

Refining the temporal relation between Large Igneous Provinces and carbon cycle perturbations: not every LIP triggers environmental crises, not every crisis is due to a LIP!

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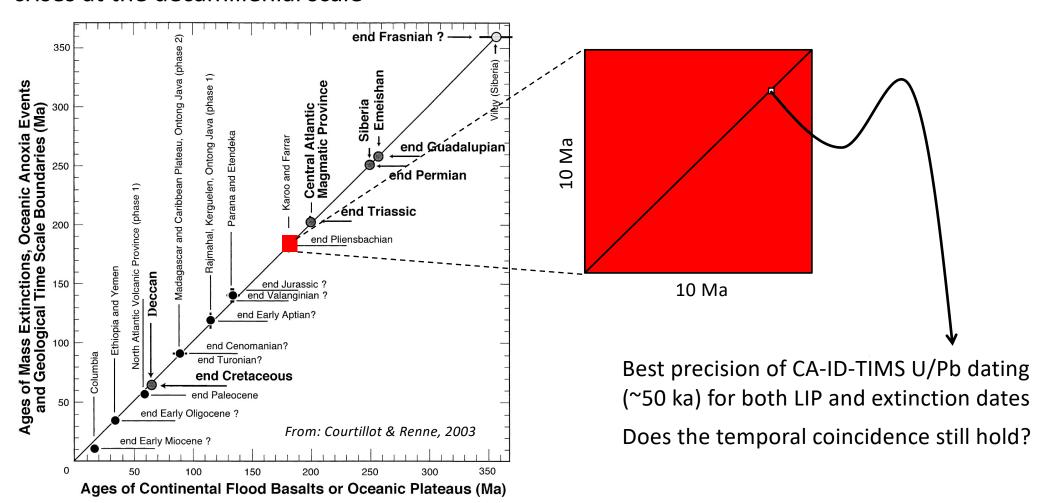
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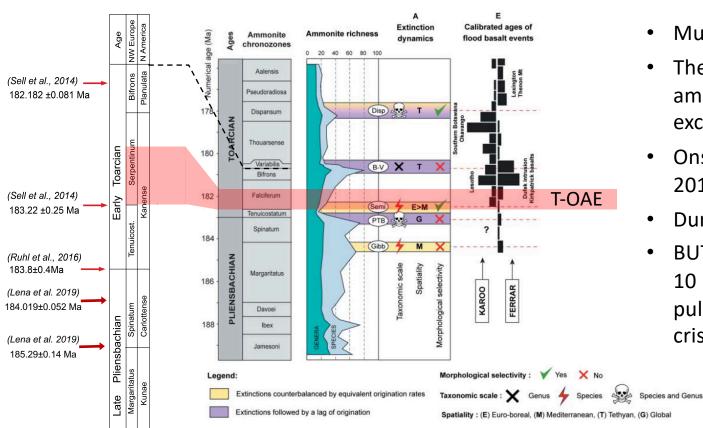




Linking volcanism of Large Igneous Provinces with environmental perturbation and biotic crises at the decamillenial scale



Lower Jurassic: Repeated biotic crises – ocean anoxia – LIP activity

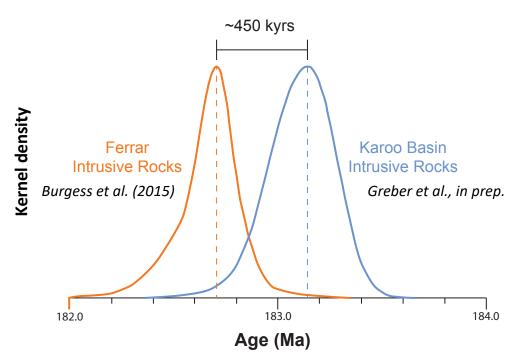


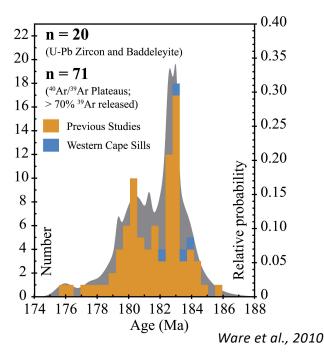
- Multi-episodic biotic crises in the Toarcian
- The T-OAE is defined by the Falciferum ammonite zone negative carbon isotope excursion (CIE)
- Onset at 183.22 ± 0.25 Ma (Sell et al., 2014)
- Duration 300-500 kyr (Boulila et al., 2014)
- BUT: known Karoo LIP dates scatter over 10 Ma from 186 to 176 Ma – is just one pulse responsible for the environmental crisis?

Dera et al., 2010

Karoo sill/dyke complex coincides with onset of T-OAE

- Age of Karoo basin dyke-sill complex is ca. 183 ± 0.1 Ma and correlates with onset of Falciferum ammonite zone negative carbon isotope excursion (Svensen et al., 2012; Corfu et al., 2016; Greber et al., manuscript in prep.)
- Continued volcanic activity (Ferrar and beyond?) may be responsible for environmental perturbation
- No volcanic activity pre-183 Ma
- Previously published age scatter (40Ar/39Ar and U/Pb) is at least partly an analytical artefact





Conclusion: No volcanic driver for the Late Pliensbachian cooling event

Cooling event recorded by C, O, Os isotope time series is older than oldest highprecision age from Karoo LIP

→ non-volcanic driver for a global C-cycle disturbance

(Korte & Hesselbo 2011 Ma (Burgess et al., 2015) Hesselbo et al., 2007 Storm Peak Lava Littler et al., 2010; 182.182 ±0.081 Ma (Sell et al., 2014) 182.430±.0.036 Ma Jenkyns et al., 1997 and this study) Toarcian Sepentinum $\delta^{13}C$ Karoo-Ferrar Kanense New Amalfi Sheet warming 183.243±.0.045 Ma~ 800 ka Early Tenuicost. 183.22 ±0.25 Ma (Sell et al., 2014) 183.8 Ma (Ruhl et al., 2016) 184.019±0.052 Ma (Lena et al., 2019) Spinatum Pliensbachian 184.5 Cooling 185.0 Late Pliensbachian event 185.29±0.14 Ma (Korte & Hesselbo, 2011) 185.5 Margaritatus 186.0-186.5

Lena et al., 2019, Sci. Rep.

Conclusions – hypotheses – outlook

- High-precision U/Pb geochronology allows correlation of volcanism with environmental and biotic crises in Earth's history at the ~50'000 years' level of resolution
- Intrusive phases of LIP correlate with mass extinctions
- Periods of massive C-cycle disturbance may have non-volcanic triggers
- Pre- and post-intrusive phase flood basalt phases have less prominent impact on environmental change and biotic crises
- Sill-dyke intrusions into sediments rich in organic matter produce thermogenic CO₂ through from contact metamorphism of sediments (e.g., Svensen et al., 2004, 2008; Heimdal et al., 2017), which seems to be the main driver of environmental change