





The baroclinic instability of an initially stratified fluid layer

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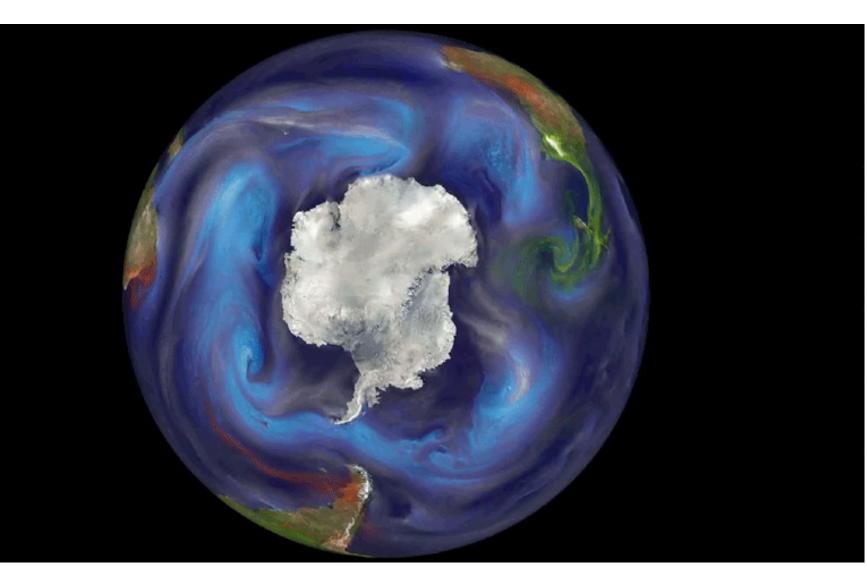
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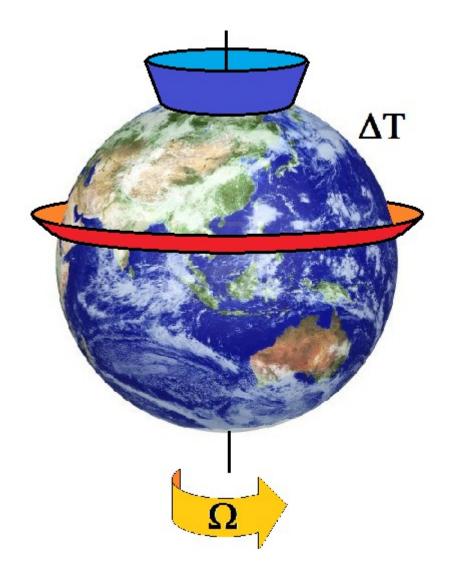
General motivation



Credit NASA: Aerosol transport patterns from the GOES-5 satellite data.

A minimal model of weather-like dynamics

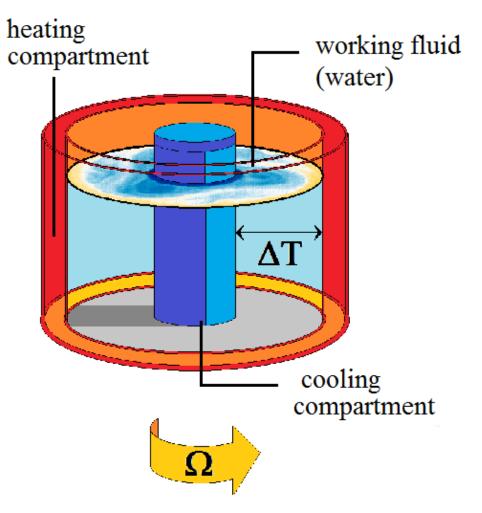
- A large variety of the typical atmospheric phenomena of the midlatitudes are primarily driven by two factors only.
- Rotation + meridional temperature difference ≈ weather





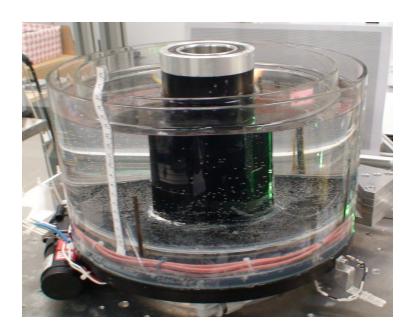
A minimal model of weather-like dynamics

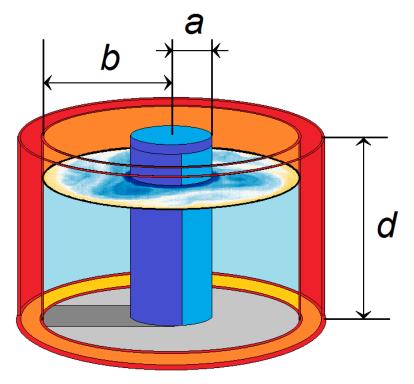
Laboratory experiments On the baroclinic instability





CoGeoF1 plateform Brandenburg University of Technology (BTU) Cottbus



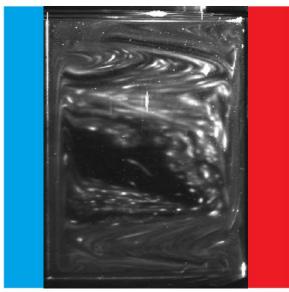


Geometric parameters:

- a = 45 mm
- b = 120 mm
- d = 135 mm



Basics: baroclinic instability



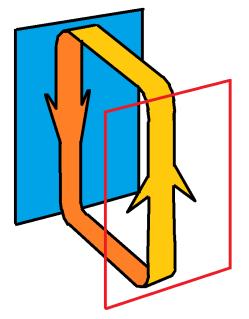
cooling

heating

"Sideways convection" – no threshold in ΔT (i.e. No 'critical Rayleigh number')

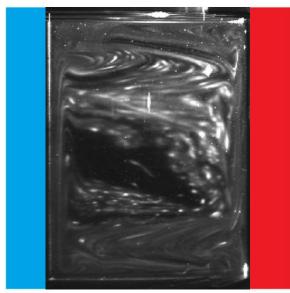
Any temperature difference can initia the flow

No Rotation





Baroclinic instability



cooling

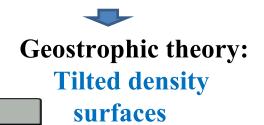
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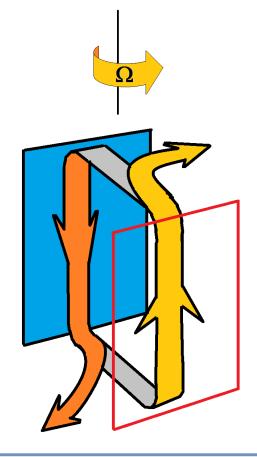
CC

heating



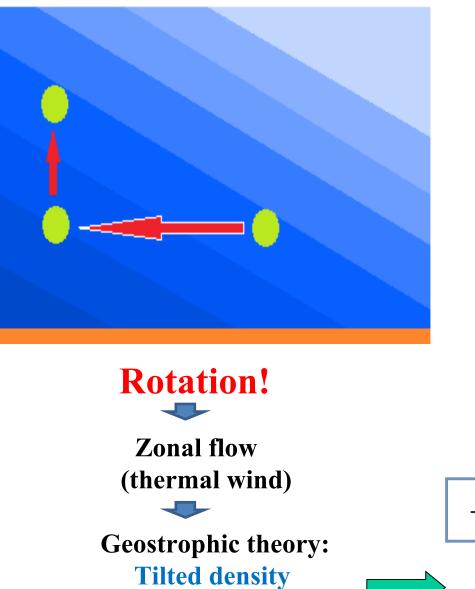
(thermal wind)





 $-2\Omega \vec{e_z} \times \vec{u}$

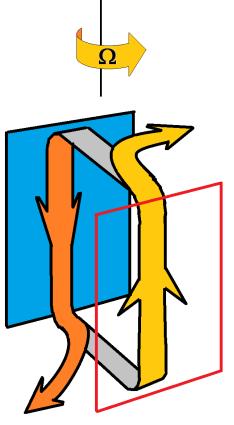
Baroclinic instability

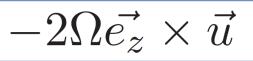


surfaces

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CC







Experimental baroclinic instability set-up mimics:
Convective cells, jets and zonal winds,
Rossby waves of planetary atmospheres

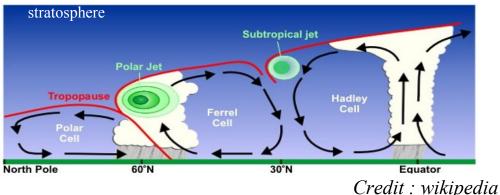
- control parameters:

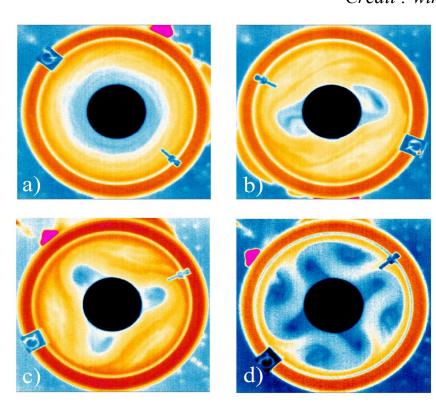
rotation rate, radial temperature difference

- Different planetary atmospheres can be modelled

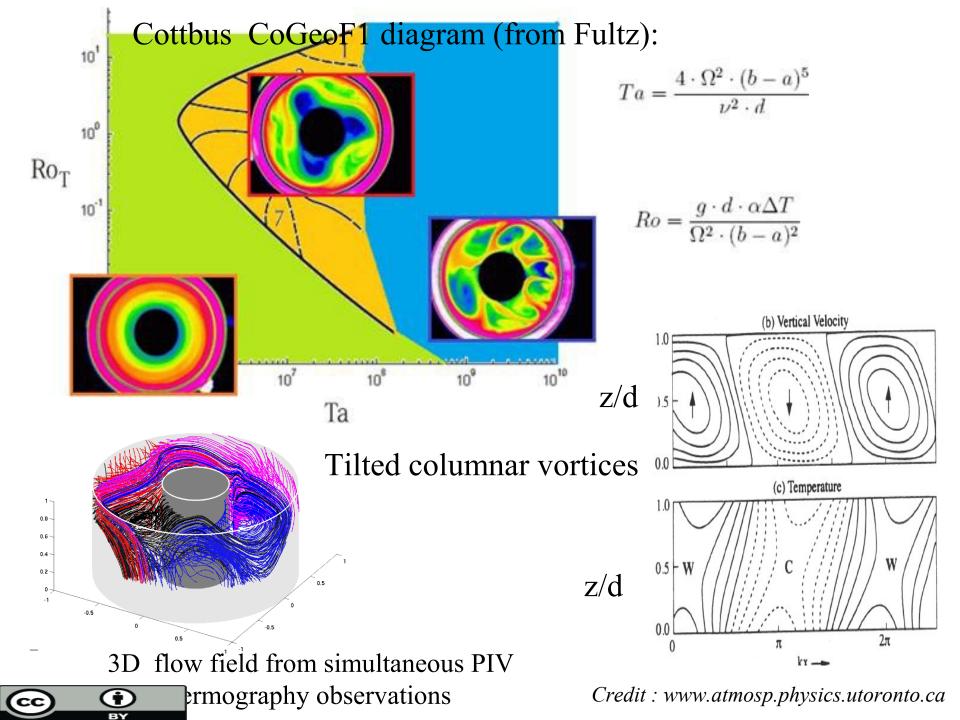
<u>Venus:</u> slow rotation, zonal flow

<u>Earth:</u> fast rotation \rightarrow Coriolis effect \rightarrow vortices ("weather")









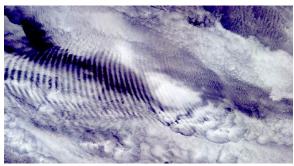
Can we add a stratosphere ?

- **D** To get the « right » horizontal stress free Boundary Condition
- To generate « pancake » vortices in rotating and stratied flows : Jupiter Great Red Spot

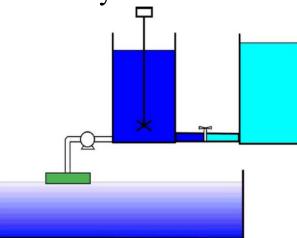


Credit:NASA

• To excite Gravity Waves from eventual unbalanced dynamics



Credit:http://www.planetpals.com





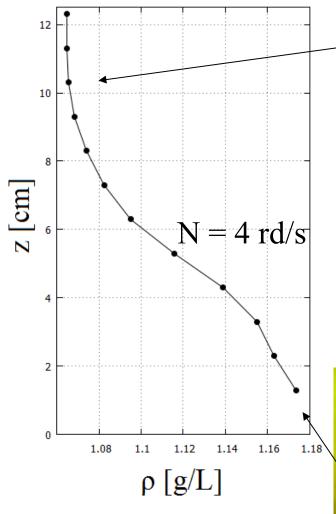
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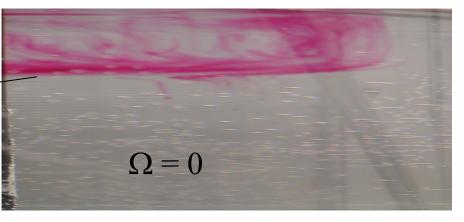
The baroclinic instability of an initially stratified fluid layer :

The BAROSTRAT Instability

Salinity stratification: "Double-bucket" method

Stratification will oppose to convection !



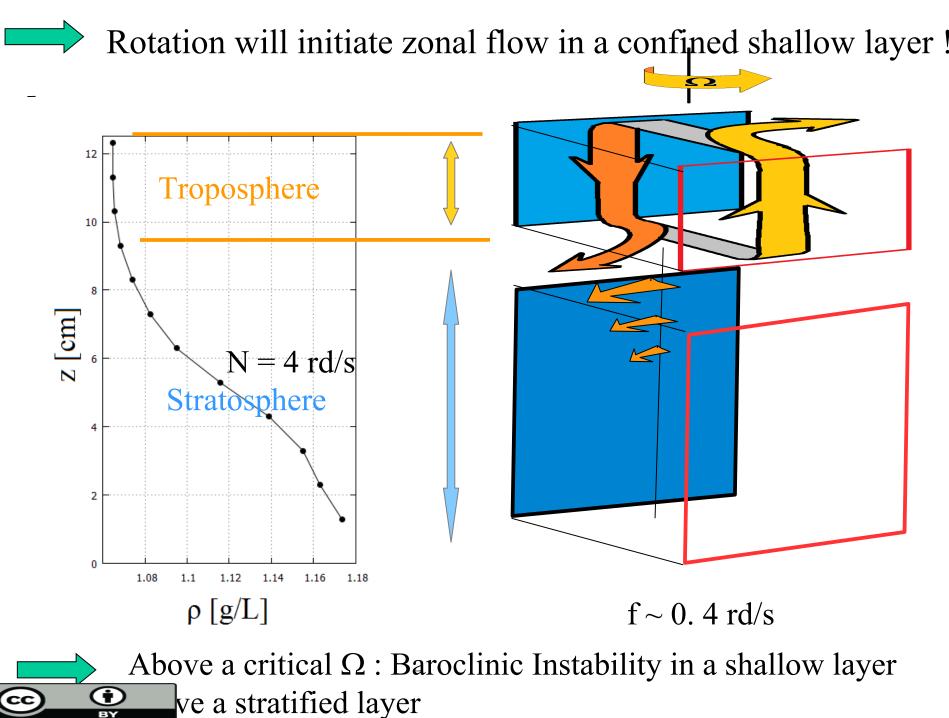


Above a temperature threshold, convection confined in weakly stratified shallow layers (Lee, Kang, Son 1999)





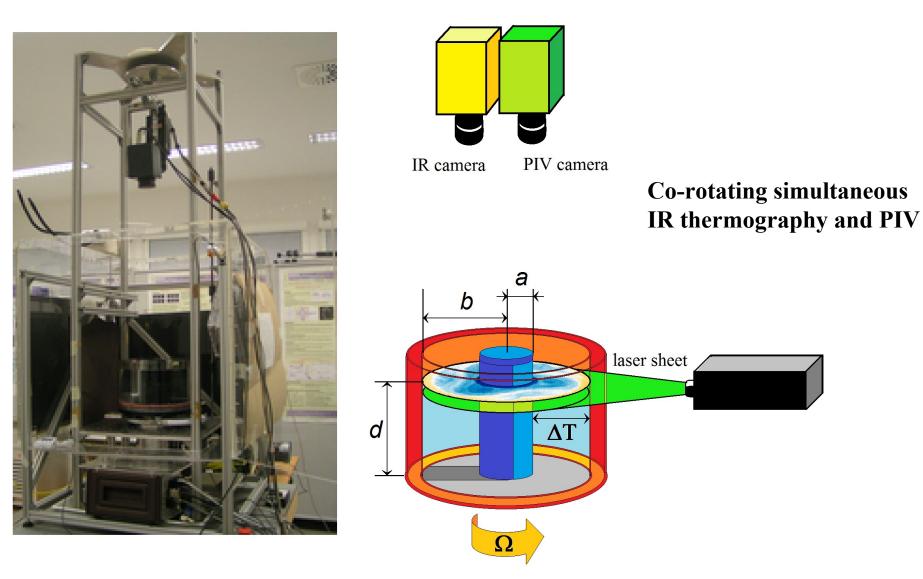
Double diffusive staircase cells at bottom



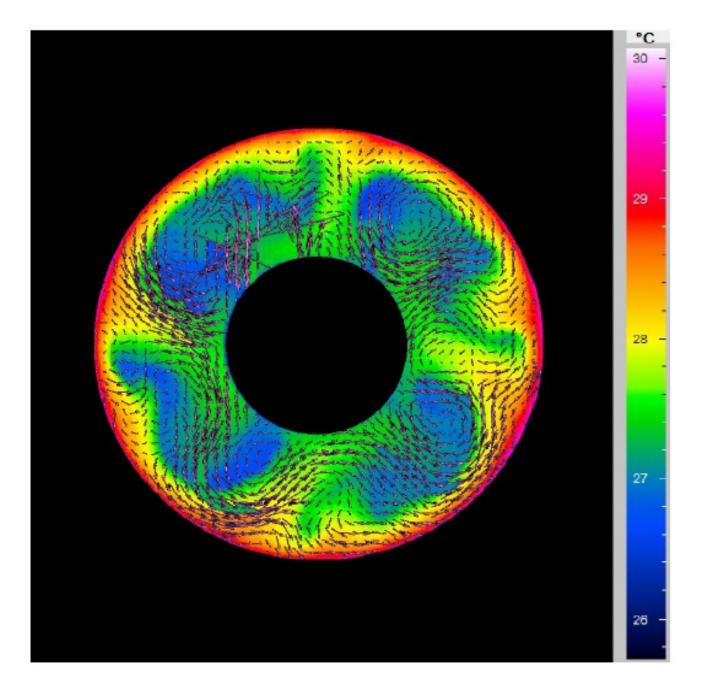
ve a stratified layer

(cc)

Measurement techniques







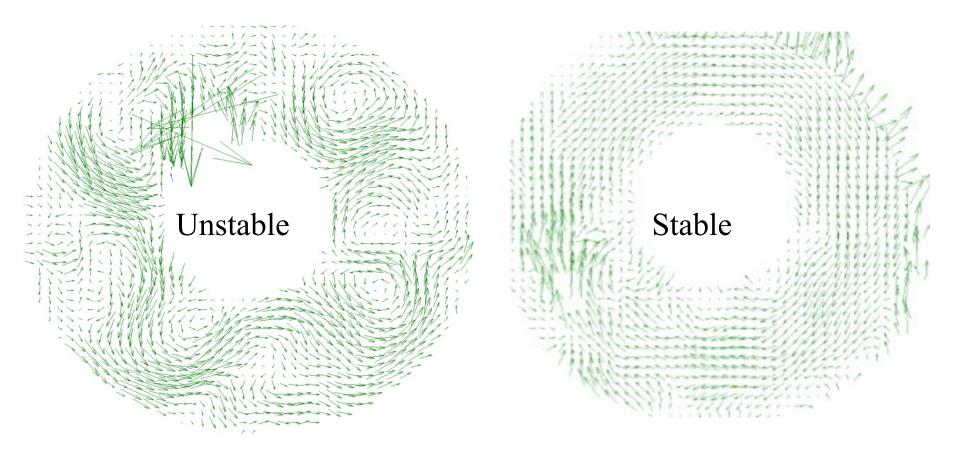
$$\Delta T = 6 \text{ K}$$

 $\Omega = 2.7 \text{ rpm}$



PIV field near surface

PIV field near bottom



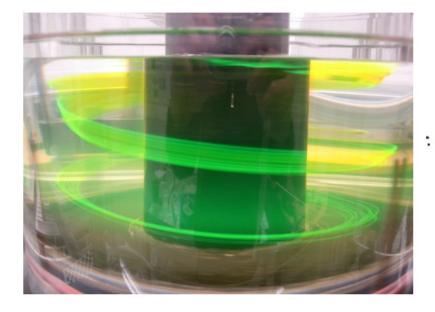
Zonal flow only



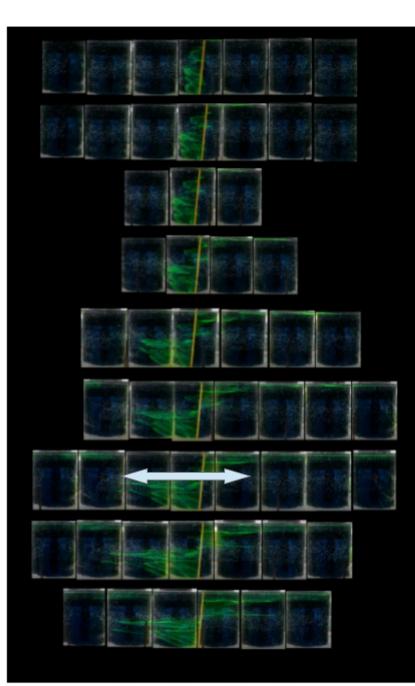
t=0



Fluoresceine visualization of thermal wind

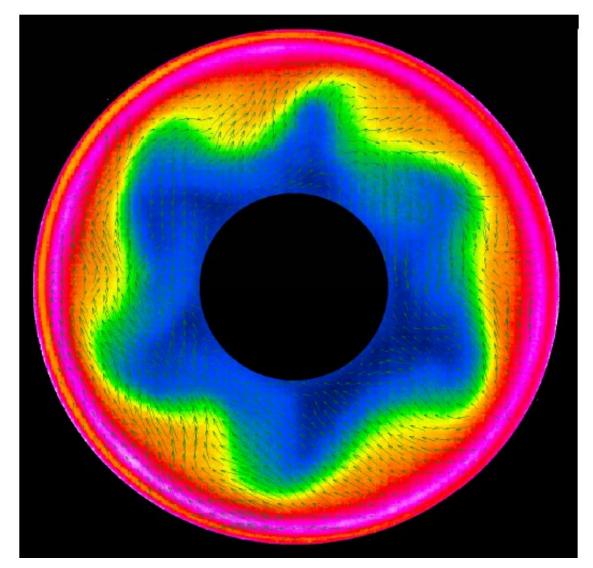


Note the stretching of dye filaments between top and bottom of the convective cell



azimuthal direction

Sucking top mixed layer : decreases its thickness and changes azimuthal wave number



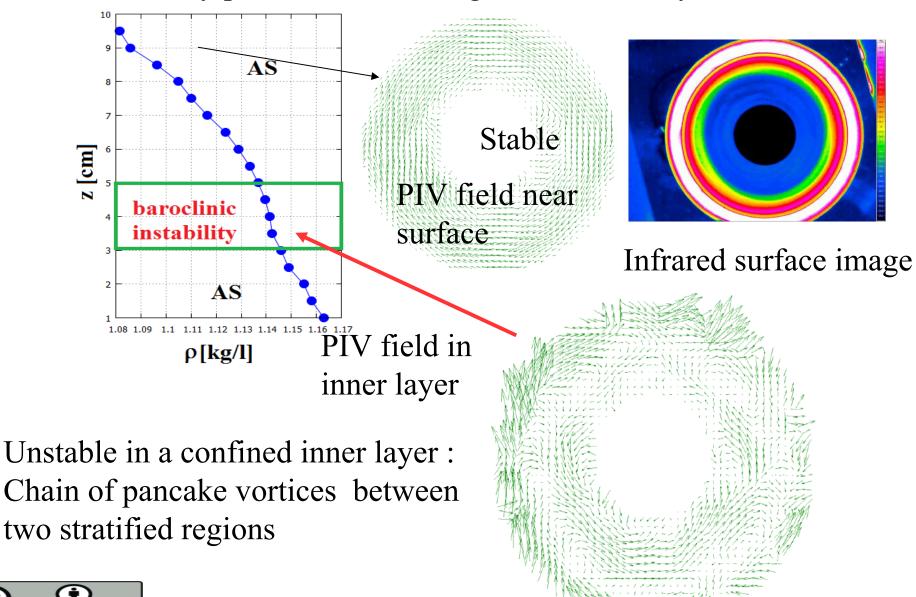
 $\Delta T = 6 \text{ K}$ $\Omega = 1.7 \text{ rpm}$

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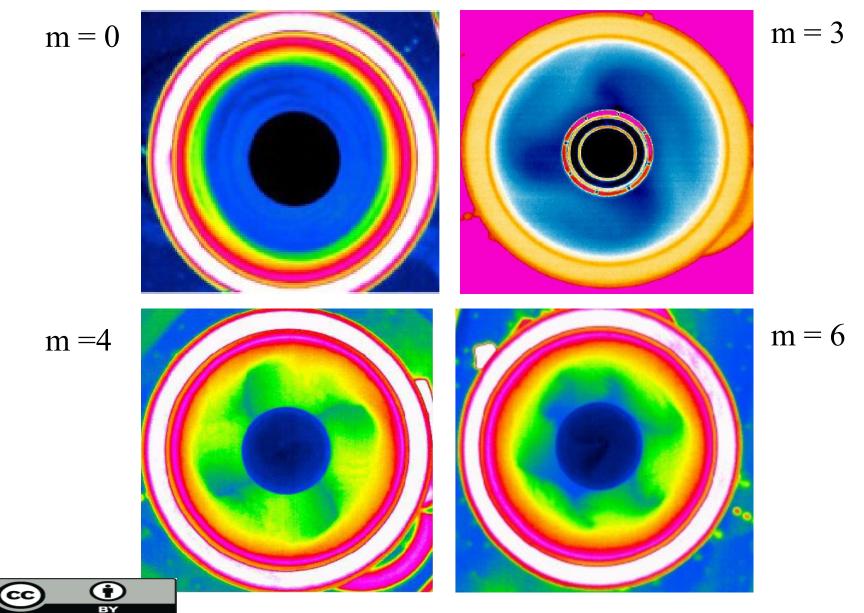
CC

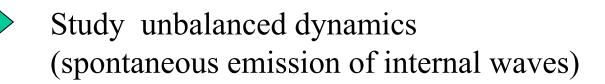
mbined thermographic/PIV image (surface temperature + PIV data from h = 12 cm)

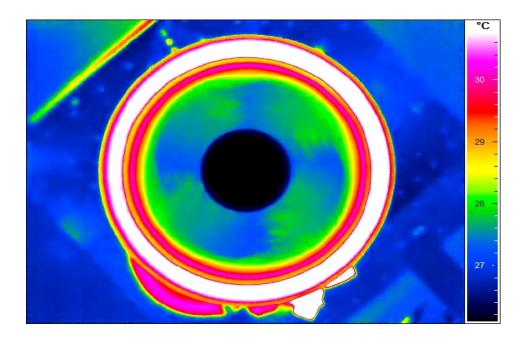
Sucking out completely top mixed layer on a vertical Density profile with a zero gradient inner layer



Zoology of different patterns of the « Barostrat intability »







$\Delta T \approx 6.5 \text{ K}$ $\Omega = 2.5 \text{ rpm}$

Waves have been observed in a 2 immiscible fluid layers experiment (Williams, Haine, Read, 2008)

