

# **Downscaling Ocean Conditions: Initial Results using a Quasigeostrophic and Realistic Ocean Model**

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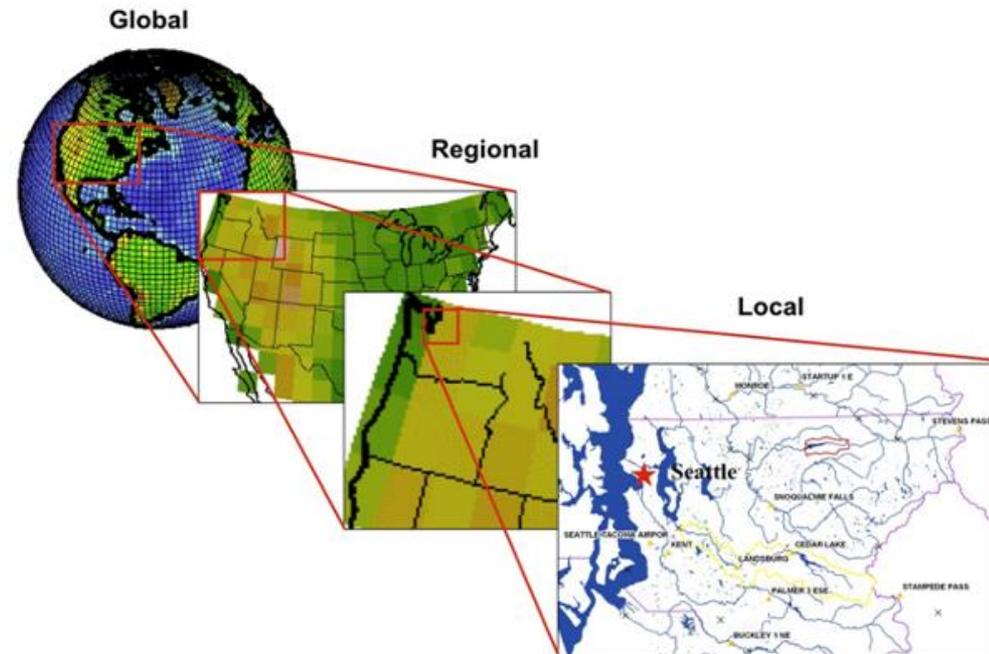
# Outline of Talk

- **Background and Motivation**
- **Approach**
- **Idealized Application**
- **Northwest Atlantic Application**

# Motivation

Need for historical reconstructions, predictions and projections of physical oceanographic conditions on **regional scales**

- The **local scale variability** is often **controlled** by processes operating on **large scales**
- Model large parts of the ocean with regions of increased resolution (**Downscaling**)
- Several **problems** with **nesting**



# Coupling of Large and Small Scales

- **Henshaw et al., (2003):** Accurate reconstruction of the small scale variability from the time history of the large scales.
- **Large scales** can guide reconstruction, prediction and projections over limited areas.

**Atmospheric Spectral Nudging** (e.g., von Storch et al., 2000): Coarser resolution fields are used to drive the higher resolution model in the **interior** and not just the **boundaries**.

$$\psi_a = \psi_f + \underbrace{\gamma \langle \psi_i - \psi_f \rangle_L}_{\text{Spectral nudging term}}$$

**Spectral nudging term**

$\psi_f$ : regional model forecast

$\psi_i$ : independent estimate of the true field (coarser resolution model)

$\langle \rangle_L$ : spectrum that corresponds to large scales

$\gamma$ : Nudging coefficient

# Downscaling Ocean Conditions

**Will it work in the ocean?**

**The different ocean length scales can be coupled through:**

- **Non-linearity** of the governing equations
- **Surface forcing**
- Reflection of the **Rossby waves** by the western boundary

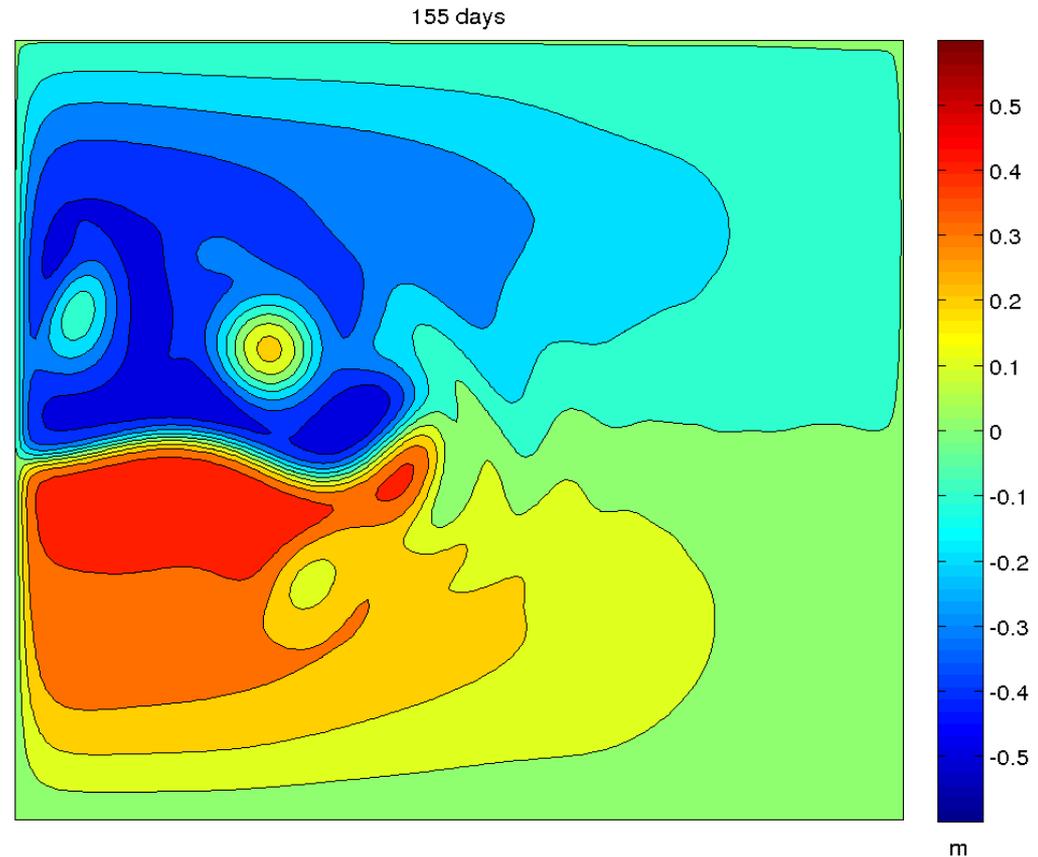
**Important Atmosphere and Ocean differences:**

- Internal **Rossby radius** of deformation
- **Lateral boundaries**
- Relative **few observations** of the ocean

**One step further:** Assimilate available sparse point observations along with the large scale information .

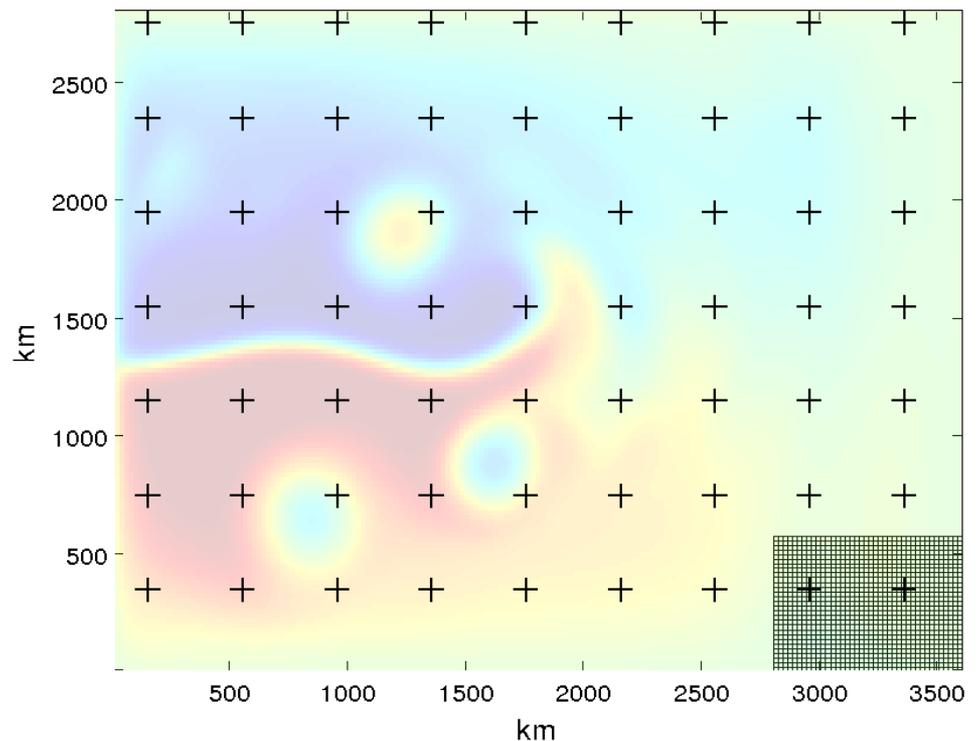
# QG Model Application

- Rectangular ocean (3600x2800 km with resolution 20km) on **beta plane**
- **1/2 layer** model
- The model is forced by a **steady zonal** wind distribution that generates two gyres separated by a meandering ocean jet
- Scale separated using 2D discrete **Fourier transform**

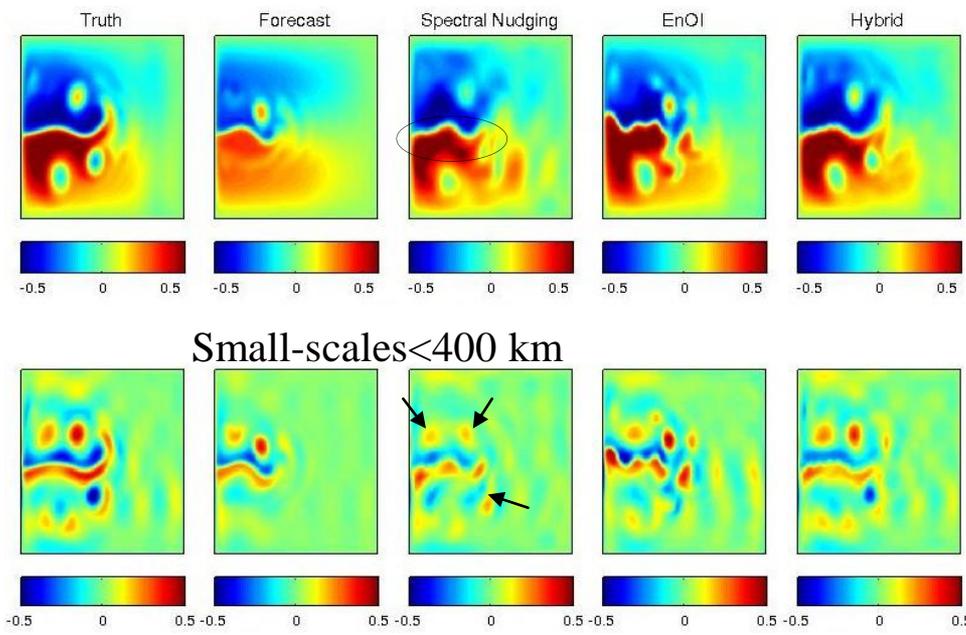


Description	Run	Period	Initial Condition	Assimilation
Spinup	SP	several decades	Rest	No
Forecast	FC	1000 days	End of run SP	No
Ensemble	EN	18 years	End of run FC	No
Truth	TR	1000 days	End of run EN	No
Spectral nudging	A1	1000 days	End of run SP	Large-scales from run TR
EnOI	A2	1000 days	End of run SP	Point obs from run TR
Hybrid	A3	1000 days	End of run SP	Large-scales and point obs

- Several twin experiments performed
- **Large scales** and **pseudo-observations** are generated by the model using different initial conditions

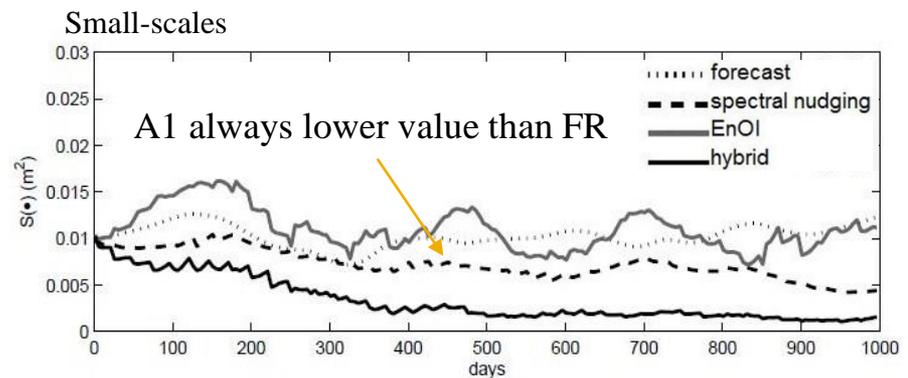
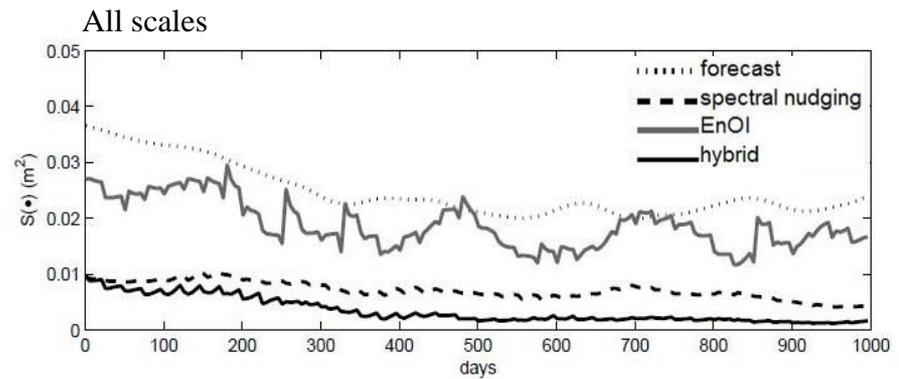


# QG Model Results

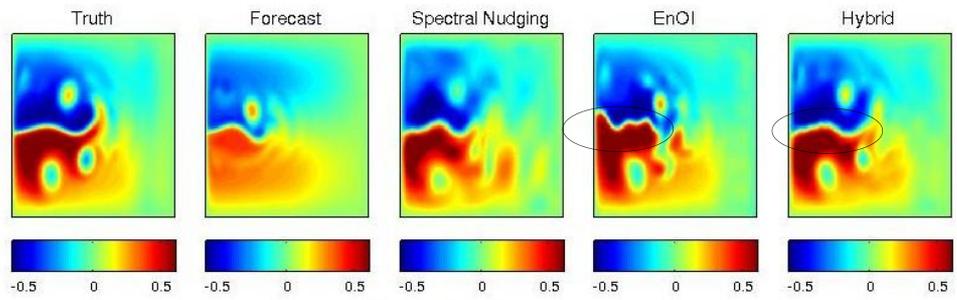


Spectral nudging removed unrealistic jet meanders and placed the **eddies** at the **correct location** (however eddies too weak)

Spectral nudging also **reduced the small scale error**

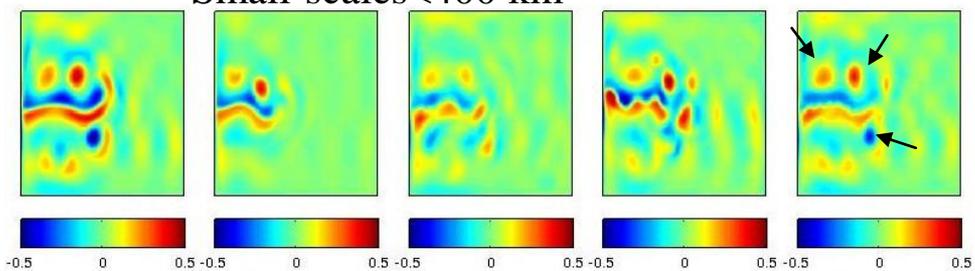


# QG Model Results



- **EnOI** : unrealistic jet and eddies position
- **Hybrid** : eddies and meanders have the correct location and intensity

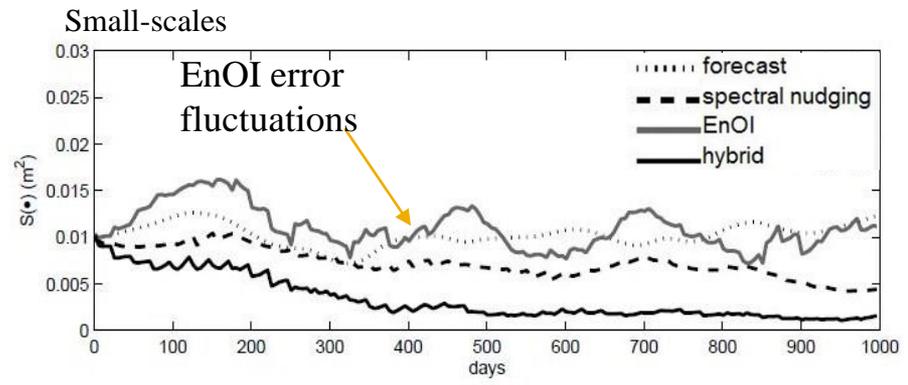
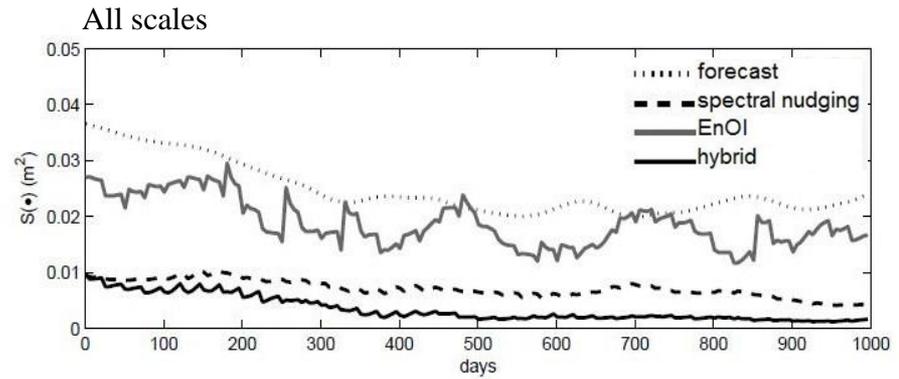
Small-scales < 400 km



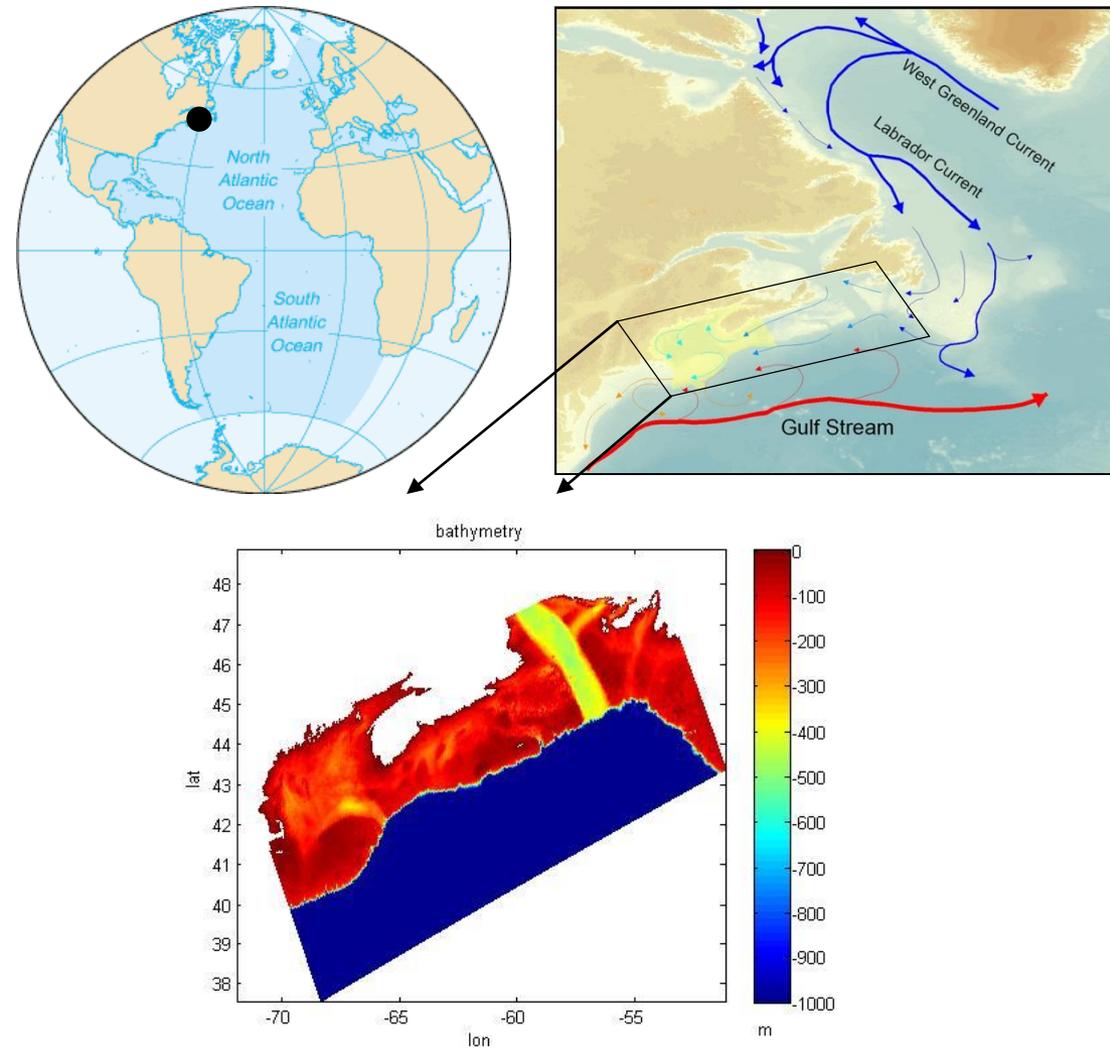
• **EnOI error fluctuation**: error inflated when the observation array cannot detect the eddies

• **Hybrid outperforms** spectral nudging and EnOI.

• Limited observations have **large impact** when used in combination with spectral nudging

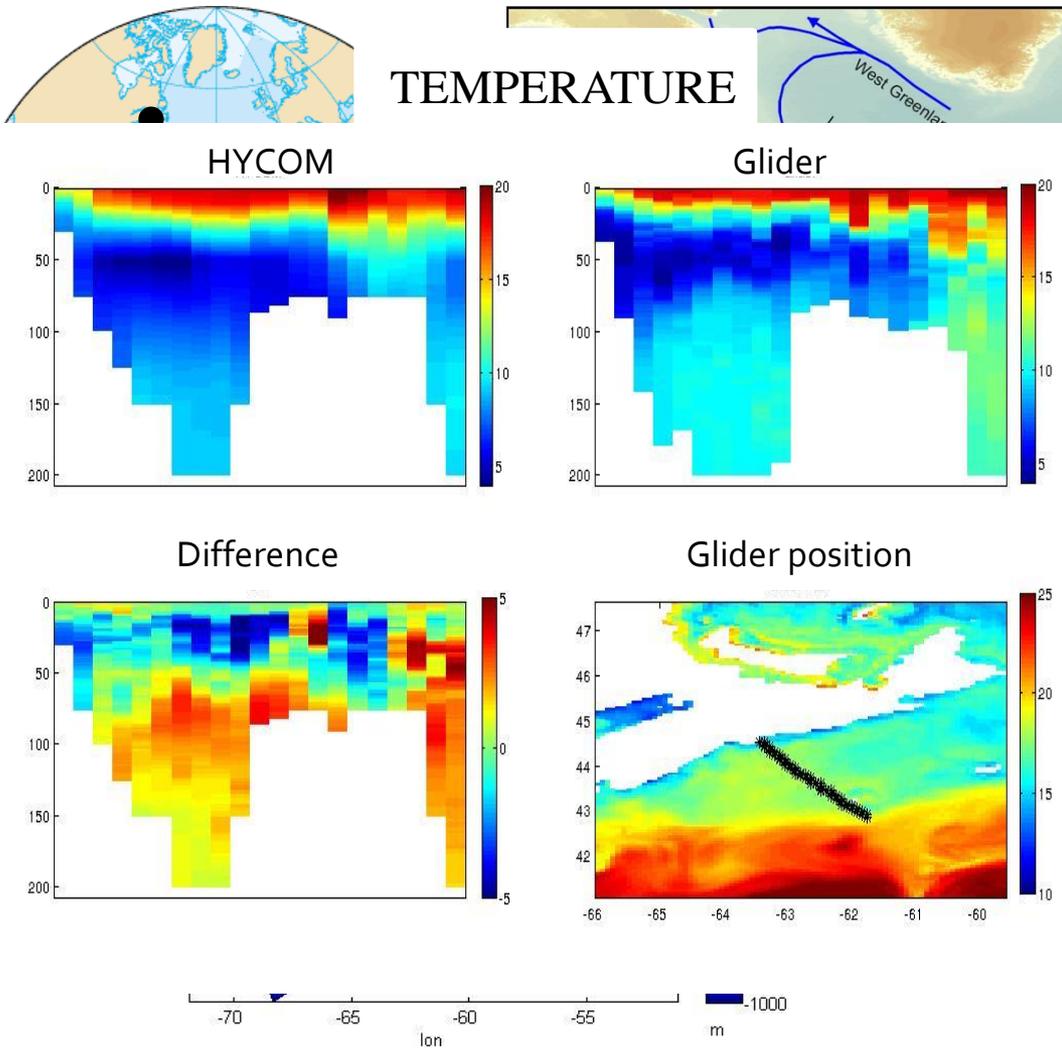


# Realistic Application



- Complex mean circulation and very high tides
- NEMO 3.1 with resolution  $1/36^\circ$  and 52 vertical levels
- Atmospheric forcing: **NCEP Climate Forecast System Reanalysis (CFSR)**
- Tides (M2,S2,N2,K1,O2) from LEGOS (Laboratoire d'Etudes en Géophysique et Océanographie Spatiales)
- Initial conditions and open lateral boundaries: **HYCOM** (HYbrid Coordinate Ocean Model) + **NCODA** (Navy Coupled Ocean Data Assimilation) global  $1/12^\circ$  analysis product

# Realistic Application



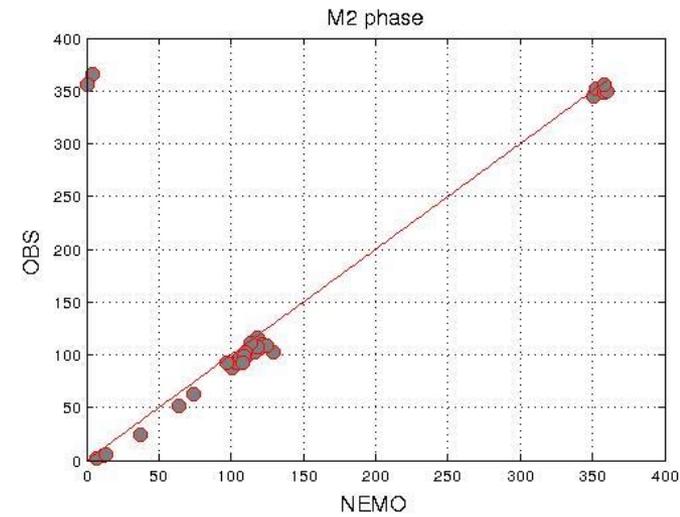
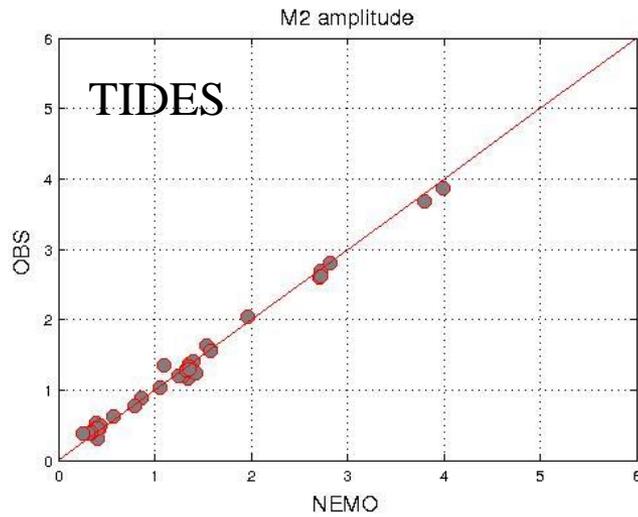
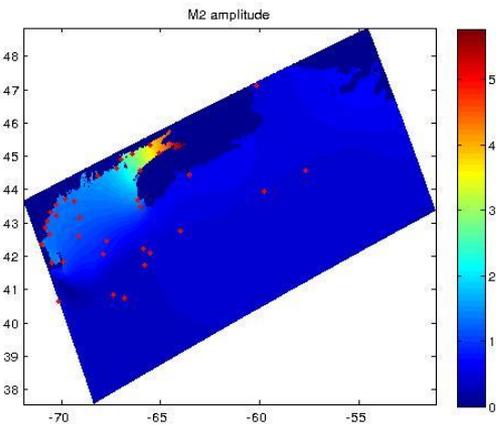
- Complex mean circulation and very high tides

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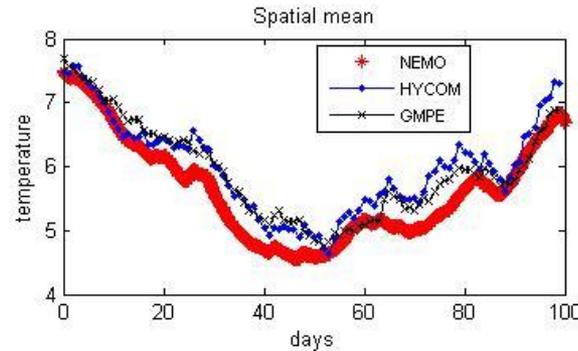
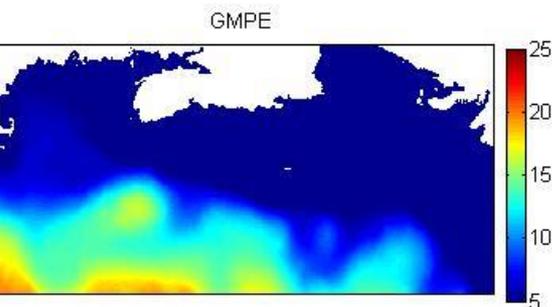
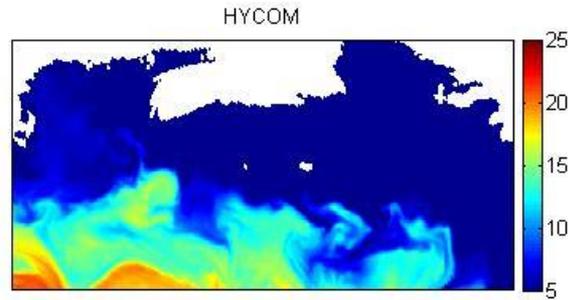
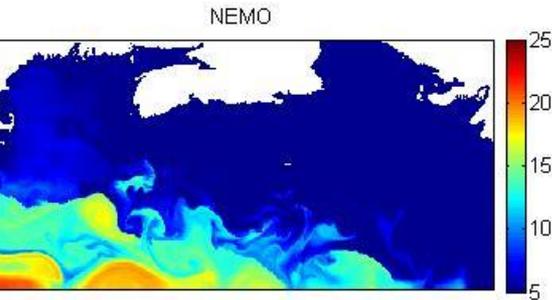
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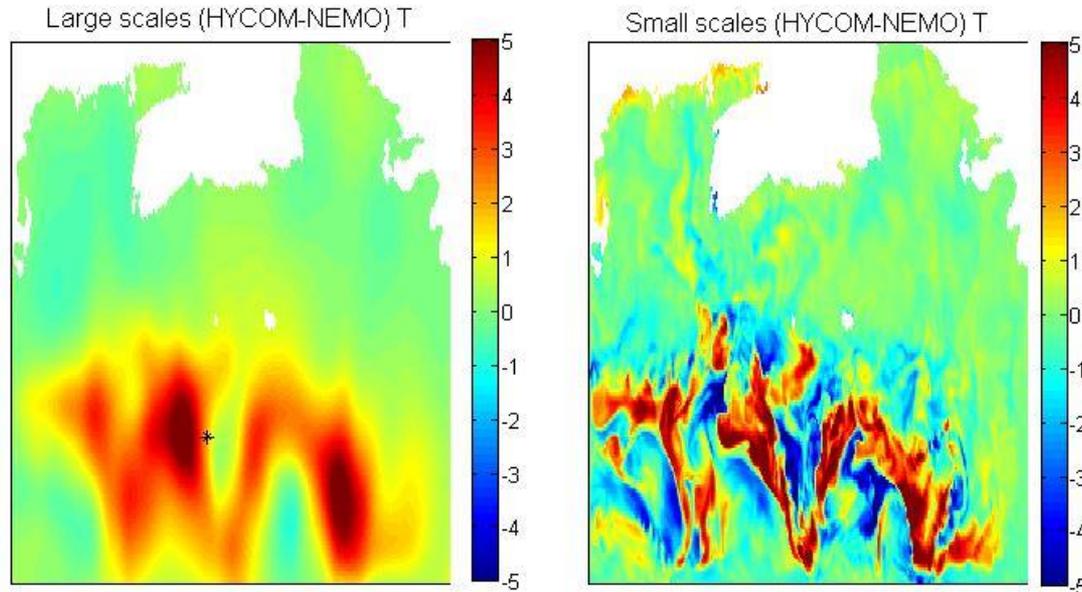
## TEMPERATURE



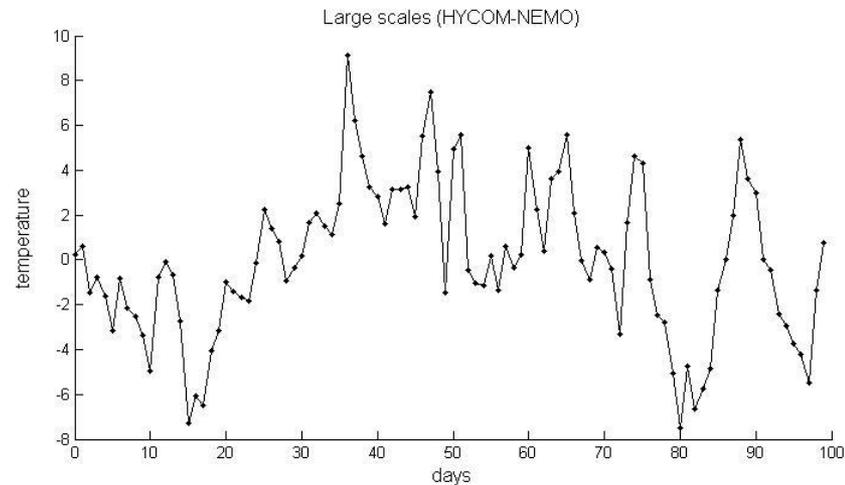
- Realistic **tides** and **general circulation**.
- Divergence of HYCOM and NEMO solutions.
- We want the model to:
  - evolve **freely** for **length scales not resolved** by HYCOM
  - stay close to HYCOM on the **large scales**.

# Large and small scales are defined using a **Butterworth filter**

TEMPERATURE



Cut off lengthscale ~ 500 km



Nudging coefficient  $\gamma$  **varying** in horizontal and vertical:

- Near the **coast** more model freedom.
- Near the **surface** stronger relaxation.

# Summary

- The time history of ocean large scales can guide the reconstruction of the small scales
- Assimilation of sparse observations more effective with spectral nudging of large scales
- High resolution nested NEMO diverges from global HYCOM on both large and small scales
- Optimistic that constraining NEMO's large scales will improve its prediction of the small scales

Thanks: Ocean Tracking Network (OTN), Marine Environmental Observation Prediction and Response Network (MEOPAR), Dr. Y. Lu, Dr. J-P. Paucin, Dr. V. Korabel

# References

Henshaw, W., Kreiss, H.-O., Ystrom, J., 2003. Numerical experiments on the interaction between the larger- and the small-scale motion of the Navier–Stokes equations. *Multiscale Modeling and Simulation* 1, 119–149.

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Martin, Matthew, et al. "Group for High Resolution Sea Surface temperature (GHRSSST) analysis fields inter-comparisons, 2012. Part 1: A GHRSSST multi-product ensemble (GMPE)." *Deep Sea Research Part II: Topical Studies in Oceanography* 77, 21-30.

# Definition and Variability of Scales

**Discrete Fourier transform:**

$$a_{k_1, k_2} = \frac{1}{N_1 N_2} \sum_{n_1=0}^{N_1-1} \sum_{n_2=0}^{N_2-1} e^{-i\varphi} \eta_{n_1, n_2}$$

$$\varphi = 2\pi \left( \frac{k_1 n_1}{N_1} + \frac{k_2 n_2}{N_2} \right)$$

**Large scales:**

$$\eta_{n_1, n_2}^L = \sum_{(k_1, k_2) \in K^c} e^{i\varphi} a_{k_1, k_2}$$

**Small scales:**

$$\eta^S = \eta - \eta^L$$

Most of the variance is captured from the first 13 wavenumbers in the zonal and 12 in the meridional ( $k^c = 13, 130$  km)

