.5-0.65 and may contribute to reduce the risk of shear reactivation of faults and fractures. The maximum sustainable fluid pressure that would not cause opening of vertical tensile fractures during CO₂ operations is about 18.5-20 MPa at a depth of 1 km.

Angle Θ between fault and σ₁ in σ₁-σ₃ plan

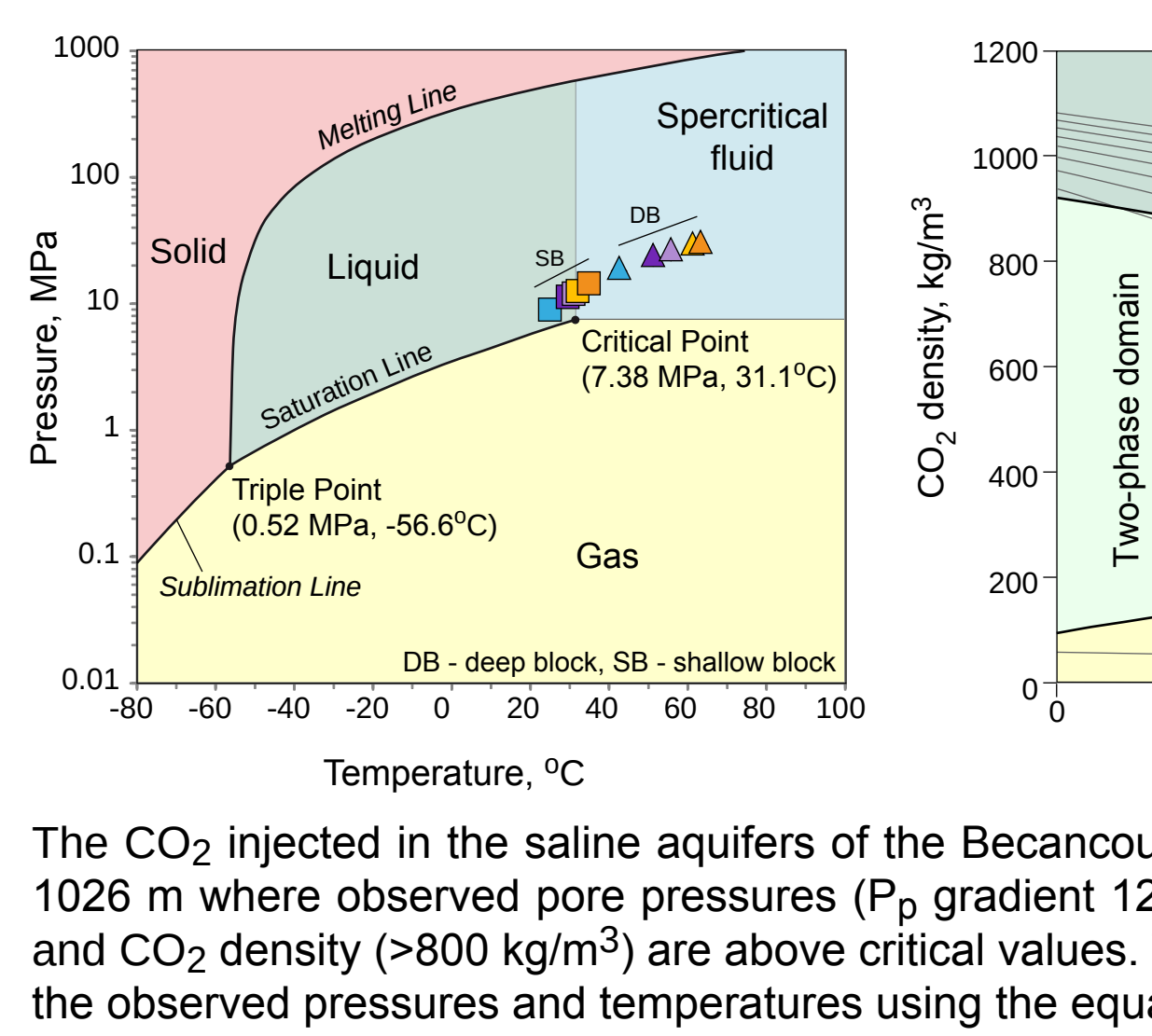
10.6° to 22°
22° to 36.6°

Slip tendency τ/σ₁n

0.30-0.47
0.56-0.58

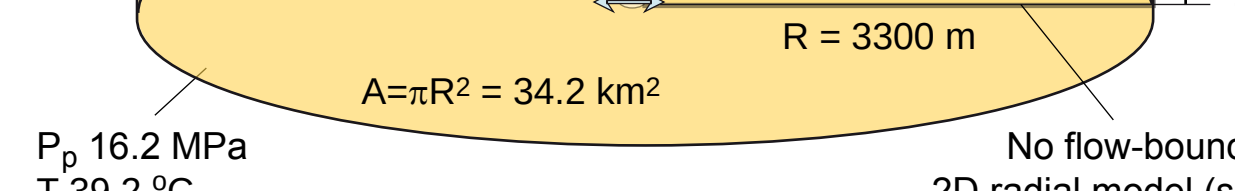


Predicted CO₂ state in the Becancour saline aquifers

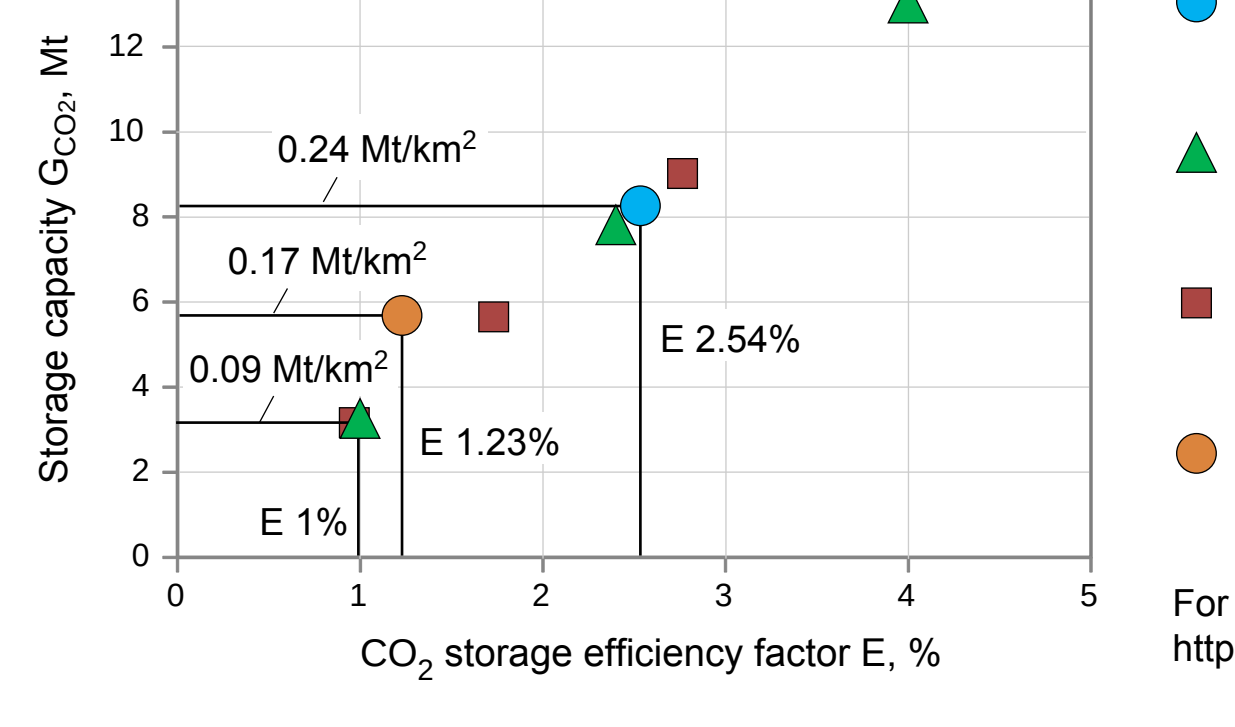


The CO₂ injected in the saline aquifers of the Becancour area is predicted to behave as supercritical fluid at depths below 1026 m where observed pore pressures (P_o gradient 12.2 MPa/km), temperatures (T_o surface 8°C, T gradient 23.5°C/km) and CO₂ density (>800 kg/m³) are above critical values. The predicted CO₂ density is calculated in the saline aquifers under the observed pressures and temperatures using the equation of Span and Wagner (1996).

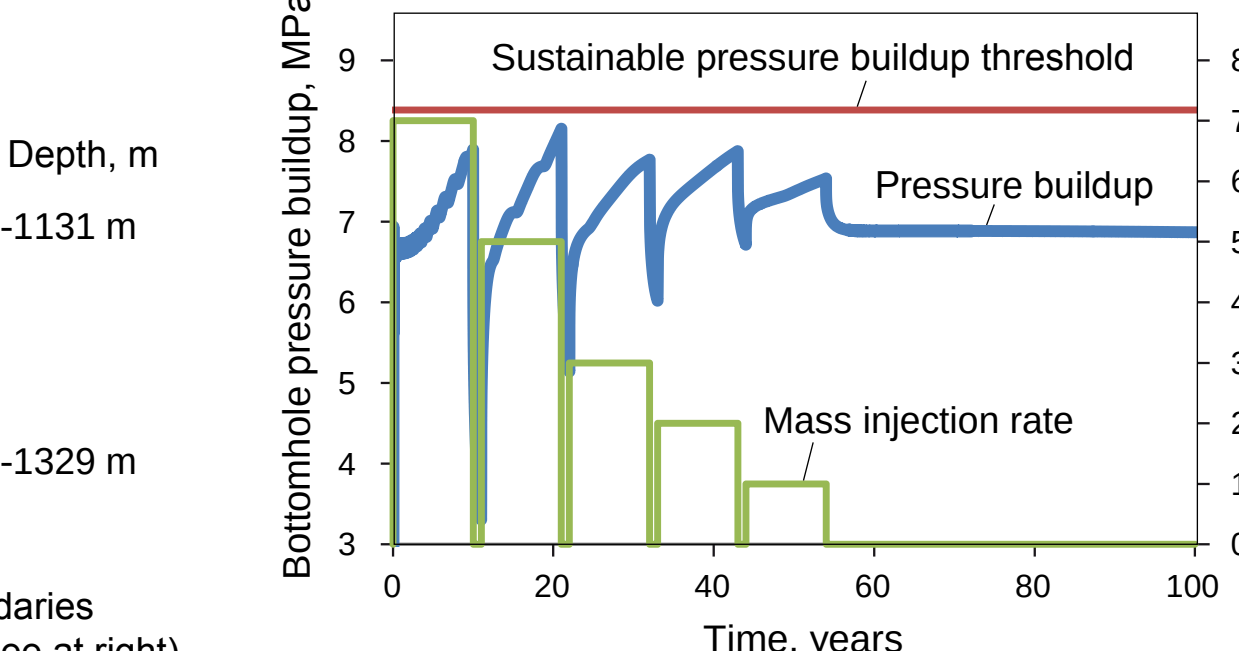
Conceptual radial model of CO₂ injection



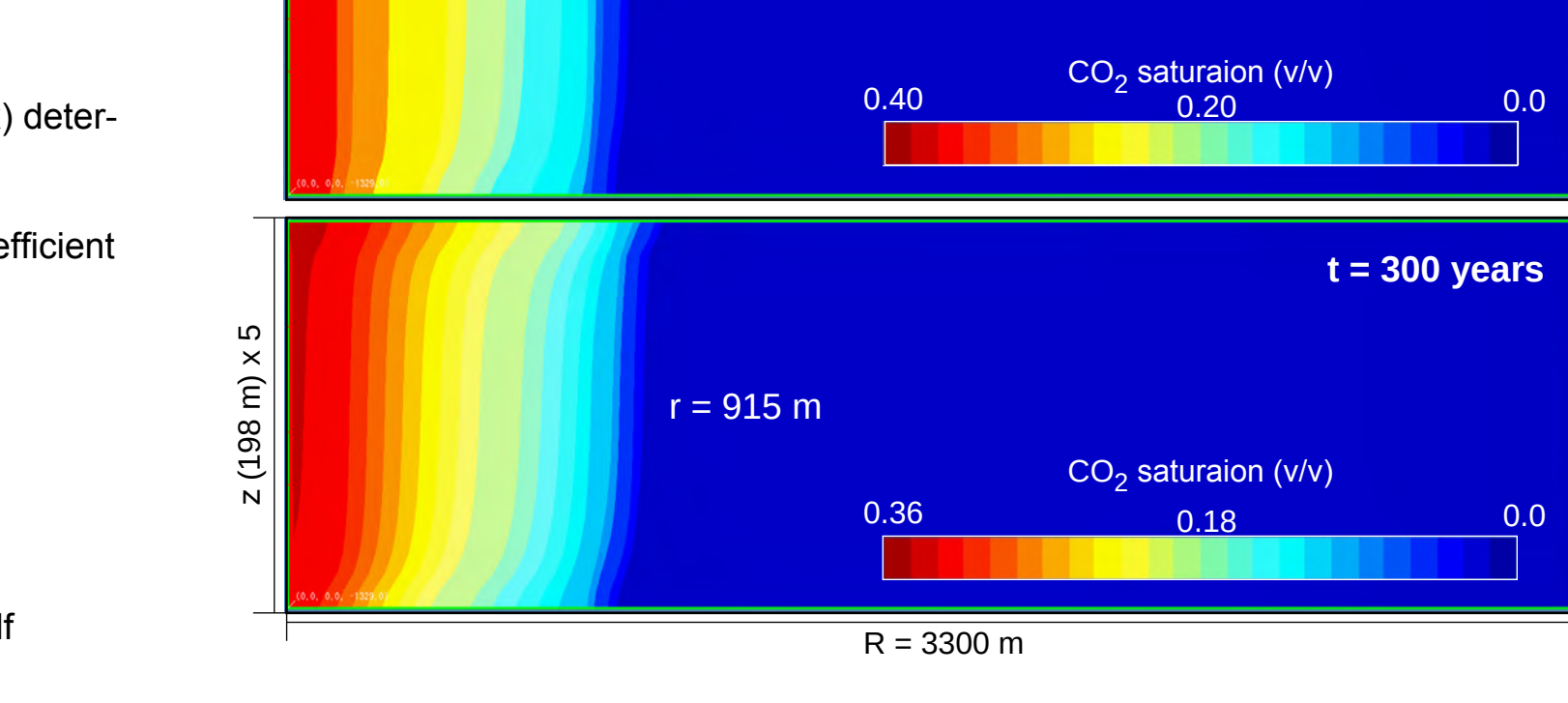
Estimation of CO₂ storage capacity in the Covey Hill sandstones for the area A=34.2 km²



An intermittent injection scenario

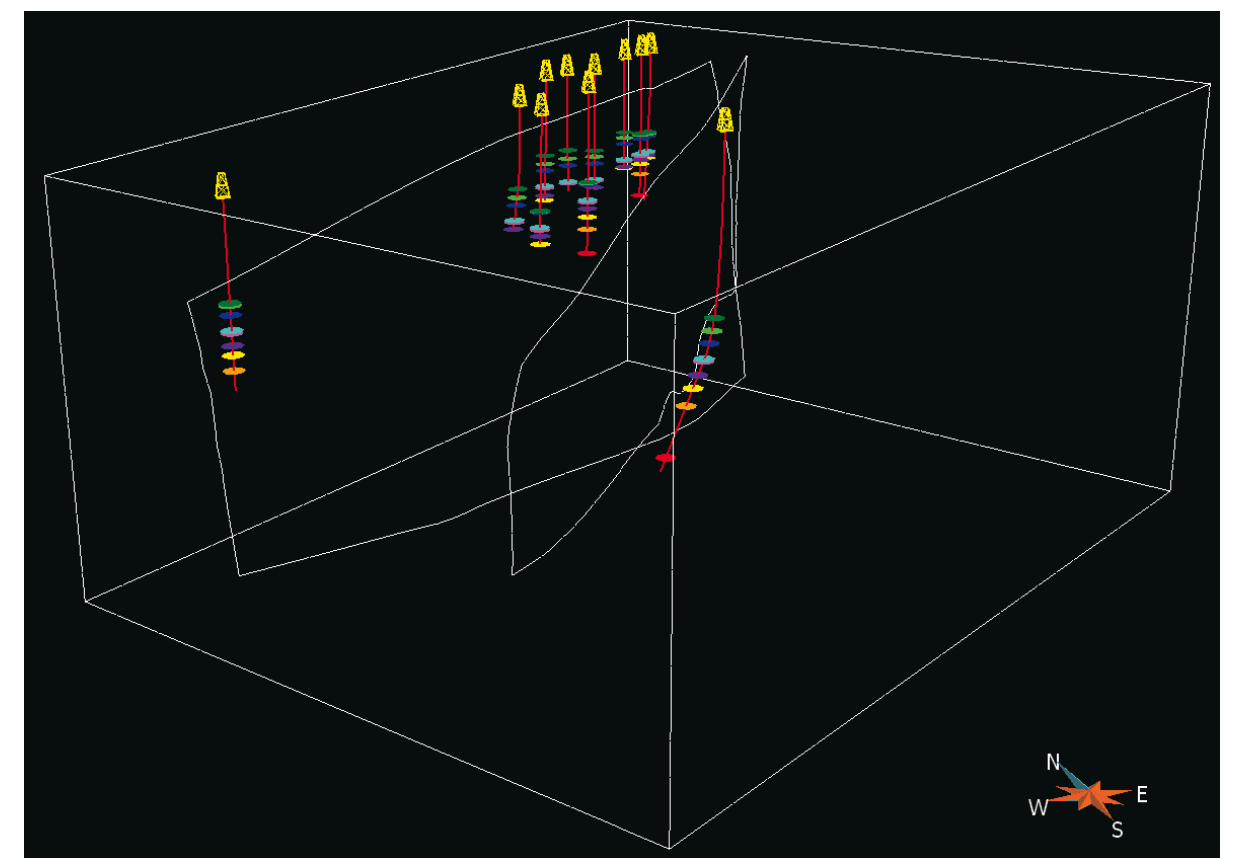


5- 3D geological model of the Becancour area in depth

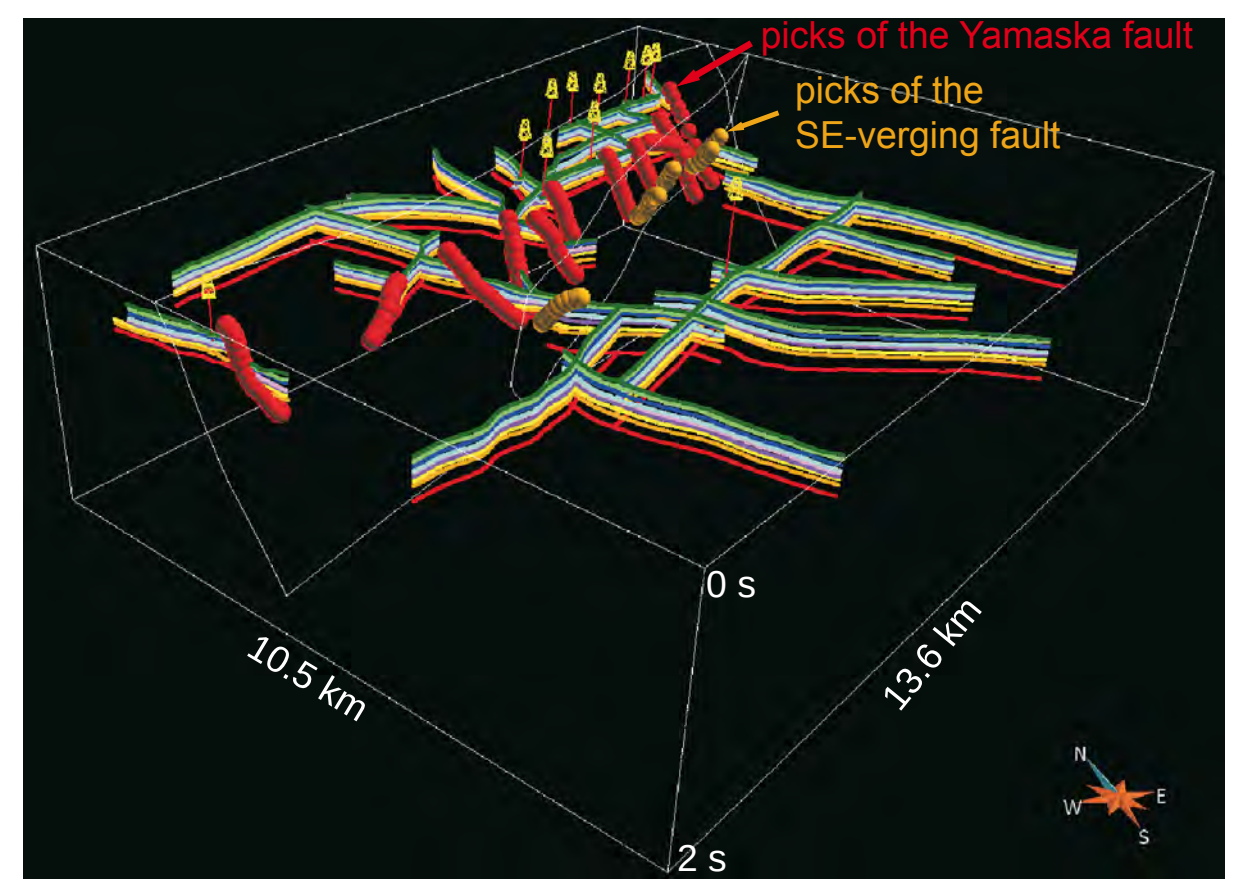


MAIN STEPS TO BUILD 3D GEOLOGICAL MODEL OF THE BECANCOUR AREA

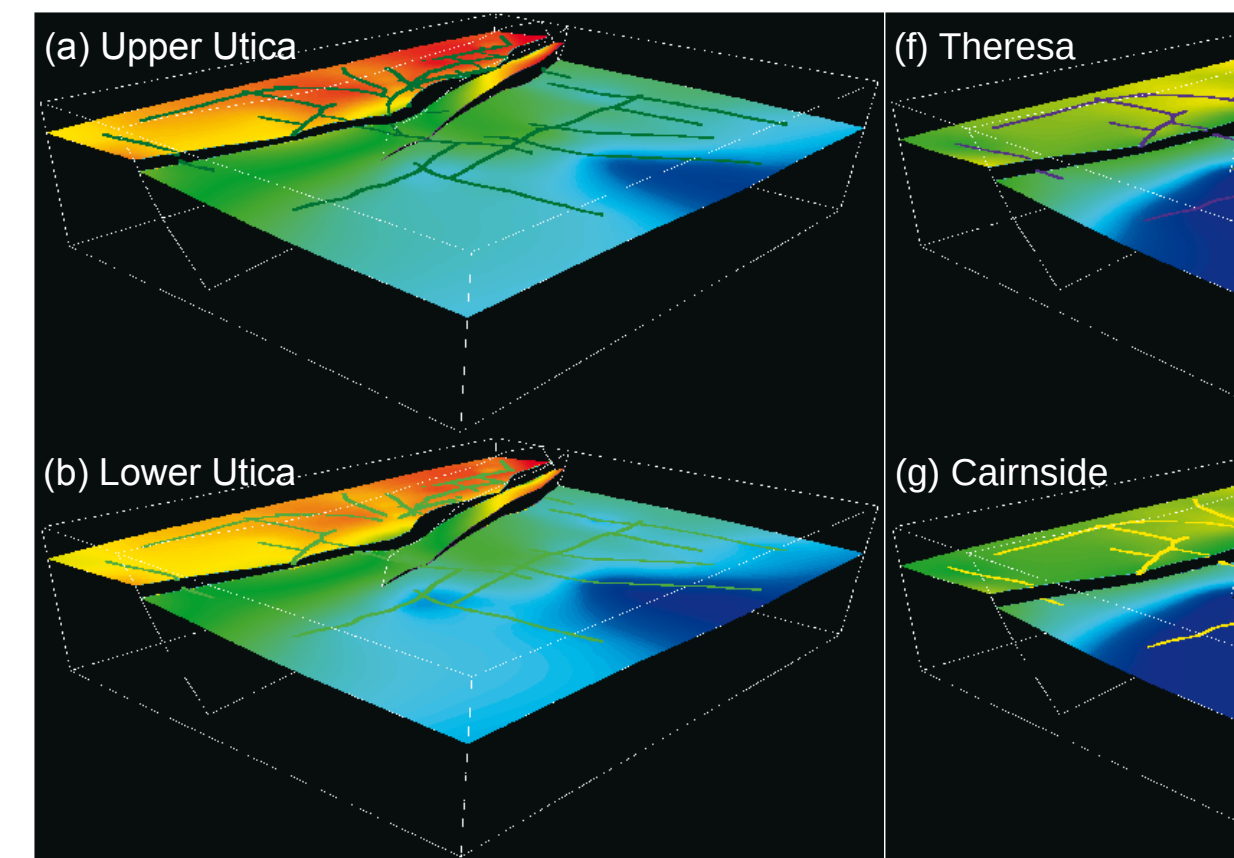
1- Identification of formation tops in depth (m)



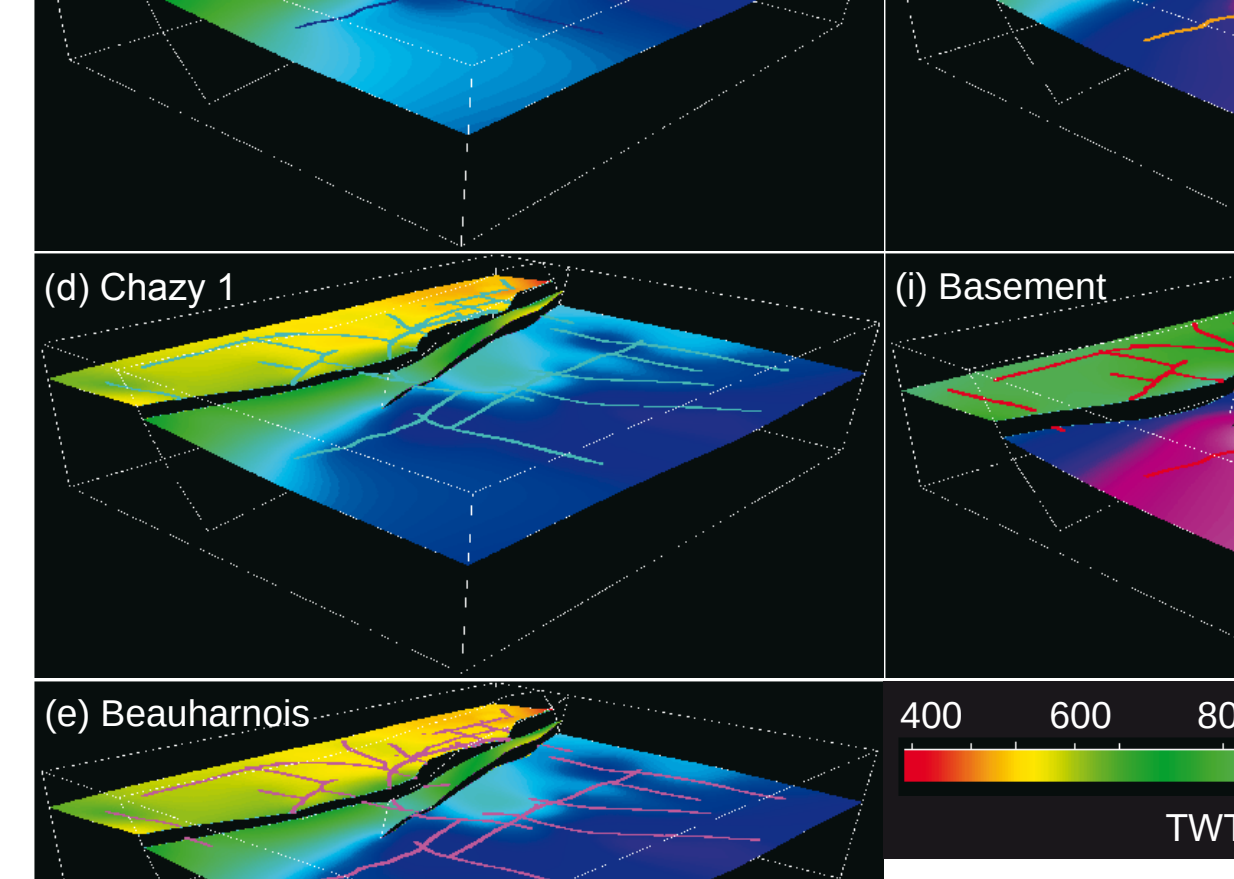
2- Picking of seismic horizons in TWT (s)



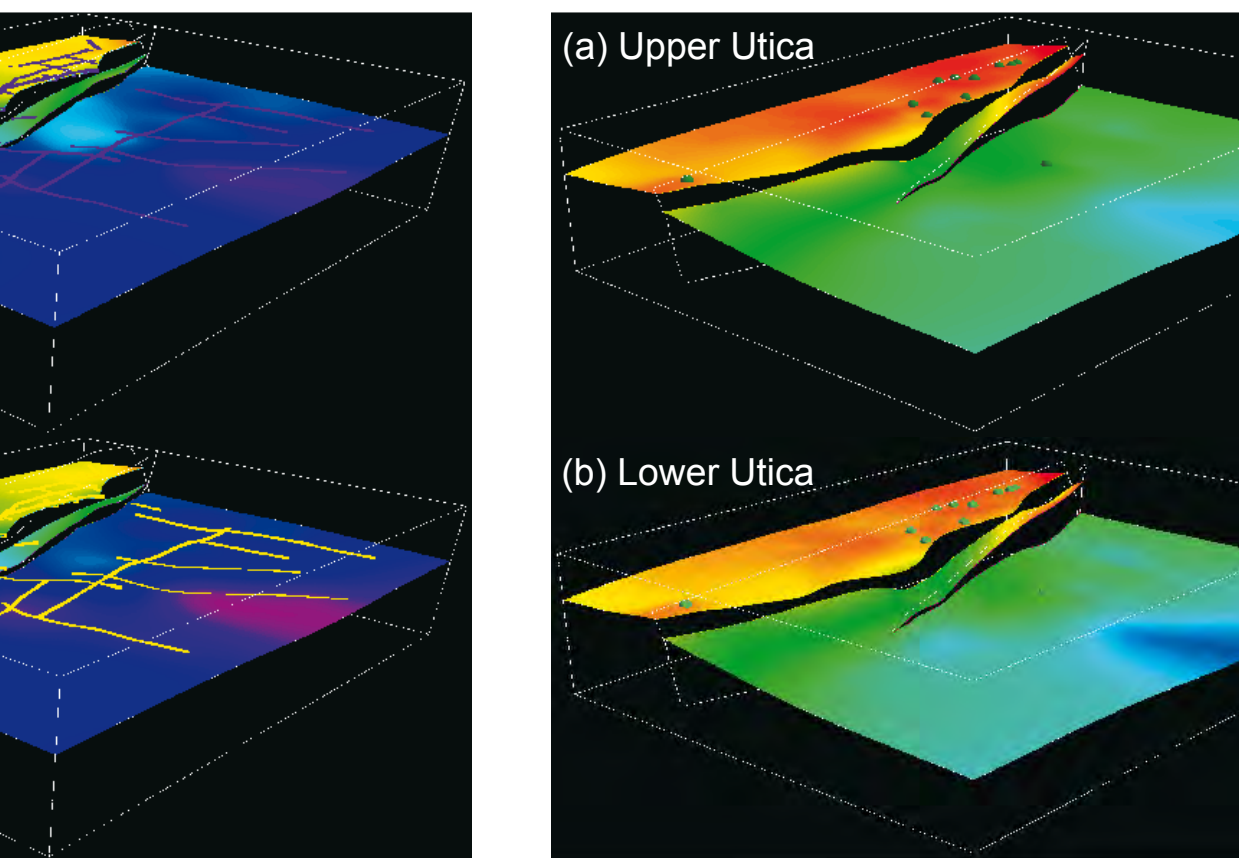
3 - Modeled horizons in TWT from seismic picks



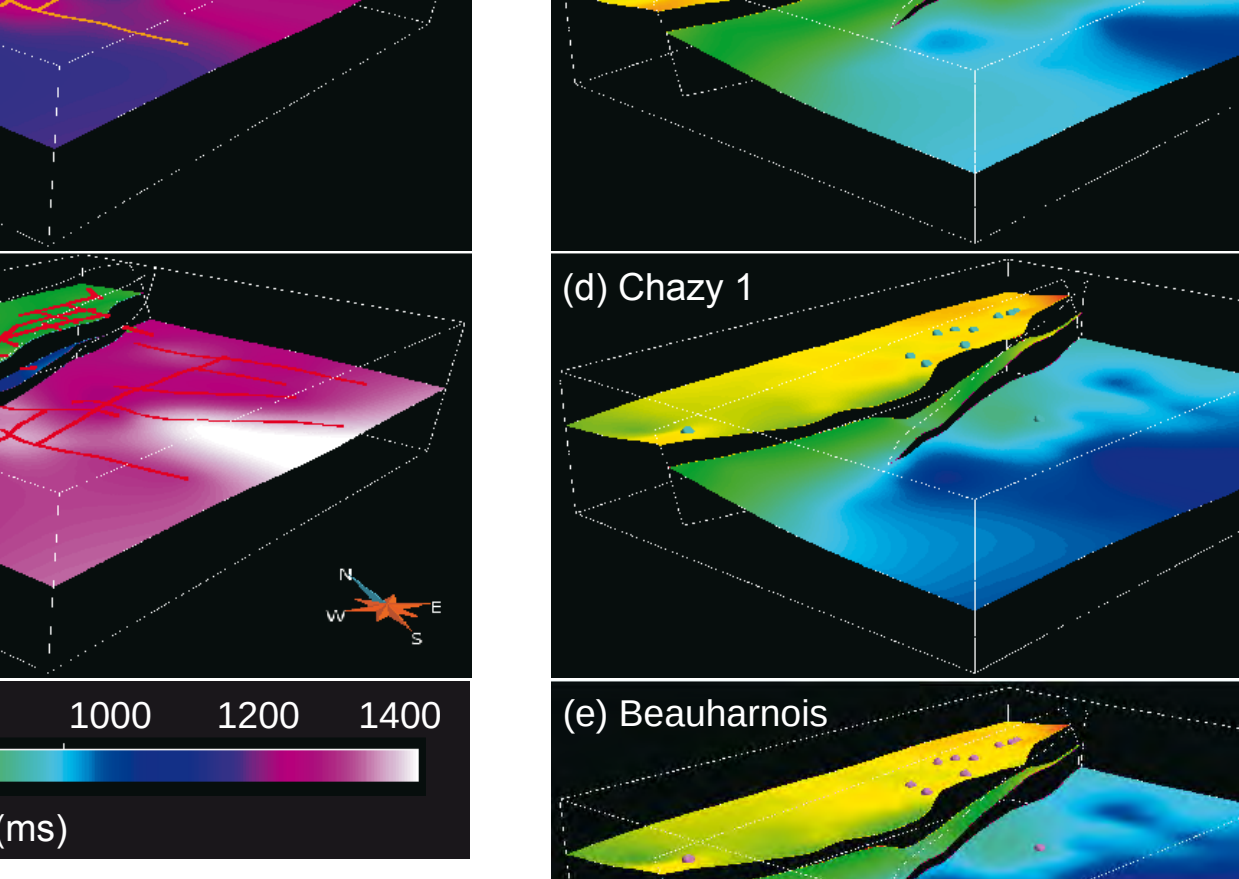
4- Kriged horizons in depth by KED



5- 3D geological model of the Becancour area in depth



6- 3D geological model of the Becancour area in depth



The 3D geological model of the Becancour area (143 km²) is based on structural interpretation of 2D seismic lines (100 km) and well log analysis (11 wells). The main steps that were followed to build the model include:

- (1) identification of formation tops in depth (m) after well log analysis
- (2a) projection of formation tops on 2D seismic lines and identification of seismic horizons by picking travel times in TWT
- (2b) structural interpretation of seismic lines and localisation of faults
- (3) kriging with the modeled horizons in TWT as the external drift
- (4) kriging with the modeled horizons in depth (m)
- (5) build-up of 3D geological model

The 3D geological model is the frame for 1) geostatistical modeling to characterize the spatial distribution of petrophysical properties (porosity, permeability) in the saline aquifers and 2) numerical simulations for CO₂ injection into the Covey Hill Formation.

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