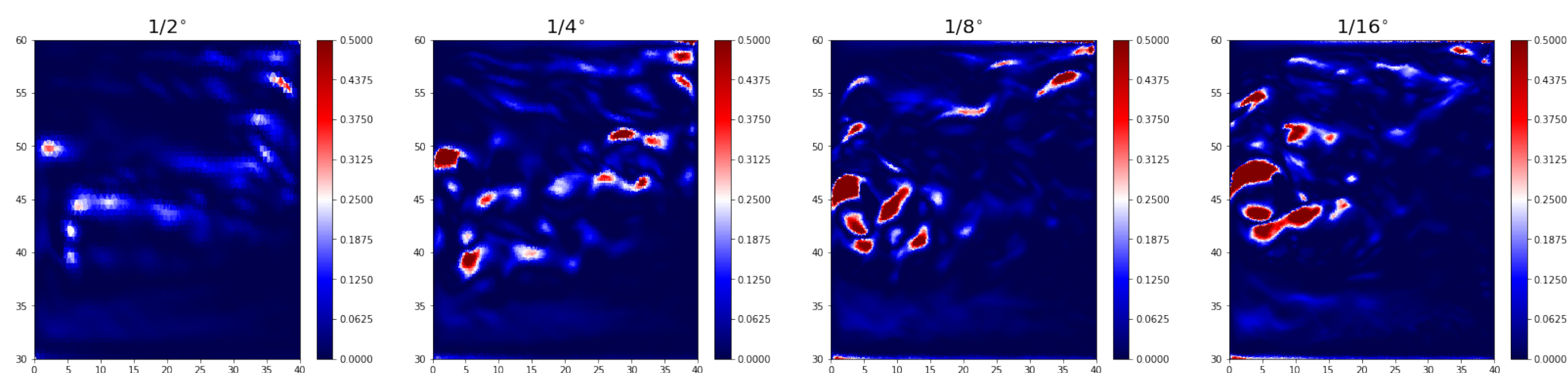


## Background

- Global ocean models incur exceptional computational expense.
- Therefore models must be run at low resolutions, and effects of small scales parameterised.
- We want to use a stochastic parameterisation added in such a way that does not destroy the fundamental physics of fluid dynamics.
- SALT (stochastic advection by Lie transport) is a method that does this (see [Hol15], [SC20]).
- We consider the ocean model FESOM2.0, see [DSWJ16].

## Differences in flow solution as we change resolution

- One field in which we can see clear differences when we change the resolution is kinetic energy. Below we plot kinetic energy field snapshots for resolutions of  $1/2^\circ$ ,  $1/4^\circ$ ,  $1/8^\circ$  and  $1/16^\circ$ .



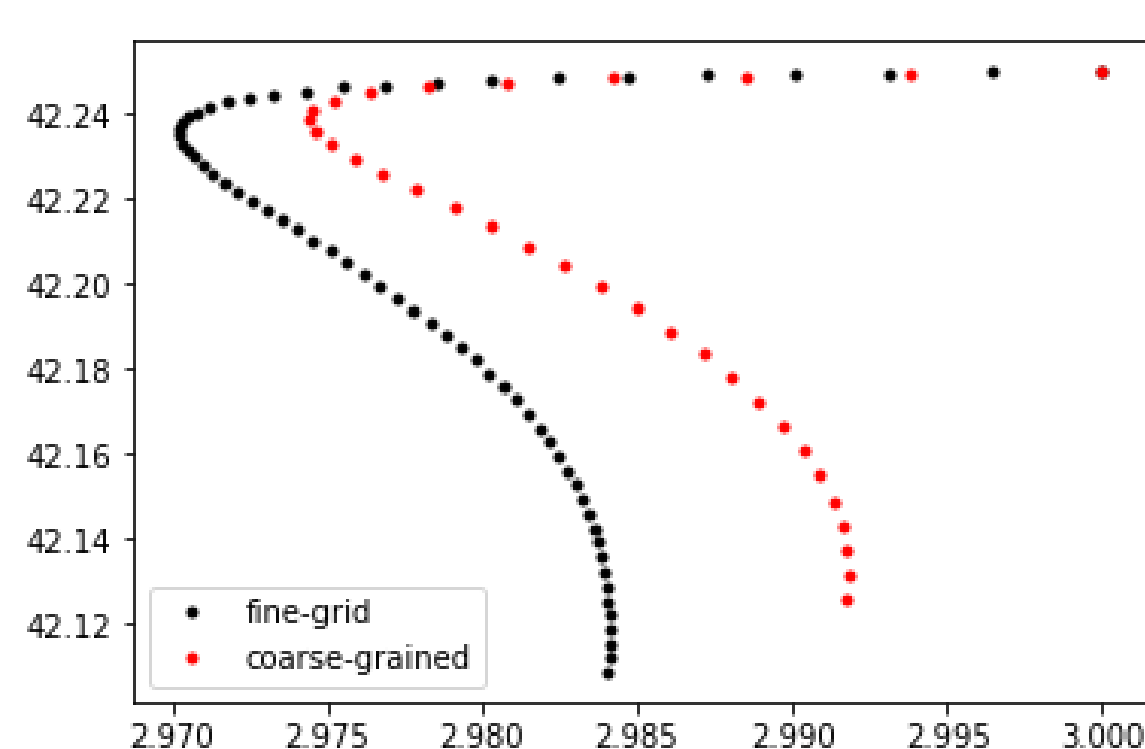
- As we increase the resolution, we see a stronger kinetic energy jet emerging from the western boundary. This kind of feature is what an effective parameterisation scheme should capture.

## Lagrangian Paths

- A fluid is made up of a continuum of particles, each following a trajectory determined by the velocity field. To capture uncertainty we add a stochastic part to the particle trajectory:

$$dx_t = u(x_t, t)dt + \sum_i \xi_i(x_t) \circ dW_t^i \quad (1)$$

- We determine the  $\xi_i$  by taking the velocity field of a simulation on the fine grid,  $u_f$ . We then coarse-grain this solution by filtering, to get a smooth field  $\bar{u}$ . We then calculate the Lagrangian trajectories for each of the fields and look at the difference (see [CCH<sup>+</sup>18]):



- We use these trajectories to calculate:

$$\Delta X = x_f(t) - \bar{x}(t) \approx \sum_n u(x_f^n, t_n) \Delta t - \sum_m \bar{u}(\bar{x}^m, T_m) \Delta T \quad (2)$$

- The  $\xi_i$  are calculated as re-scaled EOFs of the field  $\Delta X$ .
- An alternative simple approach is to take the difference in velocity fields:

$$\Delta X = (u_f(x, t) - \bar{u}(x, t)) \Delta T \quad (3)$$

## References

- [CCH<sup>+</sup>18] C. Cotter, D. Crisan, D. D. Holm, W. Pan, and I. Shevchenko. Numerically Modelling Stochastic Lie Transport in Fluid Dynamics. *arXiv preprint arXiv:1801.09729v2*, 2018.
- [DSWJ16] S. Danilov, D. Sidorenko, Q. Wang, and T. Jung. The Finite-volume Sea ice-Ocean Model (FESOM2). *Geoscientific Model Development Discussions*, pages 1–44, 2016.
- [Hol15] Darryl D. Holm. Variational principles for stochastic fluid dynamics. *Proceedings of the Royal Society A: Mathematical, Physical and Engineering Sciences*, 471(2176):20140963, Aug 2015.
- [SC20] Oliver D. Street and Dan Crisan. Semi-martingale driven variational principles. *Arxiv*, 2020.

## Variational Principles for stochastic Fluids

To use the stochastic Lagrangian paths in a way that preserves the properties of the flow, we use a variational principle involving the following action:

$$S = \int_0^T \underbrace{l(u, D, T)}_{\text{incompressibility}} dt + \underbrace{\langle d_t P, D - 1 \rangle}_{\text{Lagrangian particle paths}} + \left\langle \pi, dx_t - u(x_t, t)dt - \sum_i \xi_i(x_t) \circ dW_t^i \right\rangle \quad (4)$$

The Lagrangian  $l$  contains information about the kinetic and potential energy of the system. For the primitive equations it is given by:

$$l(u, D, T) = \int_V \left( \frac{1}{2} |\mathbf{u}|^2 + \mathbf{R} \cdot \mathbf{u} - \int_{z_0}^z (1 + B(T(\mathbf{x}, z, t), z')) dz' \right) d^3x \quad (5)$$

This consists of kinetic energy  $\frac{1}{2} |\mathbf{u}|^2$ , rotation  $\mathbf{R} \cdot \mathbf{u}$  and potential energy  $\int_{z_0}^z (1 + B(T(\mathbf{x}, z, t), z')) dz'$ , where  $T$  is potential temperature and  $B$  comes from the equation of state,  $\rho'/\rho_0 = B(T, z)$ .

Deriving the equations in this way preserves important physical properties such as circulation and potential vorticity:

$$d_t \oint_{C(t)} (\mathbf{u} + \mathbf{R}) \cdot d\mathbf{x} = (g/\rho_0) \iint_{S(t)} \hat{\mathbf{k}} \times \nabla \rho' \cdot d\mathbf{S} dt \quad (6)$$

$$d_t q + dx_t \cdot \nabla q = 0 \quad q = \nabla T \cdot (\text{curl } \mathbf{u} + f\hat{\mathbf{k}}) \quad (7)$$

## Primitive Equations with SALT

The primitive equations with the SALT method are given by:

$$d_t \mathbf{u} + \left[ u_3 \cdot \nabla_3 \mathbf{u} + f\hat{\mathbf{k}} \times \mathbf{u} + \nabla p \right] dt + \sum_i \mathbf{G}_i \circ dW_t^i = (\boldsymbol{\tau} + \mathbf{D}_u) dt \quad (8a)$$

$$\nabla_3 \cdot \mathbf{u}_3 = 0 \quad (8b)$$

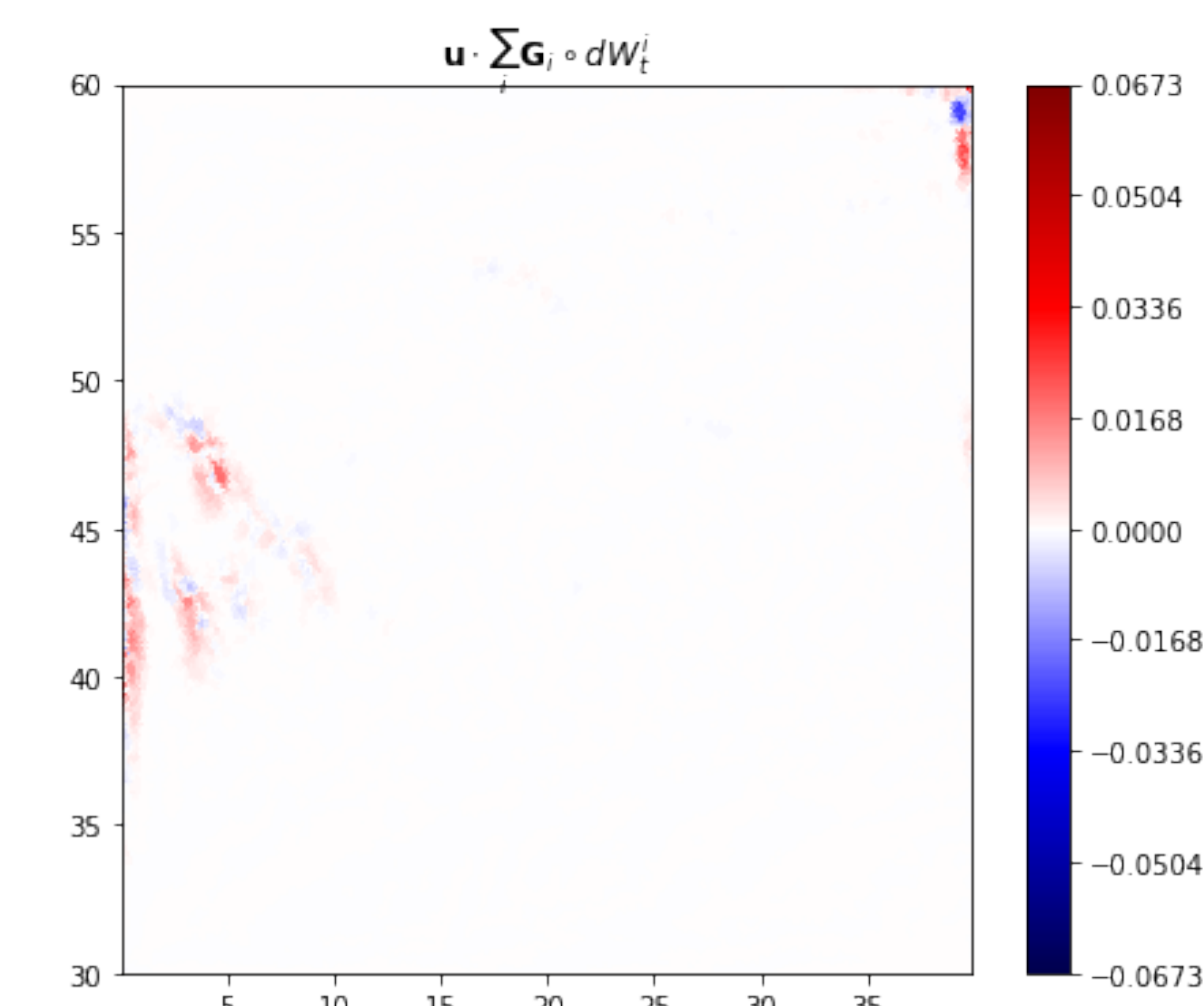
$$\frac{\partial p}{\partial z} = -g(1 + \rho'/\rho_0) \quad (8c)$$

$$d_t T + u_3 \cdot \nabla_3 T dt + \sum_i \xi_i \cdot \nabla_3 T \circ dW_t^i = F_T \quad (8d)$$

$\mathbf{G}_u$  is the stochastic forcing given by:

$$\mathbf{G}_i = \xi_i \cdot \nabla_3 \mathbf{u} + f\hat{\mathbf{k}} \times \xi_i + \nabla \xi_i \cdot \mathbf{u} + \nabla \int_z \frac{\partial \xi_i}{\partial z} \cdot \mathbf{u} dz' \quad (9)$$

This forcing adds kinetic energy to the system at a rate given by  $\mathbf{u} \cdot \sum_i \mathbf{G}_i \circ dW_t^i$ . We plot this below for a simulation on  $1/4^\circ$  grid with  $\xi_i$  calculated from a  $1/8^\circ$  simulation:



We see that the forcing is acting where needed to increase the kinetic energy.