



## Deconstructing the subtropical AMOC variability

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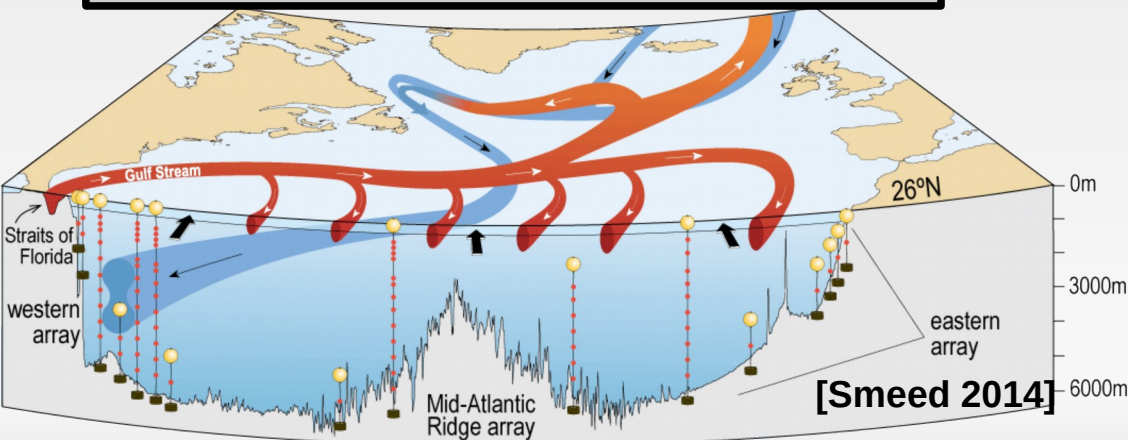
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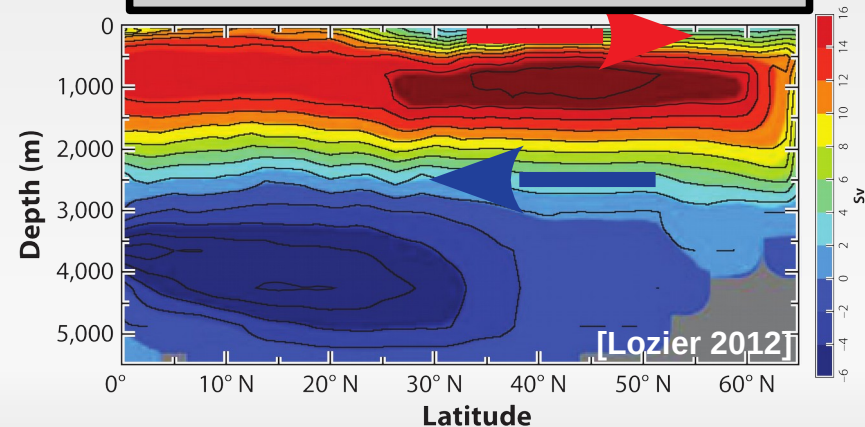


- The Atlantic Meridional Overturning Circulation (**AMOC**):
    - Synthetic representation in the y-z plan of the **complex North Atlantic ocean circulation**
    - Plays a central role in climate by **redistributing heat, freshwater and carbon**
- ➔ Long-standing interests in **understanding its variability**

Schematic of the NA ocean circulation



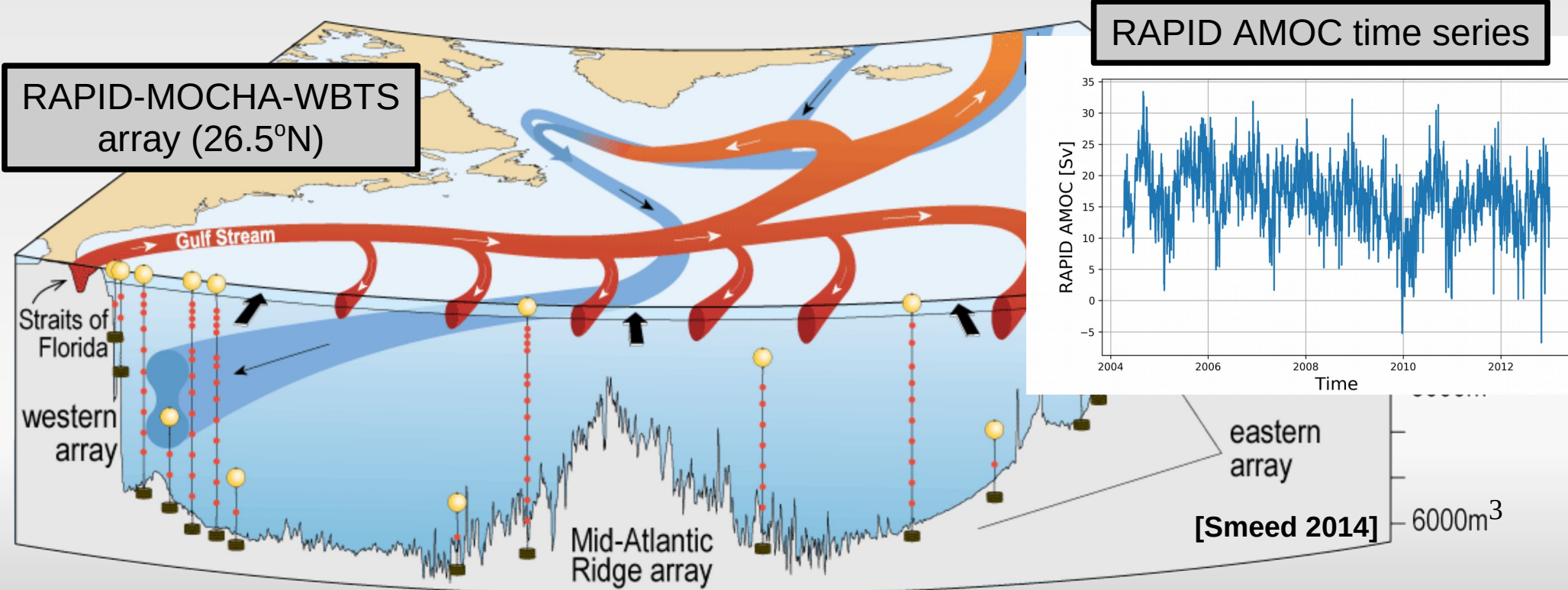
Example of time mean AMOC



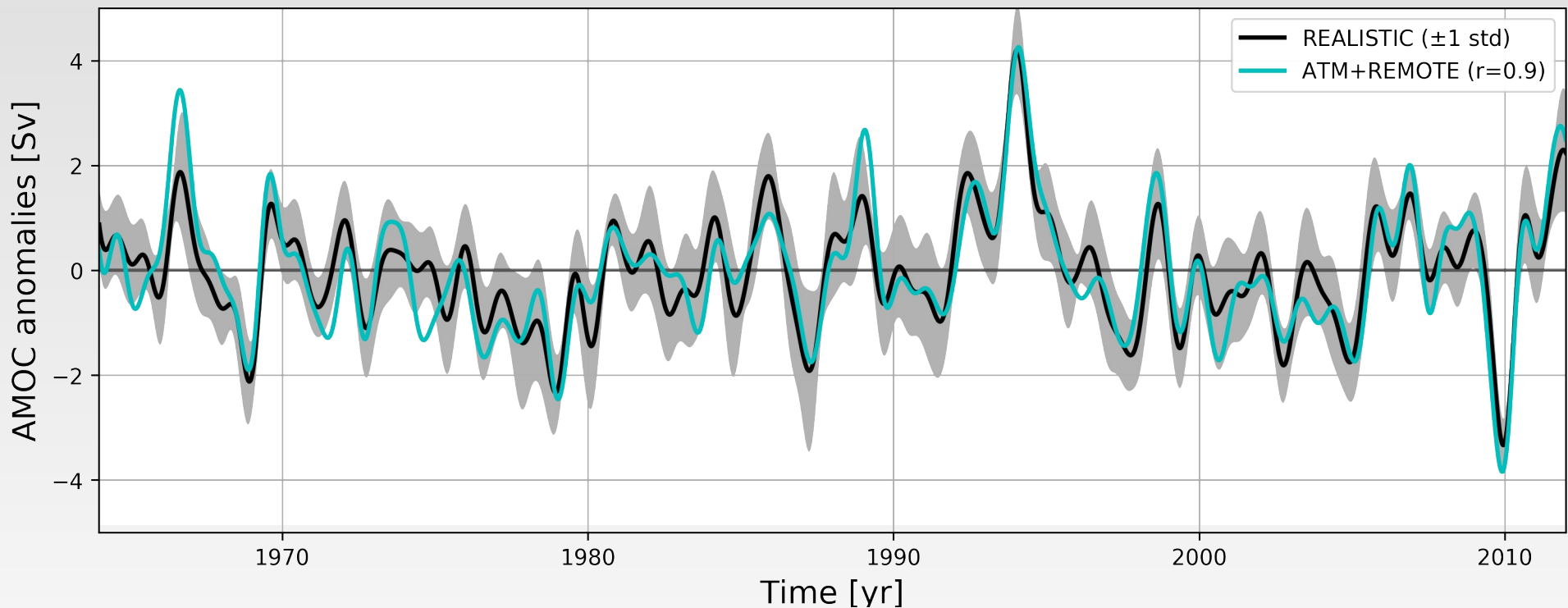
$$AMOC(y, z, t) = \int_{x_w}^{x_e} \int_{-H}^z v dz dx$$

- Sources of AMOC variability **in the NA subtropical gyre:**
  - Adjustment to signals of **remote origin** (NA subpolar gyre, South Atl.)  
[Johnson 2002; Biastoch 2008a,b; Hodson 2012; Jackson 2016]
  - **Local atmospheric forcing** [Eden 2001a,b; Hirschi 2007; Deshayes 2008; Gastineau 2012]
  - **Local intrinsic oceanic variability** [Gregorio 2015; Leroux 2018; Jamet 2019]

→ Potentially, **complex interactions** [Spall 1996a,b; Bower 2000; Zhang 2007; Andres 2016] may complicate the interpretation of observations (RAPID) ...



➔ Nonetheless, we show that AMOC can be understood as a **linear combination** of signals with different origin



$$AMOC \approx AMOC^{ATM} + AMOC^{REMOTE} + AMOC^{INTRINSIC}$$

- Isolating the NA dynamics from the rest of the world:

- Regional, **eddy-resolving (1/12°)** oceanic configuration of the MITgcm
- Partially coupled (CheapAML [Deremble 2013])
- Atmospheric forcing: DFS4.4 [Brodeau 2010]
- Boundary conditions: ORCA12 [Molines 2014]
- **50 yr long simulations**

- Isolating the forced variability from its intrinsic counterpart:

- 12 ensemble members simulations
- Micro initial conditions [Stainforth 2007]
- Ensemble spread → **intrinsic signal**

$$\sigma_I^2 = \frac{1}{N} \sum_{i=1}^N [f_i(t) - \langle f_i(t) \rangle]^2$$

- Ensemble mean → **forced signal**

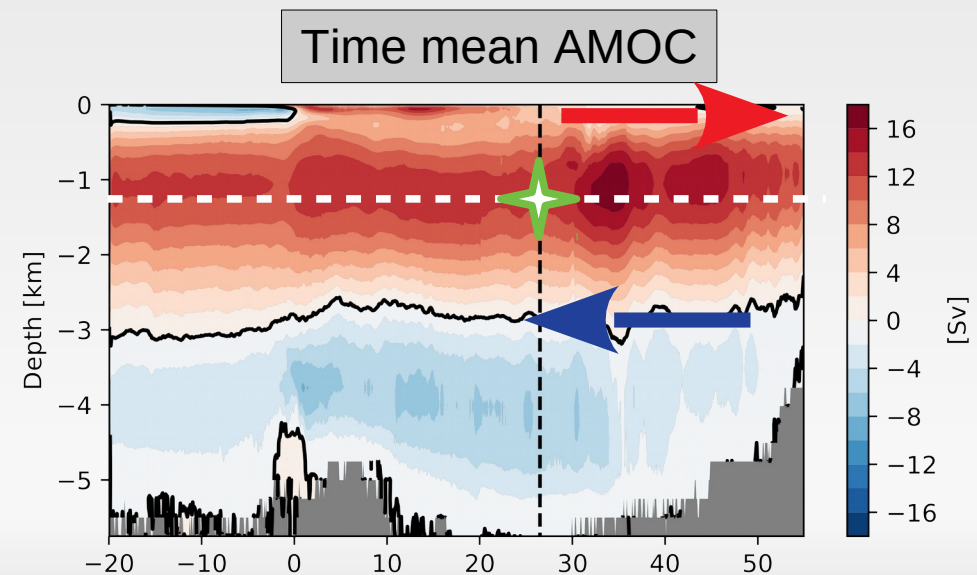
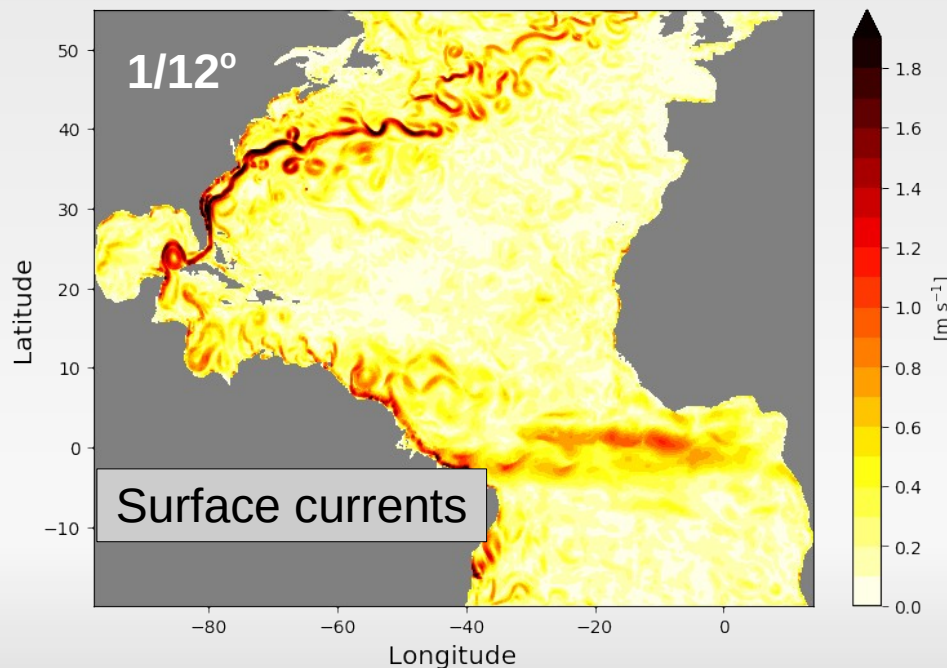
$$\sigma_F^2 = \frac{1}{T} \sum_{t=1}^T [\langle f_i(t) \rangle - \overline{\langle f_i(t) \rangle}]^2$$

- Isolating local atmospheric forcing from signals of remote origin:

Atmosphere	<i>Normal year</i> <i>Aug 2003 – July 2004</i>	<i>Fully Varying</i> <i>1963-2012</i>
Open boundaries		
<i>Climatological</i>	<b>RESIDUAL</b>	<b>ATM</b>
<i>Fully Varying</i>	<b>REMOTE</b>	<b>REALISTIC</b>



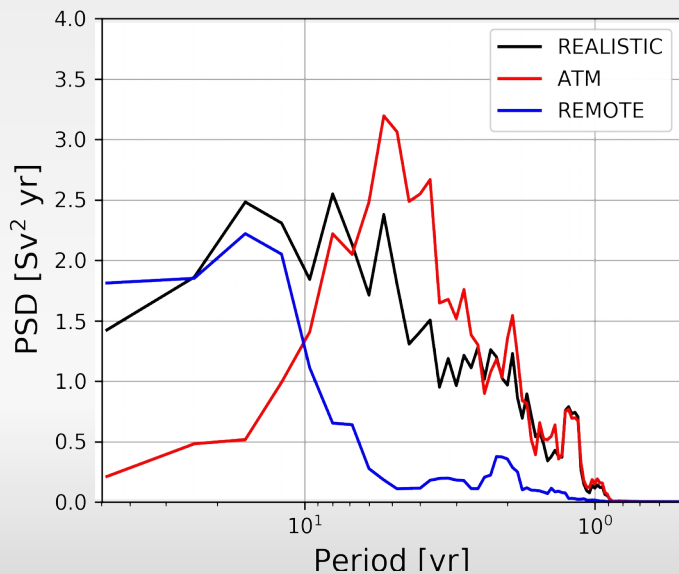
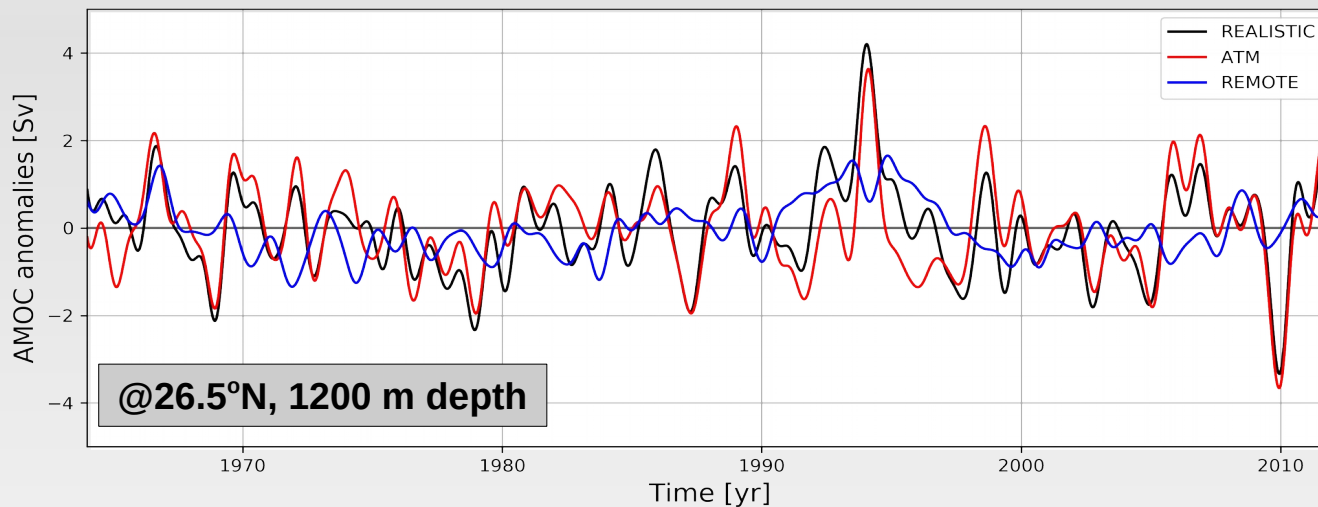
- The simulation delivers a reliable representation of the North Atlantic ocean circulation
- Ensemble production represents about 8,000,000 cph and 150 TB
- Model configuration and outputs are available at: <https://github.com/quentinjamet/chaocan>



# AMOC at 26.5°N

## Forced variability - ensemble mean

- Compare the **forced AMOC variability** simulated by the 3 ensembles **REALISTIC**, **ATM**, and **REMOTE**

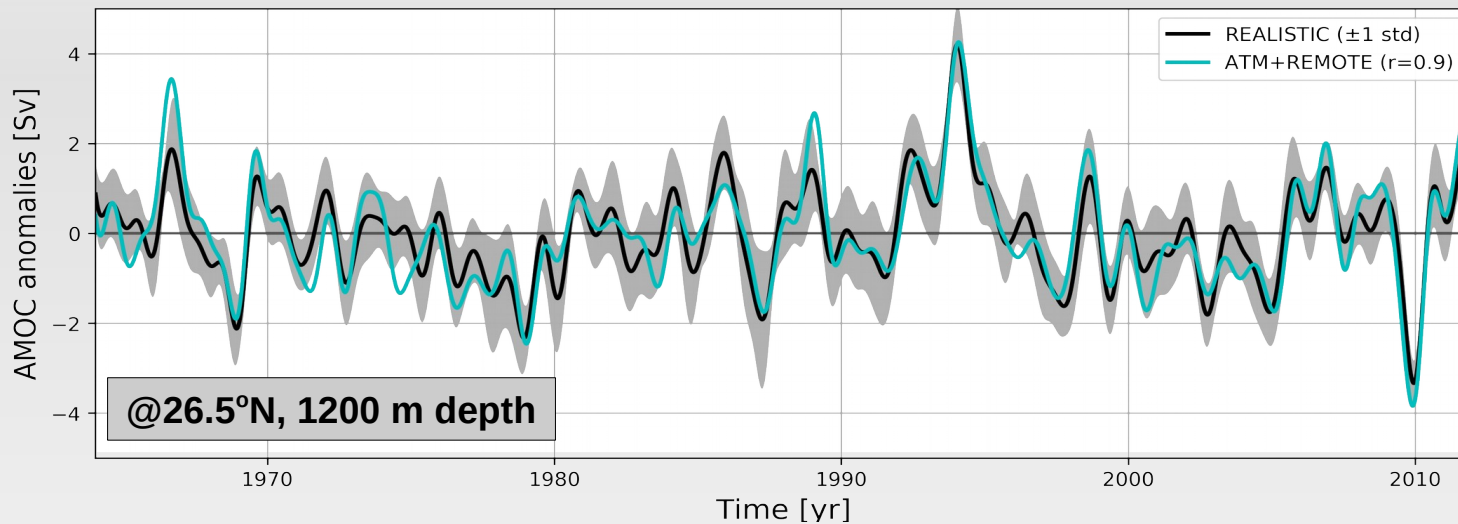


- ➔ **Marked time scales separation**
- Remote (boundary) signals:  
→ decadal time scales
  - Local (atmospheric) forcing:  
→ interannual time scales

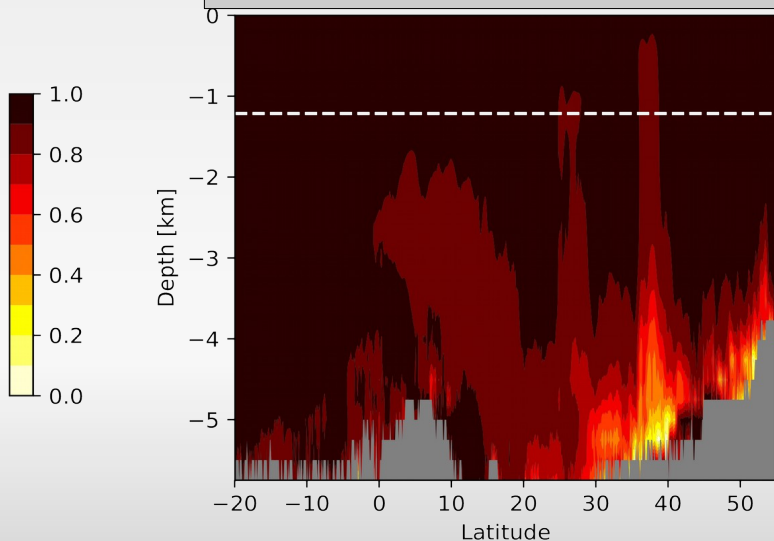
# AMOC at 26.5°N

## Forced variability - ensemble mean

- Compute a **linear reconstruction** as **ATM** + **REMOTE** and compare it with **REALISTIC**



$r(\text{ATM} + \text{REMOTE}, \text{REALISTIC})$



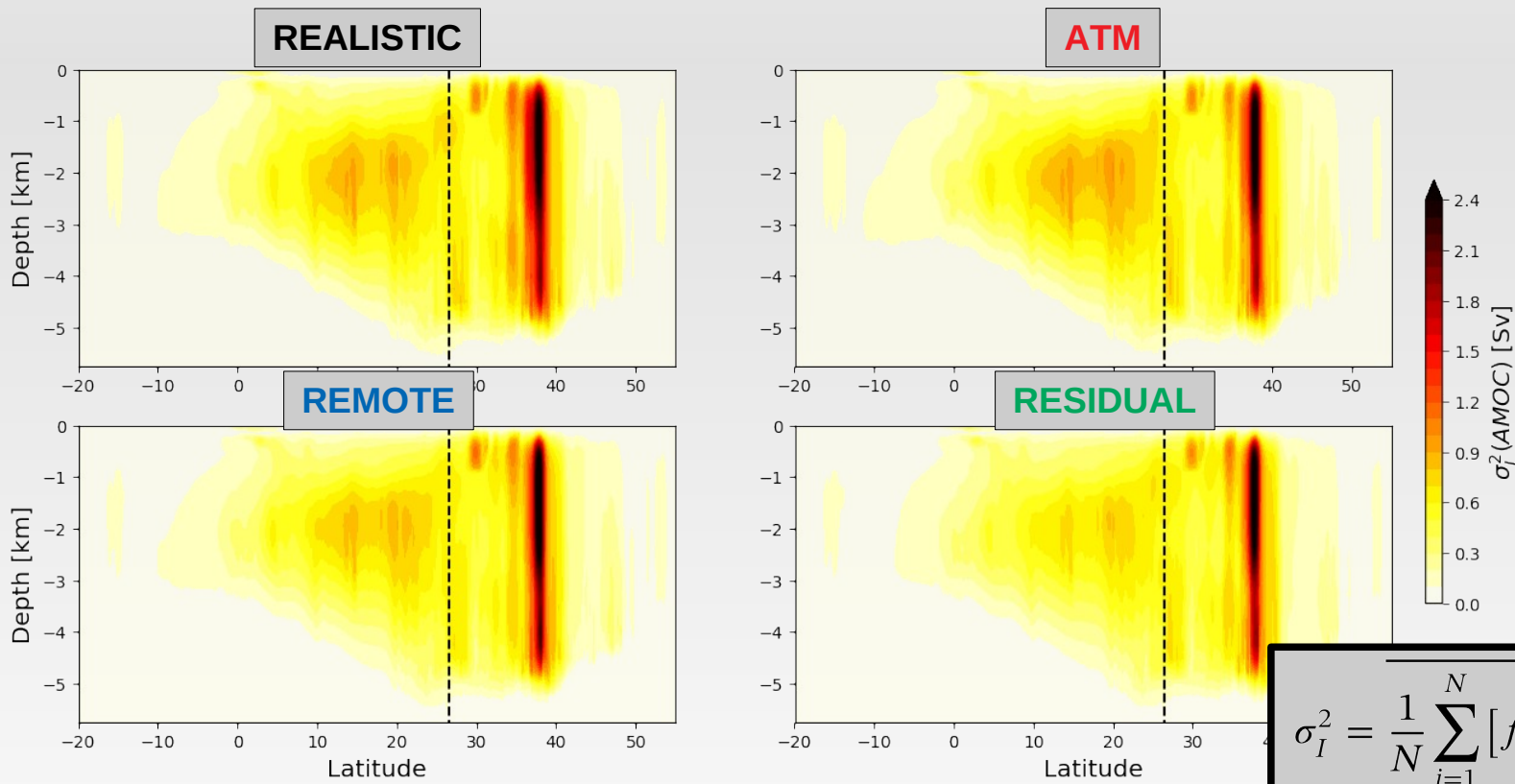
- High correlations between reconstructed and realistic AMOCs in most of the basin
- Forced AMOC variability can be understood as a linear combination:

$$\langle AMOC \rangle \approx \langle AMOC_{.<10\text{ yr}}^{ATM} \rangle + \langle AMOC_{.>10\text{ yr}}^{REMOTE} \rangle$$



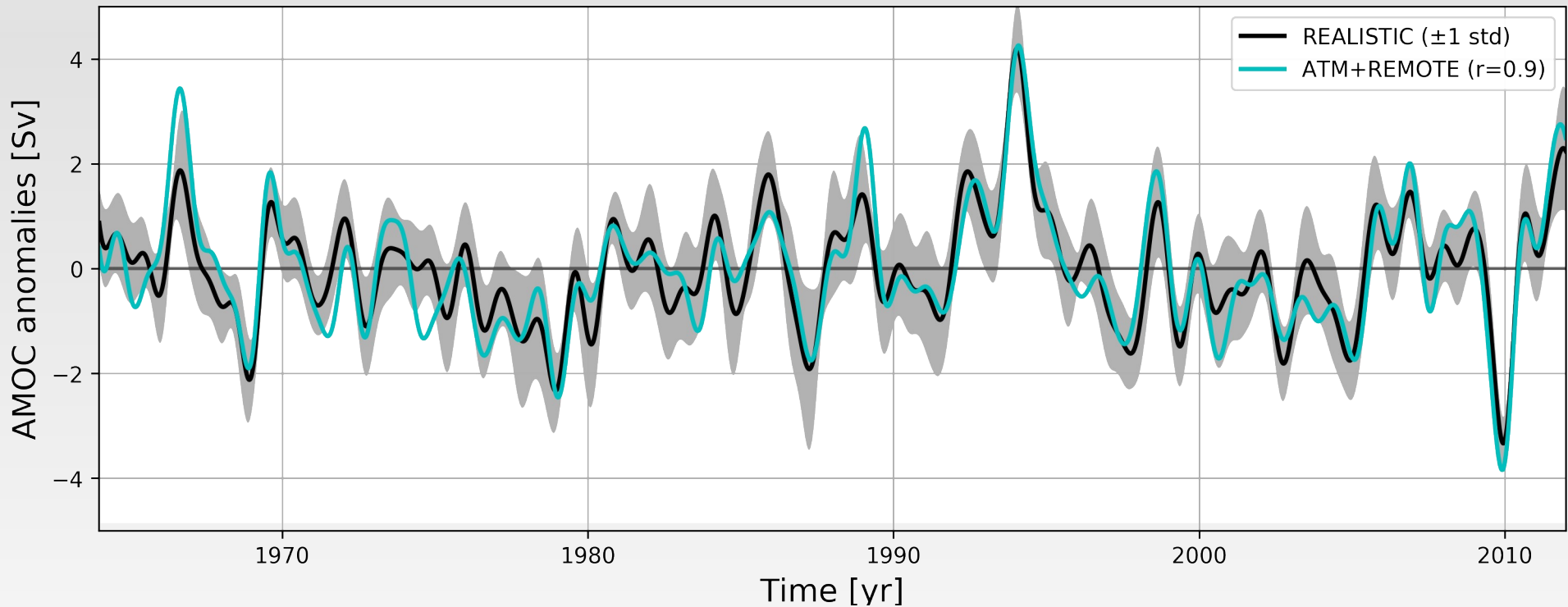
## Intrinsic variability – ensemble spread

- Compare **intrinsic AMOC variability** in each ensemble and assess its sensitivity to the surrounding forced signal



- Each ensemble exhibits specific ensemble mean AMOC variability, **BUT** they all simulate a similar ensemble spread
- ➔ **No causal relationship between intrinsic and forced AMOC**

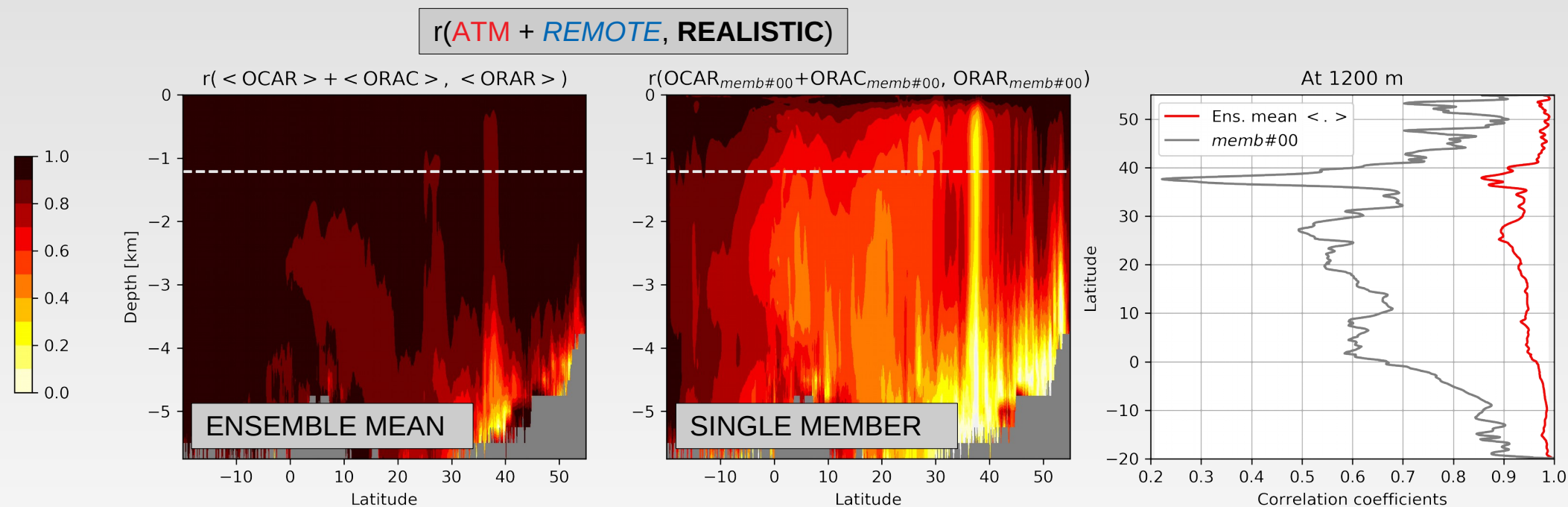
→ AMOC can be understood as a linear combination of signals with different origin



$$AMOC \approx AMOC_{<10\text{ yr}}^{ATM} + AMOC_{>10\text{ yr}}^{REMOTE} + AMOC^{INTRINSIC}$$

- For further details, see [Jamet et al. J. Clim 2020](#)

- Compare **ensemble results** with those obtained with **single simulations**, i.e. without ensemble averaging



→ Correlations decrease to  **$r=0.6$**  in the subtropical gyre, and to  **$r=0.2$**  in the eddying Gulf Stream

$$\sigma_F^2 = \frac{1}{T} \sum_{t=1}^T \left[ \langle f_i(t) \rangle - \overline{\langle f_i(t) \rangle} \right]^2$$
$$\sigma_I^2 = \frac{1}{N} \sum_{i=1}^N \left[ f_i(t) - \langle f_i(t) \rangle \right]^2$$
$$\sigma_T^2 = \sigma_F^2 + \sigma_I^2$$

