INTRODUCTION

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The main objective of our research is a comparison and validation of different models from satellite missions CHAMP, GRACE and GOCE with ground-based gravimetric measurements. Our analyses concern **Bouguer gravity anomalies**. As a reference dataset for validation of the models, we used ground-based gravity measurements for points located in Poland. These stations are divided into three groups: 1) points of POLish REFerence network POLREF (with relative gravity measurements), 2) points of Polish National Gravimetric Network POGK (with absolute gravity measurements) and 3) points located in Tatra Mountains (with relative gravity measurements). First two groups contain stations from the whole area of the country while the last one covers only a small highlands region.

In this study we compute the Bouguer's anomalies for every point in two ways: 1) with the use of different GGMs – we call them calculated anomalies and 2) with the use of ground-based gravimetric observations - we call them measured anomalies. Next we determine the differences between calculated and measured values as well as their basic statistics (minimum, maximum, mean, standard deviation). The evaluation of the accuracy of each model is based on the analysis of these differences. The use of CHAMP, GRACE and GOCE-based GGMs for modelling the gravity field over Poland is discussed here.

DATA USED

1) CHAMP, GRACE and GOCE-based GGMs – with different N_{max}, spatial resolution and input data

Mission(s)	Model	Max degree (N _{max})	Year	Type of			
CHAMP	1. ULUX_CHAMP2013s	120	2013	S (CHAMP)			
	2. EIGEN-CHAMP05s	150	2010	S (CHAMP)			
GRACE	3. ITG_GRACE2010s	180	2010	S (GRACE)			
	4. GGM05s	180	2014	S (GRACE)			
	5. ITSG_GRACE2014s	200	2014	S (GRACE)			
	6. EIGEN_GL04c	360	2006	S (GRACE, SLR) + G + A			
	7. GGM03c	360	2009	S (GRACE) + G + A			
	8. EGM2008	2190	2008	S (GRACE) + G + A			
GOCE	9. DIR-R2	240	2011	S (GOCE)			
	10. TIM-R5	280	2014	S (GOCE)			
	11. SPW-R4	280	2014	S (GOCE)			
	12. SPW-R5	330	2017	S (GOCE)			
	13. JYY_GOCE04s	230	2014	S (GOCE)			
	14. GECO	2190	2015	S (GOCE) + EGM2008			
GOCE + GRACE	15. EIGEN_6s	240	2011	S (GOCE, GRACE, SLR			
	16. EIGEN_6c	1420	2011	S (GOCE, GRACE, SLR			
	17. GOGRA04s	230	2014	S (GOCE, GRACE)			
	18. GGM05c	360	2016	S (GOCE, GRACE) + G			
CHAMP + GRACE + GOCE	19. GOCO02s	250	2011	S (GOCE, GRACE, CHA			
	20. GOCO05s	280	2015	S (GOCE, GRACE, CHA			
	21. GOCO05c	720	2016	S (GOCE, GRACE, CHA			
	22. XGM2016	719	2017	S (GOCE, GRACE, CHA			

Table 1. Global Gravity Field Models (GGMs) used for analysis

*S – satellite data, G – ground-based gravimetric measurements, A – altimetry measurements, K – kinematic orbits of satellites: Swarm A+B+C, TerraSarX, Tandem-X, CHAMP, GRACE A+B, GOCE

2) Ground-based gravity measurments for points in Poland - POLREF, POGK and TATRA



• POLish REFerence network (POLREF) - 353 points

- (POGK) 167 points
- 479 points

Total – 999 points

METHODOLOGY

Bouguer anomalies have been obtained from GGMs as follows (Barthelmes, 2013):

n _{max} p n	n	n _{max} n
$\Delta g_B(r,\varphi,\lambda) = \frac{GM}{r^2} \sum_{n=0}^{\infty} \left(\frac{R}{r}\right)^n (n-1)$	$\sum_{m=0} \bar{P}_{nm}(\sin\varphi)(\bar{C}_{nm}^T \cos m\lambda + \bar{S}_{nm}^T \sin m\lambda)$	$-2\pi G\rho R \sum_{n=0} \sum_{m=0} \bar{P}_{nm} (sin\varphi) (\bar{C}_{nm}^{top})$

where: $\bar{C}_{nm}^T i \bar{S}_{nm}^T$ – fully normalized Stokes' coefficients of gravity disturbance, $\bar{C}_{nm}^{topo} i \bar{S}_{nm}^{topo}$ – fully-normalized coefficients of topography model, P_{nm} – fully-normalized Legendre functions

Bouguer anomalies from ground-based measurements have been computed using formula: $\Delta g_B = g - \gamma_0 + R_B + R_{fa}$, where $R_B = 2\pi G \sigma H$ is a Bouguer correction and $R_{fa} = 0.3086 \cdot H$ s a free-air correction

REFERENCES

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Validation of static global gravity field models from CHAMP, GRACE and GOCE with ground data in Poland Justyna Śliwińska and Jolanta Nastula

Contact e-mail: jsliwinska@cbk.waw.pl



Polish National Gravimetric Network

Points in Tatra Mountains (TATRA) –

 $\sum_{n=1}^{\infty} \cos m\lambda + \bar{S}_{nm}^{topo} \sin m\lambda$



Fig. 2. Example of Bouguer's anomalies distribution over area of Poland - obtained from ground-based gravity measurements and three GGMs with different N_{max} (ULUX_CHAMP2013s, JYY_GOCE04s and EGM2008). The Bouguer anomalies are given in mGal.





Fig. 3. Example of Bouguer anomalies distribution for area of Tatra Mountains - obtained from ground-based gravity measurements and three GGMs with different N_{max} (ULUX_CHAMP2013s, JYY_GOCE04s and EGM2008). The Bouguer anomalies are given in mGal.

2) Differences between anomalies obtained from ground-based gravity measurements and calculated from **GGMs**

Table 2. Basic statistics of differences between calculated (from model) and measured
 (from ground-based gravity mesurements) Bouguer anomalies. The values are given in mGal.

NA - 11 - 1	min		max		mean			std				
WOdel	POGK	POLREF	TATRA	POGK	POLREF	TATRA	POGK	POLREF	TATRA	POGK	POLREF	TATRA
1. ULUX_CHAMP2013s	-59.93	-70.09	34.76	37.16	31.25	68.52	-10.88	-14.25	43.10	18.83	4.10	19.30
2. EIGEN-CHAMP05s	-45.67	-67.63	14.16	29.47	25.95	49.60	-8.44	-12.54	24.50	15.96	5.53	17.07
3. ITG_GRACE2010s	-38.87	-77.42	14.94	21.40	21.21	46.95	-9.54	-12.90	22.04	11.50	4.32	13.91
4. GGM05s	-46.34	-75.34	8.54	30.50	22.20	39.86	-9.94	-13.12	19.89	15.96	4.97	17.19
5. ITSG_GRACE2014s	-40.64	-75.23	0.76	24.42	21.91	31.70	-9.33	-12.36	11.92	14.41	5.10	16.14
6. EIGEN_GL04c	-30.17	-82.88	-15.51	14.11	8.20	15.83	-9.16	-12.60	-1.71	7.22	6.49	11.34
7. GGM03c	-30.12	-81.83	-13.55	7.82	12.43	18.20	-9.39	-12.85	-0.73	7.16	6.10	11.73
8. EGM2008	-24.51	-73.62	-8.49	10.85	5.48	39.22	-9.66	-13.34	1.56	4.38	8.23	10.25
9. DIR-R2	-38.76	-71.81	16.49	21.54	18.88	48.87	-9.04	-12.68	25.34	11.44	4.15	12.95
10. TIM-R5	-28.04	-68.95	-1.74	14.11	14.43	32.91	-9.31	-12.97	9.17	8.78	5.88	12.26
11. SPW-R4	-29.36	-69.65	-1.73	15.47	11.87	30.34	-9.31	-12.83	9.32	9.43	6.05	12.46
12. SPW-R5	-30.51	-64.92	2.33	18.82	14.85	33.03	-9.08	-12.95	14.00	10.24	6.07	12.61
13. JYY_GOCE04s	-32.93	-74.62	4.70	15.92	13.66	36.83	-9.07	-12.92	12.28	9.60	4.61	12.75
14. GECO	-24.60	-73.61	-8.62	10.61	4.69	38.68	-9.50	-13.21	1.67	4.38	8.08	10.07
15. EIGEN_6s	-38.43	-72.27	16.24	21.17	18.05	48.72	-8.96	-12.64	25.27	11.42	4.14	12.97
16. EIGEN_6c	-24.27	-80.87	-7.82	10.47	5.92	39.68	-9.57	-13.02	0.92	4.22	8.00	10.31
17. GOGRA04s	-33.11	-74.72	4.94	15.89	13.83	37.07	-9.08	-12.93	12.47	9.59	4.60	12.76
18. GGM05c	-28.97	-75.53	-4.45	12.46	10.32	28.19	-9.66	-13.25	3.74	6.80	4.88	11.29
19. GOCO02s	-33.97	-70.68	10.25	20.94	14.96	41.58	-9.40	-12.65	18.72	10.85	4.88	12.93
20. GOCO05s	-28.70	-70.11	-3.06	15.09	15.51	31.74	-9.34	-12.94	8.11	8.96	5.97	12.36
21. GOCO05c	-30.16	-71.34	-0.09	9.02	4.40	35.20	-9.95	-13.32	6.79	4.80	4.19	10.01
22. XGM2016	-31.35	-71.64	-1.40	9.21	5.15	31.89	-9.95	-13.21	5.67	4.74	4.28	10.12

3) Spatial distribution of differences between measured and calculated anomalies - Tatra Mountains





Fig. 4. Example of differences distribution for area of Tatra Mountains for four GGMs (ULUX_CHAMP2013s, JYY_GOCE04s, XGM2016 and EGM2008). The highest differences correspond to the biggest heights of the stations (southern part of the area).

ACKNOWLEDGEMENTS

This work was supported by the Polish national science foundation NCN under grant No. UMO-2013/11/B/ST10/04975 The authors would like to thanks Dr Tomasz Olszak for providing gravimetric data and for many valuable advices that improved this contribution.

Space Research Centre of the Polish Academy of Sciences, Warsaw, POLAND





5) Long wave-length components of gravity anomalies

It is obvious that the models with lower N_{max} give bigger differences of gravity anomalies than models that have higher degree of expansion. The first ones show only the long to medium wave-length components of Earth's gravity field while the second ones - also short wave-length changes. Consequently, for an objective comparison of all considered models, we computed Bouger anomalies using the same max degree/order ($N_{max} = 120$) for all GGMs.



DISCUSSION

In this study, the accuracy of GGMs determined from CHAMP, GRACE and GOCE has been estimated over the area of Poland. Bouger gravity anomalies estimated from several models with different spatial resolution and input data (Table 1) have been compared with the corresponding ones obtained from ground-based gravity measurements.

The results obtained show that the variance of modelled anomalies is bigger for the models with higher degree of expansion (Fig. 1). It is clear that higher N_{max} results in lower differences between computed and measured Bouguer anomalies. Obtaining models with high spatial resolution was possible after combining observations from satellite missions CHAMP, GRACE and GOCE with altimetric data and ground-based gravimetric measurements (Table 1).

The discrepancies between examined models are the biggest for the case of Tatra Mountains area (Fig. 2). The varied terrain elevation and geological structure of this region results in high variance of Bouguer anomalies the value of which depends e.g. on the height of the station. The satellites of gravity missions are not sensitive enough to observe ultrashortwave changes of Earth's gravity field. For this reason, standard deviations of differences between gravity functionals computed using GGMs and based on ground-based measurements are the biggest for this group of points for all considered models. This statistics is the lowest for the models: EGM2008, GECO, EIGEN 6c, GOCO05c and XGM2016 that are distinguished by the highest degree of expansion N_{max} (Table 2). The highest standard deviation of differences is observed for satellite-only models from CHAMP (ULUX CHAMP2013s and EIGEN-CHAMP05s) and GRACE (GGM05s, ITG GRACE2010s and ITSG GRACE2014s) missions. The residuals between measured and computed anomalies are bigger for the stations with higher elevations (Fig. 4).

The comparison of Bouguer anomalies for the points of POLREF and POGK networks shows that standard deviation of differences between observed and computed values are bigger for POLREF for most of combined models (satellite + ground-based measurements as input data) and lower for these stations for most of satellite-only models (only observations from satellite gravity missions) (Table 2, Fig. 5). It is worth mentioning that on the stations of POGK network the gravimetric measurements were made using absolute methods while for the case of POLREF network - with the use of relative measurements. The mean of differences is lower for POGK for all considered models (Table 2).

The maximum degree/order of the model (N_{max}) has dominant impact on the accuracy of GGMs (Fig. 5). The biggest increase in quality is observed up to the degree 300. However, further increase in N_{max} also results in decreasing standard deviation of differences between Bouguer anomalies computed using GGMs and obtained from ground-based gravity measurements. For the case of Tatra Mountains area, starting with $N_{max} = 700$, standard deviation of these residuals in-

There is no significand difference between long wave-length components of Bouguer gravity anomalies obtained from different Global Gravity Field Models. Only EIGEN-CHAMP05s, GGM05s and ITSG_GRACE2014s models have bigger standard deviation of differences between computed and measured anomalies for both POGK and TATRA points (Fig. 6).

(\mathbf{i}) (cc)

4) The impact of N_{max} on the differences between Bouguer anomalies computed from GGMs and the respective ones from ground-based gravity measurements

applied N_{max} [degree]

Fig. 5. Standard deviation of differences between Bouguer anomalies computed from GGMs and the respective ones from ground-based gravity measurements - models EGM2008 and EIGEN-6c

> Fig. 6. Standard deviation of differences between Bouguer anomalies computed from GGMs and the respective ones from ground-based gravity measurements, $N_{max} = 120$. Points of POGK (left panel) and points located in Tatra Mountains (right pa-