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Introduction

Clouds cover large areas of the Earth's surface throughout the year and thus play an important role in terms of climate, especially concerning the Earth's radiative energy budget. The radiative cloud properties themselves strongly depend on the fractions of ice crystals and cloud droplets, their size spectra, as well as ice crystal shapes and surface roughness. Hence, for quantifying cloud radiative properties, it is essential to know the

Particle Scattering Simulations



Fig. 1: Intensity ratios $(I_{0.77}/I_{1.00})$ in dependence of a and χ_{ve} (d_{ve}). Blue color: $I_{1.00} \approx I_{0.77}$; white color: $I_{1.00} > I_{0.77}$; other color: $I_{1.00} < I_{0.77}$ (good discrimination)



- (ice crystals)
- water = 1.33)



 χ_{ve} (d_{ve}). Blue color: $I_{1.00} \approx I_{1.30}$; white color: $I_{1.00} > I_{1.30}$; other color: $I_{1.00} < I_{1.30}$ (good discrimination)

Results and Outlook

1.00

- > Based on particle scattering simulations it was found that a scattering angle of 100° is best for phase discrimination when using linearly polarized light.
- Circularly polarized light in combination with an scattering angle of 142° result in an even higher sensitivity concerning the particle phase but it is restricted to particles > 2.54μ m.
- > Based on these results, a new optical particle counter called SHERLOCC was developed.
- > A detailed comparison of SHERLOCC and TOPS-Ice based on measurements using illite, kaolinite, Snomax[®], ATD and/or biological particles as CCN/IN is planned.

A New Optical Ice Particle Counter at LACIS: SHERL&CC



Fig. 4: Sketch of the particle generation, LACIS, and the optical detection system (modified from Hartmann et al., 2013.

- LACIS (7m length, 15mm \emptyset) is a laminar flow tube consisting of seven 1m sections and is connected with a particle generation system (Fig. 4).
- Due to the adjusted thermodynamic conditions, liquid droplets and/or ice The particle generation system crystals can be formed in upstream of the LACIS tube various sizes of the lower μ m produces dry quasirange. monodisperse or polydisperse All particles generated by particles.
- The thermodynamic condition (temperature, relative humidity) inside the LACIS tube are

Fig. 3: Intensity ratios $(I_{1,30}/I_{1,00})$ in dependence of a and χ_{ve} (d_{ve}). Green color: $I_{1.00} \approx I_{1.30}$; white color: $I_{1.00} > I_{1.30}$; other color: $I_{1.00} < I_{1.30}$ (good discrimination)

12

14

Abbreviations

ATD – Arizona Test Dust; **CCN** – Cloud Condensation Nuclei; CPC – Condensation Particle Counter; DMA – Differential Mobility Analyzer; **DPM** – Dew Point hygroMeter; **IN** – Ice Nuclei; **PMT** – PhotoMultiplier Tube; SHERLOCC – Single Hydrometeor Electromagnetic Response Logger for Characterization and Counting; **TH** – Thermostat; **TOPS-Ice** – Thermostabilized Optical Particle Spectrometer for the detection of Ice

5.412

2.275

1,490

0.7059

-0.8627

-0.07843



Counting) was developed similar to TOPS-Ice (Thermostabilized Optical Particle Spectrometer for the detection of Ice, Clauss et al. 2013). SHERLOCC is designed for laboratory measurements in combination with the laminar flow tube LACIS (Leipzig Aerosol Cloud Interaction Simulator, Stratmann et al., 2004). This setup aims at improving the water droplet and ice crystal discrimination in the lower micrometer range.

Particle Detection: SHERL optical fiber (400µm) **PMT A** sensitive volume beam trap air tight optical cell cylindrical lenses **100**% polarizer spherical lenses PMT B pol. beam optical fiber (200µm) splitter cube PMT C

separately controlled for each of the seven sections.

LACIS are directly fed into the optical detection system TOPS-Ice or SHERLOCC.

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Fig. 5: Sketch of SHERLOCC (top view).

- SHERLOCC (Fig. 5) is based on TOPS-Ice (Clauss et al., 2013).
- The key modification in SHERLOCC is the change of the detected scattering angle from 42.5° to 100° based on particle scattering simulations.
- SHERLOCC is designed to detect the number concentration, size (PMT A) and phase of the particles (PMT B/C) in the lower μ m range.
- This setup uses a vertically polarized 200mW laser at 532nm to measure single ice particles inside the sensitive volume.

References

Clauss, T., et al.: Application of linear polarized light for the discrimination of frozen and *liquid droplets in ice nucleation* experiments, Atmos. Meas. Tech., 6, 1041–1052, 2013.

Hartmann, S., et al.: Immersion freezing of ice nucleating active protein complexes, Atmos. Chem. Phys., 12, 5751–5766, 2013.

Stratmann, F., et al.: Laboratory studies and numerical simulations of cloud droplet formation under realistic supersaturation conditions, J. Atmos. Ocean. Tech., 21, 876–887, 2004.



 The change in polarization detected by PMT C is used for the discrimination between liquid droplets and ice crystals.

- The lower detection limit of TOPS-Ice was ~1 μ m for mixed phase and pure ice clouds, and ~ 0.5μ m for liquid clouds. Due to the modification of the scattering angle in SHERLOCC, the detection limit is supposed to decrease even further.