

3000

500

Karstified

Non karstified

Mixt

1000 1500 2000 2500 3000

P (mm/v

–W* = P

INVESTIGATING THE ROLE OF GEOLOGY ON THE SPATIAL VARIABILITY OF CATCHMENT WATER BALANCE **AND ITS IMPACT ON FLOOD PROCESSES**

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Geology is a main pattern of catchment organization, and we ask whether it impacts the hydrological response of catchments, especially during floods. The aim of our study is to assess which hydrological processes are linked to the an of our study is to assess which hydrological processes are linked to the an of our study is to assess which hydrological processes are linked to the geological and karst pattern, and to which extend the latter can affect water balance and runoff coefficient during floods. Indeed, karst catchments are generally prone to Interbasin Groundwater Flows (IGF), resulting in streamflow losses or gains. Common examples of IGF in karst are underground river losses, or springs draining areas out of the topographical catchment. The proposed approach is a combined analysis of the spatial variability of hydrological indices are multi-annual average values obtained from daily streamflow and rainfall records of 9 to 30 hydrological years. Geomorphological parameters are IDPR (see dedicated box), and drainage density. Such indicators, easily calculable, could allow the transferability of results to ungauged catchments. The study site (total of 25 000 km²) includes 12 French catchments, located in various karstic regions and equipped with 120 gauging stations.





The L'Vovich annual water balance theory, adapted to non-conservative basins thanks to independent evapotranspiration data, is applied in a first step. It shows that many catchments are prone to IGFs, and that it can often be linked to the catchment geology (differentiating hard-rock) karst, and other sedimentary formations). Results also show that both quick and slow flow components are correlated with geology. In a second step, we characterize relationships between those hydrological specificities and some simple geomorphological parameters (i.e. drainage density and IDPR). Results show that correlations exist, relatively strong depending on the study sites and geology. The impact of land use is less important within each site, but is significant when considering all 3 sites. The case of intensely karstified areas is highlighted due to the highest variability of hydrological indicators in catchments covers by such patterns. Finally, our results highlight the geological control of catchment hydrology and the specific role of groundwater flows in karst areas on flood flows. Knowing that karst areas cover 20% of Europe outcrops, this work promises interesting perspective in mapping hydrological indicators in link with catchment organization and geology.

INTRODUCTION

Model improved: more consistent values, with physical meaning

Still few unexpected values (out of « conservative » zones) A better repartition of I between S and U is under research...

RESULTS: FLOOD INDICES SPATIALISATION AND LINK WITH GEOMORPHOLOGY

1. Mapping flood runoff



2. Hydrology – Geomorphology correlations





CONCLUSIONS



S/P index mapping (quick flow component normalised by rainfall) shows the spatial variability of flood runoff. It appears to be correlated with geology, especially in Cévennes (A) and Normandy (B). Similar results are obtained with IDPR or drainage density.

S/P values are low on karstified limestone basins, and higher on hard-rock terrains. Other flood indices are not presented here.



Mapping of indicators shows consistency between physiographical and hydrological the intermediate parameters at catchment scale.

Regressions between those indicators show correlations at the topographic between S/P and catchment scale, geomorphological indices (drainage density and IDPR), especially for Cévennes and Normandy. Weak correlation for the Jura Mountains can be explained by an intense karstification, promoting strong groundwater influence, not reflected in surface geomorphological parameters.

In a similar way, several land use indicators are computed and their correlation with hydrological index studied. Within each study site, determination coefficient values are low (all R² < 0.23), showing a less important influence than geomorphology on the S/P index.

Yet, unlike with geomorphological parameters, the global regression (all 3 study sites) gives interesting results $(R^2 = 0.33)$ for the topographic catchment scale.



Continuous carbonate rocks Discontinuous carbonate rocks Study sites

Potential karst aquifers [Chen et al., 2017]



	loboBidbille	meenie
Streamflow	Q _o	Q _o - Q _i
Surface runoff	S _o	S _o - S _i
Underground runoff	U _o	U _o - U _i

network characterizes connectivity the IDPR hydrographic network. 2000 It is non dimensional and its value ranges from 0 to 2000. IDPR is available at the 25 m spatial resolution over France.



2017 IDPR grid – BRGM

This index is computed by estimating the gap between a theoretical river network corresponding to the topographic thalweg network, and the existing hydrographic network. For example, a dry thalweg will give an IDPR close to 0, and a wetland corresponds to a high IDPR value (close to 2000).



IDPR computation by crossing the hydrographic network and the thalweg network

IDPR value at each node is obtained from the following conceptual equation:

Distance to closest theoretical river IDPR = $-\times 1000$ Distance to closest existing river

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