

# Impact of the 536/540 CE double volcanic eruption event on the 6th-7th century climate using model and proxy data

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# Motivation and Background

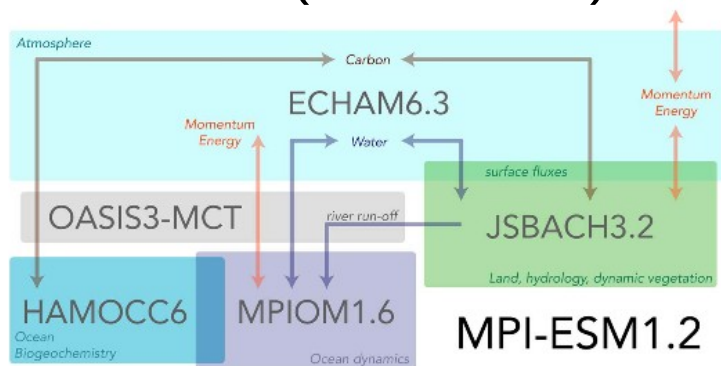
- Volcanic eruptions are important climate drivers (Crowley et al., 2000; Robock 2000)
- Very cold period after the 536/540 CE double eruption event in the mid-6<sup>th</sup> century (Larsen et al., 2008; Sigl et al., 2015)
- Evidence from multiple tree-ring records for a centennial cooling up to 660 CE (Büntgen et al., 2016)
- Previous MPI-ESM simulations show a decrease in surface temperature and an increase in Arctic sea-ice up to 15 years (Toohey et al., 2016)

Research question:

**Can we force a century long lasting cooling due to major volcanic eruptions in the 6th-7th century in earth system models?**

# Model experiment and set-up

- MPI-ESM-LR1.2 version for CMIP6/PMIP4 (Mauritsen et al., 2019)
- ECHAM6: T63 → 200x200 km, 47 vertical levels, top @ 80km
- MPIOM: GR1.5 → 150x150 km, 40 vertical levels
- 10 x 160 years → 520-680 CE
- spin up from PMIP4/Past2k run
- PMIP4 volcanic forcing  
(Toohey and Sigl 2017, Jungclaus et al., 2017)
- Anomalies calculated wrt 0-1850 CE (Past2k run)



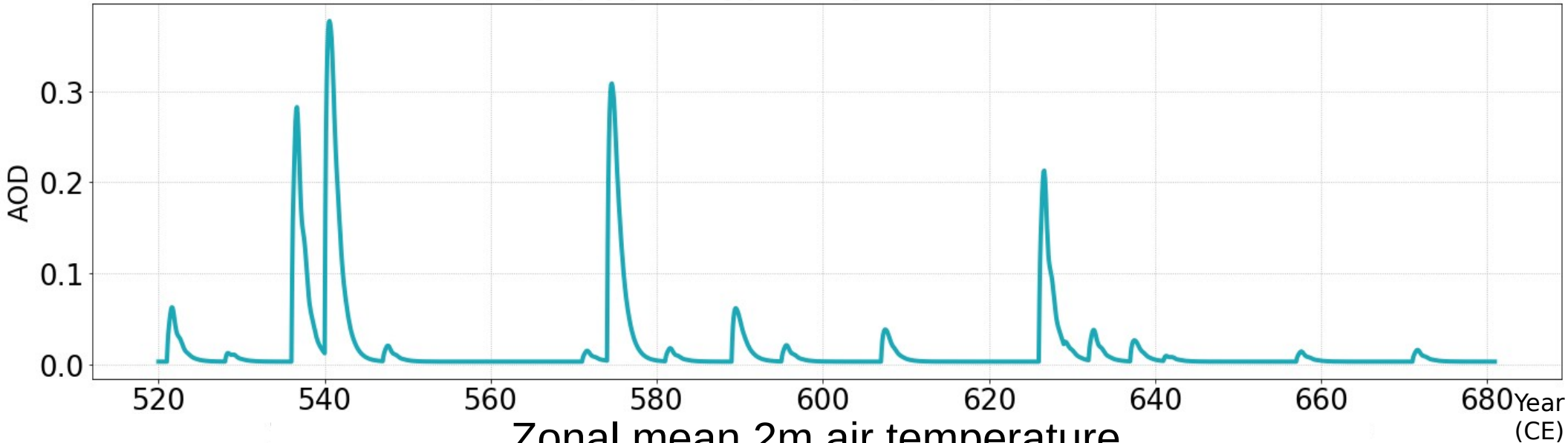
Mauritsen et al. (2019)

| Eruption year CE | Eruption month | S injected | Peak aerosol optical depth in model | Latitude of eruption |
|------------------|----------------|------------|-------------------------------------|----------------------|
| 536              | Jan            | 18.8 Tg    | 0.5                                 | NHext (~45N)         |
| 540              | Jan            | 31.8 Tg    | 0.7                                 | Tropical (~15N)      |
| 574              | Jan            | 24.2 Tg    | 0.6                                 | Tropical             |
| 626              | Jan            | 13.2 Tg    | 0.4                                 | NHext                |

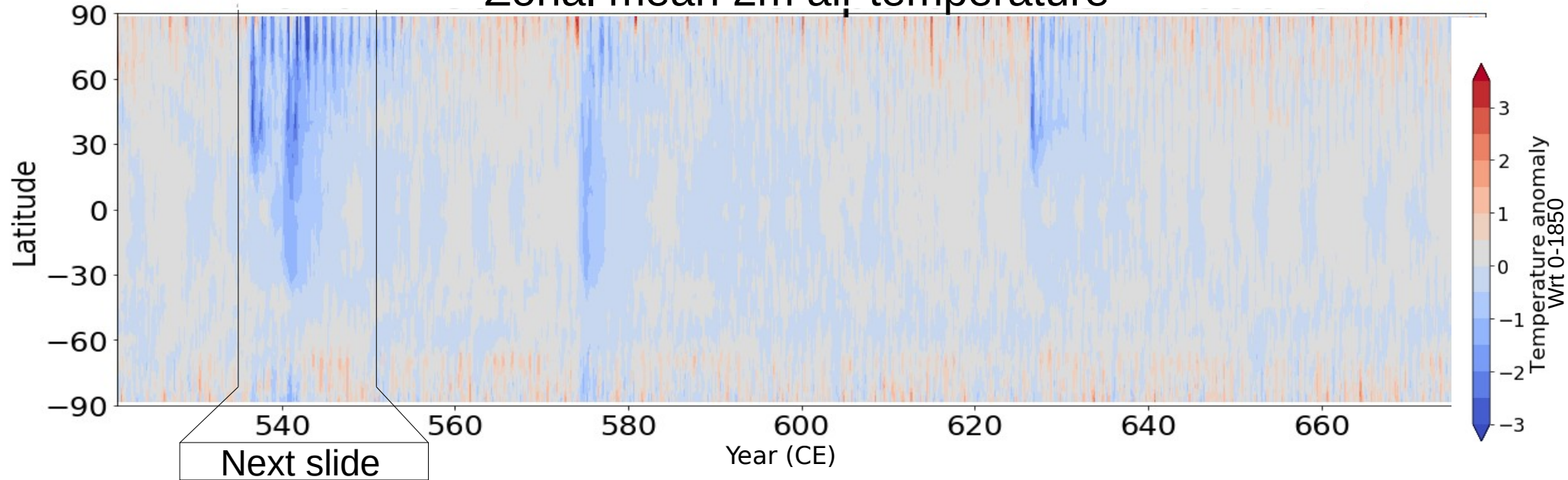
\*Toohey and Sigl (2017): eVolv2k, based on ice-core records + Easy Volcanic Aerosol model + scaling factor (Gao et al. 2006)

# Volcanic forcing - temperature response

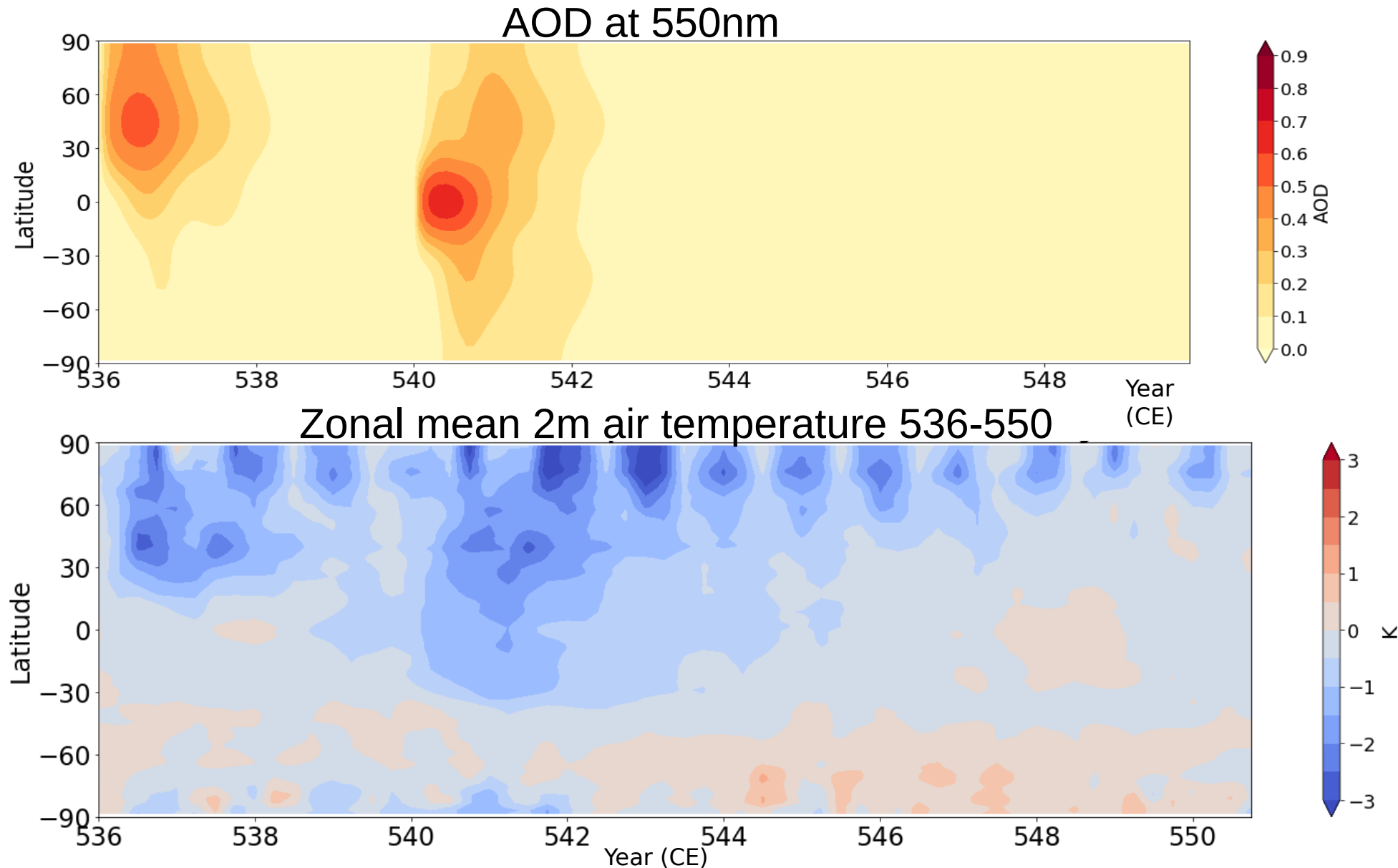
Global mean AOD at 550nm



Zonal mean 2m air temperature



# Volcanic forcing - temperature response

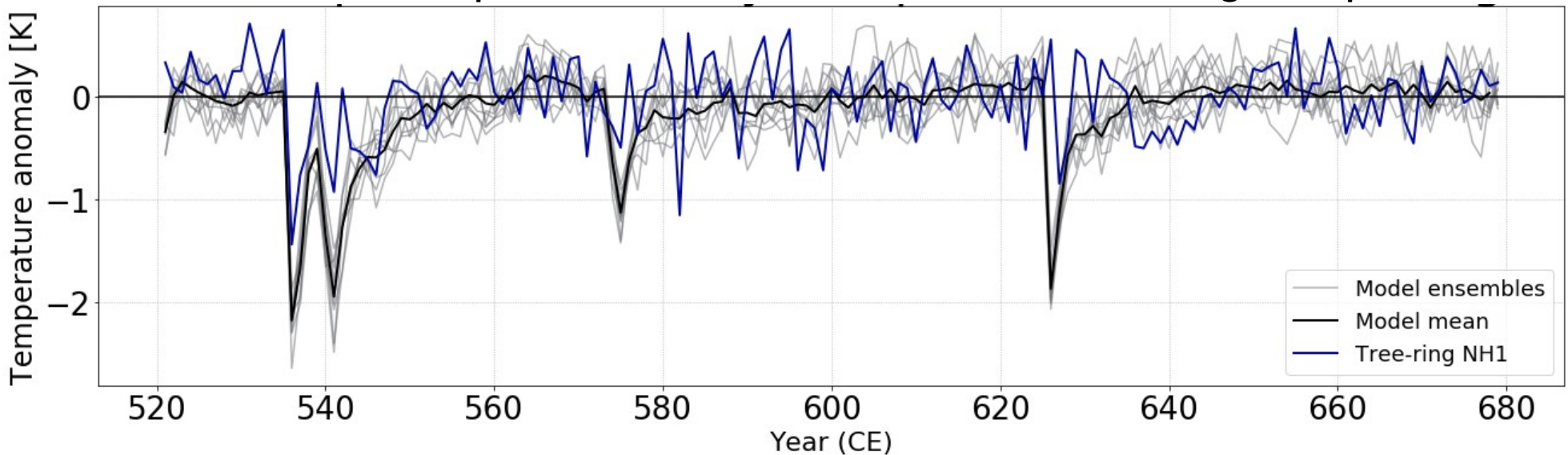


**AOD disappears after ~2 years → temperature anomaly at high latitudes lasts longer  
~20 years → ocean/sea-ice feedback?**



# Northern Hemisphere temperature: Model – Tree-ring comparison

40-75N JJA temperature anomalies Model – tree-ring comparison

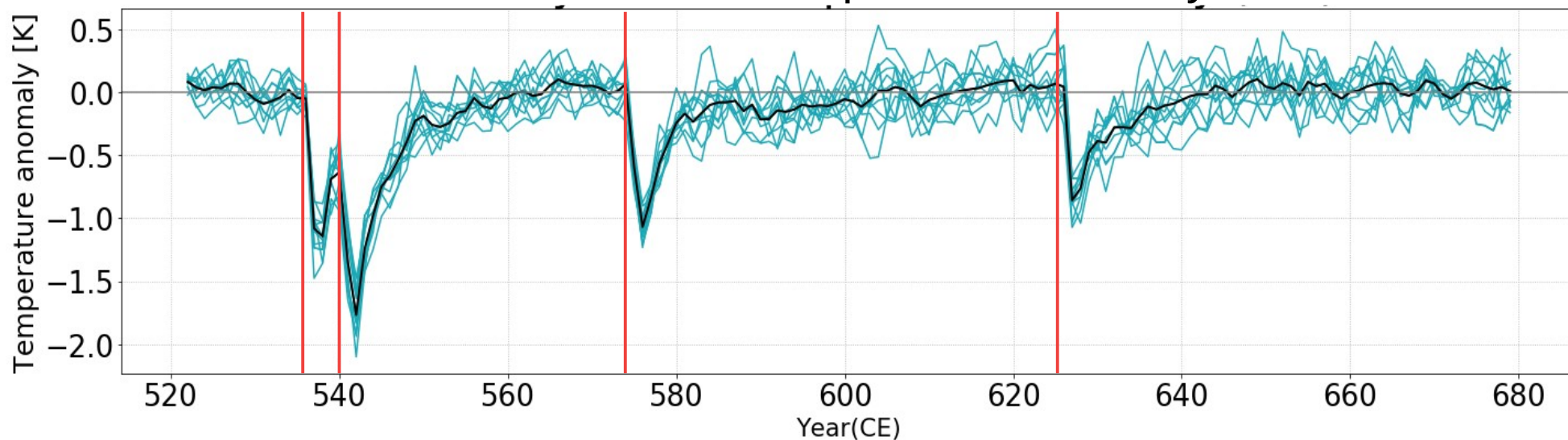


- Tree ring data from Stoffel et al. (2015)
- Anomalies wrt 1961-1990 → (NH1)
- Model data between 40-75N, land only
- JJA anomalies wrt Past2k (0-1850 CE)

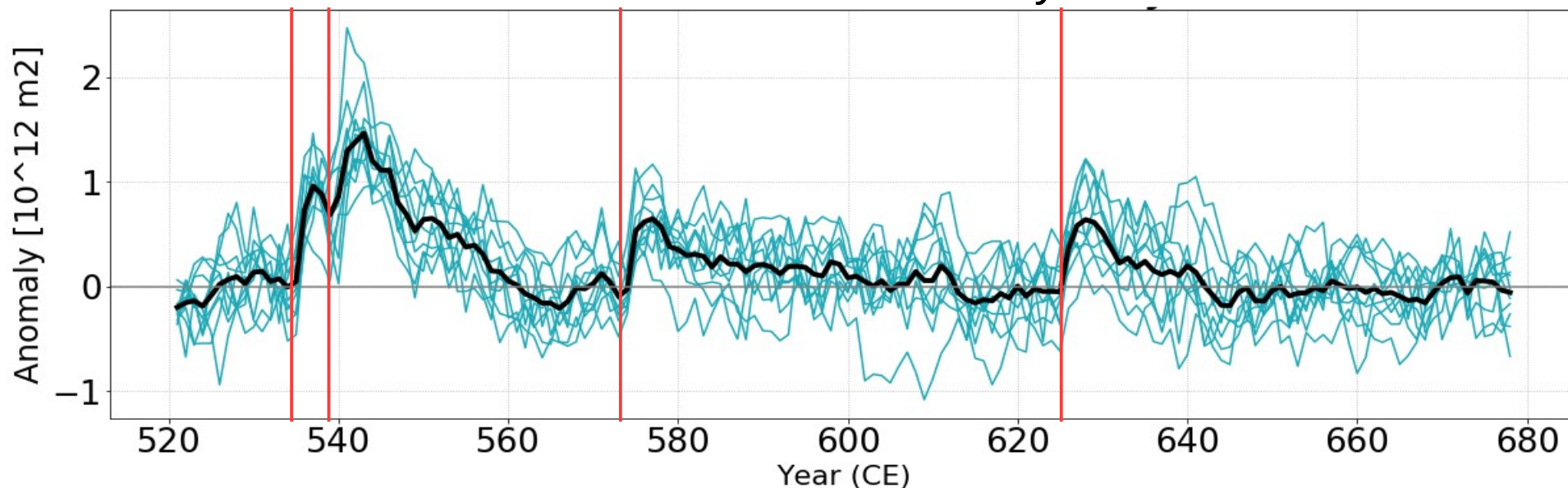
**Good agreement between model simulations and tree-ring data  
~20 year cooling after eruptions**

# NH 2m temperature and sea-ice cover

## NH mean 2m air temperature anomaly



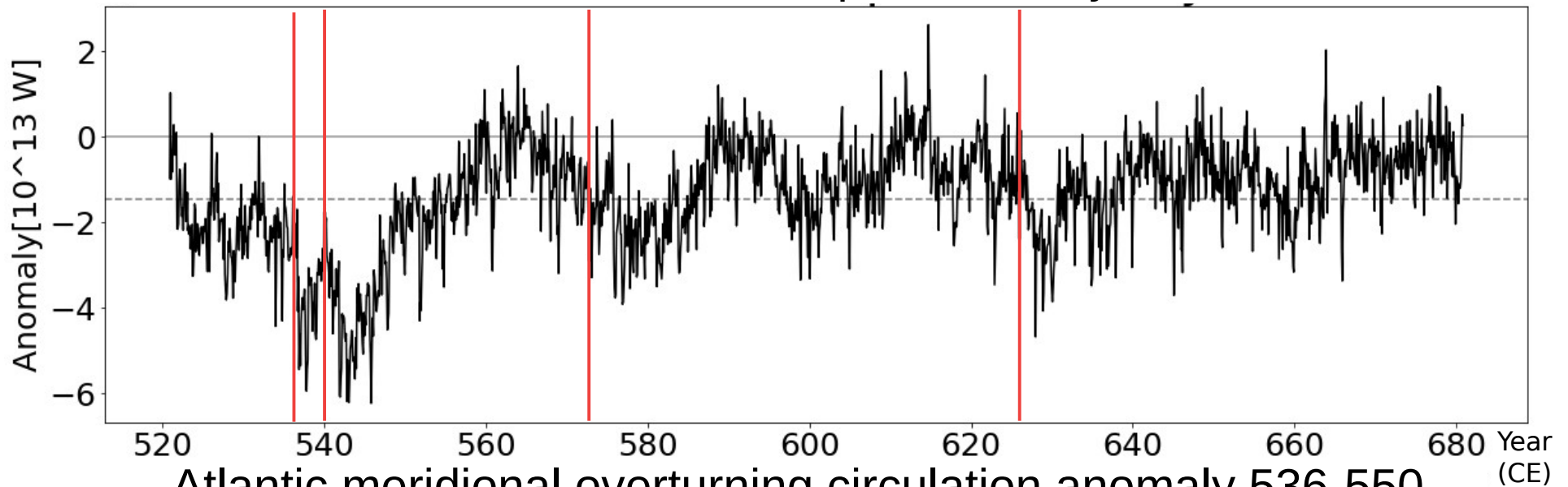
## NH sea-ice area anomaly



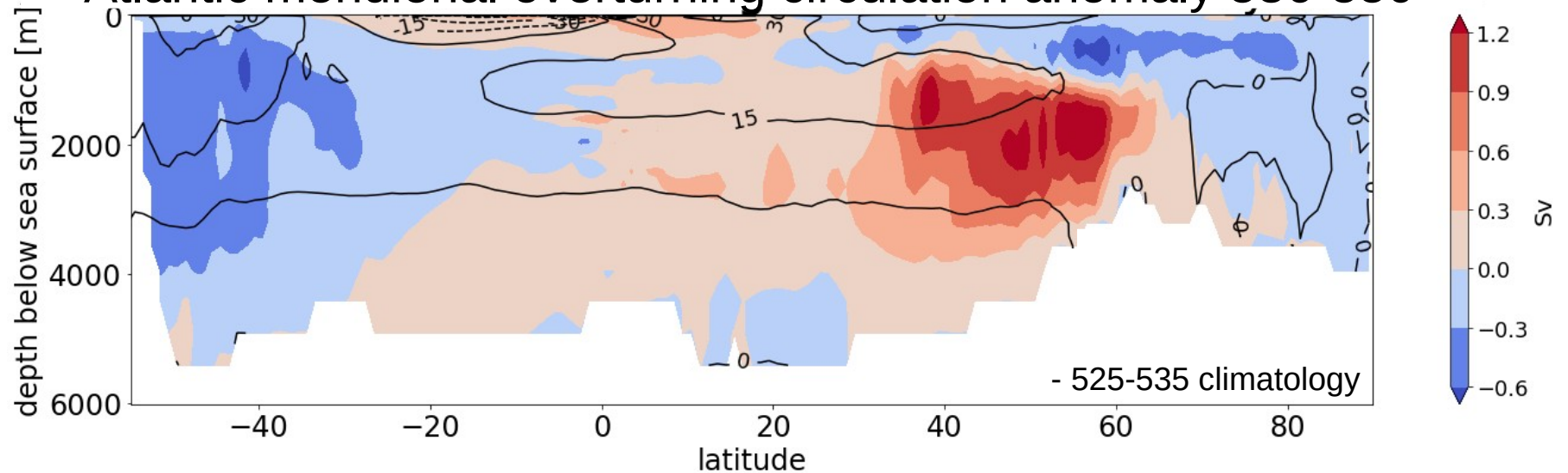
**Sea-ice and temperature anomalies correspond very well, with the same recovery time**

# Ocean heat transport and circulation

## Mean Atlantic heat transport anomaly 58-66N



## Atlantic meridional overturning circulation anomaly 536-550



**Reduced northward heat transport after the eruptions (top figure) → decadal variability**  
**Enhanced Atlantic meridional overturning circulation @ ~40-60N (bottom figure)**



# Summary

New model simulations using the MPI-ESM1.2 using the PMIP4 volcanic forcing for 520-680 CE reveal that:

- Modeled surface temperature response lasts up to 20 years, longer than the prescribed volcanic forcing (2-3 years)
- MPI-ESM runs capture the tree-ring proxy temperature anomalies well
- Ocean – sea-ice feedbacks lead to a prolonged surface climate response
- Ocean heat transport shows a decadal variability with eruptions superimposed on top
- The Atlantic meridional overturning circulation is enhanced after the 536/540 CE double volcanic eruption event