

Methane production and oxidation patterns along a hydrological gradient in Luther Bog, Ontario

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Introduction

Methane emissions from natural peatlands contribute significantly to the global budget of atmospheric CH₄ (Mikaloff Fletcher et al. 2004). In the northern hemisphere, where climate models predict rising temperatures and precipitation rates, these emissions are likely to rise (IPCC 2013). So far, little is known about the change of processes of methane production and oxidation, which influence the total amount of methane emissions, in peatland soils under warmer and wetter climate conditions. This study focuses on anaerobic CH₄ production and aerobic CH₄ oxidation processes along a hydrological gradient in an ombrotrophic bog in Ontario, Canada. We aim to quantify the amount of (a) anaerobic CH₄ and CO₂ production, (b) aerobic CH₄ oxidation and (c) their contribution to the total ecosystem respiration flux.

Study area and Methods

Luther Bog is located in the Luther Marsh Conservation Area in Southern Ontario (43.9° N, 80.4° W). We established four sites along a hydrological gradient differing in water table levels and dynamics and vegetation patterns, whereas water table fluctuations increase from site 1 to site 4 due to a reservoir at the northern end of the bog. From May to September 2015, we examined depth profiles of CO₂ and CH₄ concentrations and



Fig. 1: View of Site 1 in May

delta ¹³C isotope ratios in the peat using silicon samplers, dialysis chambers and multi-level piezometers. In addition to that, we determined carbon fluxes with chamber flux measurements using a Los Gatos Ultraportable Greenhouse Gas Analyzer.

Results and Discussion

Both CH₄ and CO₂ fluxes are overall higher at sites 3 and 4 than at the reference sites 1 and 2. CO₂ fluxes are increasing over the summer except from site 2, where they are decreasing (see Fig. 2). DIC concentration in the peat is rising over time whereas CH₄ concentrations are more stable. Altogether, DIC concentrations are lowest at site 4 and CH₄ concentrations are highest at site 4 (see Fig. 4). δ ¹³C signatures of CH₄ increase with time at all sites, δ ¹³C values for CO₂ do not show such patterns (see Fig. 3).

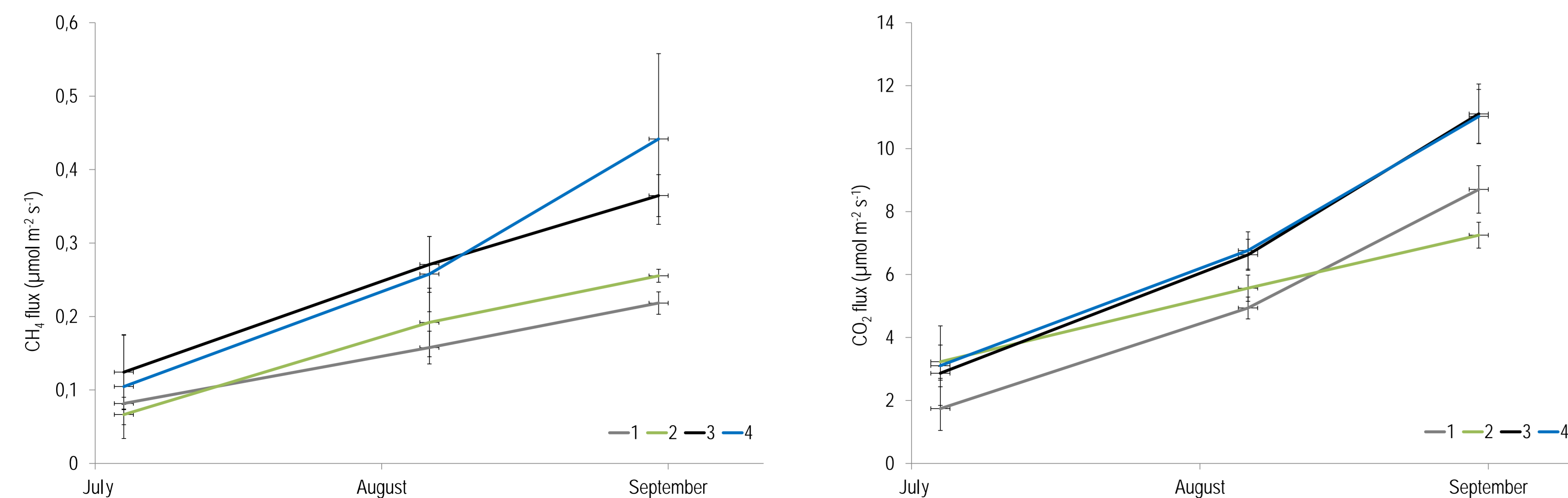


Fig. 2: Cumulated CH₄ and CO₂ fluxes at sites 1-4 from July to September.

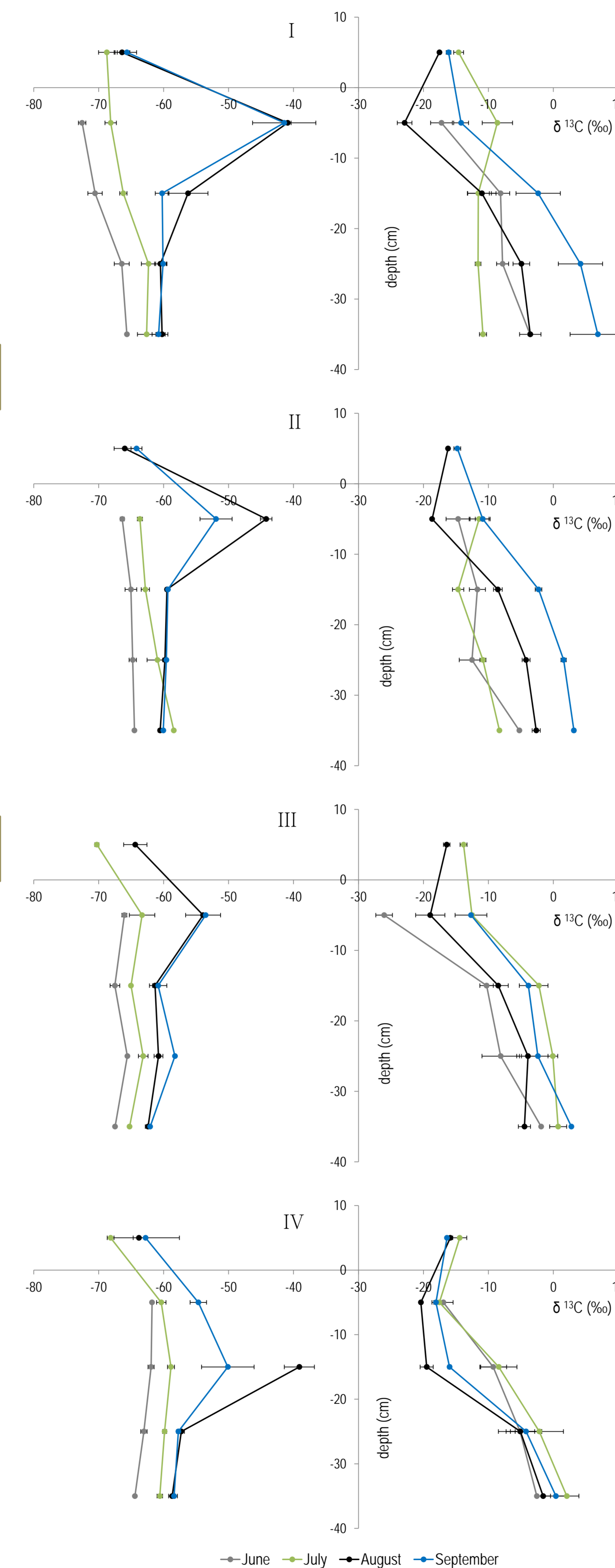


Fig. 3: δ ¹³C signatures of CH₄ (left) and CO₂ (right) in the peat and from chamber measurements at sites 1-4.

Both DIC and CH₄ concentrations in the upper soil level are reflecting water table dynamics. In case of low water tables, the emerging aerobic conditions lead to CH₄ oxidation in the upper soil, which is also reflected in the δ ¹³CH₄ values (see Fig. 3). Moreover, at all sites, the CH₄ production pathway changes with depth from aceticlastic to hydrogenotrophic methanogenesis (cf. Whiticar et al. 1986). This shift arises because with increasing depth, the organic matter becomes more hardly decomposable (Miyajima et al. 1997).

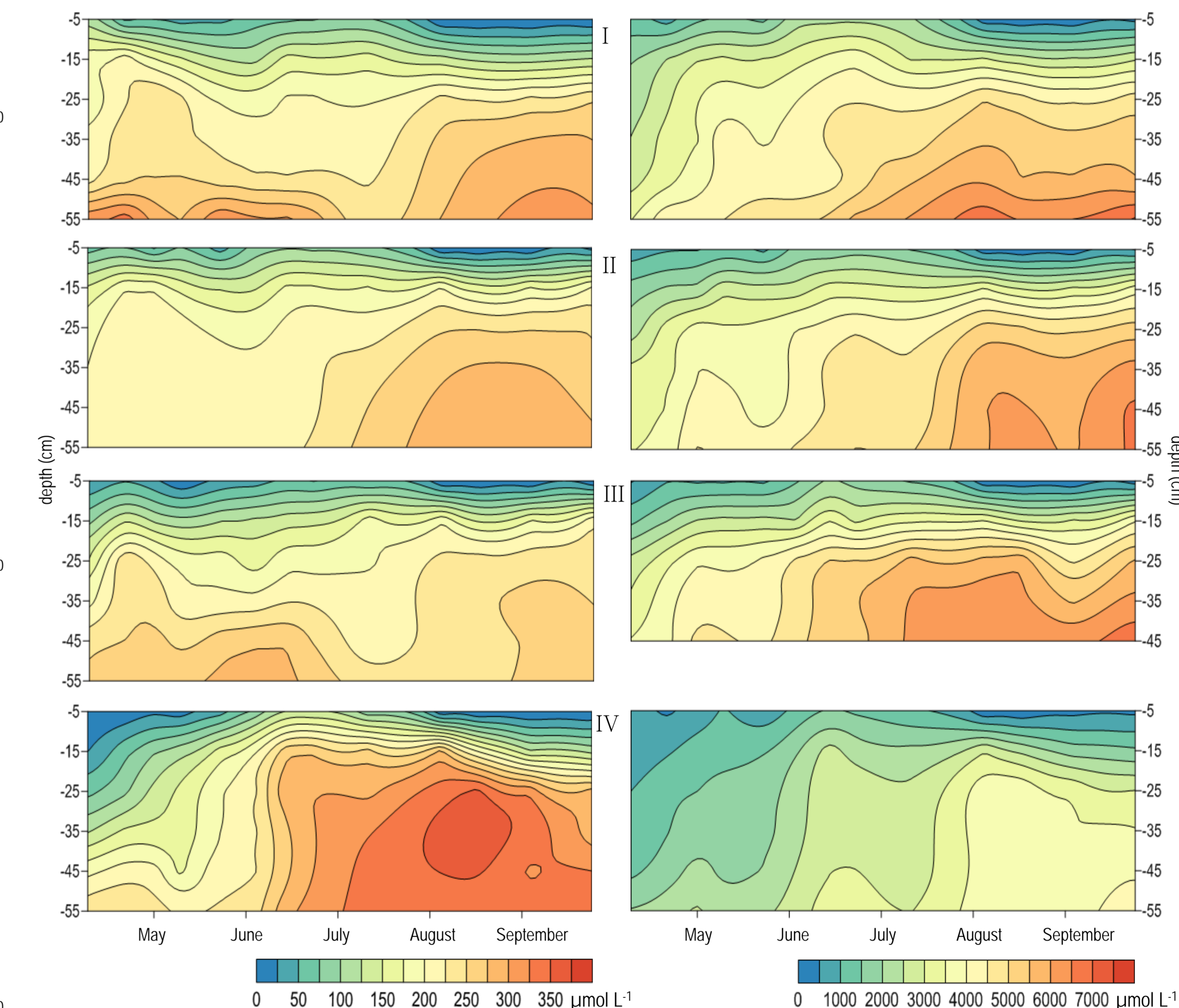


Fig. 4: Concentration profiles of CH₄ (left) and DIC (right) in the peat at sites 1-4.

Conclusion and Outlook

Water table dynamics seem to have an impact on methane and DIC dynamics and differ in-between the four sites. To quantify turnover rates, next steps to accomplish are calculating isotope mass balances as suggested in Landsdown et al. (1992) and estimating gas transport in the soil by Fickian diffusion.