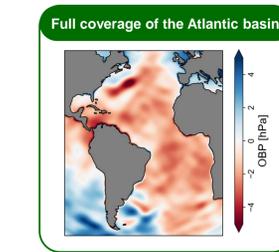
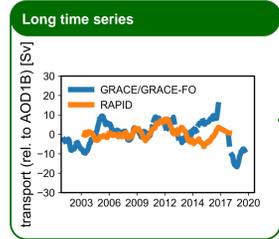


# Monitoring the AMOC with GRACE/GRACE-FO – How far can we push the spatial resolution?

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The AMOC is an important component of the Earth's climate system

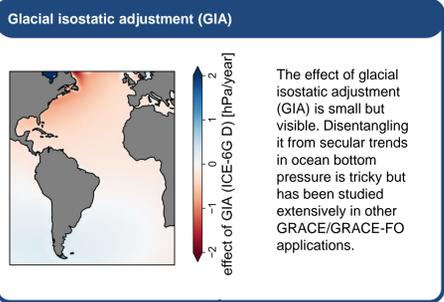
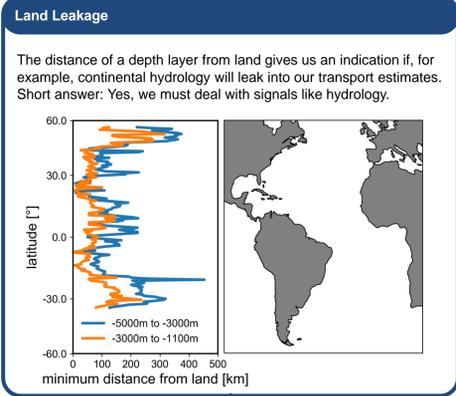
Why?

How?

## Monitoring the Atlantic meridional overturning circulation (AMOC) with GRACE/GRACE-FO

Landerer et al. (2015) have shown that satellite gravimetry has (some) skill at observing the lower limb of meridional transport.

Previous Work



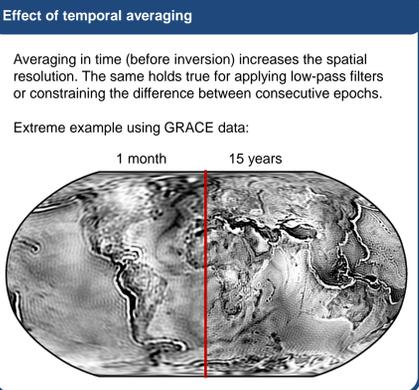
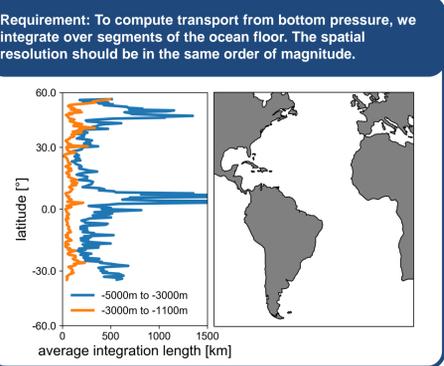
Signal separation on sensor measurement level: Why?  
Models (for example continental hydrology) can be applied in their native spatial and temporal resolution.  
Model output is passed through the GRACE/GRACE-FO processing chain and therefore smoothed in the same way.

Signal Separation

Challenges

Spatial Resolution

Resolution of GRACE and GRACE-FO



MOC is defined as the maximum of the streamfunction, it is therefore related to meridional transport.

Transport through a basin cross-section is computed by integrating the velocity over the cross-section surface

$$\psi = \iint_S v(\lambda, z) d\lambda dz$$

$\psi$  ... transport  
 $v$  ... northward velocity  
 $z$  ... elevation  
 $\lambda$  ... longitude

geostrophic flow, neglect stress terms

$$\psi = \iint_S \frac{1}{\rho_0 f} \frac{\partial p}{\partial \lambda} d\lambda dz$$

$\rho_0$  ... seawater density  
 $f$  ... Coriolis parameter  
 $p$  ... pressure

split integral into longitude/depth

$$\rho_0 f \psi = \int_{z_1}^{z_2} \sum_k \int_{\lambda_k} \frac{\partial p}{\partial \lambda} d\lambda dz$$

solve longitudinal integral

$$\rho_0 f \psi = \int_{z_1}^{z_2} \sum_k p(\lambda_k^k(z)) - p(\lambda_k^k(z)) dz$$

separate western and eastern boundary, change integration bounds

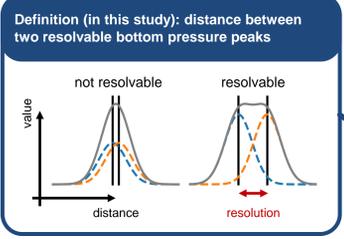
$$\rho_0 f \psi = \int_{z_1}^{z_2} \sum_k p(\lambda_w^k(z)) dz + \int_{z_1}^{z_2} \sum_k p(\lambda_e^k(z)) dz$$

substitute integration over depth with integration over longitude (assumes narrowing topography)

$$\psi = \frac{1}{\rho_0 f} \int_{\lambda_w}^{\lambda_e} p(\lambda) z'(\lambda) d\lambda$$

$z' = 0$

We can compute transport by integrating bottom pressure values from the western to the eastern depth layer boundary. Points outside the layer are excluded by setting the slope to zero.



How can we derive the spatial resolution from a (constraint) GRACE/GRACE-FO solution?

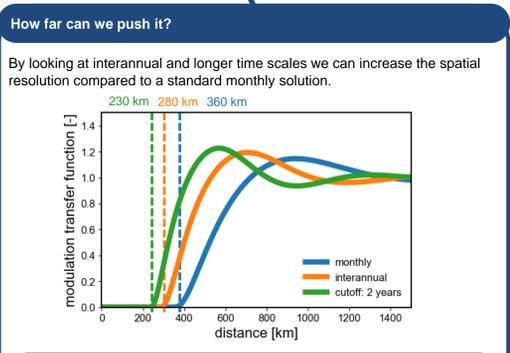
A zero-constrained OBP estimate can be rewritten as a filter:

signal covariance

$$\frac{S}{S + \Sigma}$$

estimated formal errors

This filter has a kernel function, which shows how a single OBP peak is seen by GRACE/GRACE-FO and allows us to quantify the resolution:



We should be able to compute transport in the deeper layers (< -3000m) of the Atlantic. The depth layer -1100m to -3000m with a horizontal extent of 100km to 150km might be too narrow to properly resolve.