Silicon isotopes as tracers of paleoenvironmental conditions during laterite formation in the Amazon Basin



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Lateritic soils: composition and occurrence on Earth

- 1/3 of emerged land but 85 % of soil volume on Earth
- Composition: kaolinite clay and Fe/AI (oxi)hydroxides
- Formation under tropical rainy climate



Nahon 2003

Laterite developed on sedimentary Alter do Chão Formation





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Mottled zone Differentiation between whitish (kaolinite) and reddish (Fe/Al oxides) sequences

Parent rock or sediment
Here: sedimentary Alter do Chão formation (Sandy and kaolinite-rich sediment)

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"Iron crust" - Pisolitic horizon

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- Laterites formed by successive weathering episodes (based on dating techniques)



Does a geochemical tracer exist to identify the paleoenvironmental conditions during these successive episodes of formation?

Si isotopes: fractionation of kinetic origin during chemical weathering process (Oelze et al. 2014, 2015)

Enrichment of solution in heavy isotopes





Si isotope fractionation expected during lateritization Control of climatic conditions on recorded Si isotope signal

"Test" lateritic profile: soil III North of Manaus (Lucas et al. 1996)

- Lateritic profile developed on Alter do Chão (65-45 Ma) sedimentary formation
- Already studied for dating and kaolinite crystallinity (Balan et al. 2005)



Do Si isotopes record a similar evolution over the lateritic profile ?
 Do Si isotopes record distinct conditions for the formation of the two kaolinite populations?

Modified from Eyrolle et al. 1996 and Balan et al. 2005

Focus on the clay fraction (<2 µm after CBD removal of iron oxides)</p>

 $T_{X,Nb} = (X_{soil} \times Nb_{UCC})/(X_{UCC} \times Nb_{soil}) - 1$



→

Alter do Chão sediment experienced *in situ* pedogenetic weathering before laterization (Fritsch et al. 2002, Lucas et al. 1989, Balan et al. 2005)



1/ Limited Si isotope fractionation ($\Delta^{30}Si_{sediment-UCC} = -0.40 \%$) compared with fractionation reported in literature for kaolinite formation (ca. $\Delta^{30}Si_{clay-parent rock} = -2 \%$)

2/ Silicon isotope recording kinetic fractionation (Oelze et al. 2014)
 Slow kaolinite precipitation

3/ Post-Oligocene weathering phase (25-30 Ma) under climatic conditions enabling the slow precipitation of kaolinite mineral

Limited water infiltration



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- Two possible scenarios to explain Si isotope evolution during laterite formation
 - Progressive lateritization process between 22 Ma and 6 Ma
 - Binary mixing with a short-term episode around 6-8 Ma





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Two possible scenarios to explain Si isotope evolution during laterite formation

Progressive lateritization process between 22 Ma and 6 Ma



Binary mixing with a short-term episode around 6-8 Ma

Leaching of light REE: persistence of similar chemical conditions over 14 Ma unlikely

2nd episode: lateritization process

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 Kaolinite replacement at 8 Ma independently of altitude and thus of water dynamics

- Desilication of the clay fraction associated with strong enrichment in light Si isotopes fastest kaolinite formation than during the first weathering episode
- Lateritization prone to occur with a persistent rainy season (Beauvais et al. 1999)
- Massive changes in paleogeography and water drainage during early Miocene (8-10 Ma) may have cause important changes in seasonality and water drainage in the Amazon Basin





◆ Onset of transcontinental Amazon River between 8 and 10 Ma (Hoorn et al., 2010)

Conclusions and future steps

✦ Si isotopes: a powerful tool to identify weathering episodes

- Two distinct weathering periods with distinct climatic conditions
- Current development: use of a transport reactive model
 - Constraints on water flow dynamics during Alter do Chão weathering and lateritization process

✦ Use of Si isotopes at a larger scale

- On lateritic soils with various ages
- On lateritic soils developed on crystalline rocks



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