

# Locating and characterizing seismic events in Los Humeros (Mexico) using time-reverse imaging

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# Motivation for seismicity analysis in geothermal projects

## **spatio-temporal distribution of seismic events:**

- mapping of fault networks
- assess and mitigate seismic hazard
- reservoir characterisation

## **focal mechanisms (beachballs, fault planes, ...):**

- estimate local stress regime
- kinematic processes in the subsurface
- distinction between tectonic and anthropogenically induced events?

informed understanding  
of geological and  
geophysical processes in  
the subsurface

# Why Time-Reverse Imaging (TRI)?

**TRI is a method for locating and characterising seismic events.**



- full waveform → usage of all available data
- no picking → quasi-simultaneous & low SNR events
- all scales (lab, field, ...) → non-destructive testing, micro-seismicity, tremors, ...
- no a priori assumptions about sources → may be used to infer unknown subsurface processes

well-constrained high-resolution velocity model is needed



- computational costs may be high
- constraints on station network

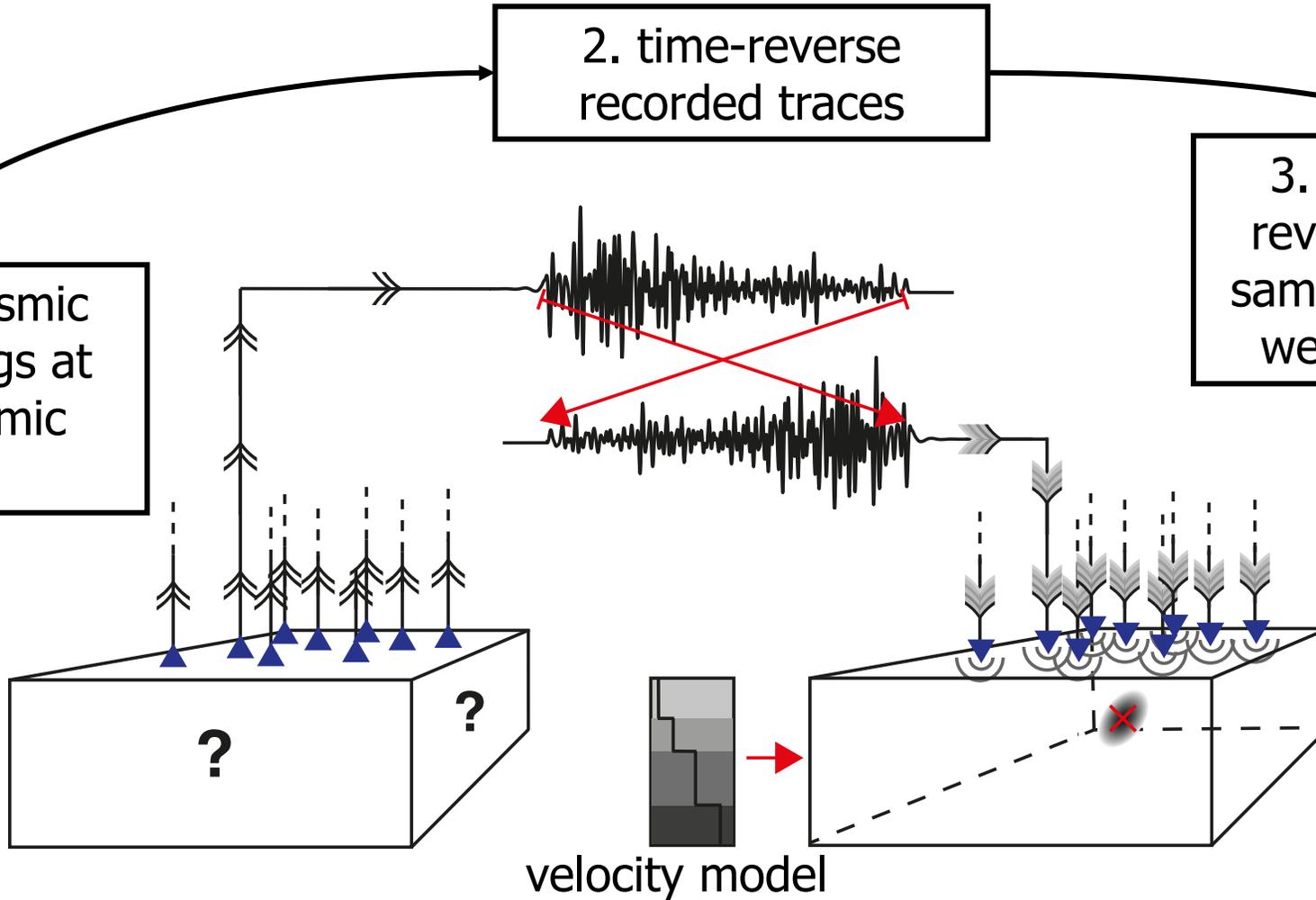
# The principle of TRI

1. passive seismic data recordings at discrete seismic stations

2. time-reverse recorded traces

3. reinsert time-reversed traces at same positions they were recorded at

4. back propagate time-reversed wave field until it focuses on the source location and application of imaging conditions



# Imaging Conditions used for TRI

*Imaging conditions are used to quantify the convergence of the time-reversed wave field. They are based on different characteristics of the wave field and are thus used simultaneously to obtain a more complete source image.*

Maximum displacement:

$$\mathbf{I}_d(\mathbf{x}) = \max_{t \in [0, T]} \|\mathbf{u}(\mathbf{x}, t)\|$$

Maximum P-wave energy density:

$$\mathbf{I}_p(\mathbf{x}) = \max_{t \in [0, T]} (\lambda + 2\mu) [\nabla \cdot \mathbf{u}(\mathbf{x}, t)]^2$$

Maximum S-wave energy density:

$$\mathbf{I}_s(\mathbf{x}) = \max_{t \in [0, T]} \mu [\nabla \times \mathbf{u}(\mathbf{x}, t)]^2$$

Maximum total energy density:

$$\mathbf{I}_e(\mathbf{x}) = \max_{t \in [0, T]} \sum_i \sum_j [\sigma_{ij}(\mathbf{x}, t) \varepsilon_{ij}(\mathbf{x}, t)]$$

$\mathbf{x}$ : space vector –  $t$ : time –  $\mathbf{u}$ : displacement –  $\lambda, \mu$ : Lamé parameters –  $\sigma_{ij}$ : stress –  $\varepsilon_{ij}$ : strain

Saenger (2011): Time reverse characterization of sources in heterogeneous media, *NDT&E International*, <https://doi.org/10.1016/j.ndteint.2011.07.011>

# Challenges when using TRI

most  
influenced

## 1. Station Distribution

*Irregular and sparse station distributions lead to under-sampled wavefields and thus incomplete reconstruction of time-reversed wave field.*

→ solvable by:

- optimised station network **SLIDE 7**
- identification of areas with low or insufficient source-location accuracy **SLIDE 10-11**

## 2. Velocity Model

*Incorrect velocity models result in incorrect travel paths of the time-reversed wave field and thus hinder the convergence at the source location.*

→ solvable by:

- identification of areas with low or insufficient source-location accuracy **SLIDE 10-11**
- improved velocity models

## 3. Ambient Seismic Noise

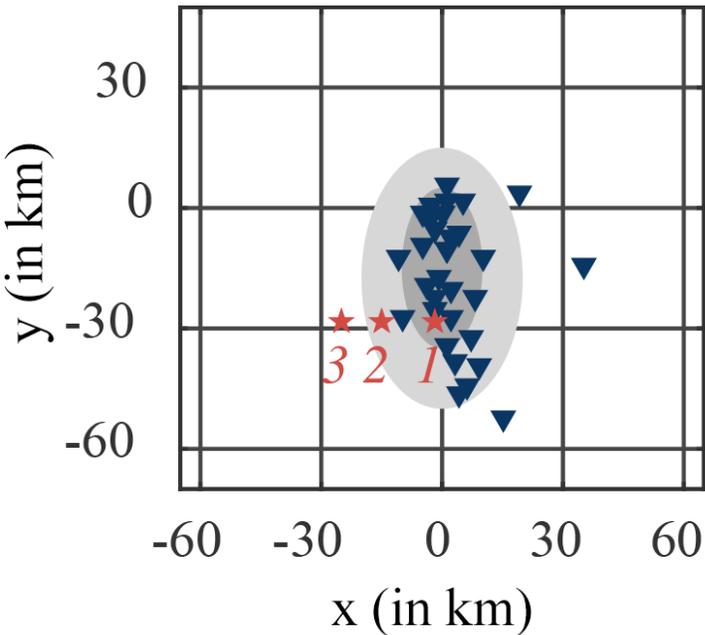
*Low SNR ( $< 1$ ) may introduce artificial convergence spots overshadowing the convergence spot at the source location.*

→ understand limits of method **SLIDE 8**

least  
influenced

# Optimise station network for TRI

## station network in Southern California



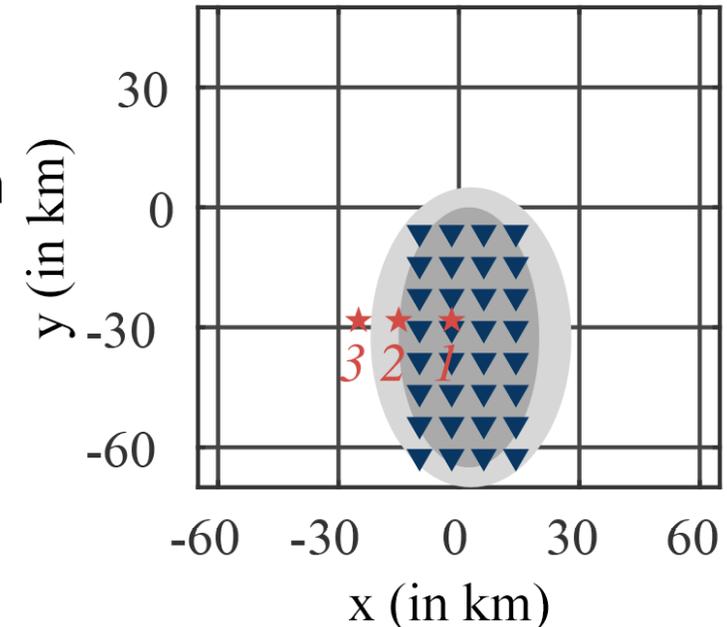
*dark grey circle: In this area sources can be located accurately in the target depths*

## guidelines for optimised station network:

- inter-station distance  $<$  source depth
- aperture of network  $>$  2 x source depth
- most accurate when sources are beneath the network
- regular station networks improve accuracy for shallow sources

*light grey circle: In this area sources may be located with decreased accuracy*

## station network optimised for TRI

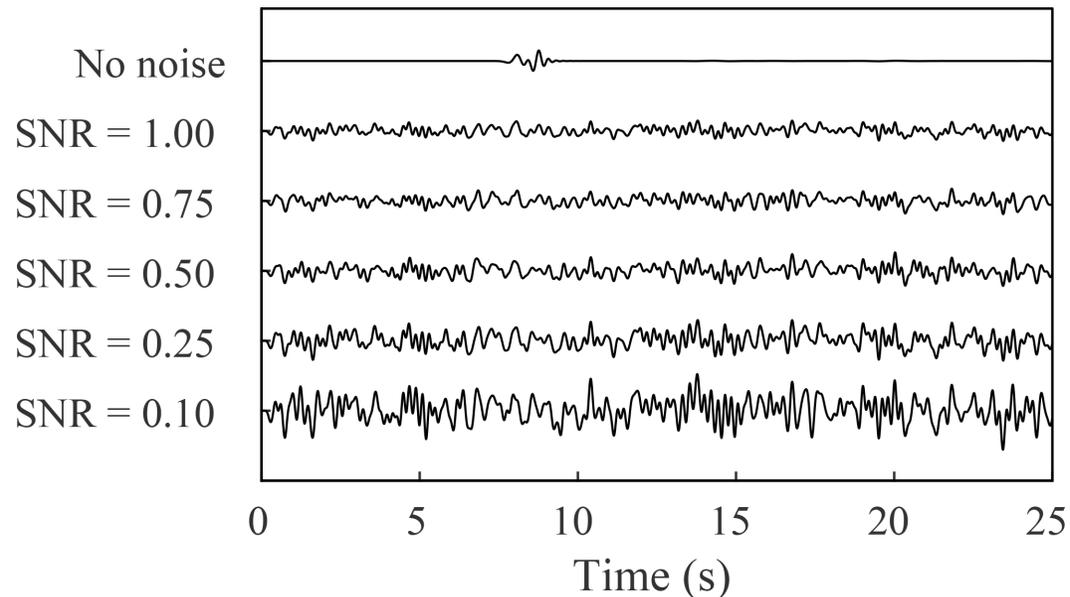


*blue triangles: stations*  
*red stars: sources*

Werner and Sanger (2018): Obtaining reliable source locations with time reverse imaging, *Solid Earth*, <https://doi.org/10.5194/se-9-1487-2018>

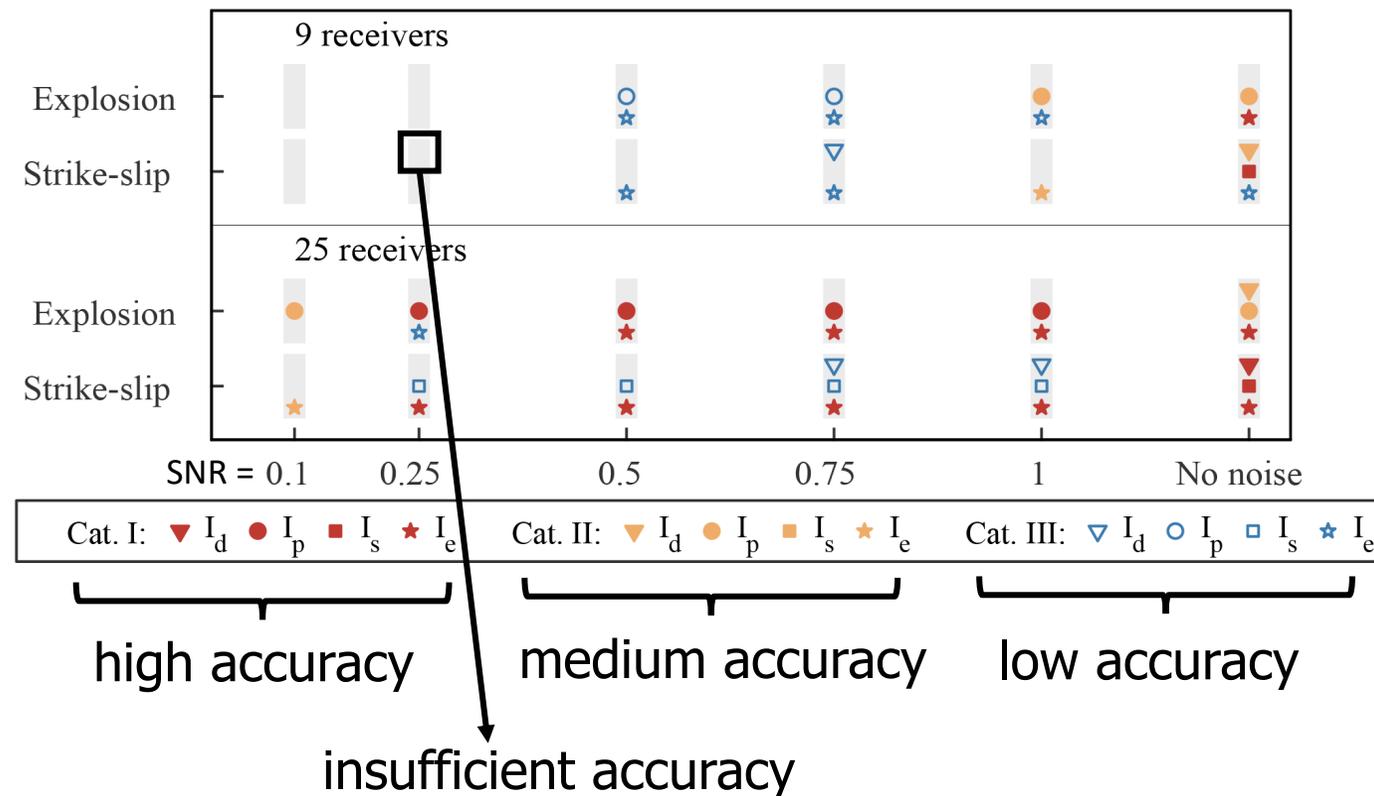
# Handling low noise levels with TRI

Example traces for a synthetic source with noise



$I_d$  : Maximum displacement imaging condition  
 $I_p$  : P-wave energy density imaging condition  
 $I_s$  : S-wave energy density imaging condition  
 $I_e$  : Total Energy density imaging condition

*TRI handles low SNR well if the station network is well enough (i.e. enough stations). Depending on the imaging condition, sources may be located with an SNR as low as 0.1.*

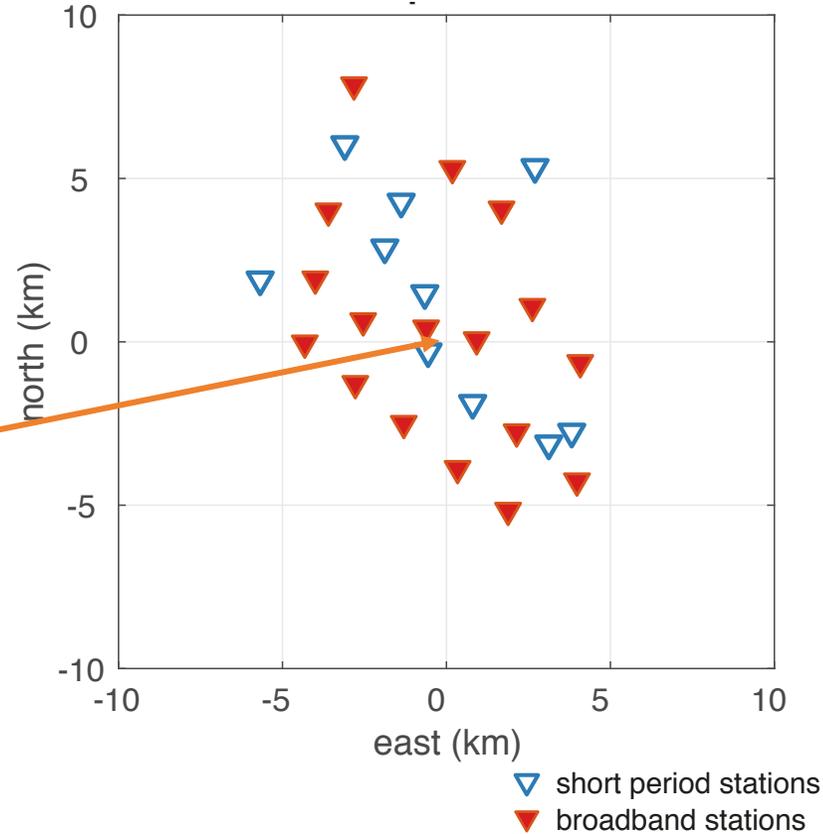


Werner and Sanger (2018): Obtaining reliable source locations with time reverse imaging, *Solid Earth*, <https://doi.org/10.5194/se-9-1487-2018>

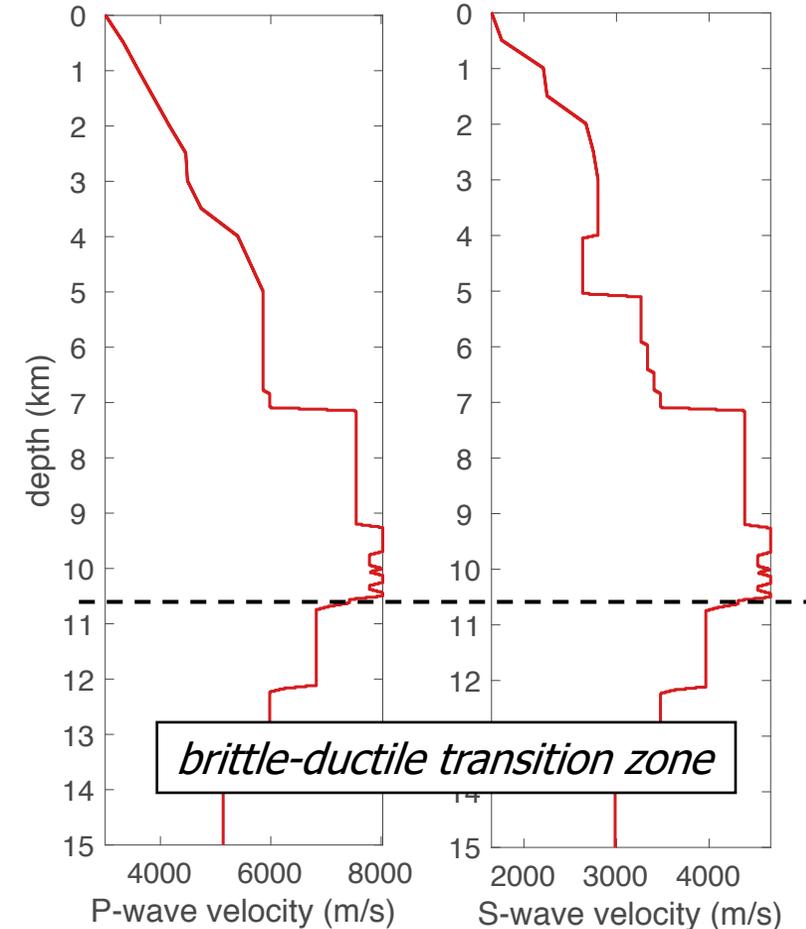
# Los Humeros geothermal field (Mexico)



27 stations (deployed 09/2017 – 09/2018)



Toledo, T. et al., 2019. Dataset of the 6G seismic network at Los Humeros, 2017-2018, *GFZ Data Serv.*, <https://doi.org/10.14470/1T7562235078>

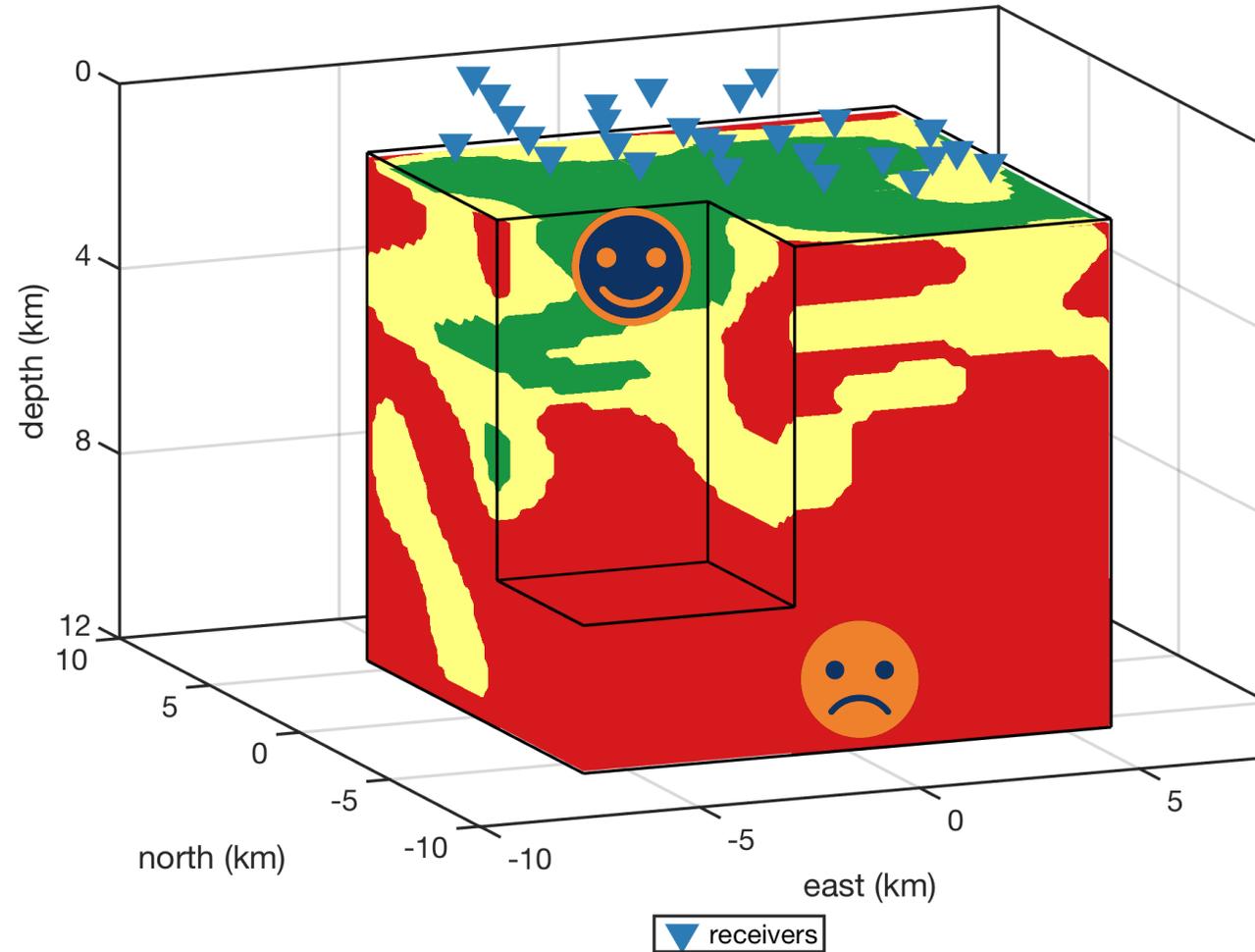
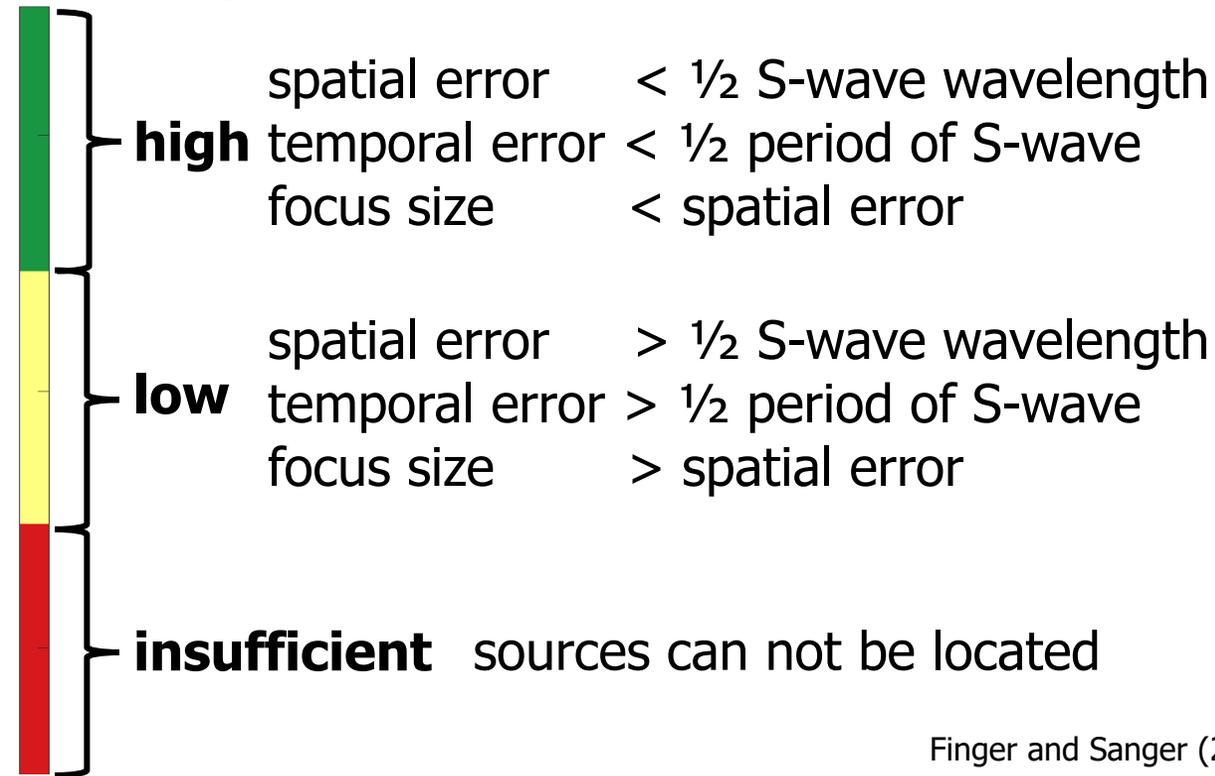


Löer, K. et al., 2020. Imaging the deep structures of the Los Humeros geothermal field, Mexico, using three-component ambient noise beamforming, *Seismological Research Letter*, *submitted*

# Sensitivity maps reveal influence of stations and velocity model

*Sensitivity maps highlight areas with low or insufficient source-location accuracy due to the combination of the station distribution and the velocity model. Convergence spots in these areas may be excluded from further analysis.*

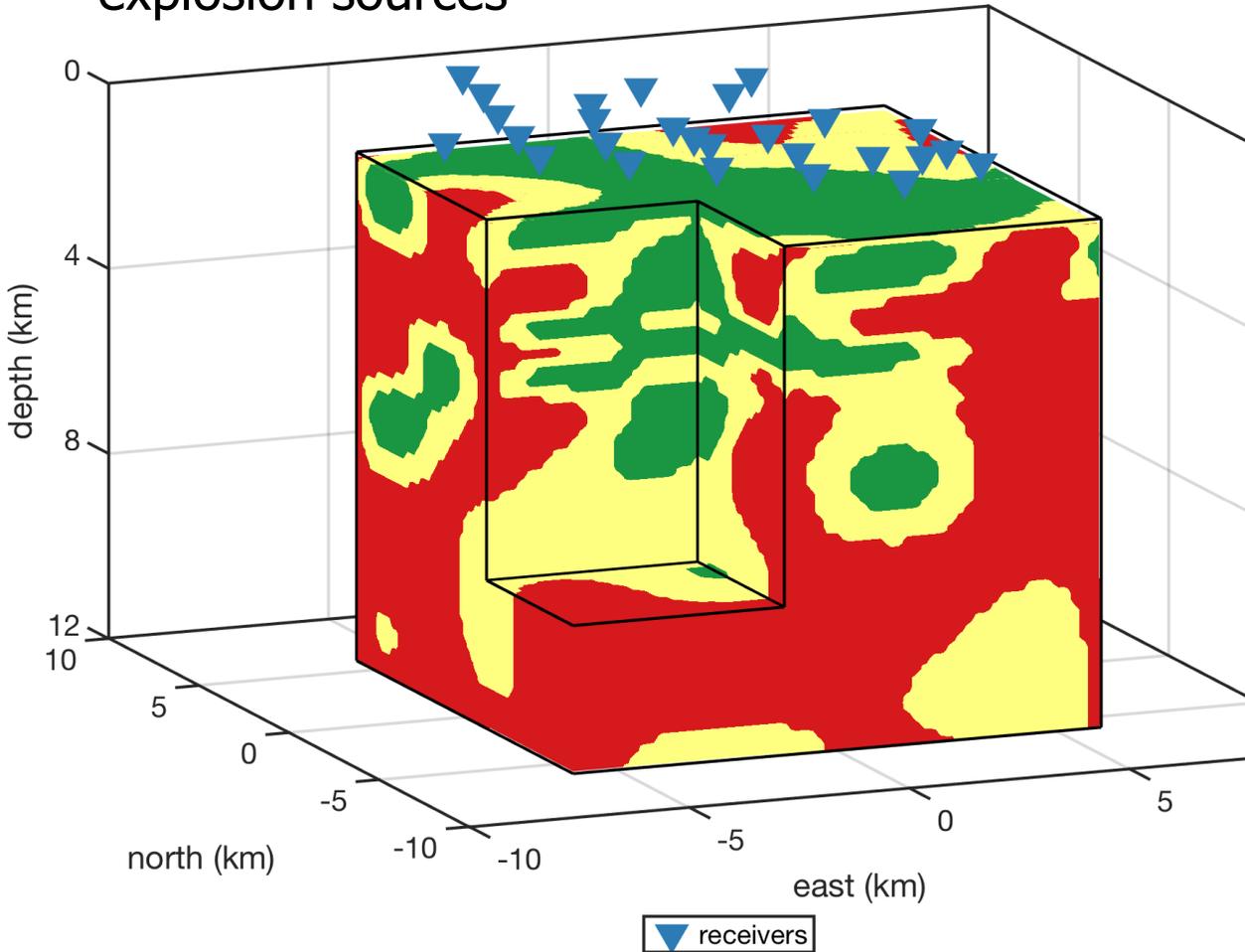
## accuracy



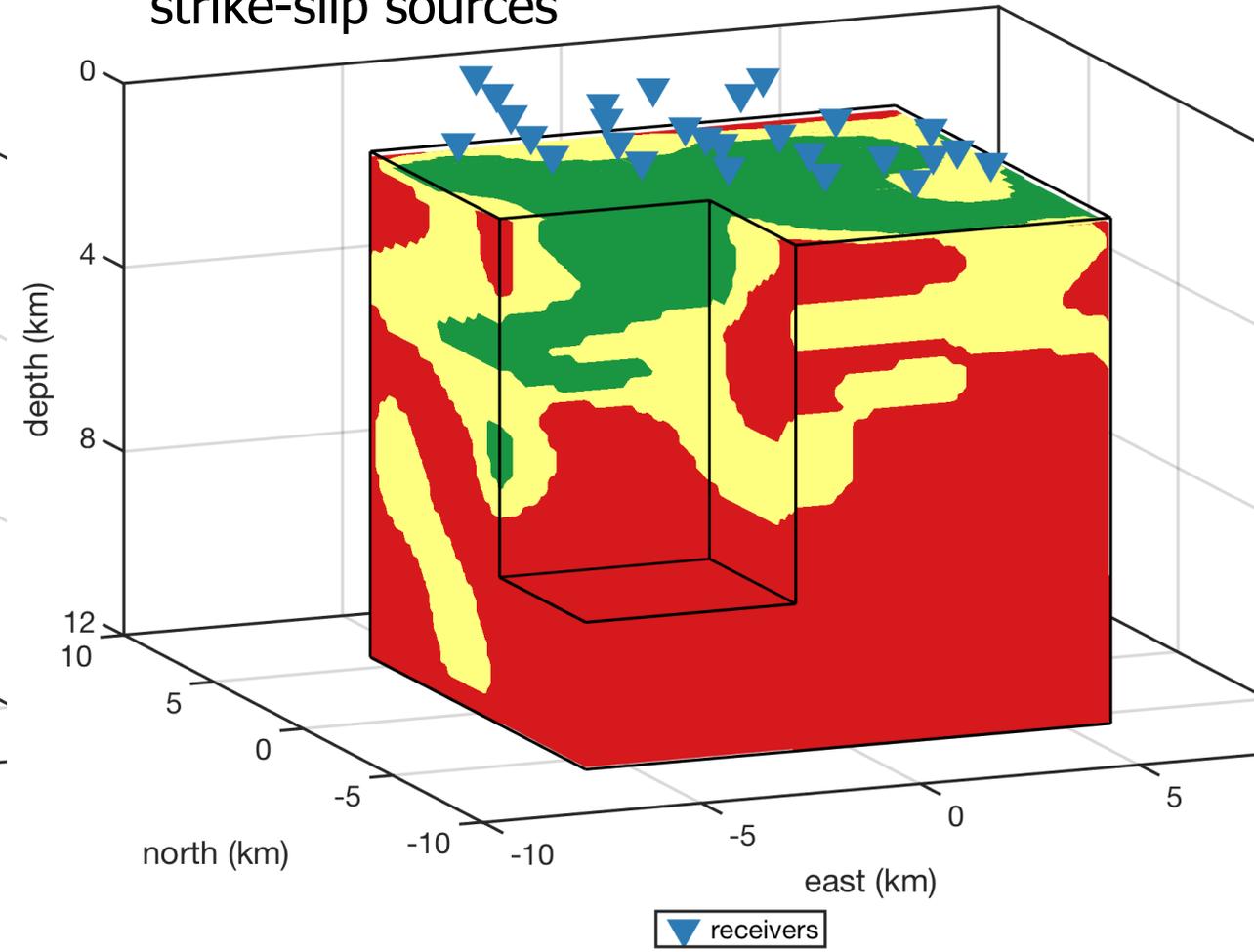
Finger and Sanger (2020): Sensitivity Maps for Time-Reverse Imaging, *GJI*, <https://doi.org/10.1093/gji/gqaa160>

# Sensitivity Maps are sensitive to the source type

explosion sources



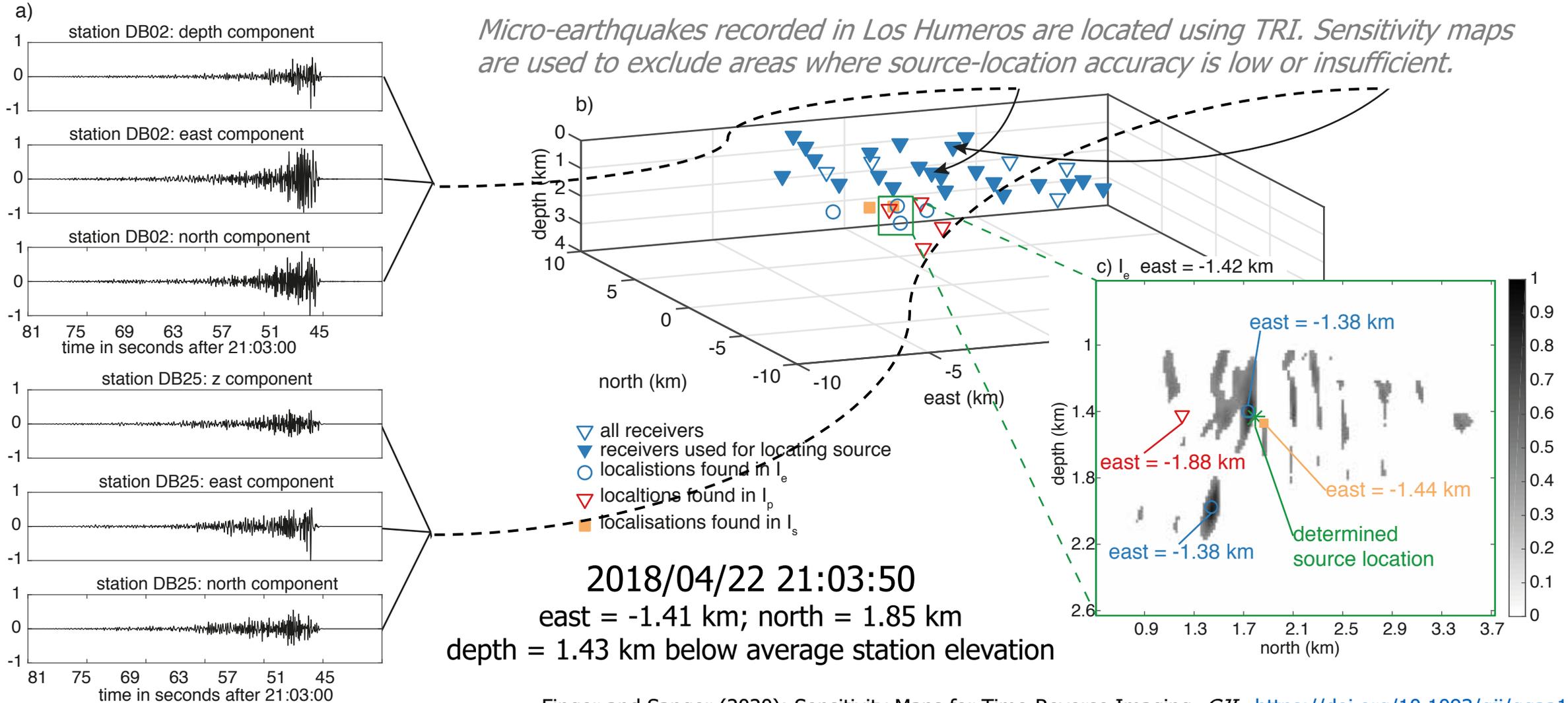
strike-slip sources



Finger and Sanger (2020): Sensitivity Maps for Time-Reverse Imaging, *GJI*, <https://doi.org/10.1093/gji/gqaa160>

# Locate exemplary event in Los Humeros with TRI

Micro-earthquakes recorded in Los Humeros are located using TRI. Sensitivity maps are used to exclude areas where source-location accuracy is low or insufficient.



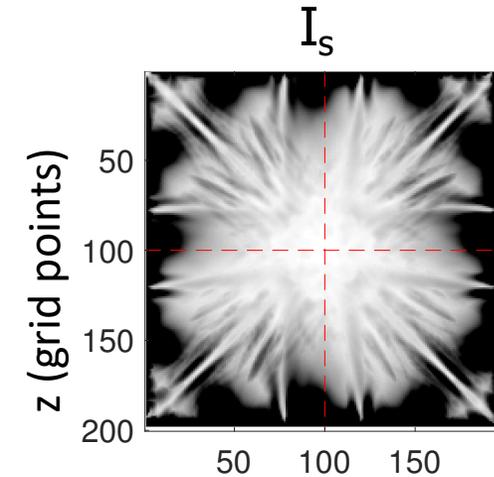
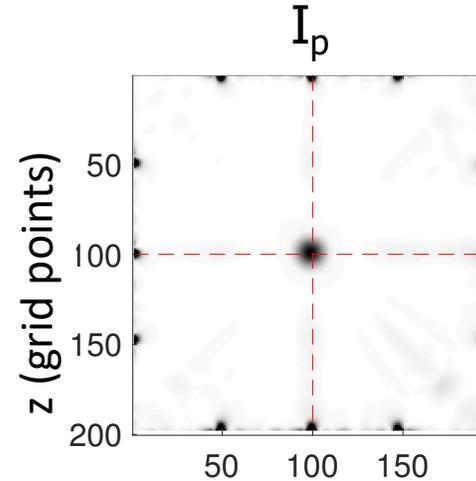
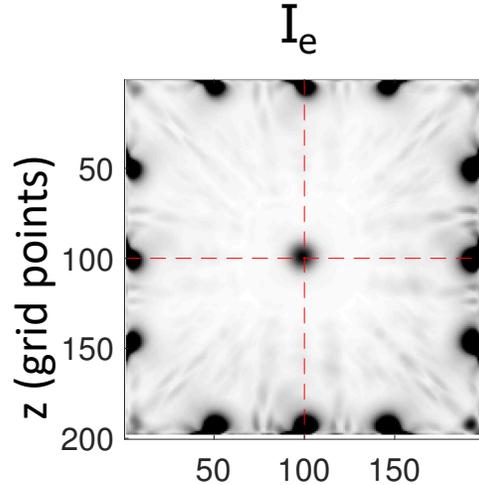
Finger and Sanger (2020): Sensitivity Maps for Time-Reverse Imaging, *GJI*, <https://doi.org/10.1093/gji/ggaa160>

# Characterise events with TRI

## synthetic explosion source in a homogeneous cube

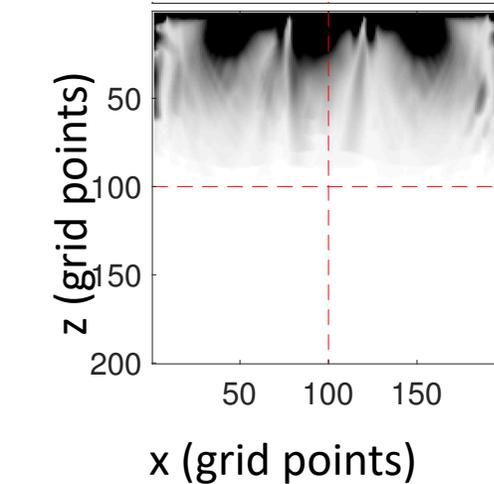
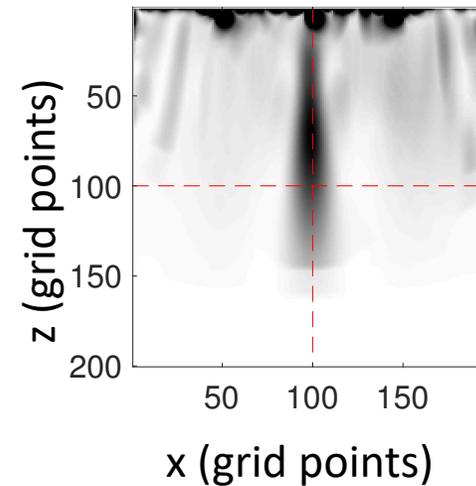
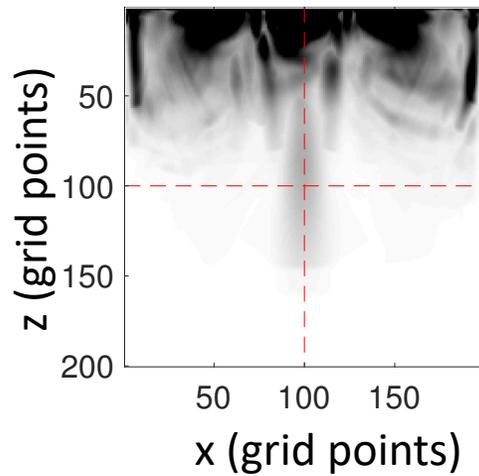
stations all around  
3D cube

*source location is  
recovered with  $I_e$   
and  $I_p$*



stations only at  
surface

*source location is  
smeared in z-  
direction*

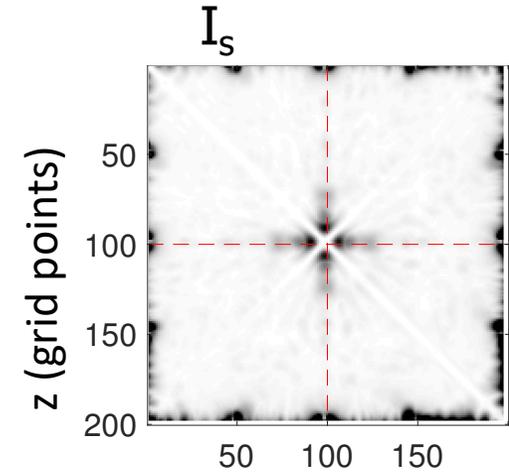
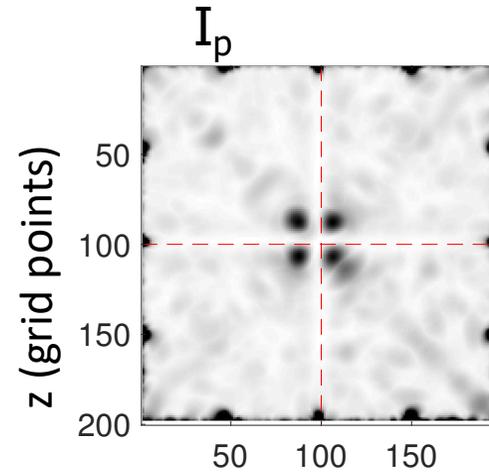
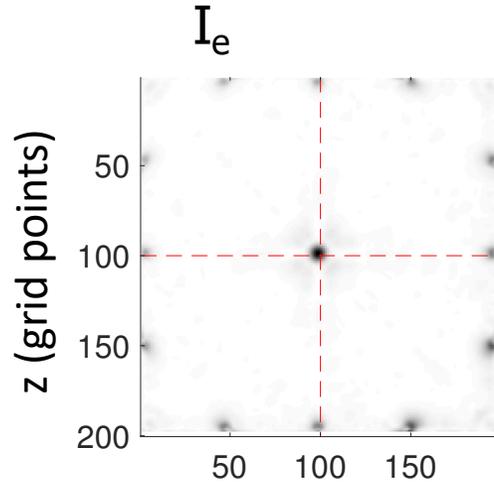


# Characterise events with TRI

## synthetic dip-slip source in a homogeneous cube

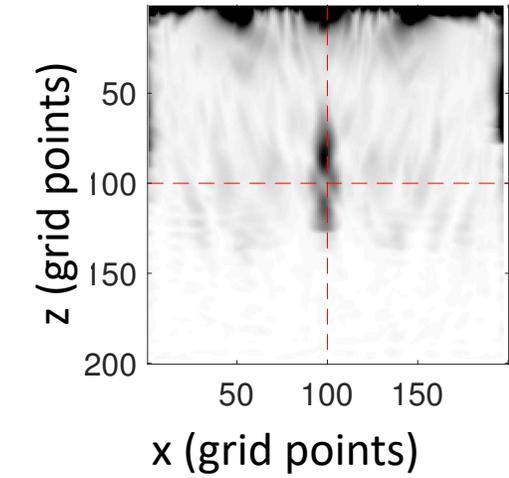
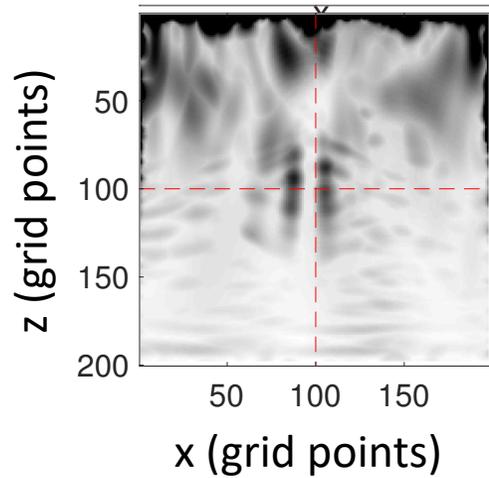
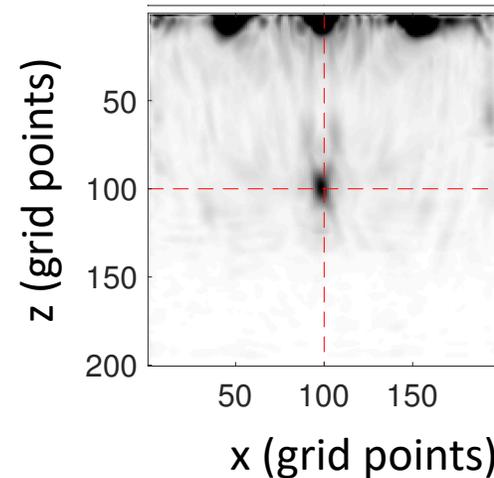
stations all around  
3D cube

*source location is  
recovered with  $I_e$ .  
Radiation pattern is  
recovered with  $I_p$   
and  $I_s$ .*



stations only at  
surface

*source location is  
recovered with  $I_e$ .  
Radiation pattern is  
not clearly visible.*



# Summary and Outlook

- TRI is able to locate and characterise seismic events in low SNR conditions and is especially well-suited for applications such as aftershock sequences, induced seismicity due to fluid injection and hydro-fracture stimulation.
- TRI is a complementary method to standard analysis routines.
- An in-depth understanding of source mechanisms is possible.
- An automatic work flow for locating and characterising seismic events is being derived (manuscript in preparation)

# Thank you for your Attention!



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[www.rockphysics.org](http://www.rockphysics.org)

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