

The Caribbean conundrum of Holocene sea level ¹ School of Earth & Environment, University of Leeds, Leeds LS2 9JT National Oceanography Centre Luke Jackson^{1,2} and Jon Mound¹ ² National Oceanography Centre, 6 Brownlow Street, Liverpool L3 5DA

Introduction

Modern sea-level rise is a hotly debated subject and is of pressing societal concern. A fuller understanding of prehistoric sea-level rise will add valuable insight to this discussion. We consider Holocene sea-level change in the Caribbean guided by three questions about data treatment, resolution and model representation. First though, the data.

Two Data sets: One Past, One Present

There are two main types of palaeo-sea-level indicator in the Caribbean: mangrove peat and fossil coral. Our 'past' catalogue of 565 published sea-level indicators has wide spatial coverage and most sub-regional sub-sets extend in time across the Holocene (12000 cal yr BP - present).



Our 'present' catalogue contains published growth ratedepth and abundance-depth information of 15 modern coral species. Sampling is discontinuous, so we calculate a normalised median depth (z_{med}) and an optimum range (σ).



Can we improve Holocene sea-level reconstructions using existing data?

The habitable range of a certain species can be a proxy for Yes, the curves show distinct spatio-temporal variability. For sea-level position. Previously, this range hasn't included example, relative sea level was up to 10 m higher in Florida natural variability. We use our 'present' catalogue to decompared to the Lesser Antilles around 9000 cal yr BP, yet scribe the probability of a mangrove or coral species occursea-level in the Lesser Antilles reached present day levels at ring at some position (z) relative to sea level. We use these least 500 years earlier than Florida, around 2000 cal yr BP. distributions for each indicator in our 'past' catalogue to make 1000 'sea-level points'.

Normal Distribution (Mangrove) $\rho^{-(z-\sigma)^2/2\sigma^2}$ $P_n =$ $\sigma\sqrt{2\pi}$



Mangroves lie within a narrower range around sea level. Tidal range is a strong control, hence Gaussian distribution.

Scaled Growth Distribution (Coral) $P_g = \tanh 8e^{-2z/z_{med}}$

Adapted from Chalker (1981), which relates growth rate to light intensity (i.e., depth). Good for corals that dominate breaker to fore-reef environments.

Logistical Distribution (Coral)

$$P_l = \frac{e^{z - z_{med}/\sigma}}{\sigma(1 + e^{z - z_{med}/\sigma})^2}$$

Designed to account for corals with preference down fore-reef slope.



largely explained by melt-In each sub-region there are water from the last deglacia-1000 'sea-level realisations'. For tion, loading the ocean basins each realisation, we derive pocausing subsidence, in accord $\frac{10}{2}$ sitions (•) within successive, with Milne & Peros (2013). overlapping, 3000 year time windows. We calculate the References Chalker, B. E. (1981), Simulating light-saturation curves for photosynthesis and -25 mean (•) and standard deviacalcification by reef-building corals, Mar Biol, 63, 135-141. Milne, G. A., and M. Peros (2013), Data-model comparison of Holocene sea-level change in the circum-Caribbean region, Global Planet Change, 107, 119-131. tion (- -) to give a probabilistic, Mitrovica, J. X., and G. A. Milne (2003), On post-glacial sea level: I. General Theory, Geophys J Int, 154 (2), 253-267. *relative* palaeo sea-level curve. Peltier, W. R. (2004), Global Glacial Isostacy and the Surface of the Ice-age Earth: The ICE-5G (VM2) Model and GRACE, Annu Rev Earth Pl Sc, 32, 111-149.

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Do the reconstructions show valid changes across the Caribbean?



What causes the sea-level rise?

We correct the reconstructed curves (---) for the difference between the optimum glacial isostatic adjustment model (Mitrovica & Milne, 2003; Peltier, 2004) derived sealevel curves (—) and the eustatic rise (- -).

Despite the variable tectonic setting in the Caribbean, the Holocene sea-level rise is

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Toscano, M. A., and I. G. Macintyre (2003), Corrected western Atlantic sea-level curve for the last 11,000 years based on calibrated 14C dates from Acropora palmata framework and intertidal mangrove peat, Coral Reefs, 22, 257-270.

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