

A new approach to modelling differentiation (with particular focus on granitic magmatism): Equilibrated Major Element Assimilation with Fractional Crystallisation (EME-AFC)

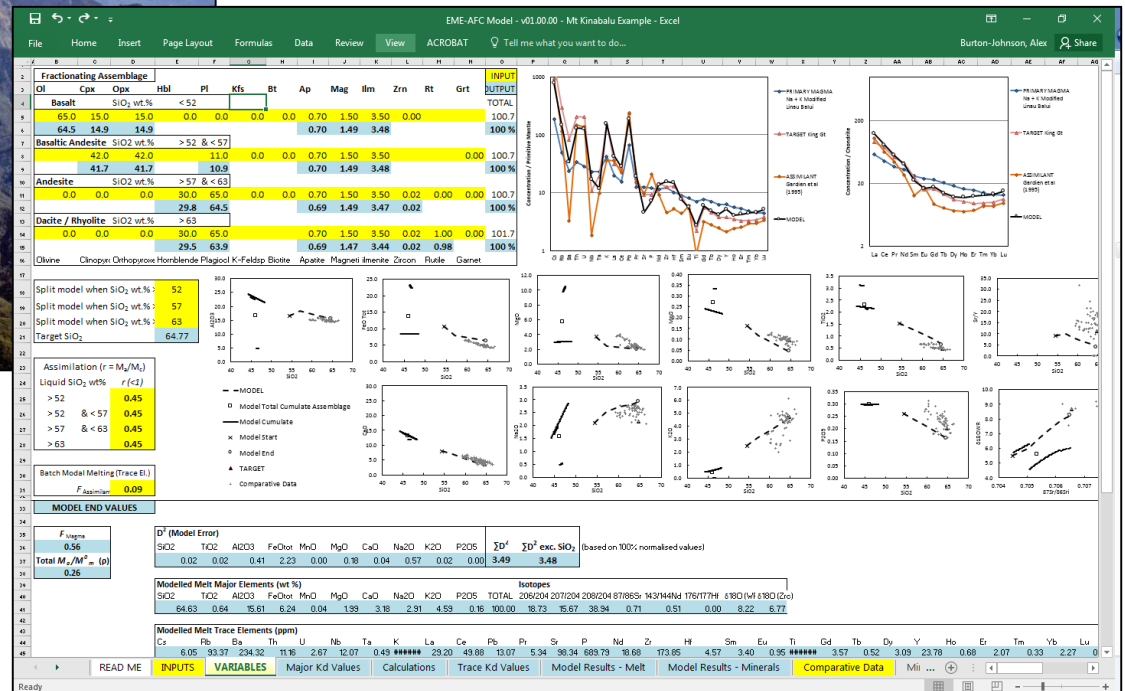
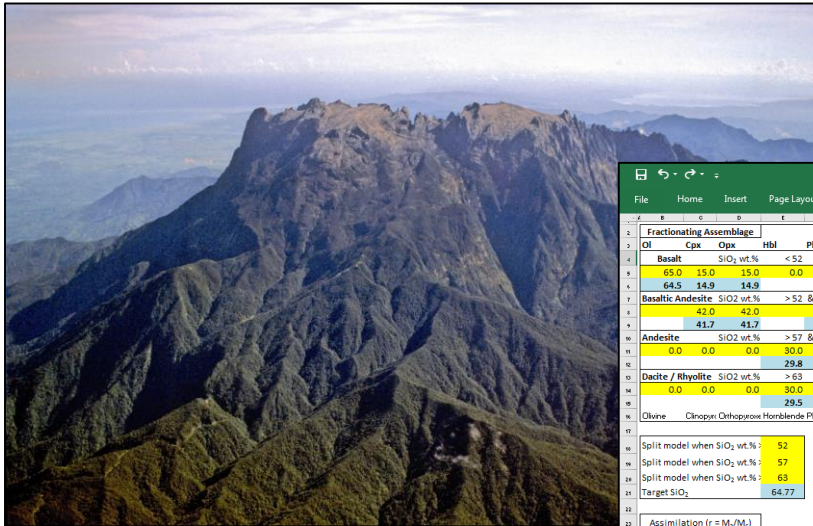


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The Challenge

Modelling AFC differentiation from basaltic to silicic magma; particularly to derive granite.

The Problems

- 1) Current AFC models (MELTS, Rhyolite-MELTS, Magma Chamber Simulator, alphaMELTS) cannot model hydrous magmatic systems involving substantial hornblende or biotite fractionation. Given evidence for “cryptic amphibole fractionation”, this applies to most/all arc magmas.
- 2) Trace element partition coefficients are notoriously variable.
- 3) Modelling different elemental and isotopic systems produce different conclusions, so ideally all systems should be modelled simultaneously.

The Solution

A **new method** (Equilibrated Major Element Assimilation with Fractional Crystallisation, **EME-AFC**), using two-component major element partition coefficients and simultaneously modelling of major and trace elements, radiogenic isotopes, and stable isotopes.

[Burton-Johnson, A., Macpherson, C. G., Ottley, C. J., Nowell, G. M. & Boyce, A. J. \(2019\). Generation of the Mt Kinabalu granite by crustal contamination of intraplate magma modelled by Equilibrated Major Element Assimilation with Fractional Crystallisation \(EME-AFC\). *Journal of Petrology* **60**, 1461–1487.](#)

EME-AFC: Major Elements

Equilibrium major element compositions of rock-building minerals can be calculated from experimentally derived, two-component partition coefficients (K_D). Minor elements are calculated from their empirical relationship to the majors (*below*).

Unlike trace element partition coefficients, major element K_D values are far less variable (*right*).

Two-Component Partition Coefficients		Mean	SD	n
Plagioclase	$(Al^{Plag} \times Si^{Liq}) / (Si^{Plag} \times Al^{Liq})$	2.46	0.51	36
Plagioclase	$(K^{Plag} \times Na^{Liq}) / (Na^{Plag} \times K^{Liq})$	0.09	0.04	36
Hornblende	$(Fe^{Hbl} \times Mg^{Liq}) / (Mg^{Hbl} \times Fe^{Liq})$	0.36	0.07	28
Hornblende	$(Al^{Hbl} \times Si^{Liq}) / (Si^{Hbl} \times Al^{Liq})$	1.13	0.27	28
Clinopyroxene	$(Fe^{Cpx} \times Mg^{Liq}) / (Mg^{Cpx} \times Fe^{Liq})$	0.28	0.07	23
Clinopyroxene	$(Al^{Cpx} \times Si^{Liq}) / (Si^{Cpx} \times Al^{Liq})$	0.27	0.18	23
Orthopyroxene	$(Fe^{Opx} \times Mg^{Liq}) / (Mg^{Opx} \times Fe^{Liq})$	0.28	0.08	15
Orthopyroxene	$(Al^{Opx} \times Si^{Liq}) / (Si^{Opx} \times Al^{Liq})$	0.19	0.16	15
Olivine	$(Fe^{Ol} \times Mg^{Liq}) / (Mg^{Ol} \times Fe^{Liq})$	0.28	0.03	14

Calculations used for major element modelling of plagioclase

Composition $((K,Na)_{1-(x-1)}(Mg,Mn,Ca)_{x-1})_1((Fe,Al)_x(Si_{4-x})_4O_8$

Al-Si Kd $K_D = (Al^{Plag} \times Si^{Liq}) / (Si^{Plag} \times Al^{Liq})$

K-Na Kd $K_D = (K^{Plag} \times Na^{Liq}) / (Na^{Plag} \times K^{Liq})$

Fe $= a \times (Fe + Al)$

Mg $= b \times (Mg + Mn + Ca)$

Mn $= c \times (Mg + Mn + Ca)$

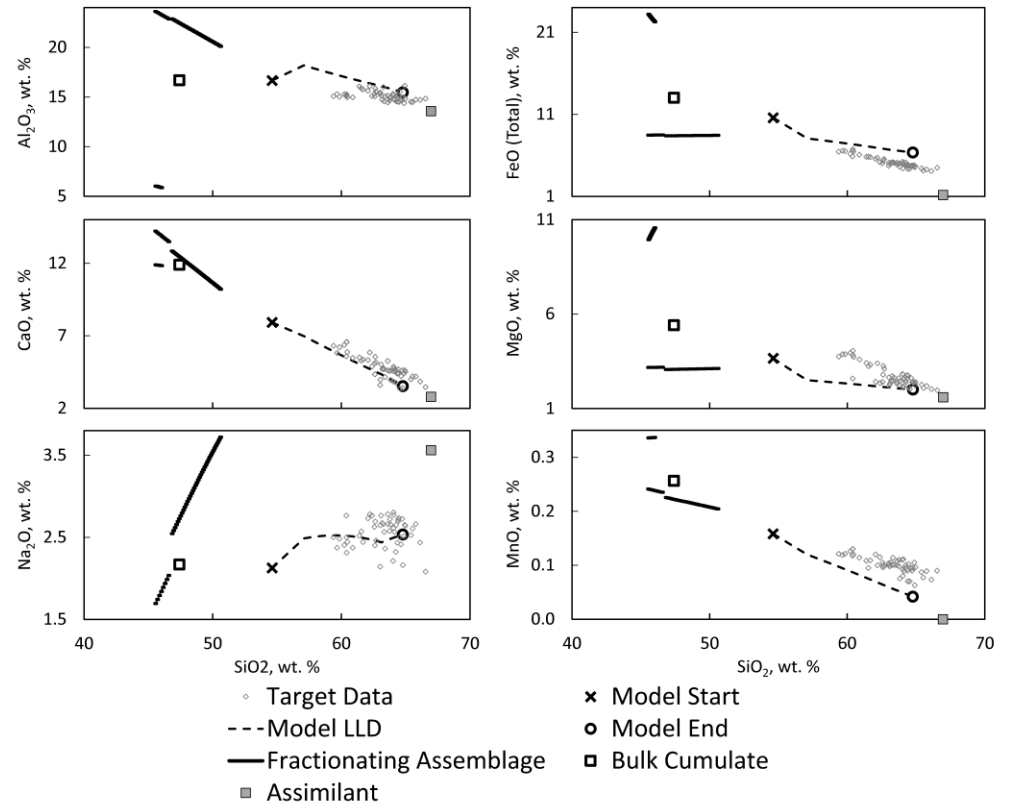
EME-AFC methodology for plagioclase.

K_D equations derive two-component major element partition coefficients from experimental data

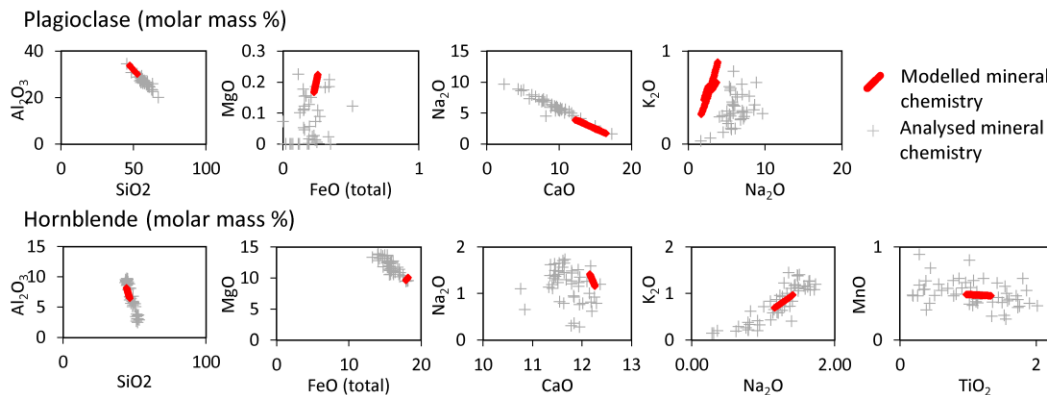
'a', 'b', and 'c' for the minor elements are calculated from mean mineral chemistry of the rocks being modelled

EME-AFC: Major Elements

The equilibrium mineral compositions are calculated and fractionated from the magma by mass balance for successive increments of fractionation, “*F*”. Assimilation is modelled at each increment by binary mixing for the major elements according to a user-determined rate of assimilation, “*r*” (the mass assimilated / mass crystallised). The instantaneous and bulk cumulate composition is also calculated.



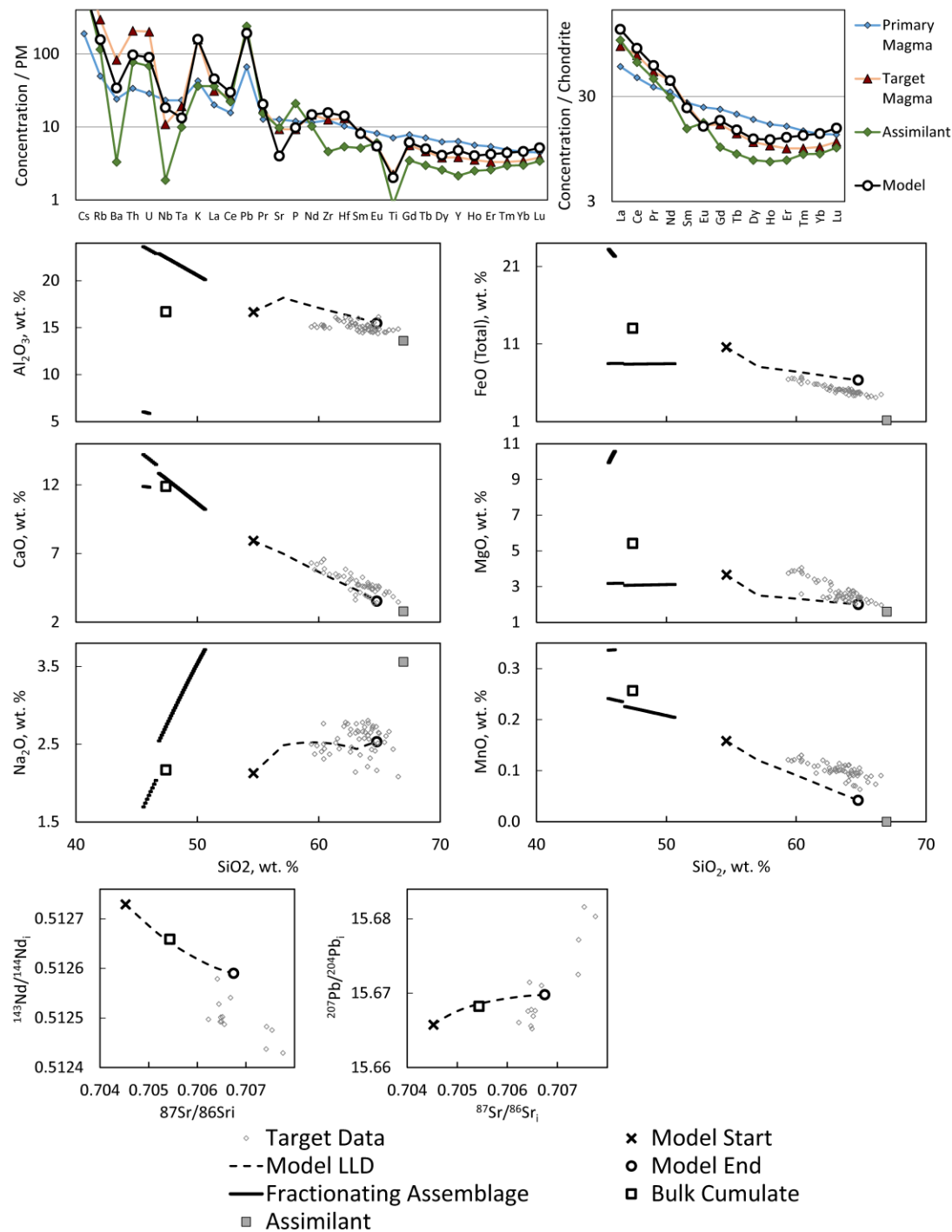
EME-AFC modelling of the Mt Kinabalu granite, Borneo



Analysed and modelled mineral chemistries of the Mt Kinabalu granite, Borneo

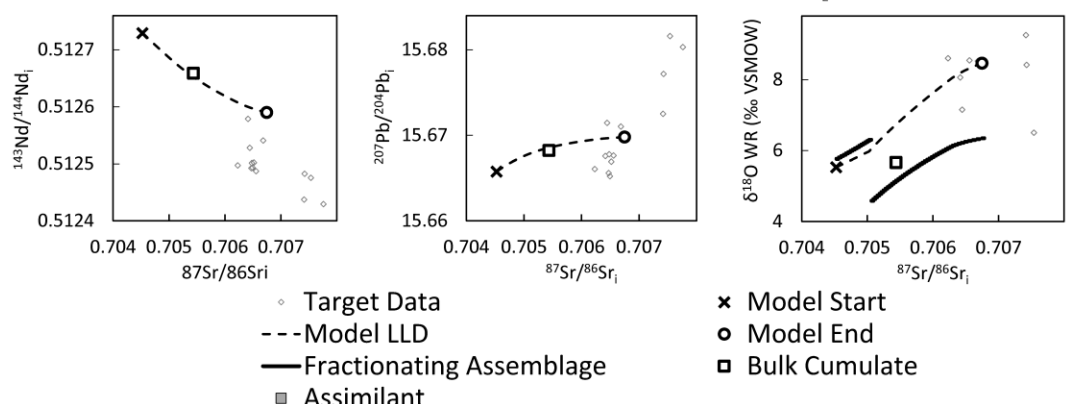
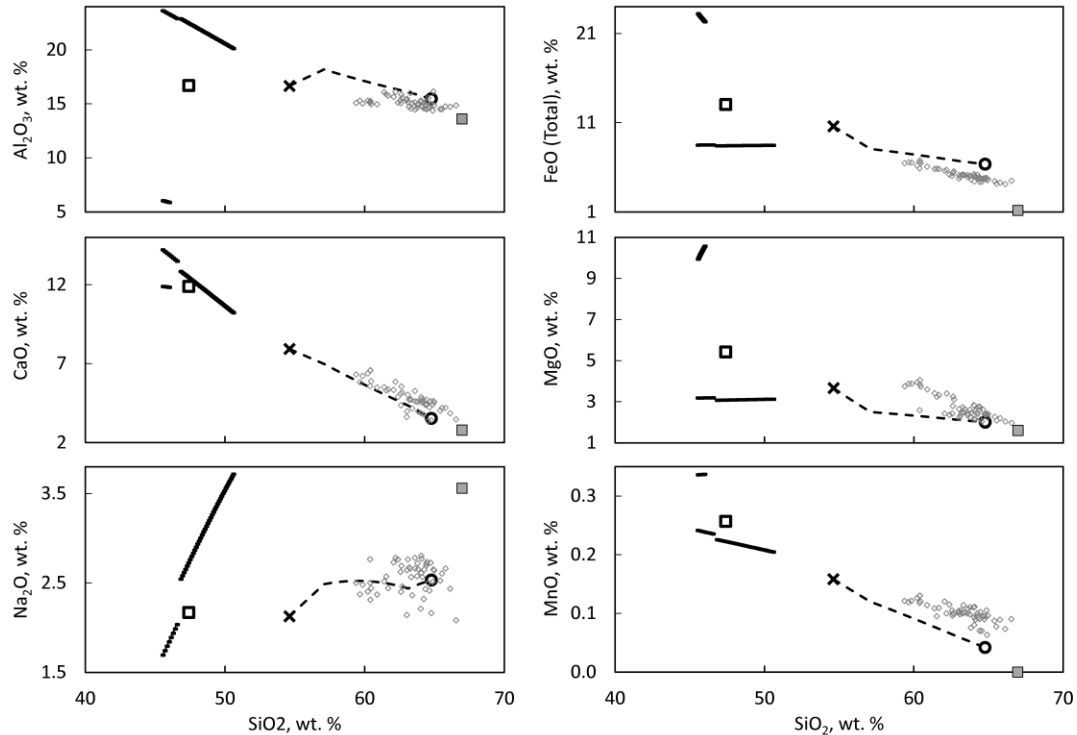
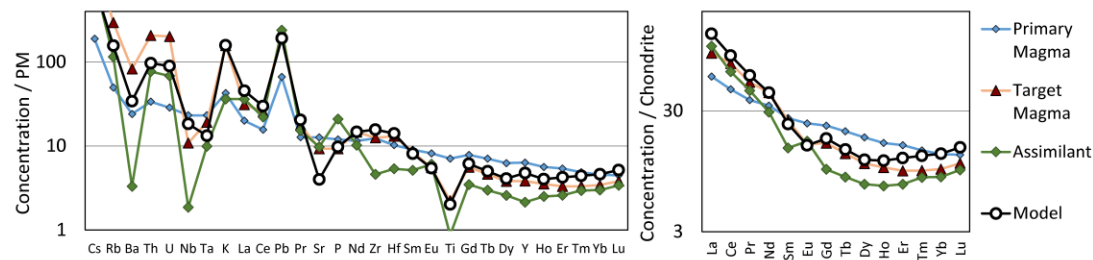
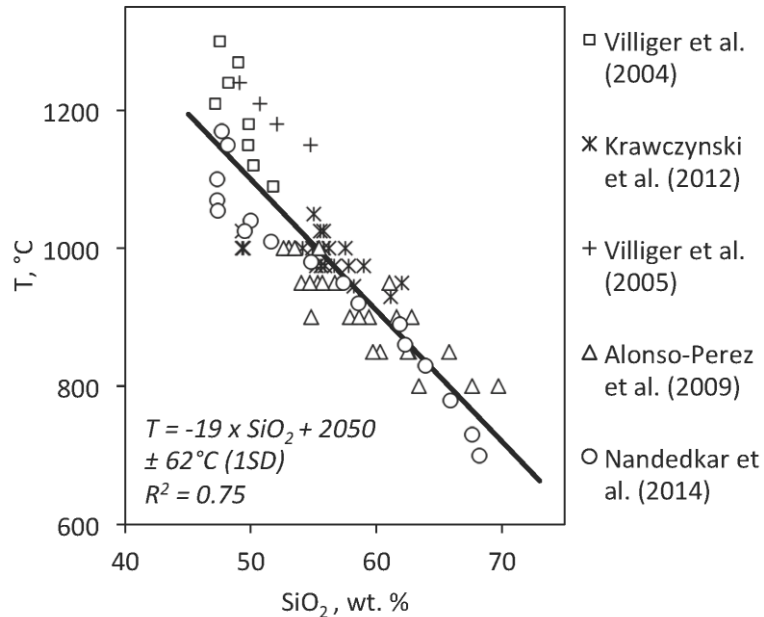
EME-AFC: Trace Elements and Radiogenic Isotopes

Trace element and isotopic compositions are calculated at each increment using the AFC equations of DePaolo (1981).



EME-AFC: Oxygen Isotopes

Oxygen isotope fractionation between coexisting phases varies with temperature. Experimental melt SiO_2 correlates strongly with temperature (*below*), allowing oxygen isotope AFC modelling at each increment (*bottom right*).



EME-AFC Excel Spreadsheet

EME-AFC can be applied using a spreadsheet-based model (with instructions).

Inputs:

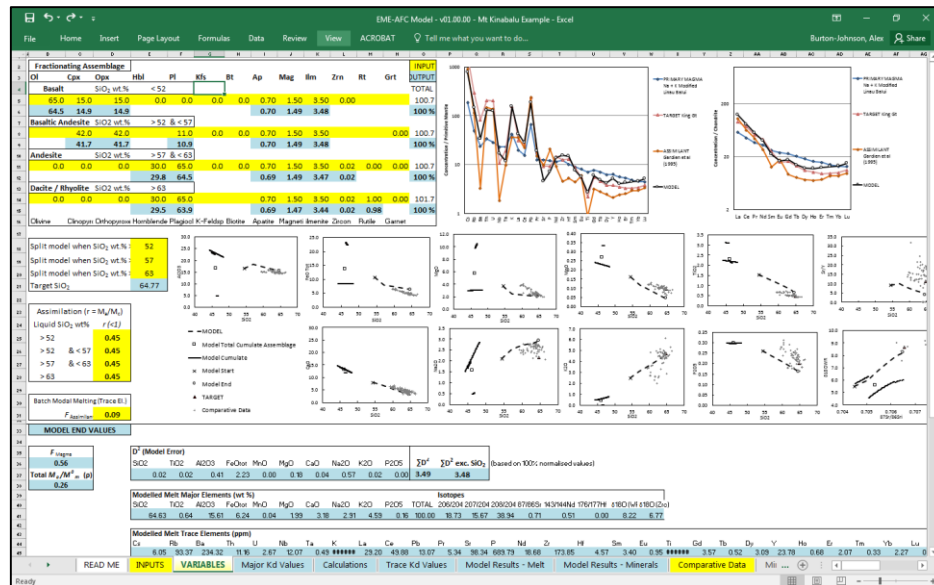
- Primary melt composition
- Target magma composition
- Assimilant composition
- Mineral separate data (optional, but ideally)

Variables:

- Rate of assimilation
- Fractionating assemblage
- Partition coefficients

Outputs:

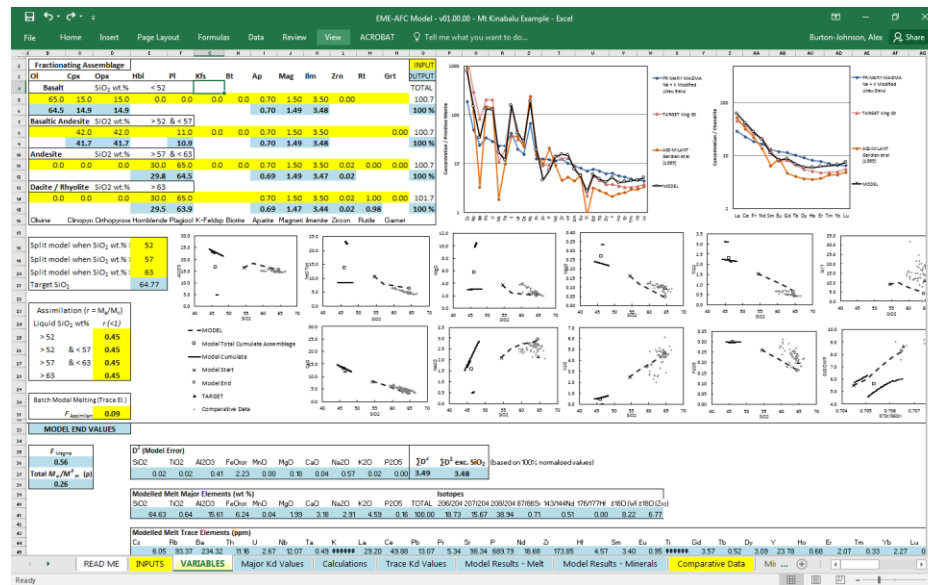
- Elemental and isotopic composition of magma and cumulates
- Bulk cumulate mineralogy
- Sum of the squares, to evaluate model agreement
- Bivariate and spidergram plots



<https://github.com/Alex-Burton-Johnson/EME-AFC-Modelling>

Summary

- EME-AFC allows modelling of magmatic differentiation even in hydrous systems
- Explore the effects of inputs and variables of the magmatic system of interest
- EME-AFC can be freely applied using a spreadsheet-based model, latest version available: <https://github.com/Alex-Burton-Johnson/EME-AFC-Modelling>



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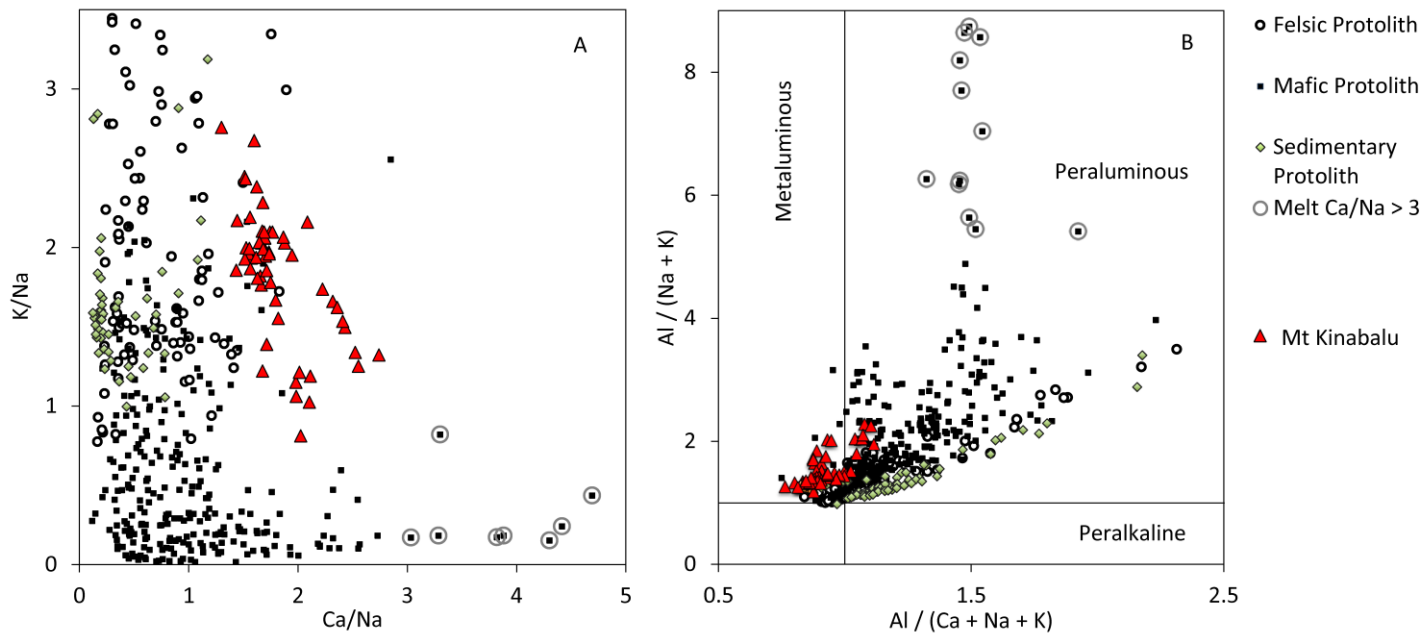
[Burton-Johnson, A., Macpherson, C. G., Millar, I. L., Whitehouse, M. J., Ottley, C. J. & Nowell, G. M. \(2020\). A Triassic to Jurassic arc in north Borneo: Geochronology, geochemistry, and genesis of the Segama Valley Felsic Intrusions and Sabah ophiolite. *Gondwana Research*.](#)

- <- Introduces EME-AFC and applies it to Mt Kinabalu in Borneo (J. Pet. Editors choice, July 2019, Open access)
- <- Further example using EME-AFC

Case Study 1: The Mt Kinabalu granitic pluton, Borneo

Whole rock isotopes indicated a continental signature, but compiled experimental melts of all rock types, melt degrees, and water contents cannot reproduce the pluton's chemistry (*below*).

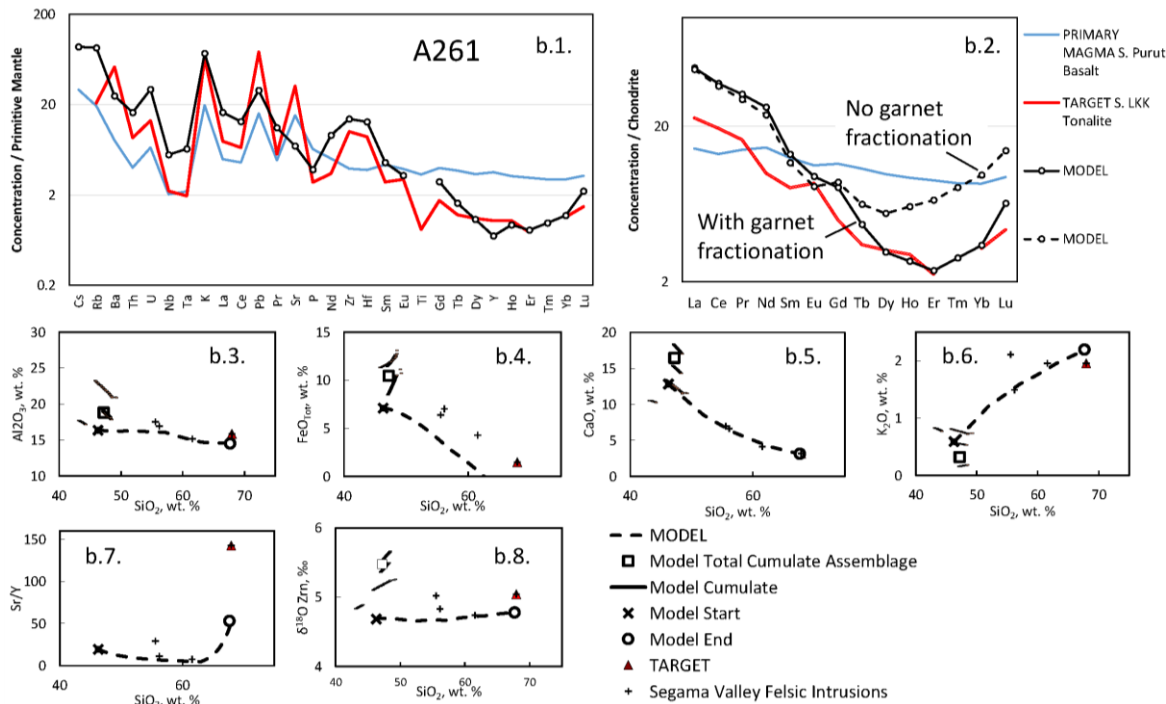
EME-AFC allowed the derivation of Mt Kinabalu from basaltic AFC to be tested with different primary melts and assimilants. The primary melt was a low-degree, K-rich, extensional mantle melt, that assimilated the sedimentary country rock.



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Case Study 2: The Segama Valley granitoids, Borneo

The granitoids were proposed to be windows or partial melts of continental crust beneath Borneo. EME-AFC showed how they could be derived from primary mantle melts without a continental crustal contribution, and how garnet fractionation was variable.



[Burton-Johnson, A., Macpherson, C. G., Millar, I. L., Whitehouse, M. J., Ottley, C. J. & Nowell, G. M. \(2020\). A Triassic to Jurassic arc in north Borneo: Geochronology, geochemistry, and genesis of the Segama Valley Felsic Intrusions and Sabah ophiolite. *Gondwana Research*.](#)