

“Changes in extreme storm surge events in Southern Europe”

***Alba Cid, Sonia Castanedo, Ana J. Abascal, Melisa Menéndez,
Fernando Méndez and Raúl Medina***

University of Cantabria

MOTIVATION OF THIS STUDY?



- ✓ For many **coastal** engineering applications (flooding, coastal management, etc...), knowledge of climate **variability** and **changes** in sea level is important.
- ✓ Storm surge is a key component of sea level.
- ✓ **Extreme** events have immediate **impact** on the coast.

- How can we **characterize** the **extreme** events?
- Are the **extremes** storm surge events **changing**?
- Do these changes have a **spatial pattern**?
- Are they changing in **frequency** or **magnitude**?



Foto: Roberto Barcellos

MAIN GOALS

- ❑ Development of a **methodology** to estimate extremes of storm surge
- ❑ Assessment of the **50-year return level**
- ❑ Analyse significant **trends** in **frequency** and **magnitude** of severe storm surges over Southern Europe

High resolution storm surge hindcast for Southern Europe. *GOS* (*Global Ocean Surge*).

MODEL DOMAIN

- ❑ Southern Europe
- ❑ Horizontal resolution of $1/8^\circ$ (~ 13 Km).
- ❑ The bathymetry was extracted from the **ETOPO 2** database.

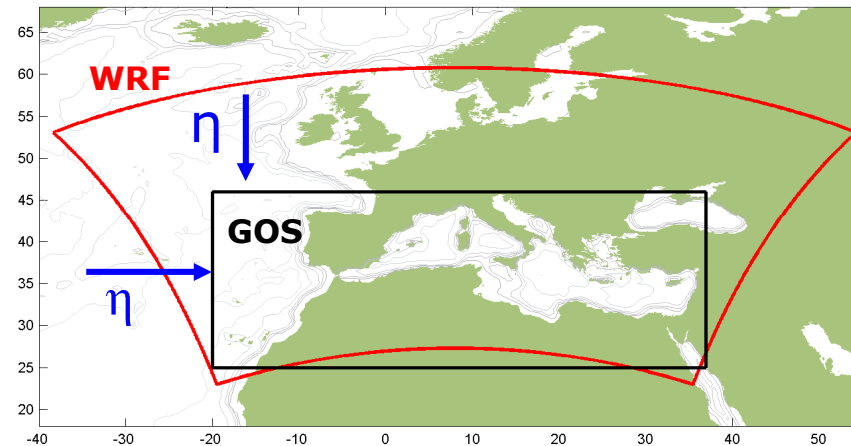
NUMERICAL MODEL

- ❑ The storm surge simulation has been performed using the **Regional Ocean Model System (ROMS)** (Shchepetkin and McWilliams, 2005).
- ❑ Barotropic mode (2D)



BOUNDARY CONDITIONS

Inverse barometer effect



ATMOSPHERIC FORCING

- ❑ **Hourly** meteorological data of **wind** and atmospheric **pressure**
- ❑ ROMS model was forced with an atmospheric dynamical downscaling (Seawind Project, (Menéndez et al., 2011))
- ❑ SeaWind: WRF forced with NCEP reanalysis (1948-2009, $\Delta x \sim 30$ km)

High resolution storm surge hindcast for Southern Europe. *GOS* (*Global Ocean Surge*).

MODEL DOMAIN

- ❑ Southern Europe
- ❑ Horizontal resolution of $1/8^\circ$ (~ 13 Km).
- ❑ The bathymetry was extracted from the **ETOPO 2** database.

Hourly dataset of **62-years** of storm surge in Southern Europe with a horizontal resolution of $1/8^\circ$ (~ 13 km)

NUMERICAL MODEL

- ❑ The storm surge simulation has been performed using the **Regional Ocean Model System (ROMS)** (Shchepetkin and McWilliams, 2005).
- ❑ Barotropic mode (2D)



BOUNDARY CONDITIONS

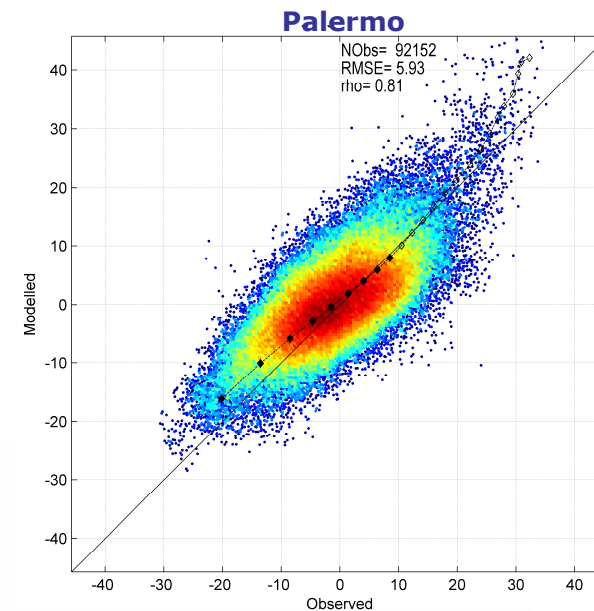
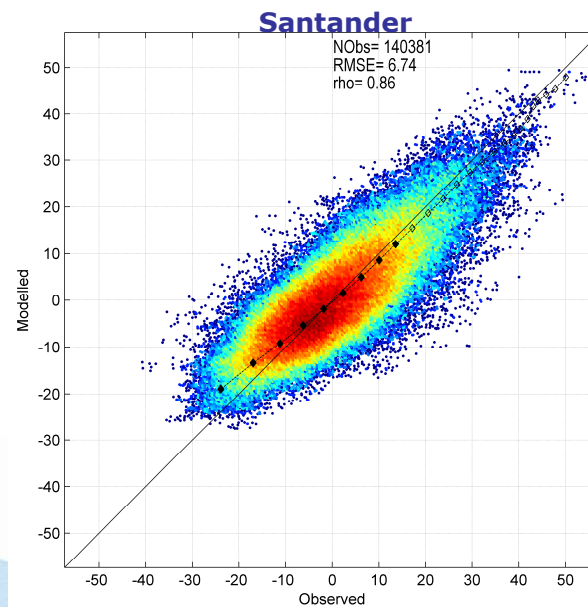
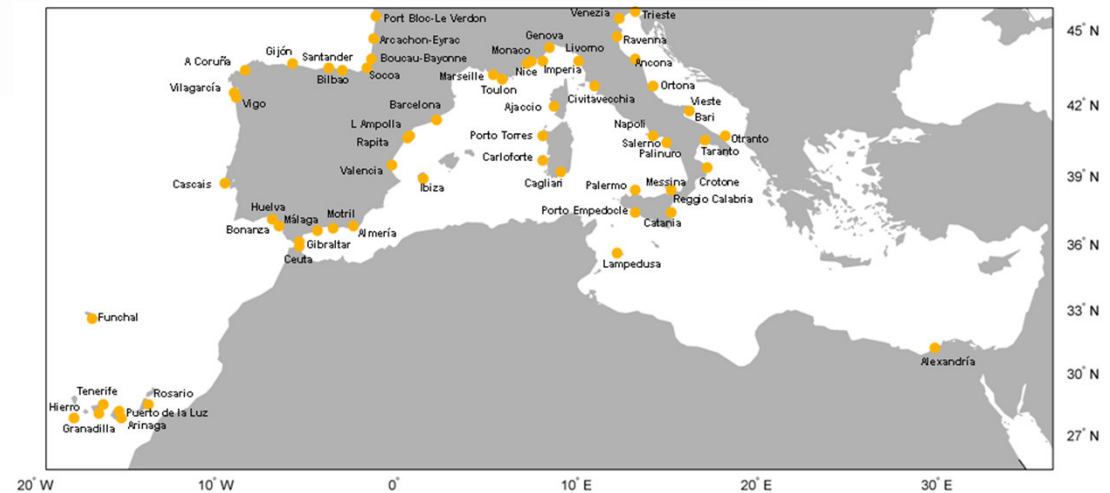
Inverse barometer effect

ATMOSPHERIC FORCING

- ❑ **Hourly** meteorological data of **wind** and atmospheric **pressure**
- ❑ ROMS model was forced with an atmospheric dynamical downscaling (Seawind Project, (Menéndez et al., 2011))
- ❑ SeaWind: WRF forced with NCEP reanalysis (1948-2009, $\Delta x \sim 30$ km)

VALIDATION

- ❑ Tide gauges (PdE, SONEL, ISPRA, UHSLC)
- ❑ Satellite altimetry data (Ssalto/Duacs products, distributed by **Aviso**, with support from Cnes)



VALID



T

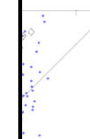
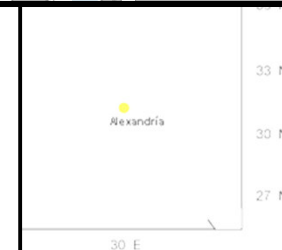
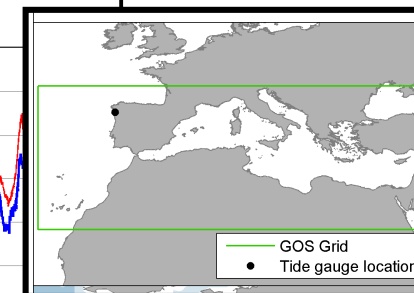
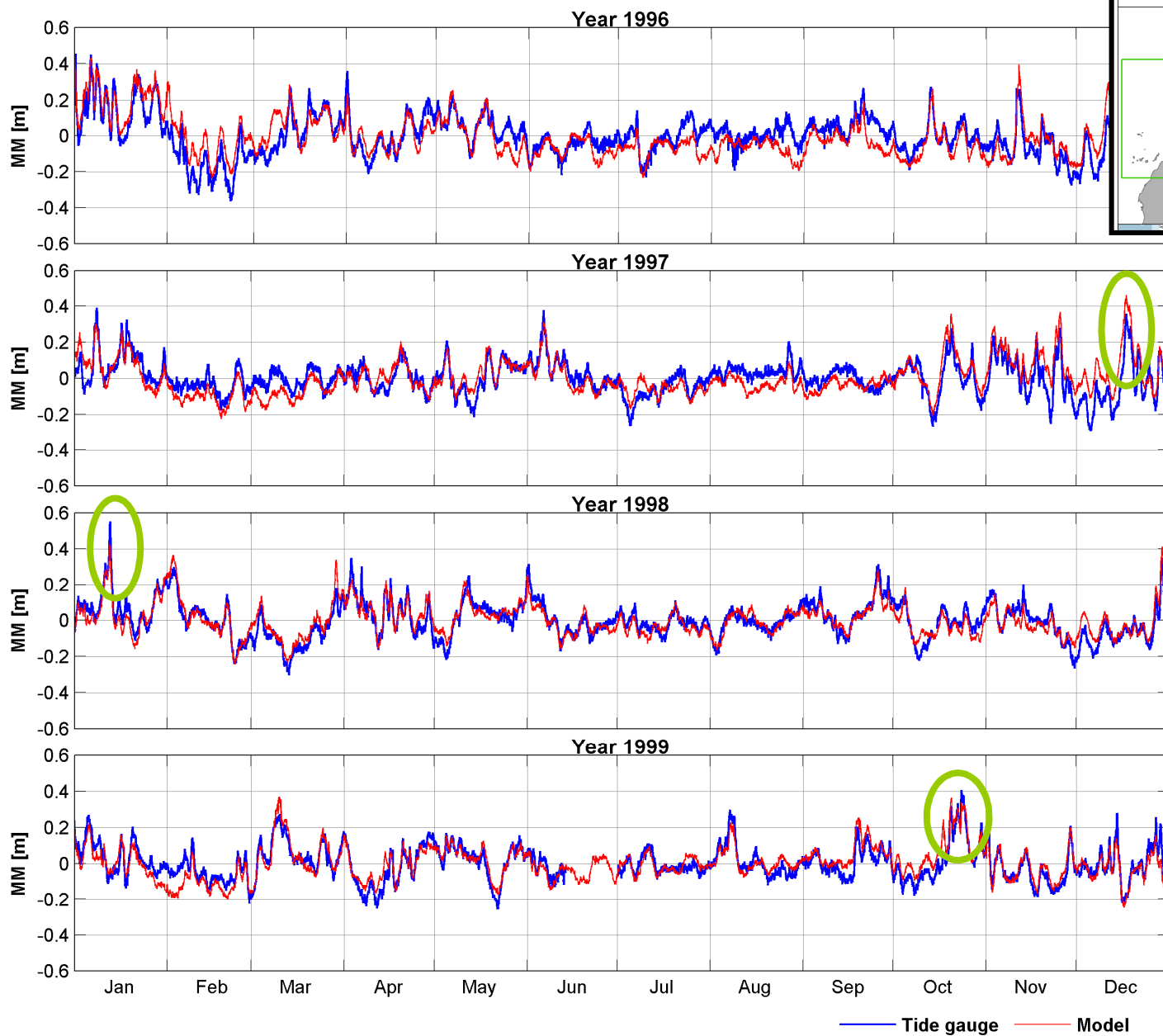


S

(

d

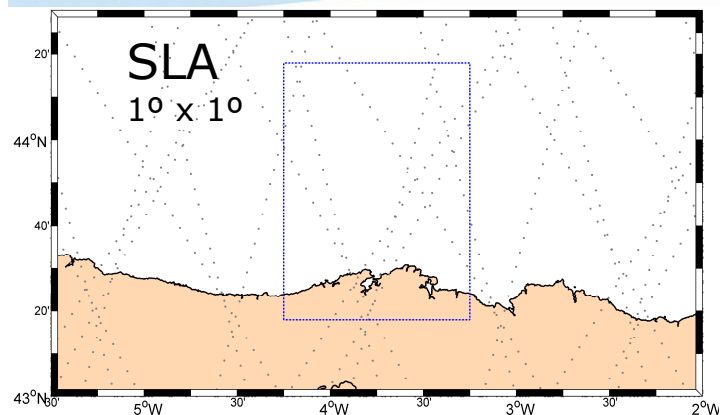
S



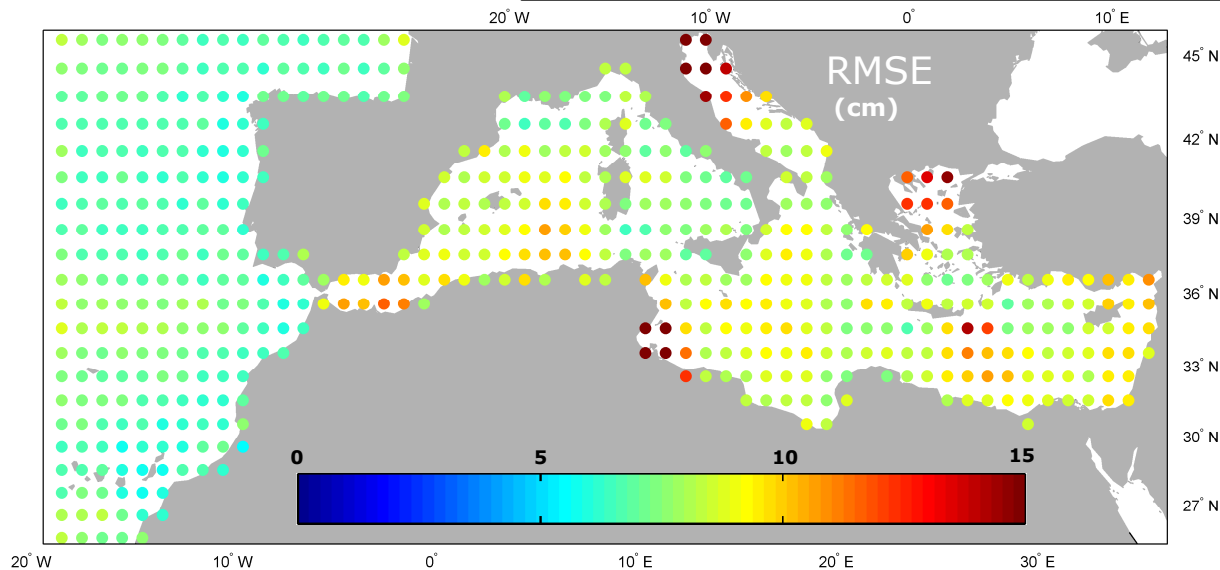
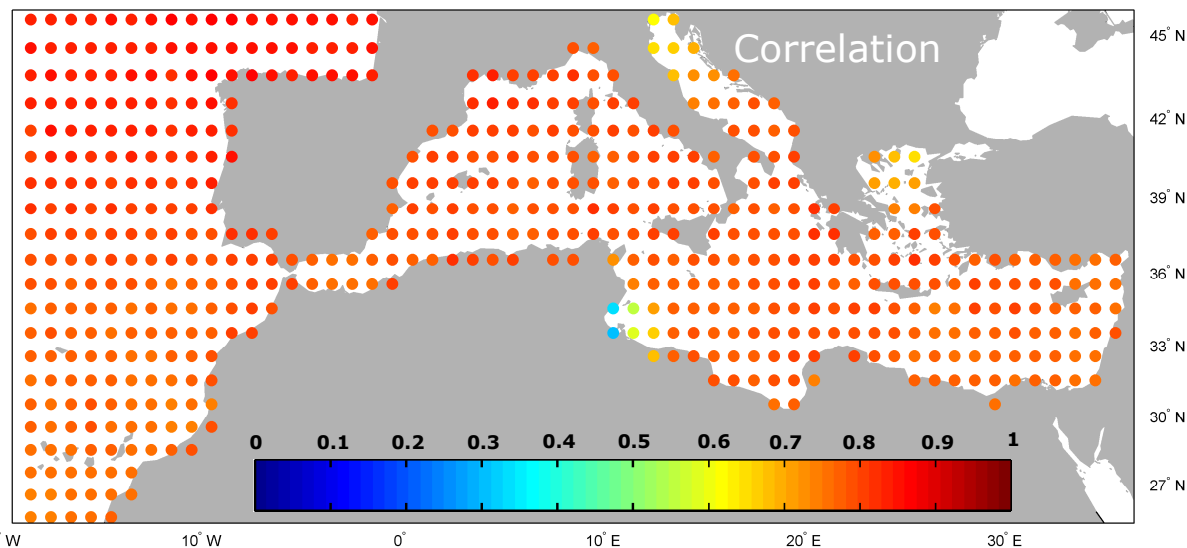
40

VALIDATION Satellite data (1992-2009)

SLA (7 different satellites, ~8 days period)
DAC (regular grid $0.25^\circ \times 0.25^\circ$, 6-hour interval)



Satellite (**SLA + DAC**) vs. **GOS**
(closest grid point in time and space)

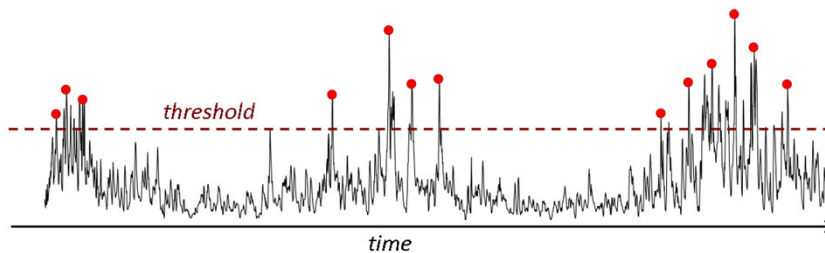


Identifying extreme events: POT

Threshold: 99.5%

Time span : 3 days

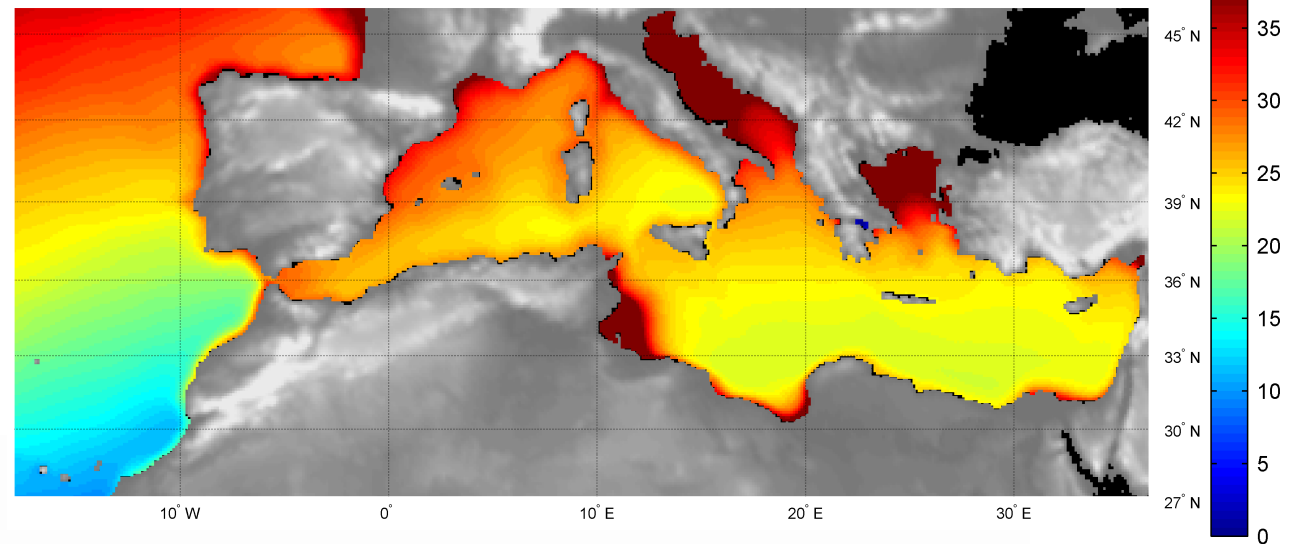
PEAK OVER THRESHOLD (POT)



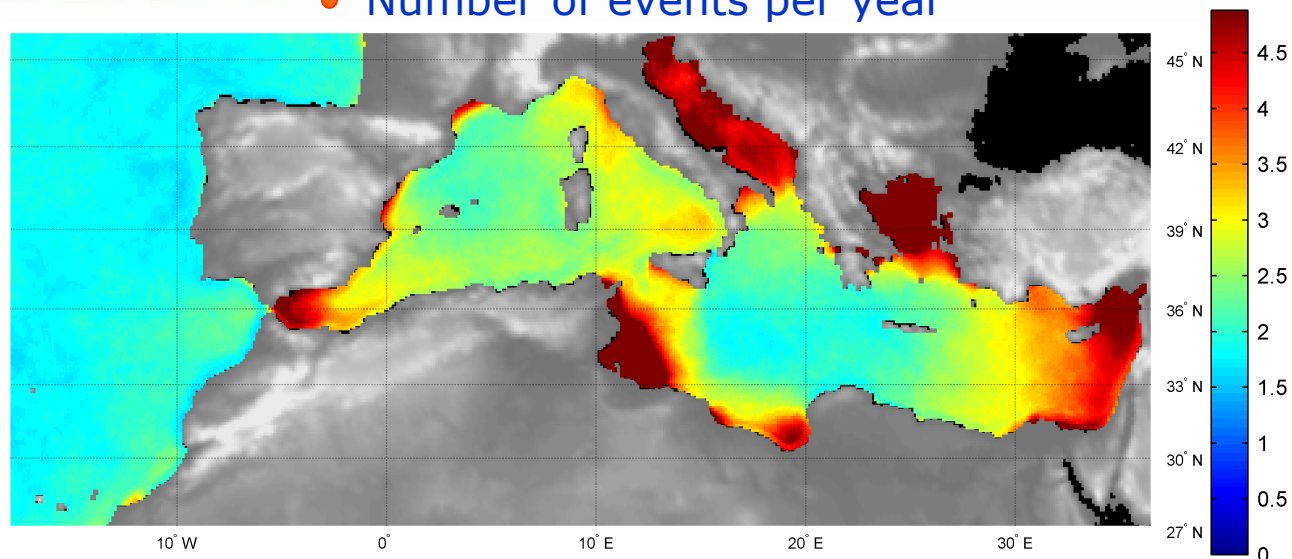
Extreme values description

• Threshold value (cm)

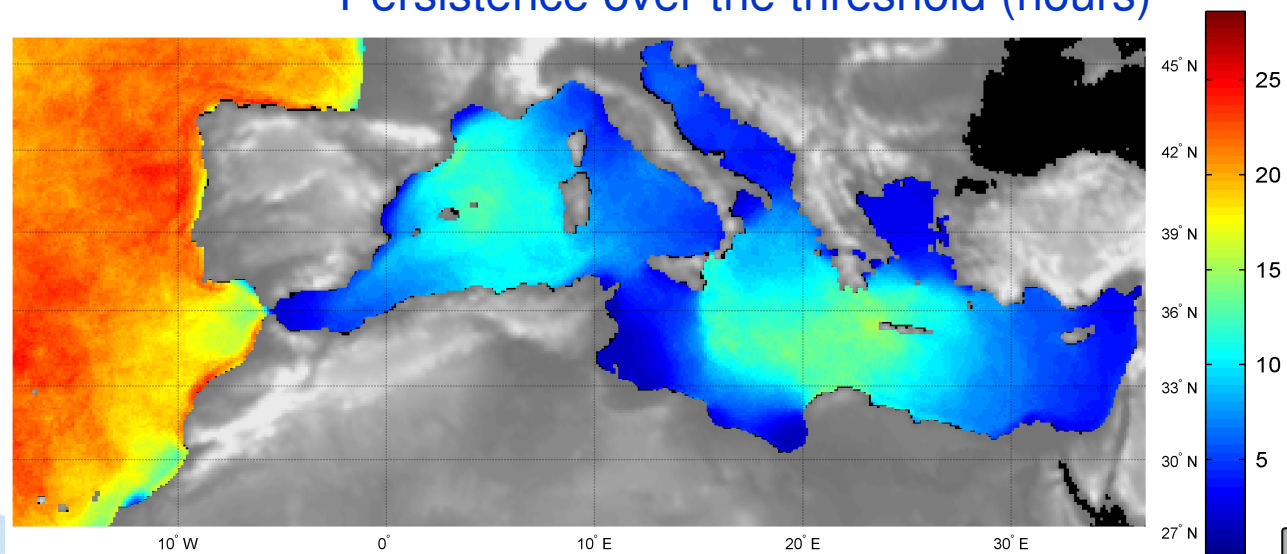
- Number of events per year
- Persistence over the threshold



- Number of events per year



- Persistence over the threshold (hours)

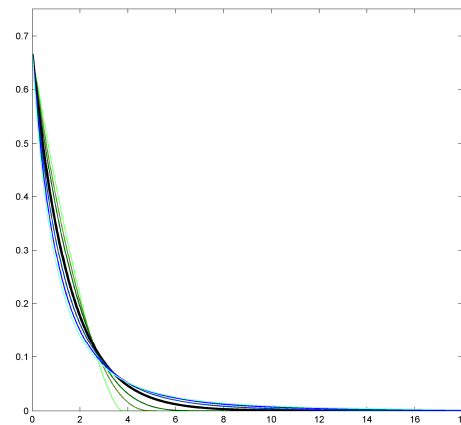


Extreme Value Theory

Statistical discipline that develops a set of techniques and methods to quantify and model the stochastic behavior of extreme events, either in magnitude or frequency.

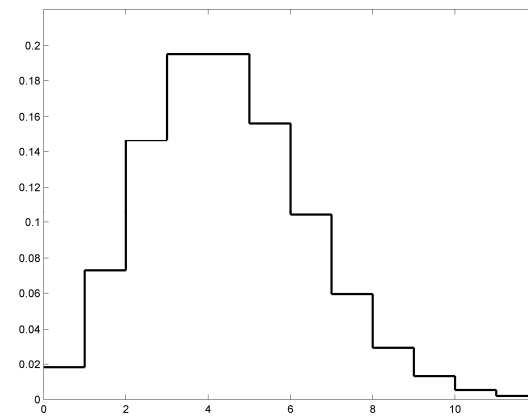
Pareto

(excess values over threshold)



Poisson

(exceedance times over threshold)



Pareto – Poisson model

$$F(x; \theta) = 1 - \left(1 + \xi \frac{x}{\sigma} \right)^{-1/\xi}$$

$$p(x; \theta) = e^{-\lambda} \frac{\lambda^x}{x!}$$

Shape parameter

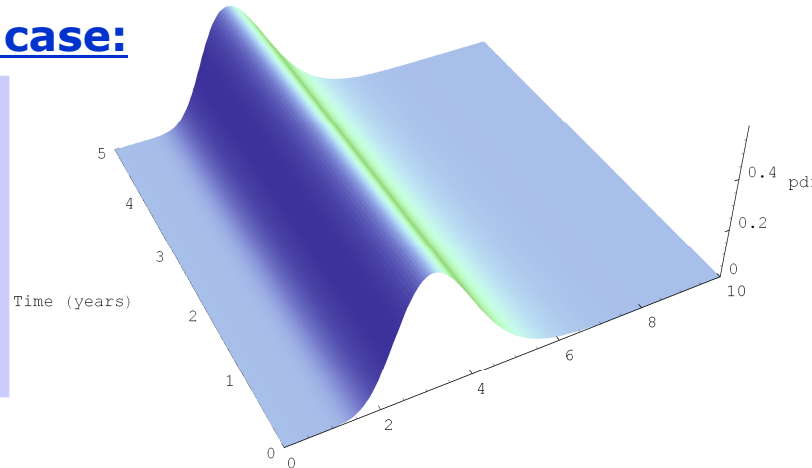
Scale parameter

Stationary case:

$$\sigma = \sigma_0$$

$$\xi = \xi_0$$

$$\lambda = \lambda_0$$



Maximum likelihood estimation

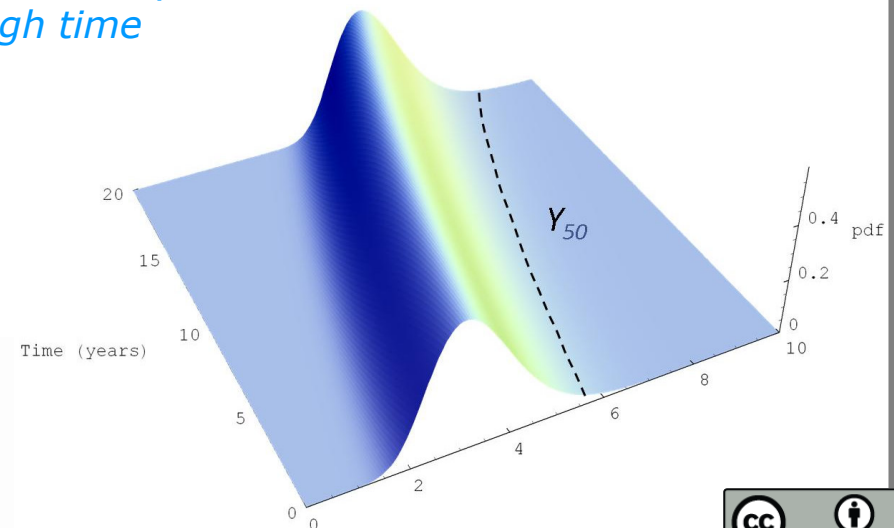
Time-dependent case:

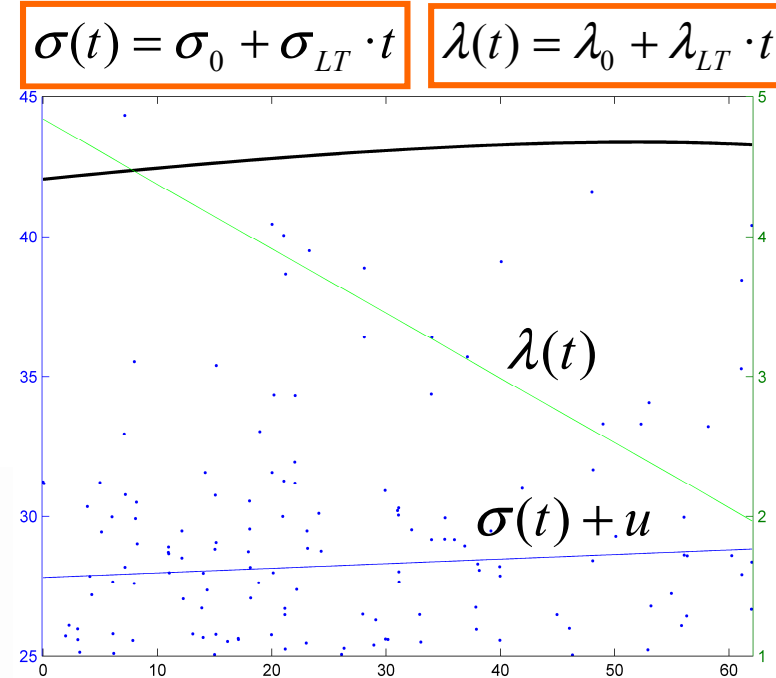
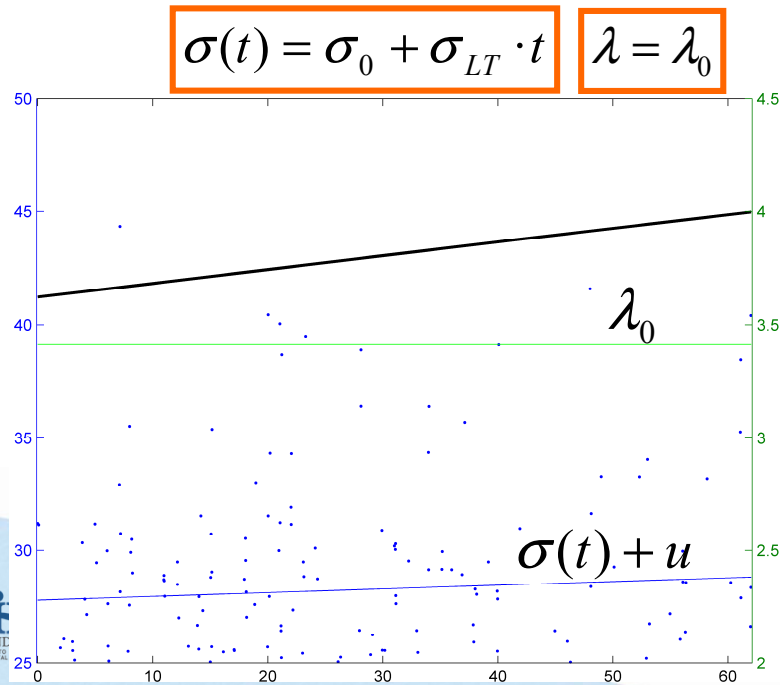
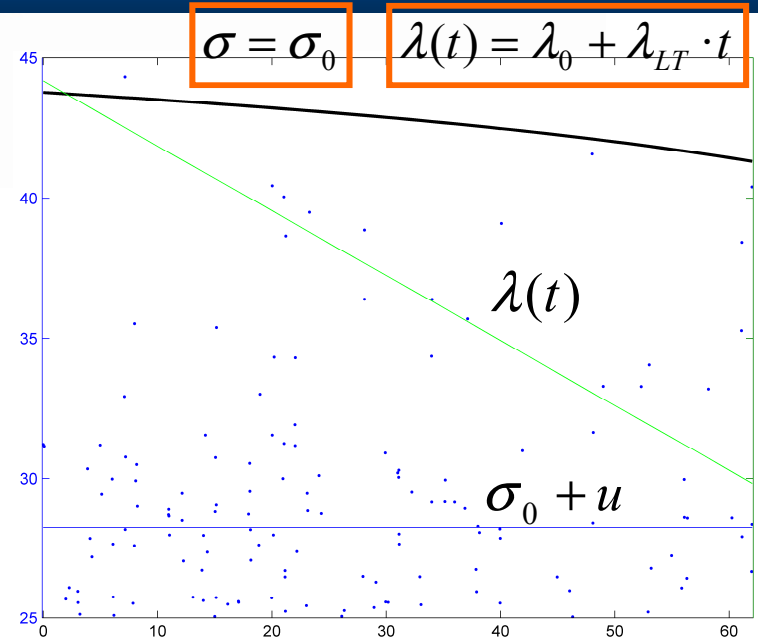
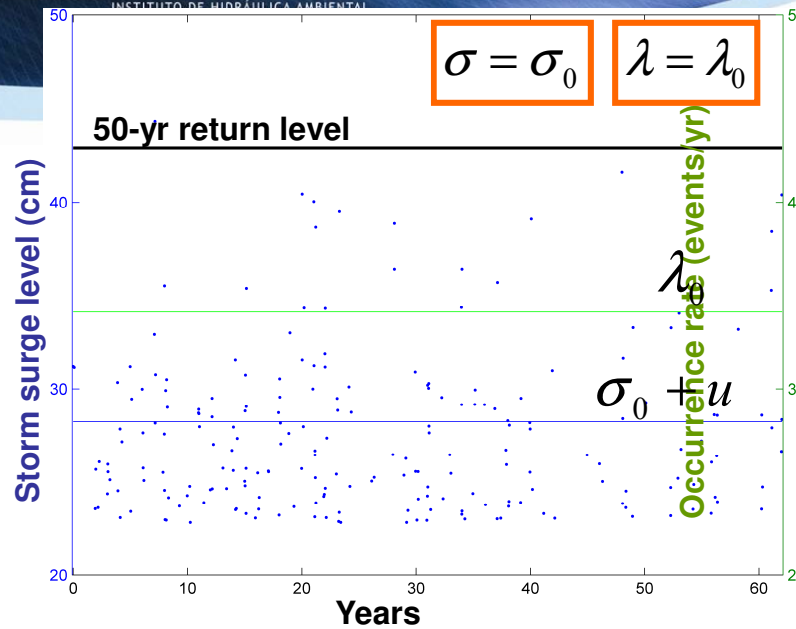
$$\sigma(t) = \sigma_0 + \sigma_{LT} \cdot t$$

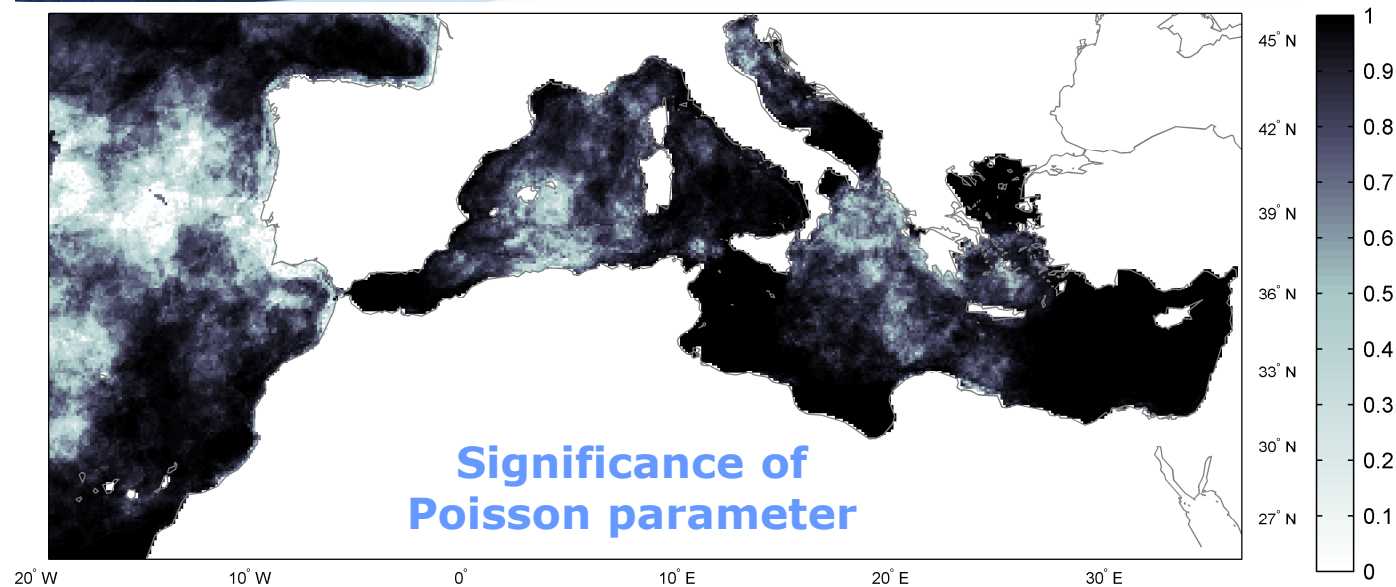
$$\xi = \xi_0$$

$$\lambda(t) = \lambda_0 + \lambda_{LT} \cdot t$$

The probability of an extreme sea level varies through time



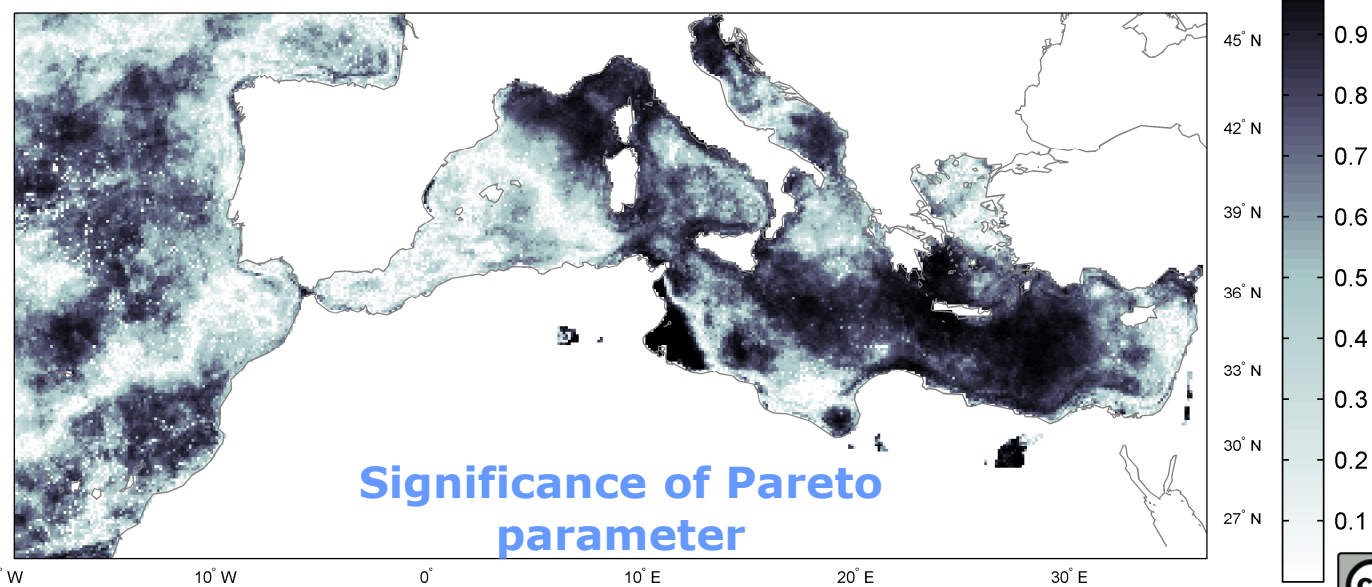


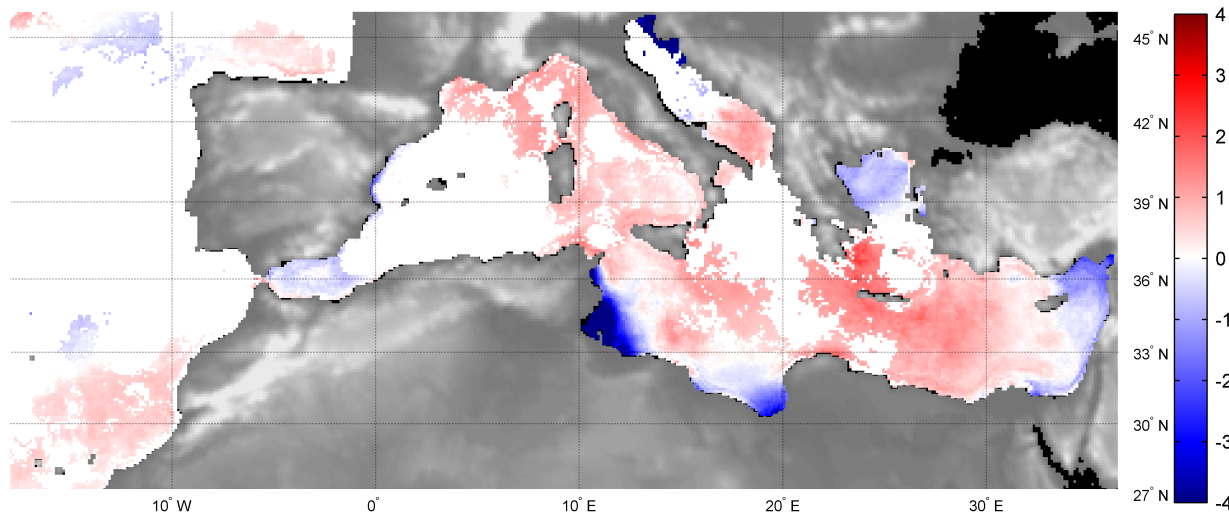


Stationary case

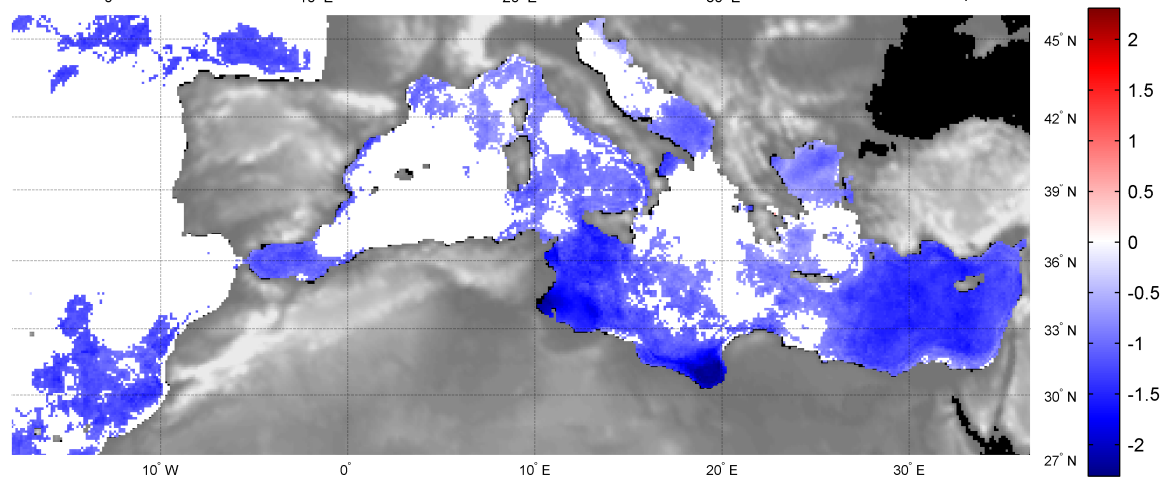
Vs.

Time-dependent case



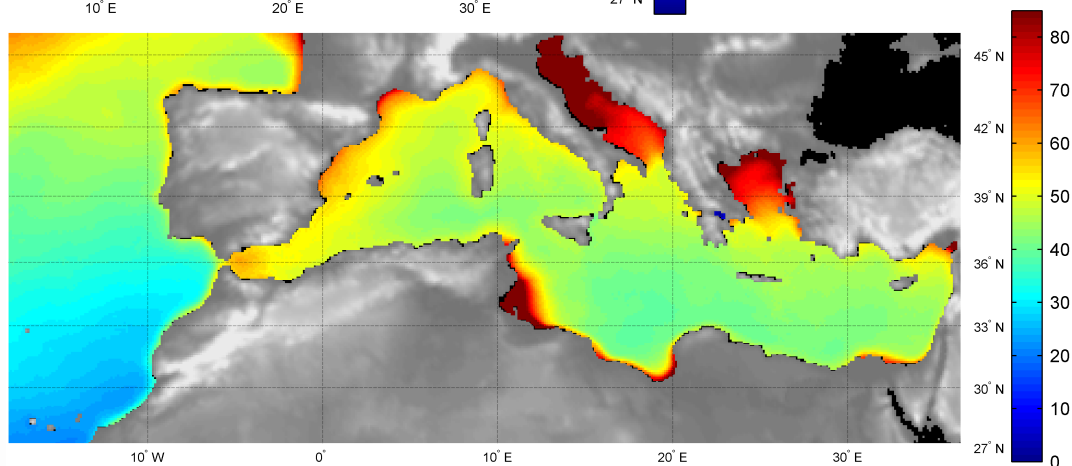


Trends statistically significant at 90%



Fifty-year return level (cm)

$$Y_{50}(t) = u - \left(\left(\frac{\sigma_t}{\xi_0} \right) \cdot \left(1 - (\lambda_t \cdot 50)^{\xi_0} \right) \right)$$



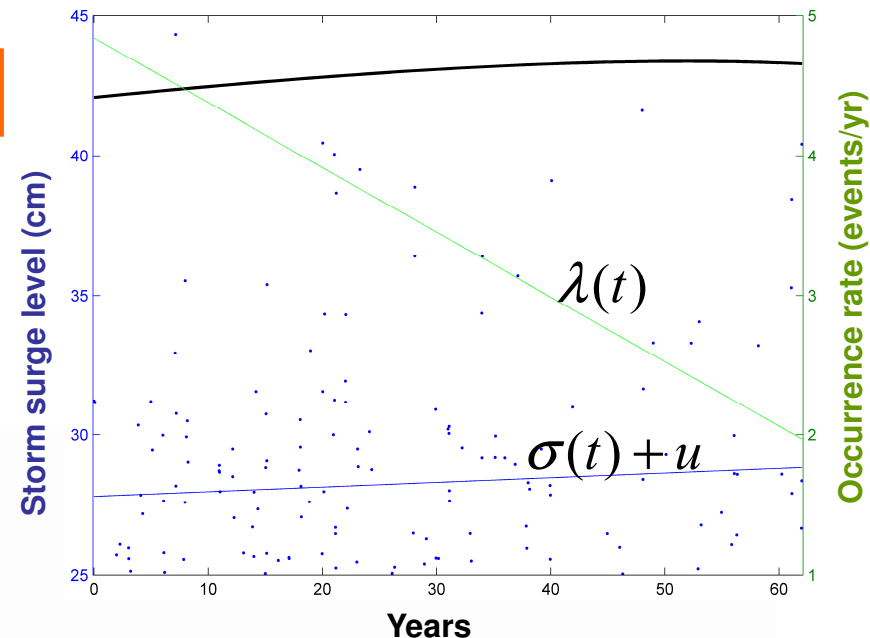
Conclusions

- ❑ Trends in the **magnitude** of extreme values are around ± 2 mm/yr
- ❑ Positive trends are found in the majority of the domain
- ❑ Trends in the **frequency** of extreme events are around -2%
- ❑ Frequency trends are negative all over the domain
- ❑ Trends in **frequency** (related to Poisson) are more important than in **magnitude** (related to Pareto).

$$\lambda(t) = \lambda_0 + \lambda_{LT} \cdot t$$

$$\sigma(t) = \sigma_0 + \sigma_{LT} \cdot t$$

- ❑ Statistical framework presented to study extreme events



Thanks for your attention

Comments? Questions?

alba.cid@unican.es