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# Processing steps for the compilation of AAGRГ gravity maps

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Online presentation

# **Basic info:**

**First unified complete Bouguer anomaly (CBA) map of AlpArray area compiled from terrestrial gravity data is currently in preparation, consisting of the following steps:**

- 1. Unification of different spatial, height and gravity systems**
- 2. Getting available detailed elevation models (mainly LiDAR-based)**
- 3. Quality control of input data**
- 4. Calculation of mass corrections (gravity effect of the topography between the surface and ellipsoid level)**
- 5. Calculation of bathymetric corrections for water masses below the ellipsoid**
- 6. Calculation of lake corrections for great Alpine lakes**
- 7. Atmospheric correction – comparison of different approaches**
- 8. Calculation of the final CBA, innovative concept of ellipsoidal heights used**
- 9. Merging individual databases into a single map (with the addition of Global Geopotential Models to fill data gaps)**

# **1. Unification of different spatial, height and gravity systems**

**Position:** Local – National Positioning Reference Systems transformed to European Terrestrial Reference System 1989 (ETRS89) and Universal Transversal Mercator (UTM)

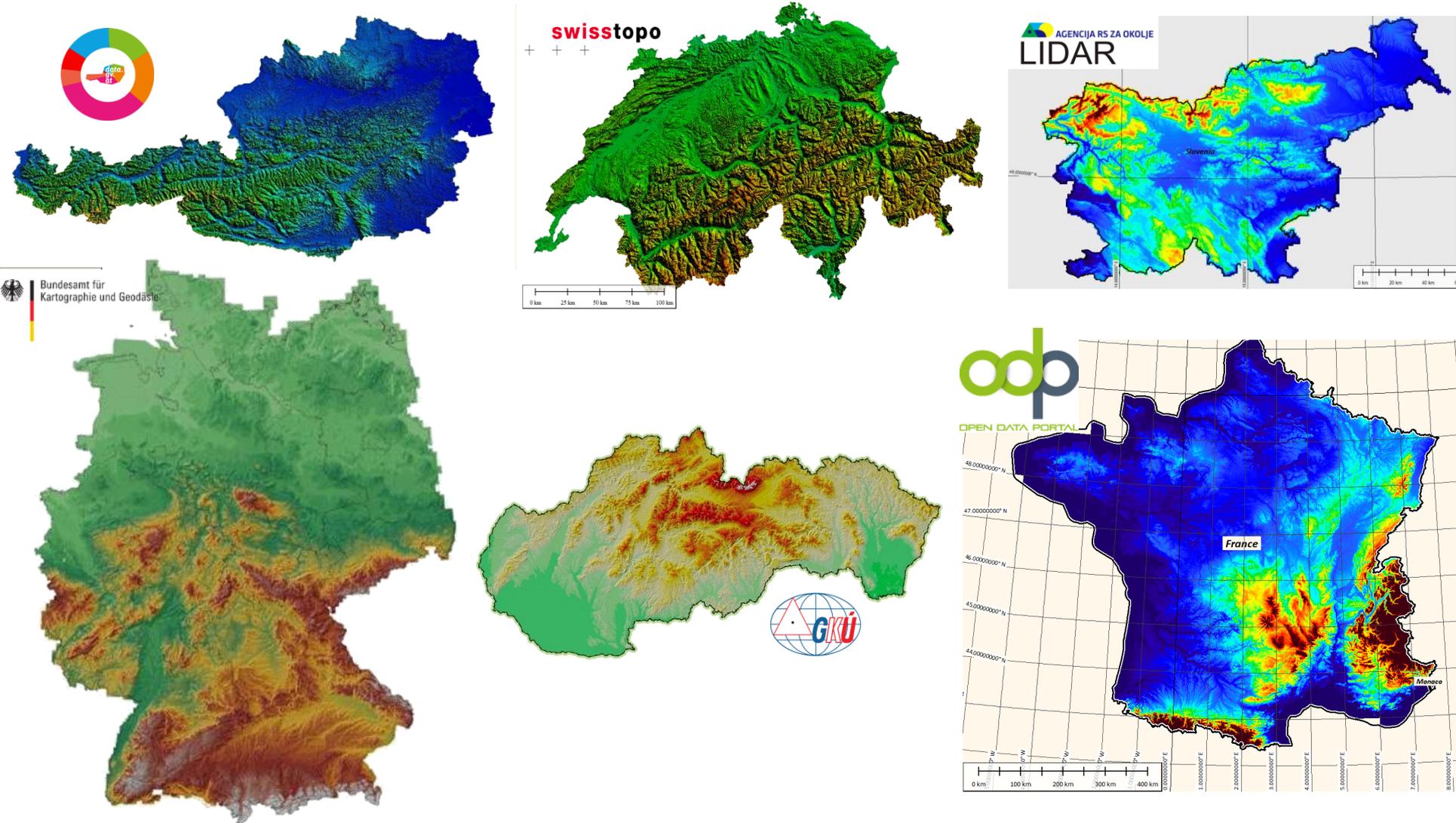
**Height:** Local – National Height Systems transformed to Ellipsoidal Heights (ETRS89, ellipsoid GRS80) using local geoid models

**Marine areas – transformation to Ellipsoidal Heights using EIGEN-6C4 model**

**Gravity:** old Potsdam Gravity System transformed to Absolute Gravity Systems (a few countries)

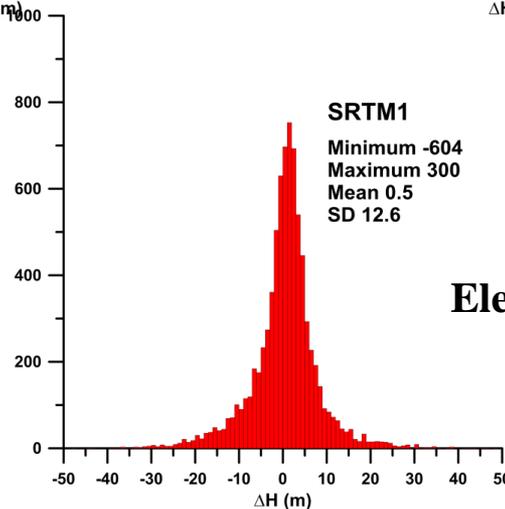
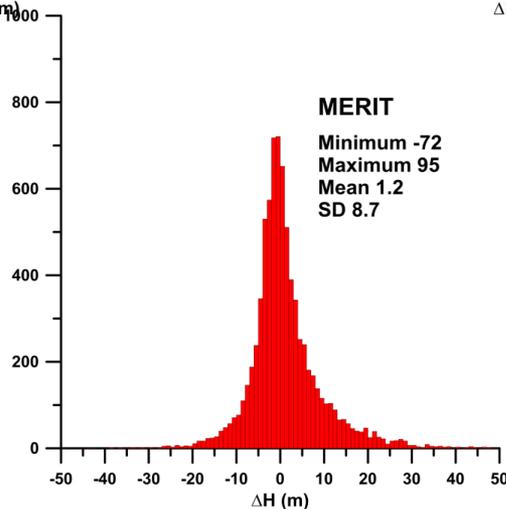
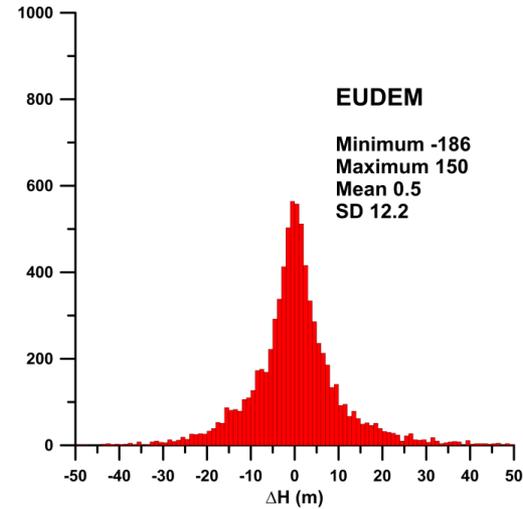
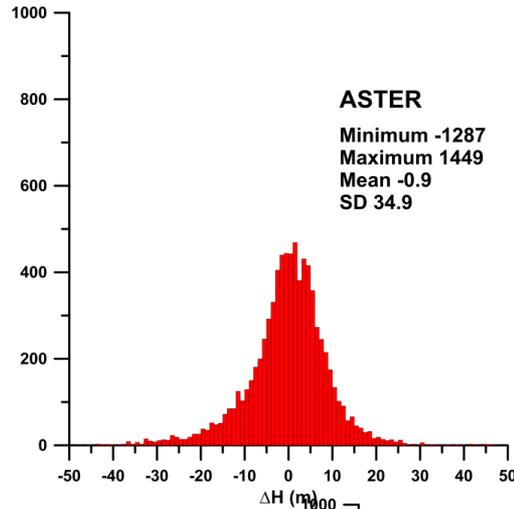
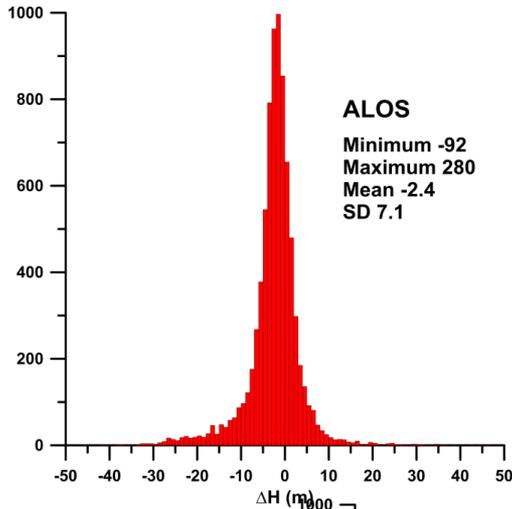
## 2. Getting available detailed elevation models

For most countries we were able to use local detailed DEMs with the resolution of 10 – 20 m for the nearest area.



## 2. Getting available detailed elevation models

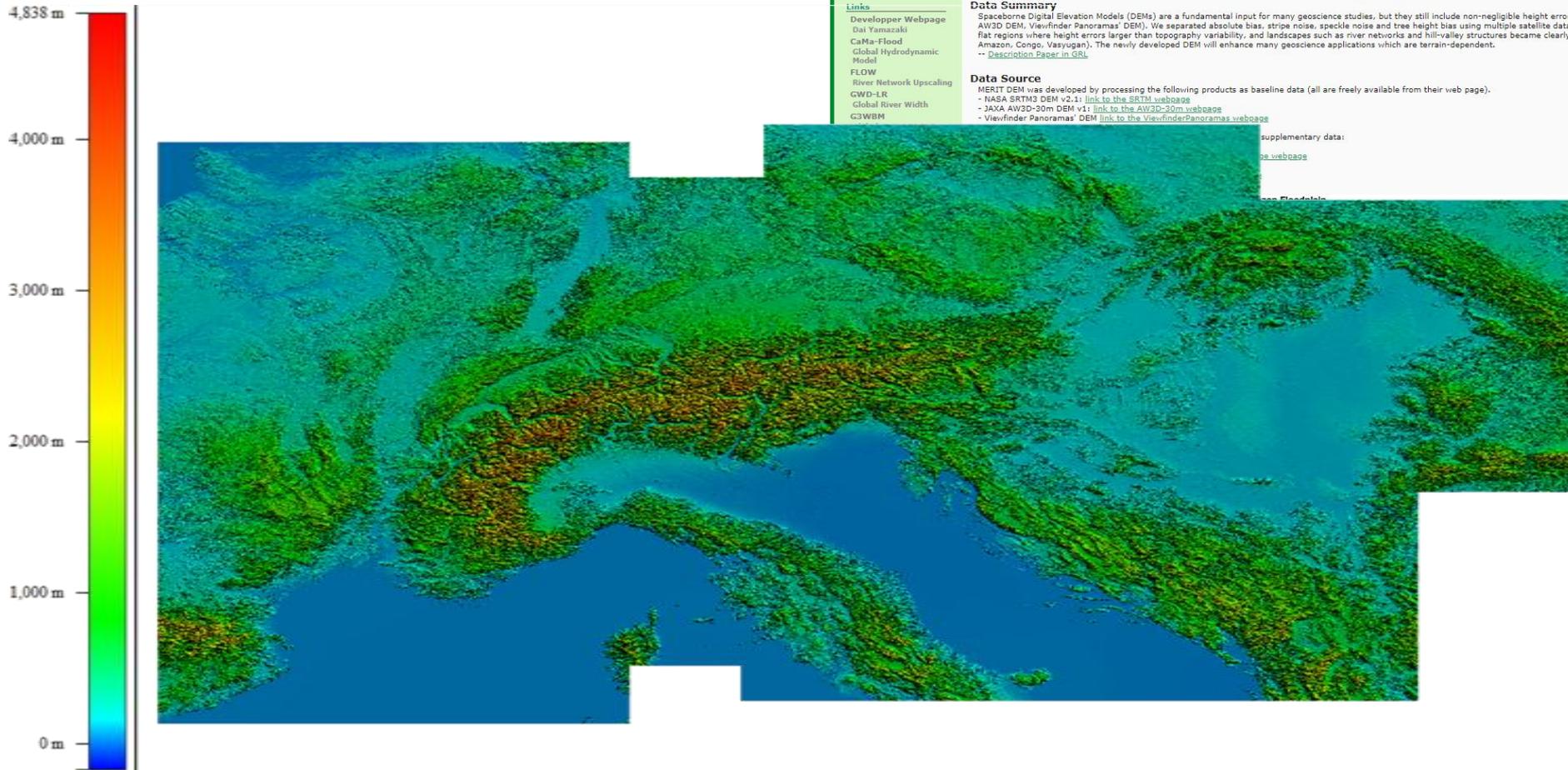
In the case of unavailability of local models, we had to choose the best available global DEM. MERIT model was selected based on the analysis.



Elevation test - Swiss database  
(7962 points)

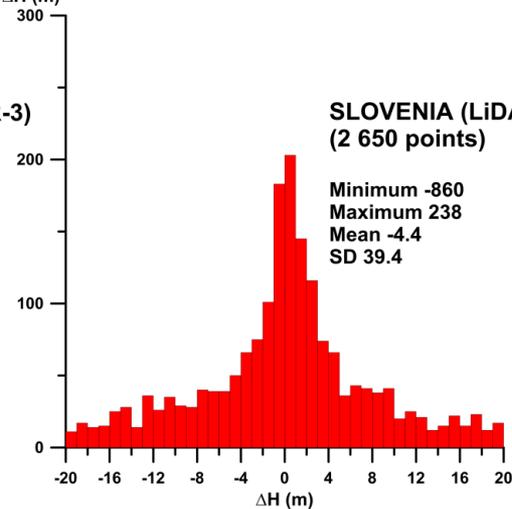
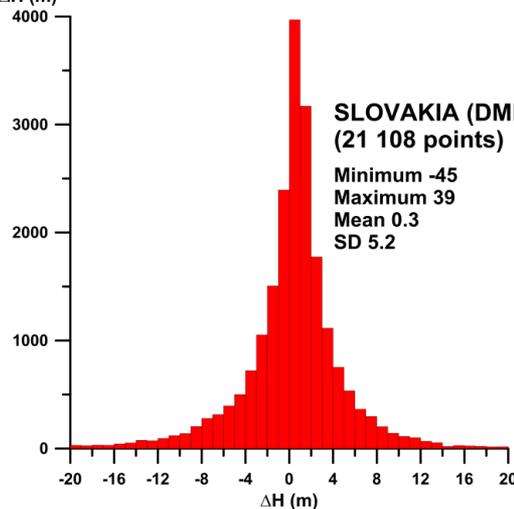
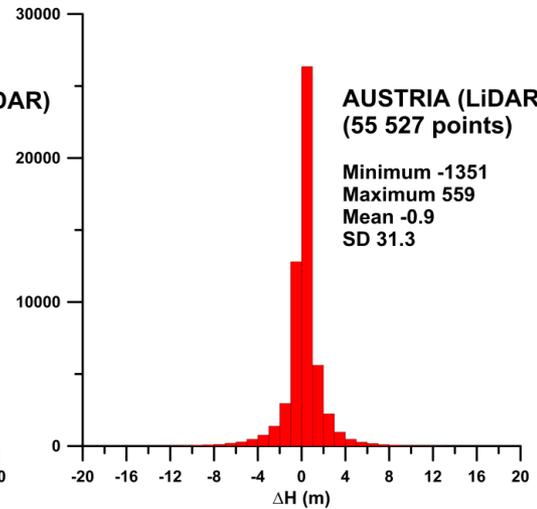
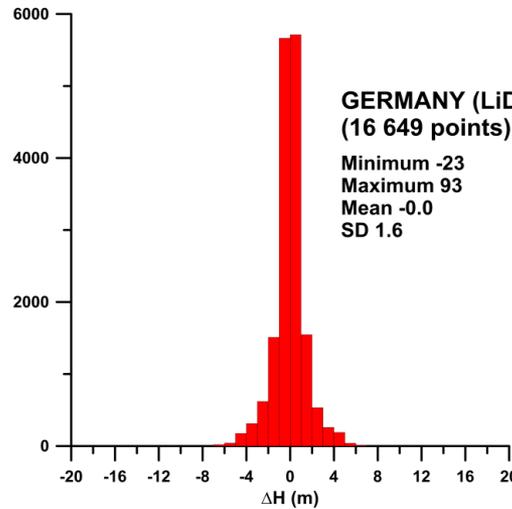
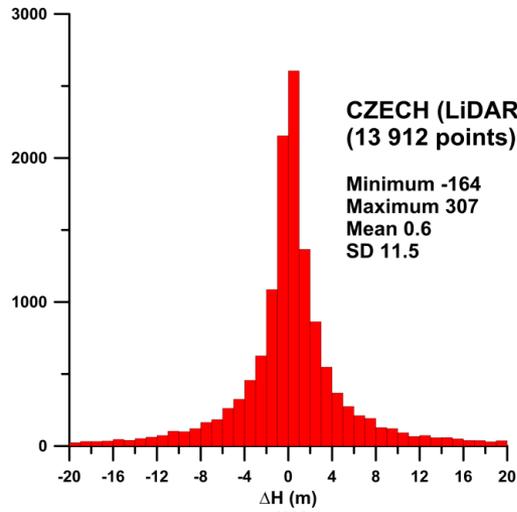
## 2. Getting available detailed elevation models

### MERIT DEM – 3 sec (resampled to 1 sec / 25 m)



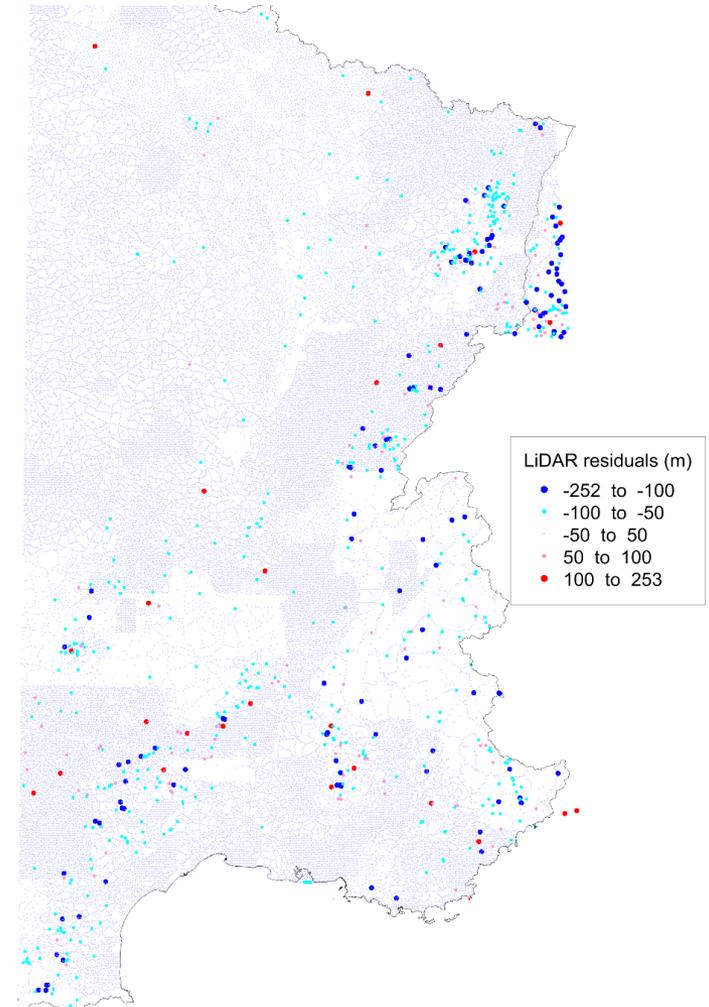
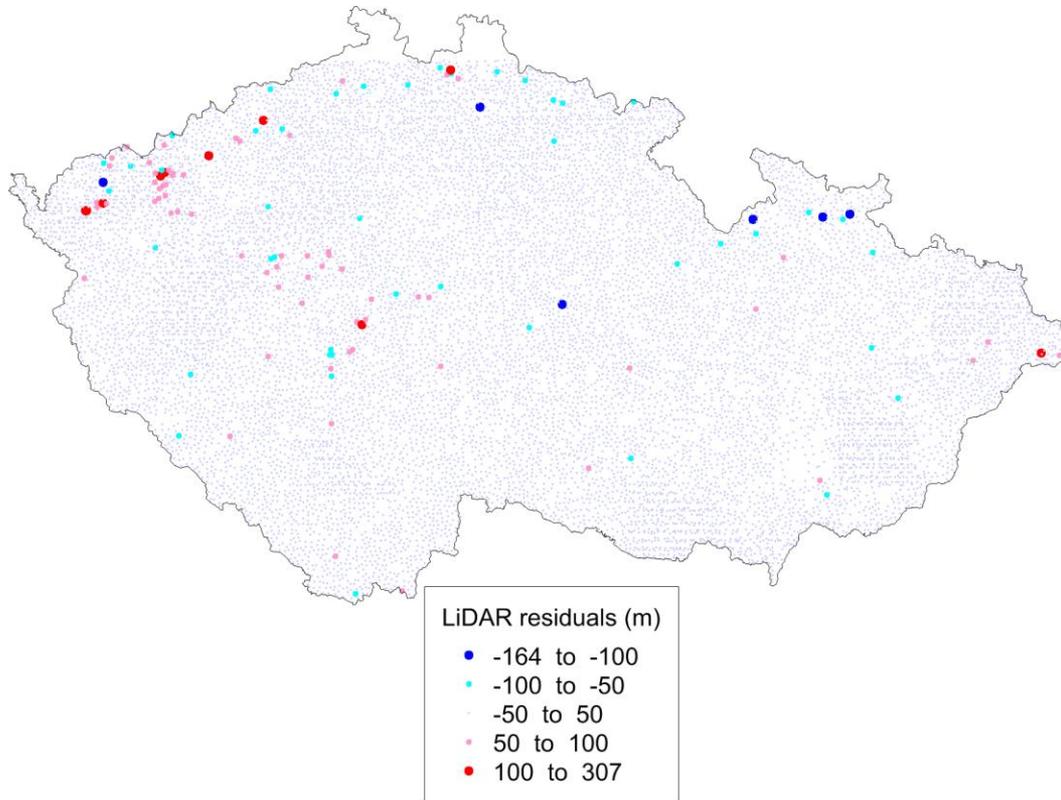
### 3. Quality control of input data

Quality control was performed based on the height differences between the point data and particular local elevation model.



### 3. Quality control of input data

Several thousand points with height residuals higher than chosen threshold ( $\pm 50$  m) were excluded.



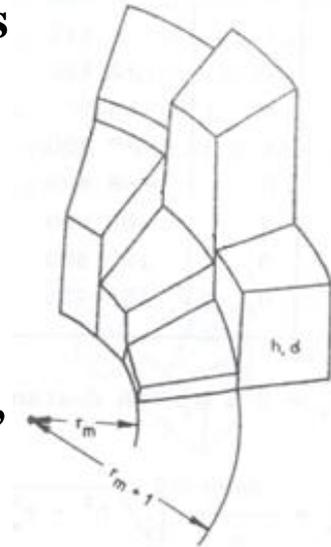
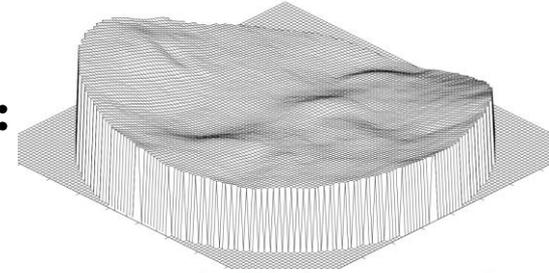
Quality control of input data –  
examples from the Czech and French  
databases

## 4. Mass correction calculation

Mass correction (topographic effect) calculated up to 166.7 km, density 2670 kg/m<sup>3</sup>, ellipsoidal heights and ellipsoidal DEMs used

Program Toposk (Slovak group) – four concentric zones:

- inner zone T1 (0-250m) – 3D polyhedral body
- intermediate zone T2 (250-5240 m) – vertical cylinder segments
- outer zones T31 (5.24-28.8km) and T32 (28.8-166.7km) – spherical layer segments

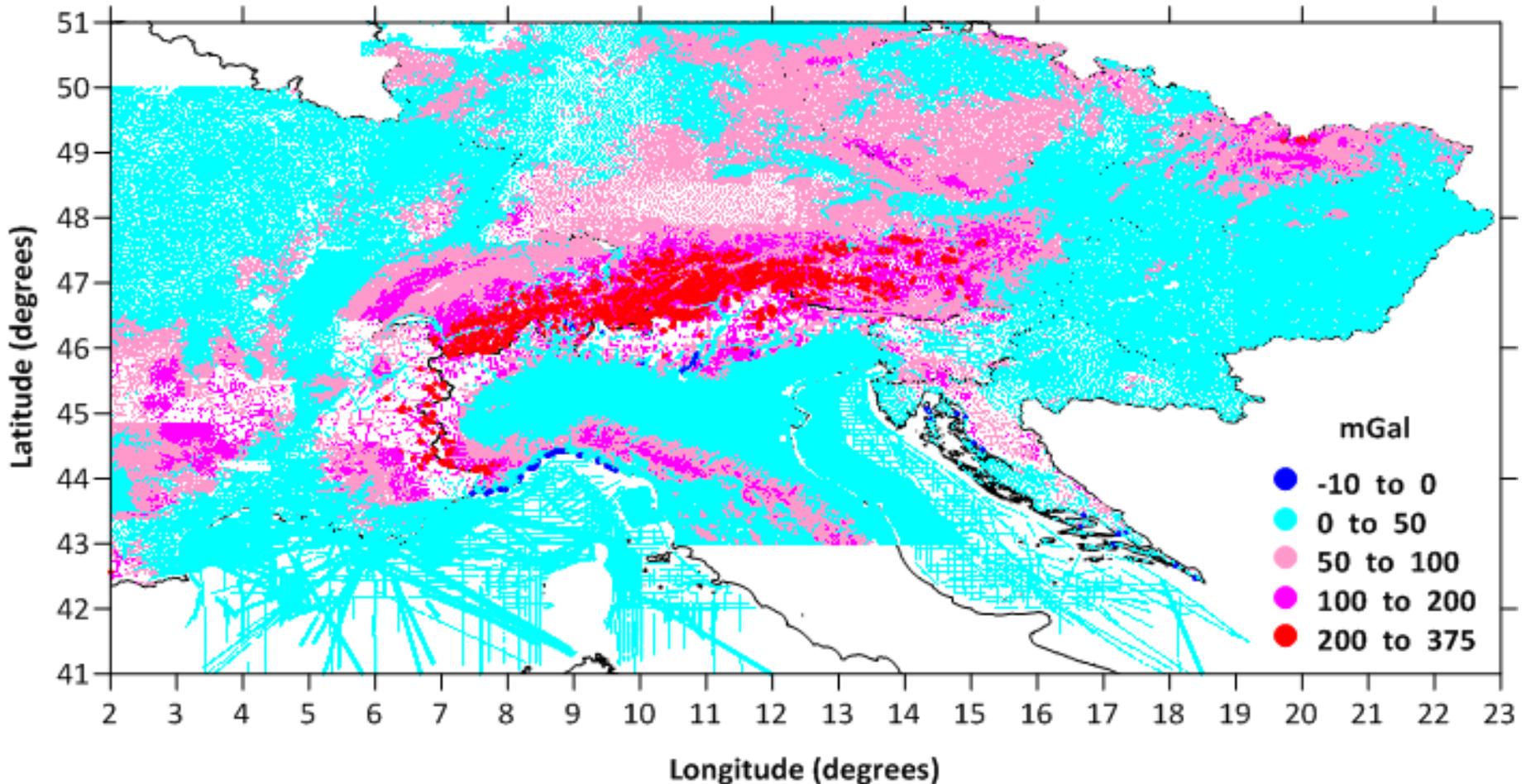


Inner zone (0-250m) - local DEMs (LiDAR) or MERIT (1x1 sec),  
outer zones – MERIT (3x3 sec), SRTM (30x30 sec)

Independent comparison with TriTop software (German group) was performed in selected areas. A good coincidence was achieved.

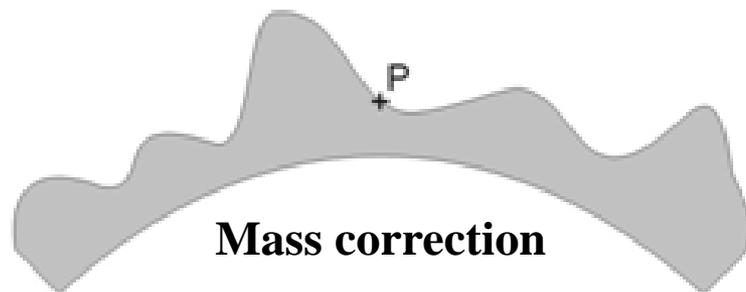
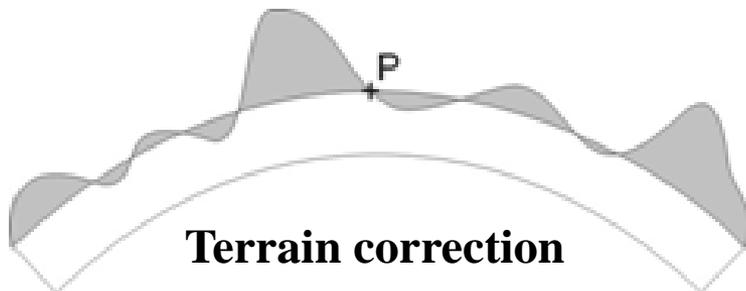
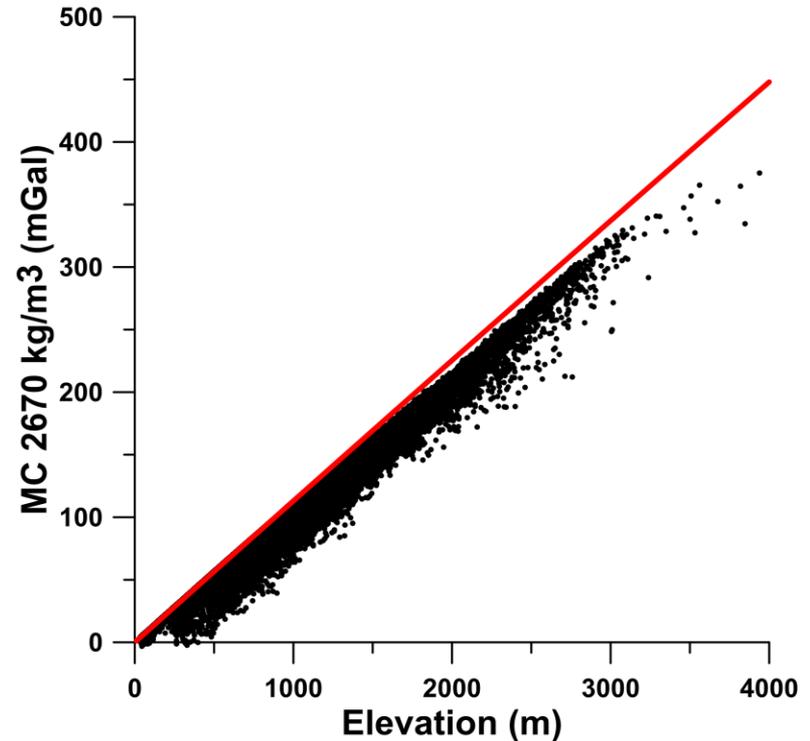
## 4. Mass correction calculation

Mass corrections (density  $2670 \text{ kg/m}^3$ ) reach values up to  $375 \text{ mGal}$ , with small negative values appearing along the coast or in deep valleys



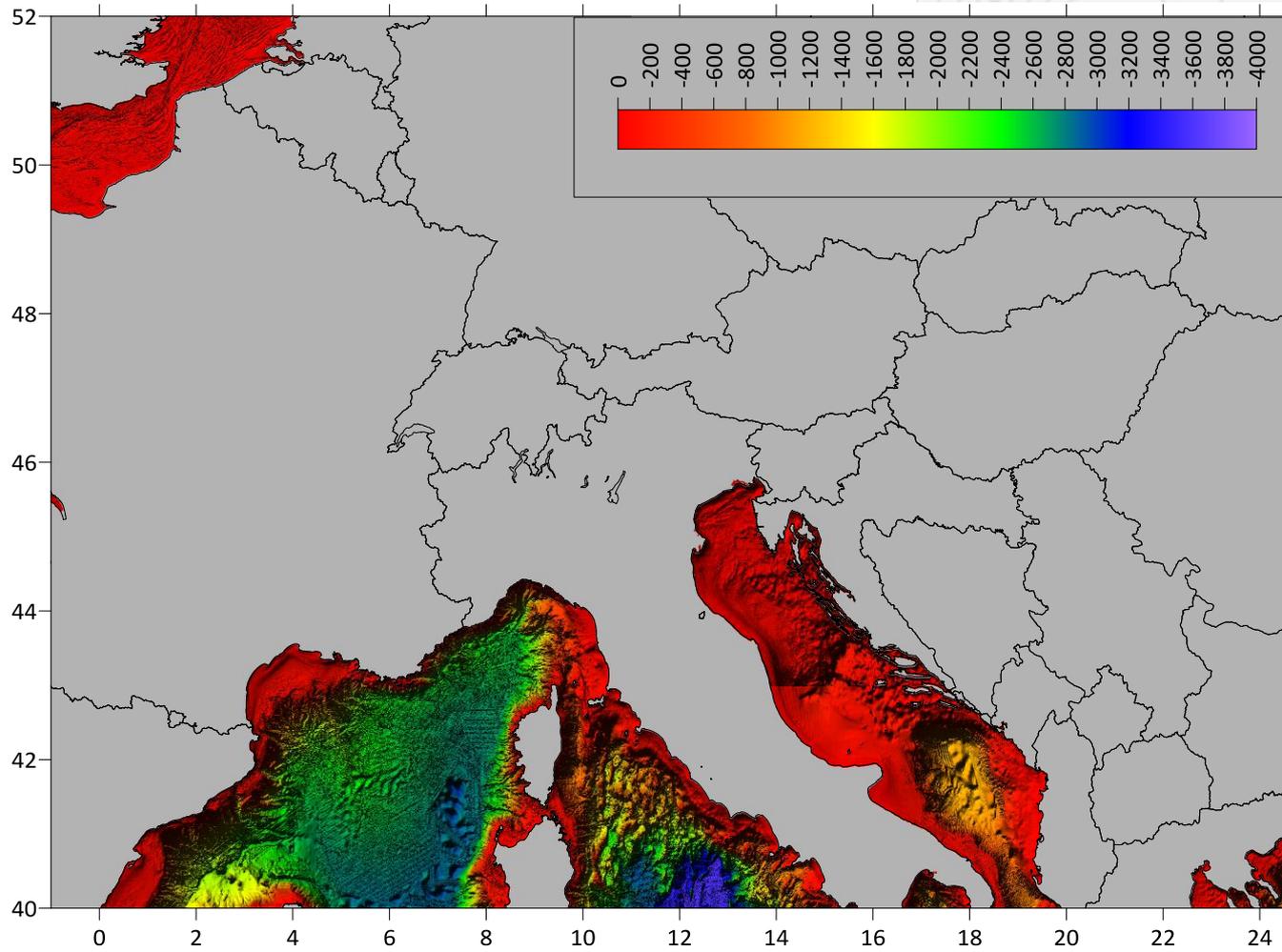
## 4. Mass correction vs. terrain correction

Mass corrections show obvious height dependence. Their differences from the gravitational effect of truncated spherical layer (= classic terrain correction) reach almost 100 mGal



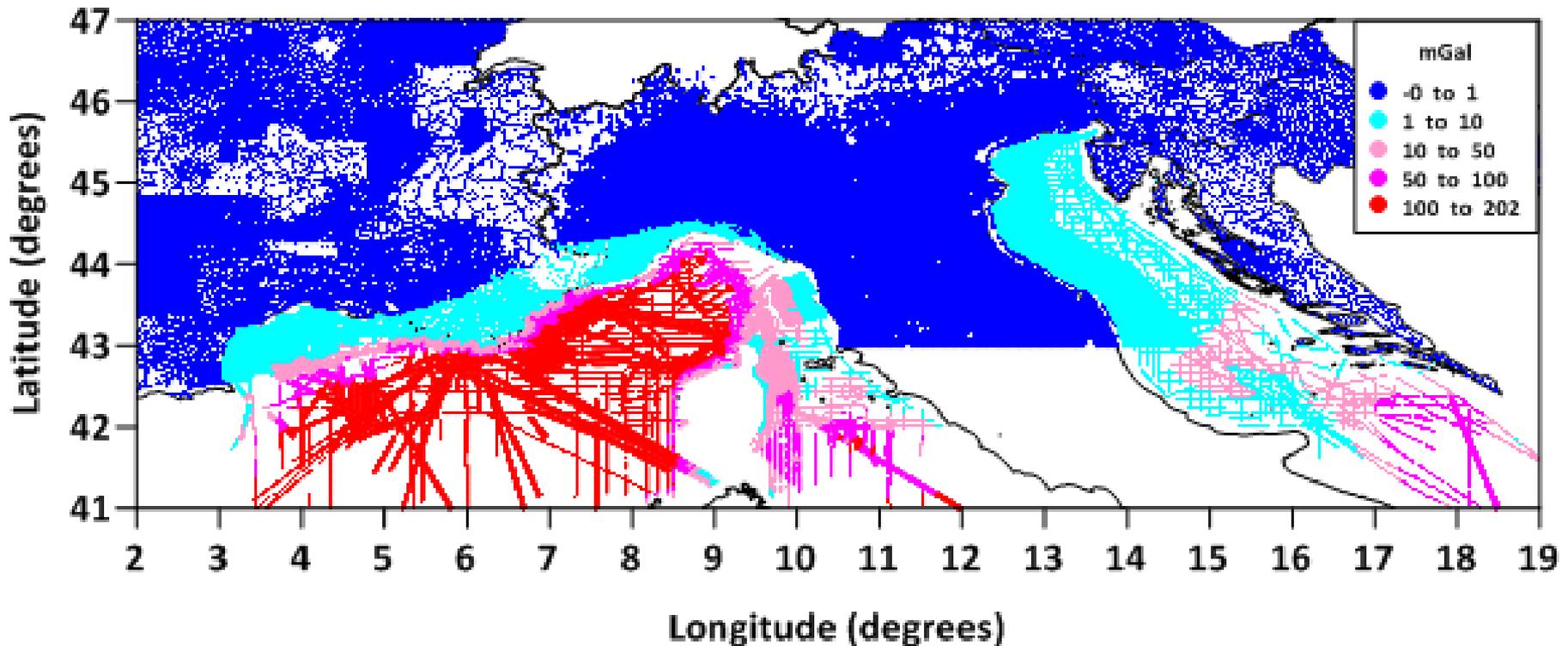
# 5. Bathymetric correction calculation

Bathymetric correction calculated up to 166.7 km, density -1640 kg/m<sup>3</sup>, bathymetry model EMODnet - 3.75 sec used



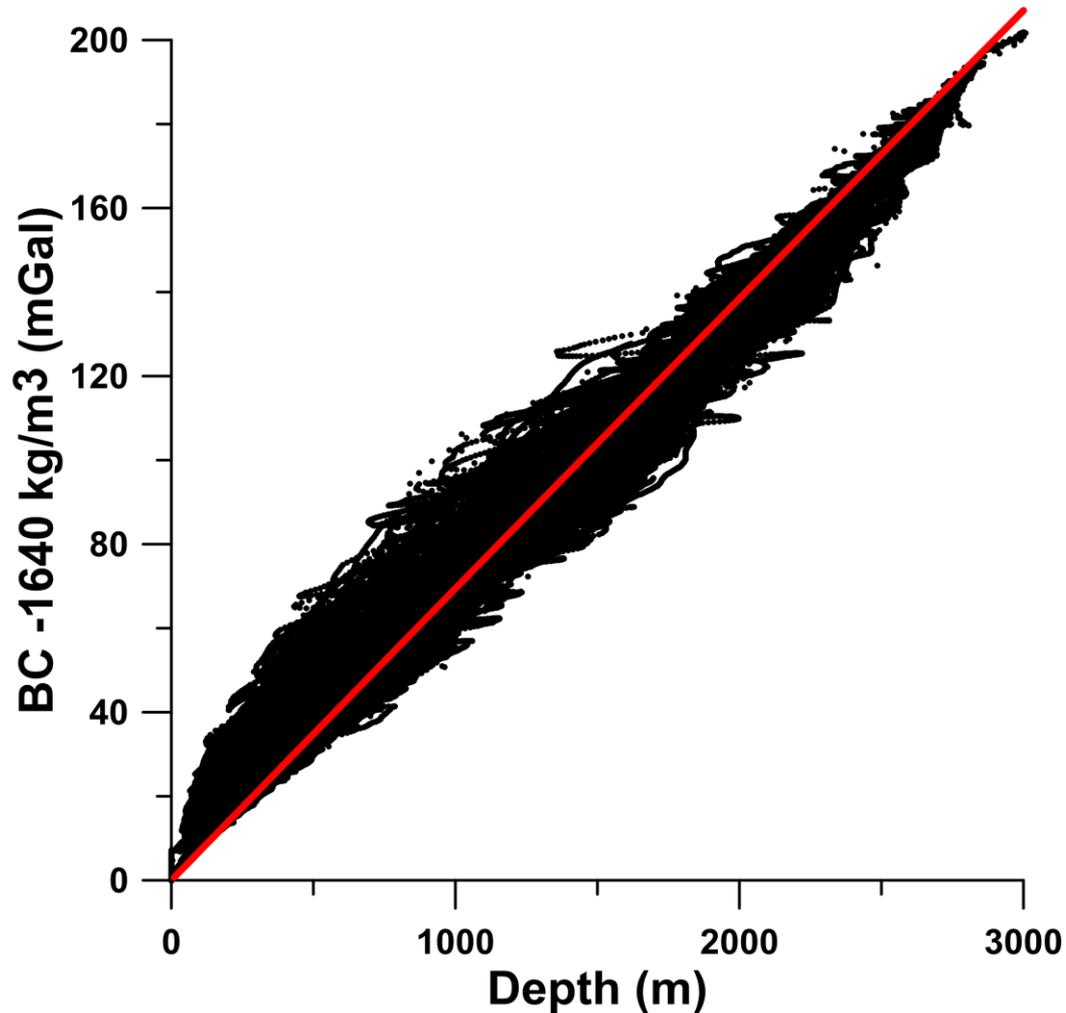
## 5. Bathymetric correction calculation

Bathymetric corrections (density  $-1640 \text{ kg/m}^3$ ) reach values up to 202 mGal (marine areas)



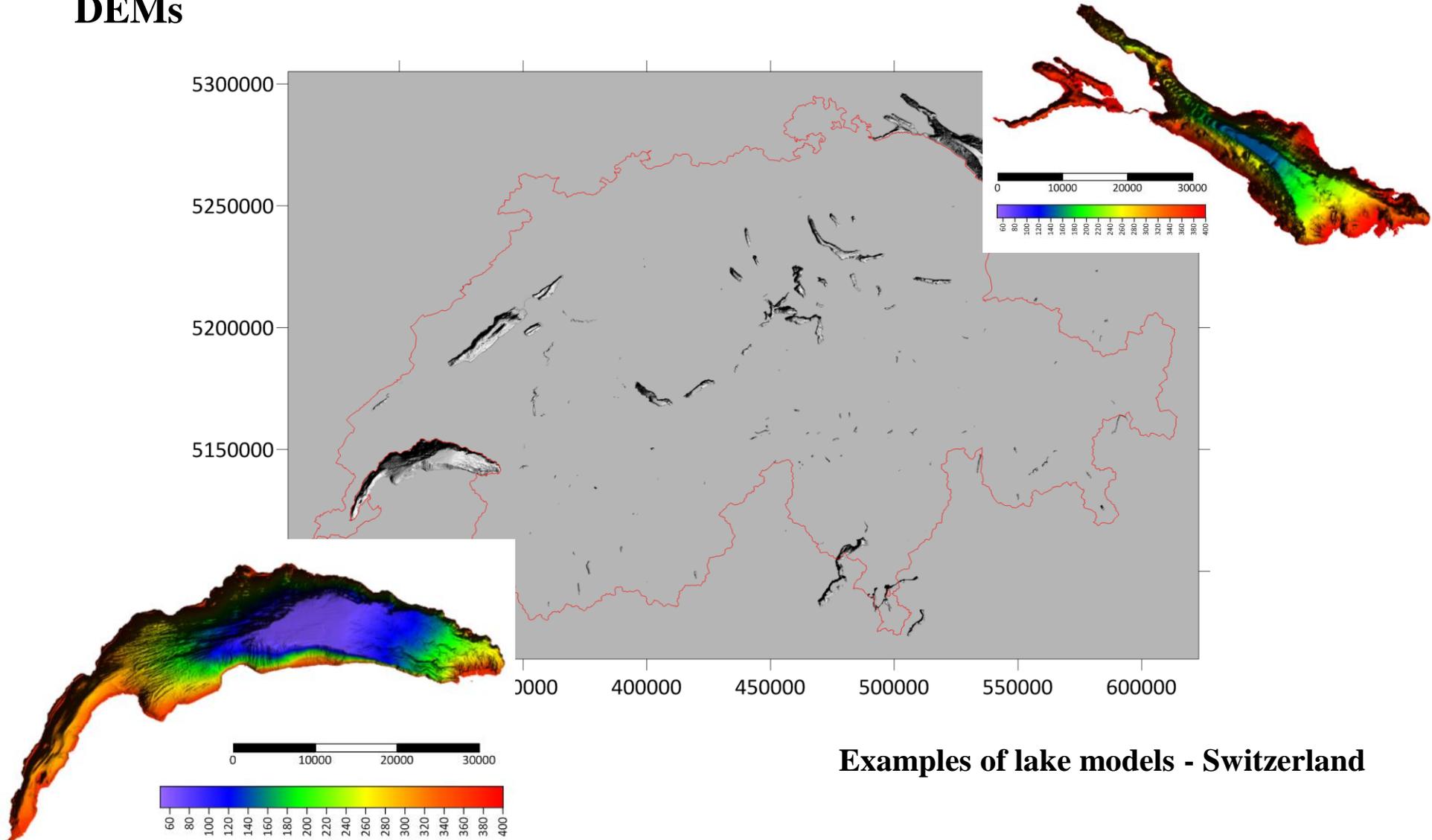
## 5. Bathymetric correction – Toposk vs. “Bouguer” approximation

Comparison of Toposk-calculated bathymetric correction and its simple “Bouguer” approximation shows relatively significant differences ( $\pm 30$  mGal)



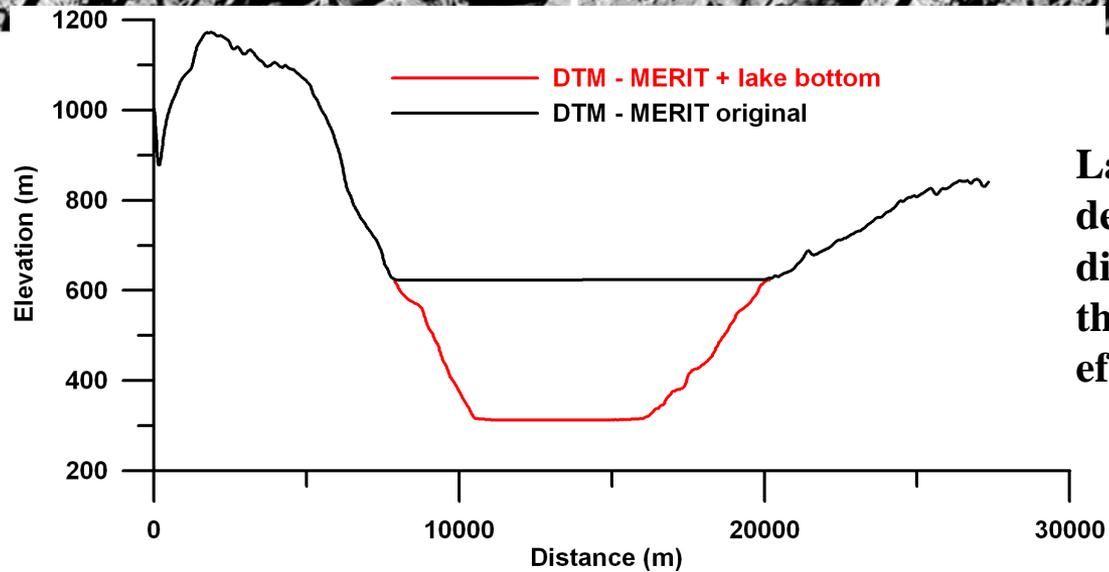
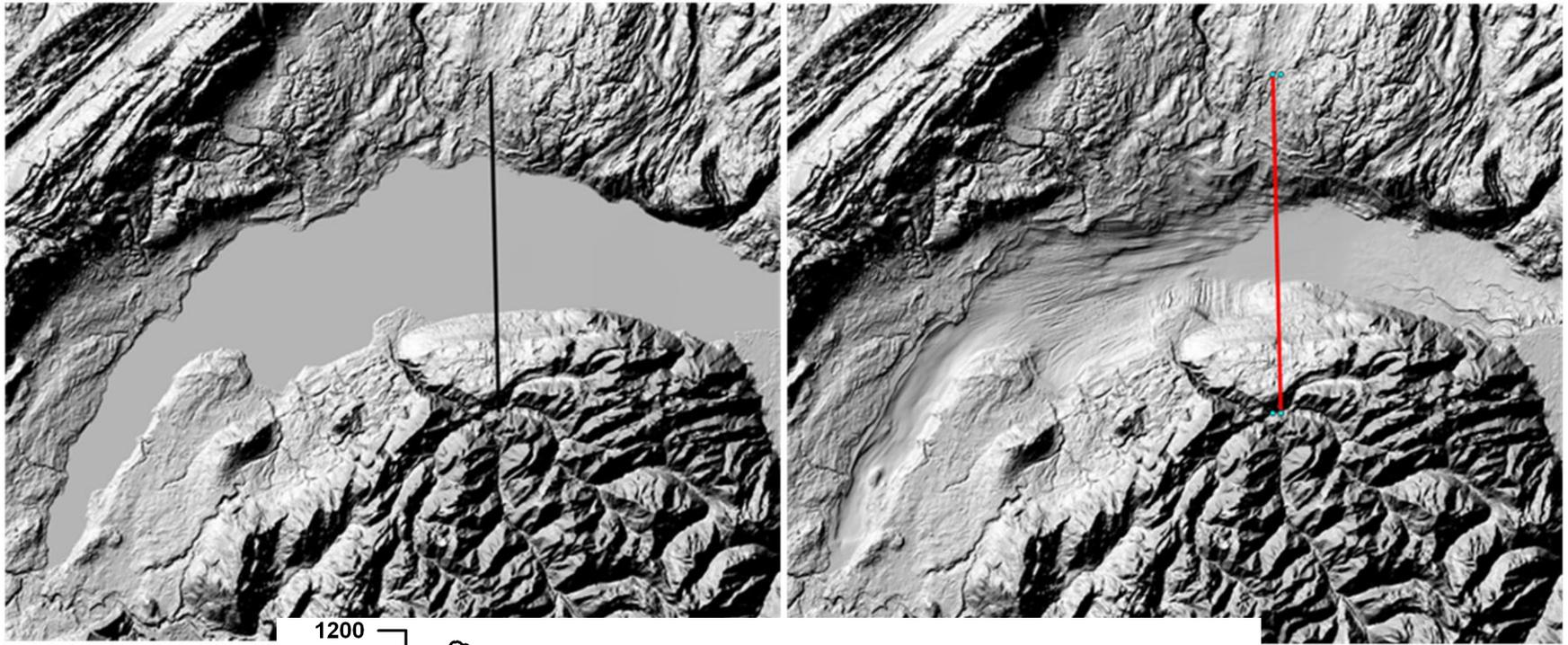
# 6. Lake corrections

Local models of large alpine lakes were used in combination with existing DEMs



Examples of lake models - Switzerland

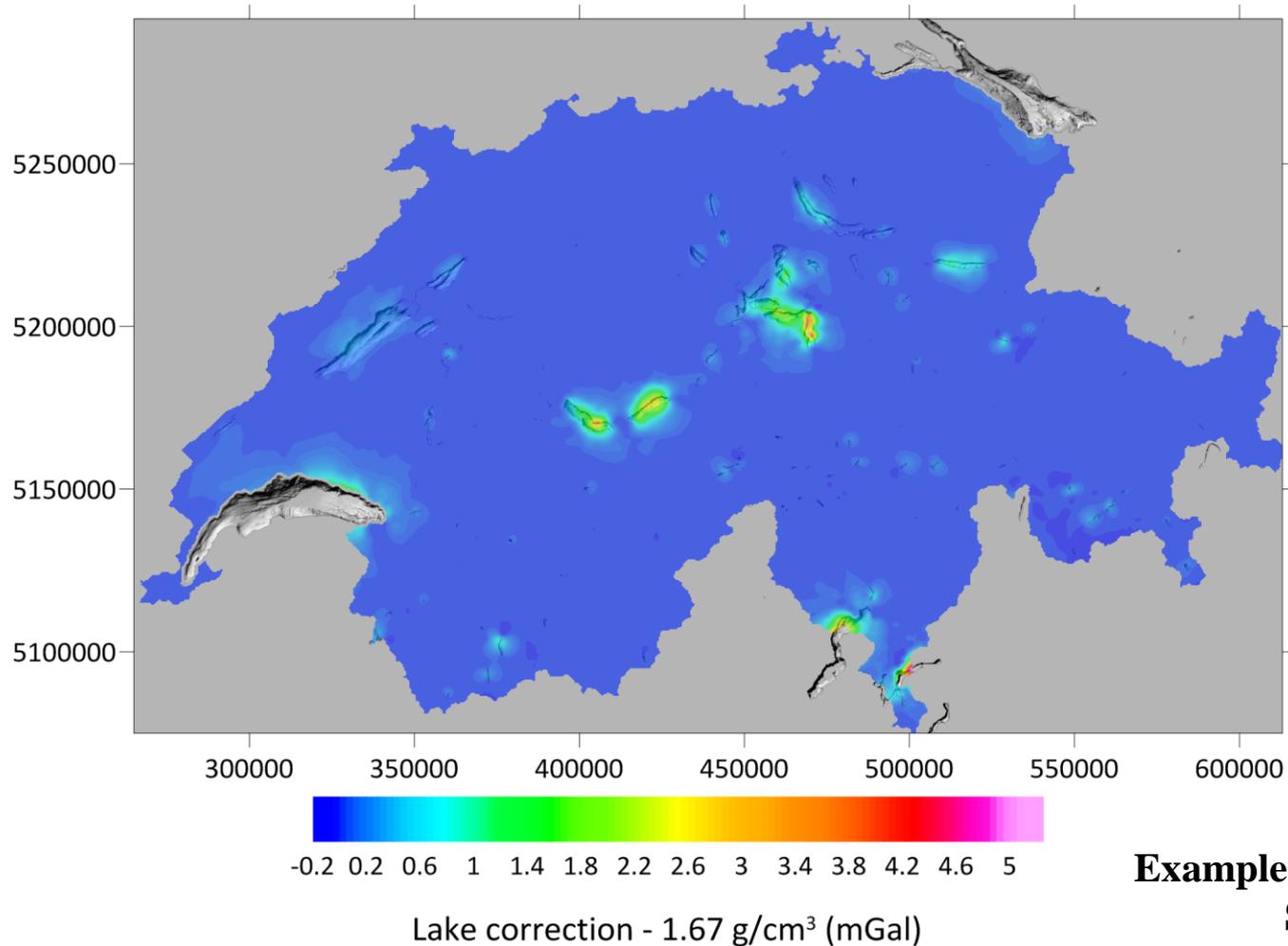
## 6. Lake corrections – calculation method



**Lake correction is defined as the difference between the gravitational effect of two models**

## 6. Lake corrections

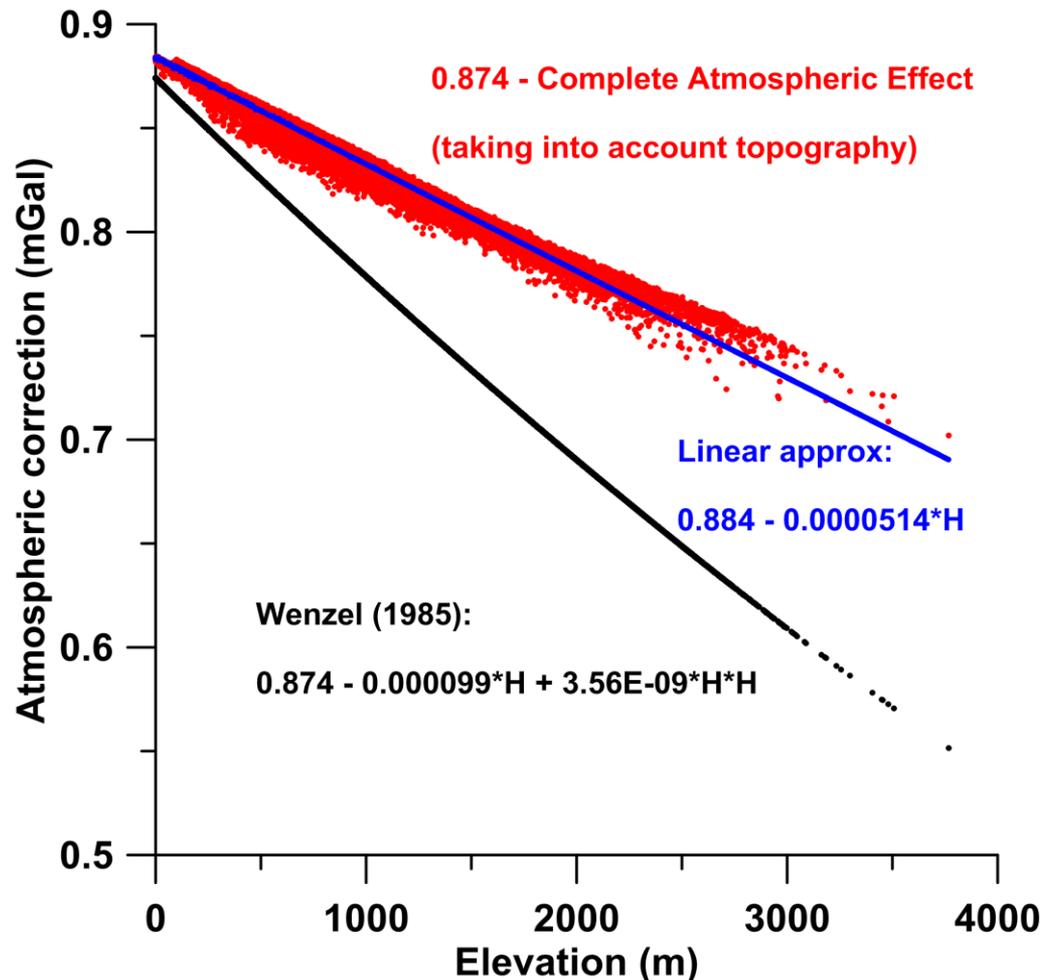
**Lake corrections (density  $-1670 \text{ kg/m}^3$ ) reach values up to approx. 5 mGal locally, with small negative values appearing in deep valleys (below lake level)**



**Example of lake corrections -  
Switzerland**

## 7. Atmospheric correction – comparison of different approaches

Atmospheric correction based on the real atmosphere model taking into account real topography (after Mikuška et al. 2008); maximum differences of about 0.15 mGal compared to the classical approach (Wenzel, 1985).



## 8. Complete Bouguer anomaly calculation

Basic formula is adopted from the paper Meurers et al. (2001), with small changes:

$$BA(\lambda, \varphi, h_E) = g(\lambda, \varphi, h_E) - \gamma(\varphi, h_E) - \delta g_M(\lambda, \varphi, h_E) + \delta g_B(\lambda, \varphi, H) + \delta g_A(\lambda, \varphi, H), \quad (1)$$

where:  $g()$  – observed absolute gravity acceleration value,  $\gamma()$  – normal gravity acceleration at height  $h_E$ ,  $\delta g_M()$  – mass correction (gravitational effect of topographical masses) for density  $2670 \text{ kg/m}^3$ ,  $\delta g_B()$  – bathymetric correction for density  $1640 \text{ kg/m}^3$ ,  $\delta g_A()$  – simplified atmospheric correction,  $\lambda$  – longitude (ETRS89),  $\varphi$  – latitude (ETRS89),  $h_E$  – ellipsoidal height (GRS80),  $H$  – physical height.

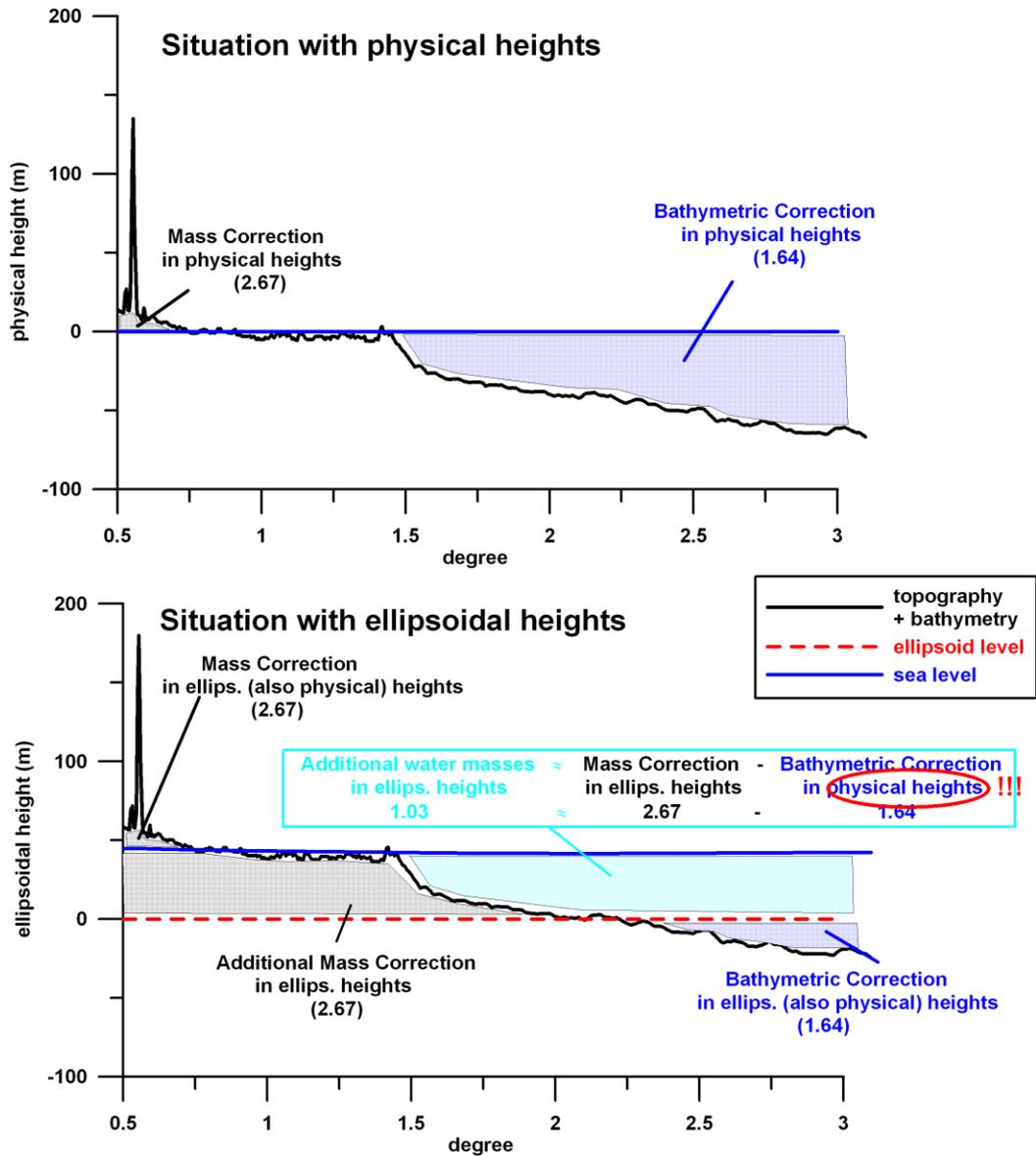
Normal gravity acceleration  $\gamma(\lambda, \varphi, h_E)$  results from a Taylor series expansion to the 2<sup>nd</sup> order both in height and also geometrical flattening (Wenzel 1985, Li and Götze 2001):

$$\gamma(\varphi, h_E) = \gamma_0(\varphi) + \left. \frac{\partial \gamma}{\partial h_E} \right|_0 h_E + \frac{1}{2} \left. \frac{\partial^2 \gamma}{\partial h_E^2} \right|_0 h_E^2, \quad (2)$$

where  $\gamma_0()$  is the well-known Somigliana formula for the normal gravity acceleration of a rotational ellipsoid at its surface (Somigliana 1929. Heiskanen and Moritz 1967):

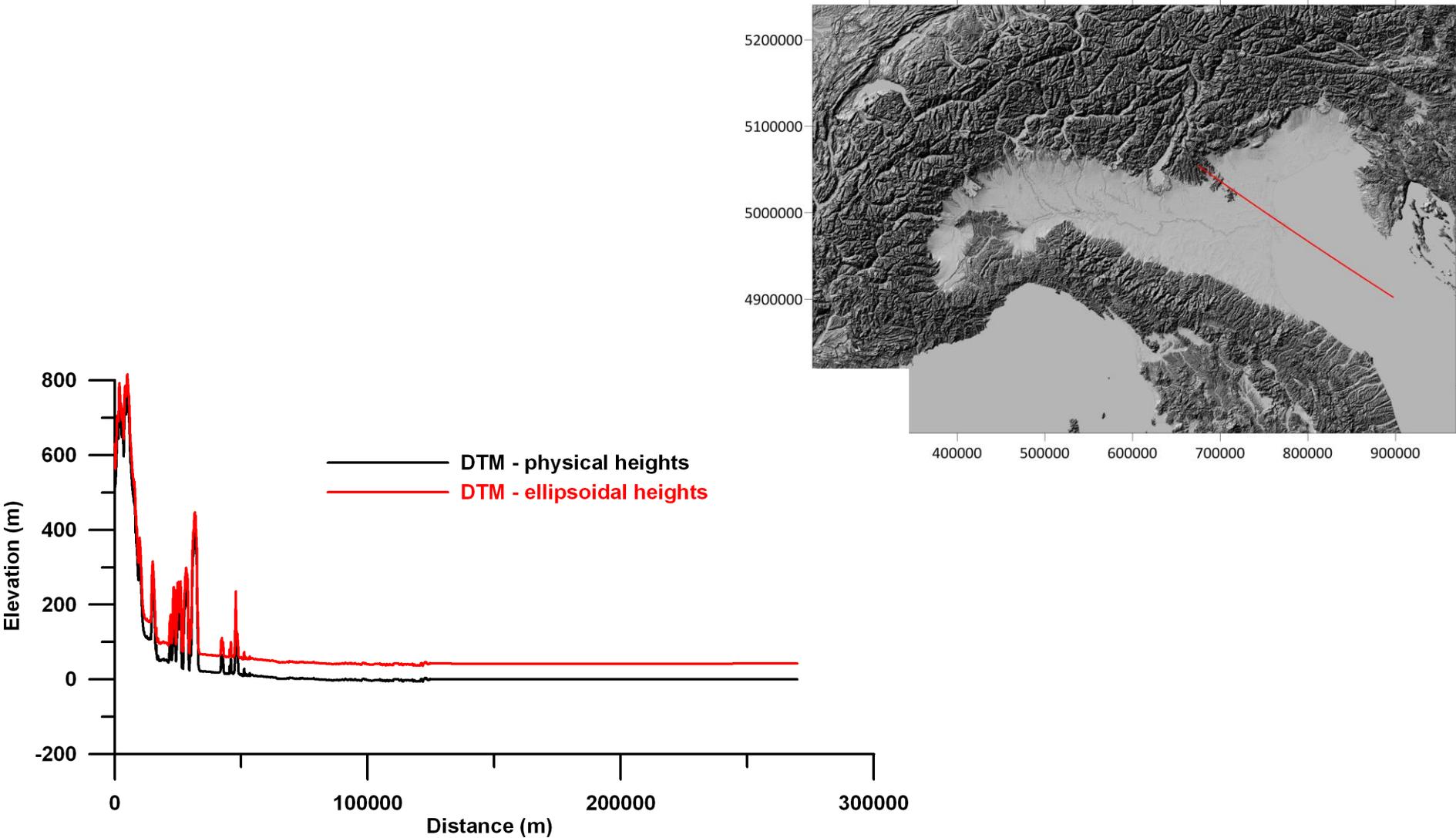
**Innovative concept of ellipsoidal heights was used**

# 8. Ellipsoidal vs. physical heights concept



## 8. Ellipsoidal vs. physical heights concept

Additional topography/water masses need to be carefully considered (with the right density), depending on DEMs used

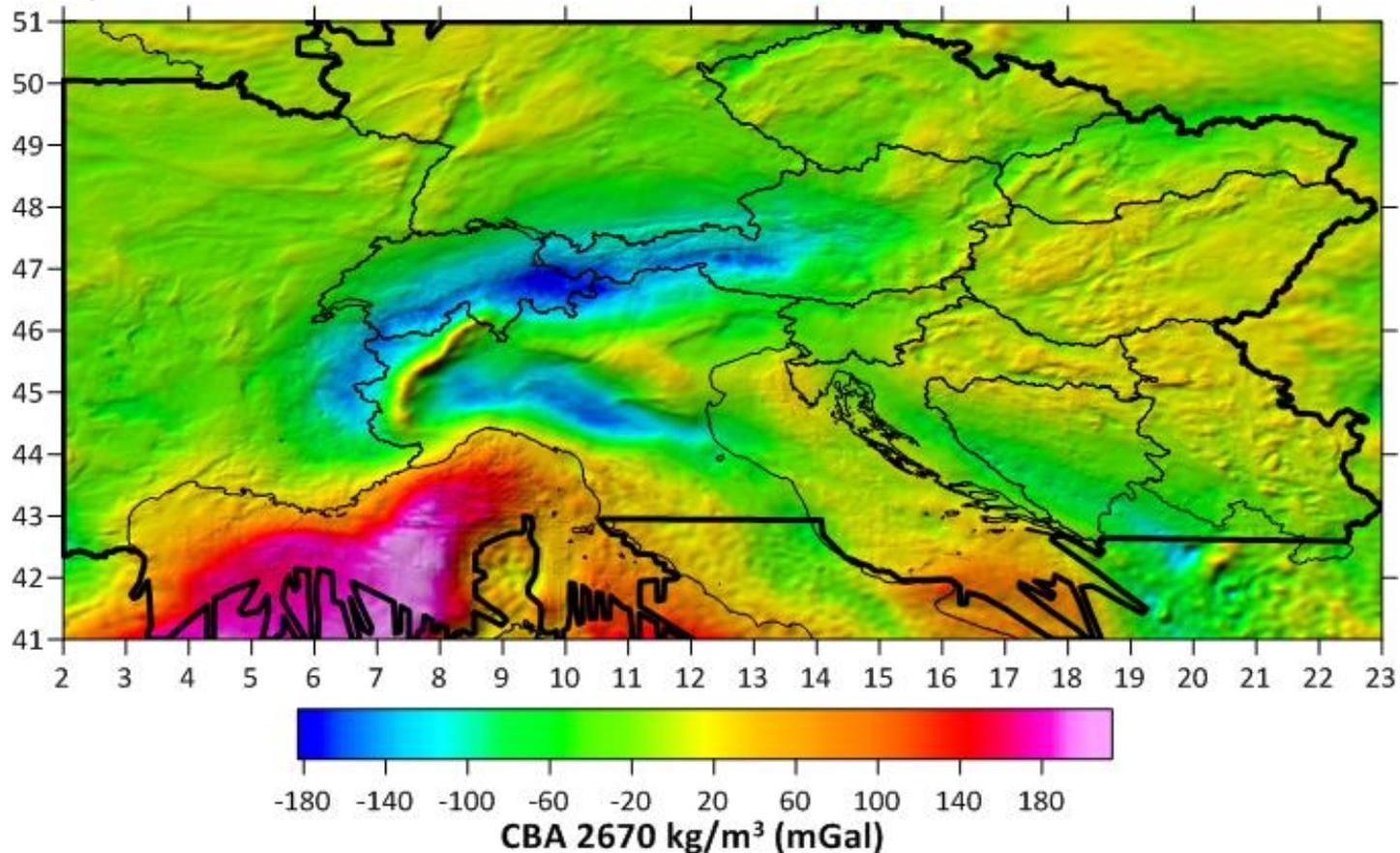


## 9. Merging individual databases into a single map

First version of the merged CBA map prepared.

Part of the former Yugoslavia was filled by digitizing an old CBA map

Peripheral areas (outside the thick black line) filled with data derived from EIGEN-6C4



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**Andrej Gosar – Slovenian Environment Agency, Slovenia**

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**Pavel Novák – University of West Bohemia, Czech republic**

**Eszter Szűcs – Geodetic and Geophysical Institute, Hungary**

**Matej Varga – University of Zagreb, Croatia**

**Thank you for your attention!**