

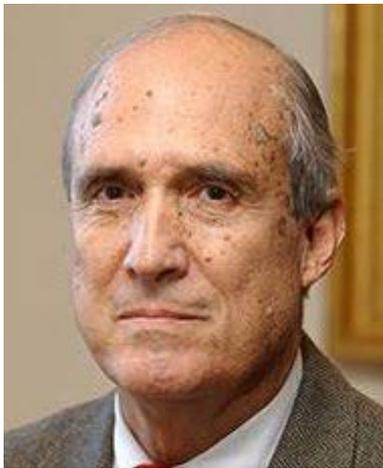


EGU2020 SSP2.13
Asian Climate and Tectonics
D840 | EGU2020-4042

**A new negative feedback mechanism for balancing
Tibet uplift-driven CO₂ drop:
Evidence from Paleogene chemical weathering records
in the northern Tibetan Plateau**

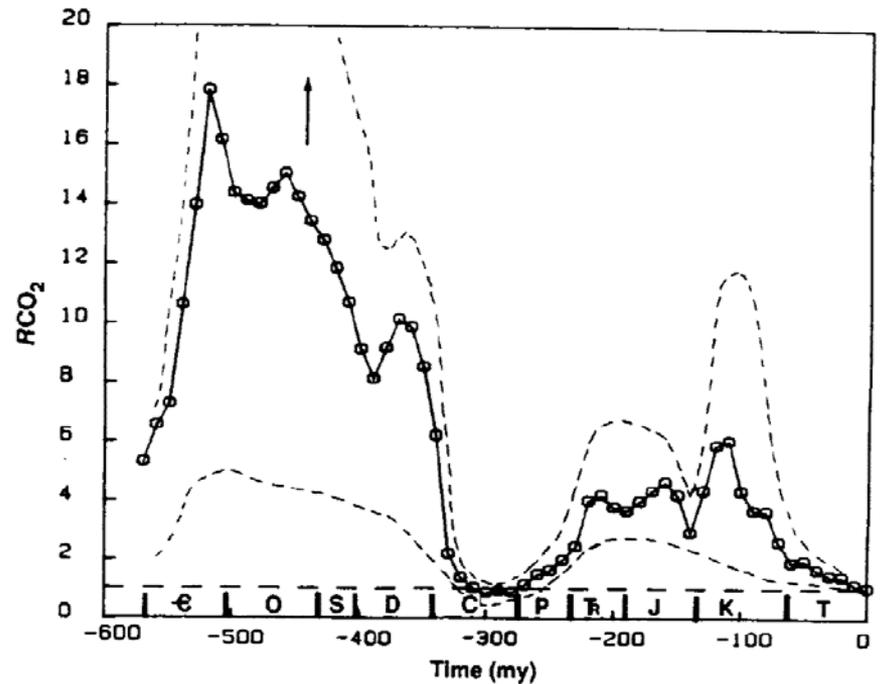
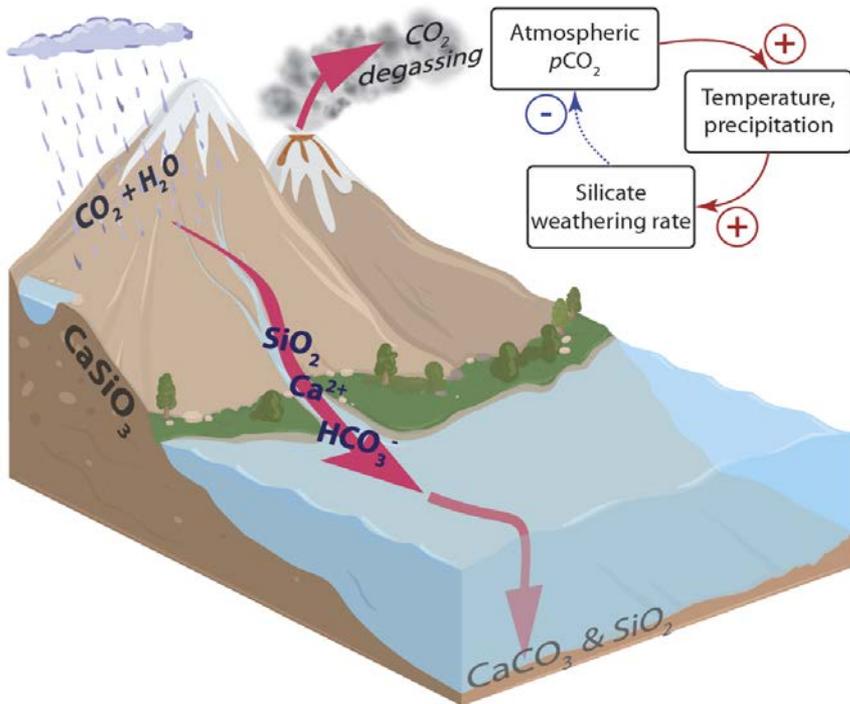
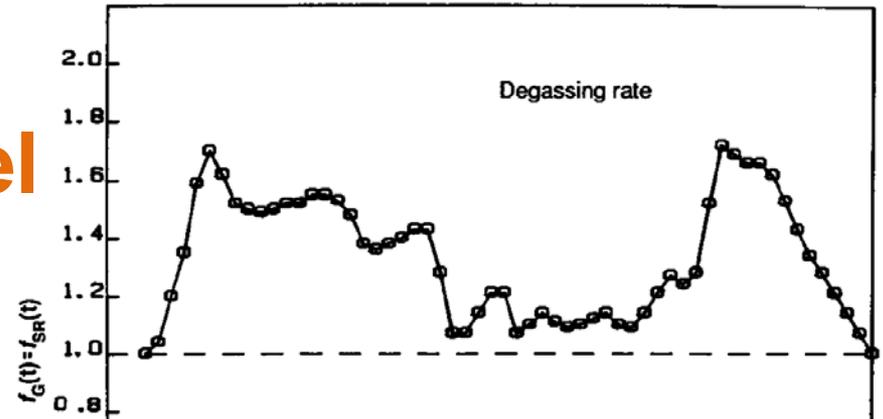
**Xiaomin Fang¹, Albert Galy², Yibo Yang¹, Weilin Zhang¹,
Chengcheng Ye¹, Chunhui Song³**

- 1. Institute of Tibetan Plateau Research, CAS**
- 2. CRPG-CNRS-University of Lorraine**
- 3. Lanzhou University**



BLAG model

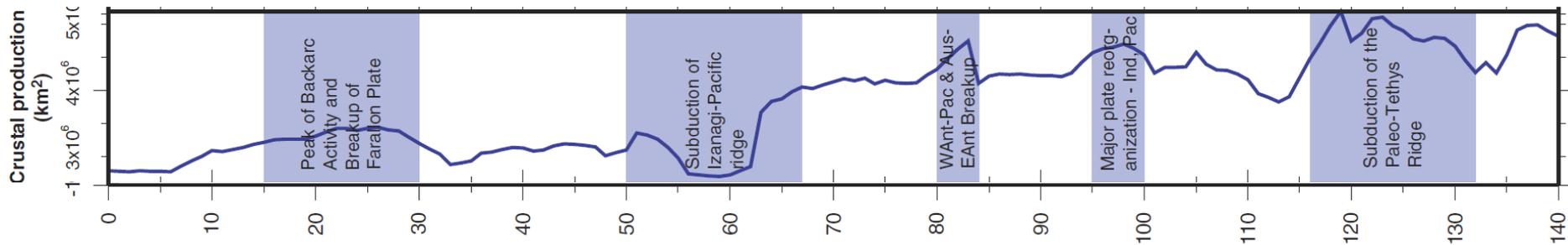
Robert A. Berner



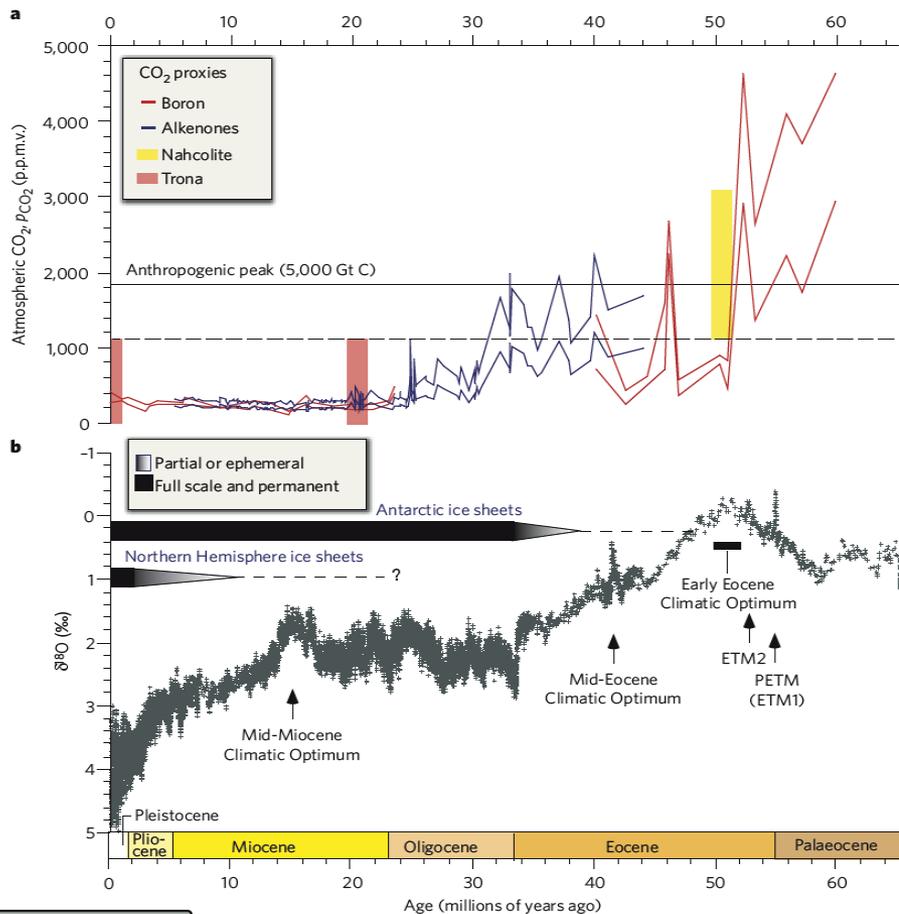
Frings, 2019

Berner, 1990





Muller et al., 2008

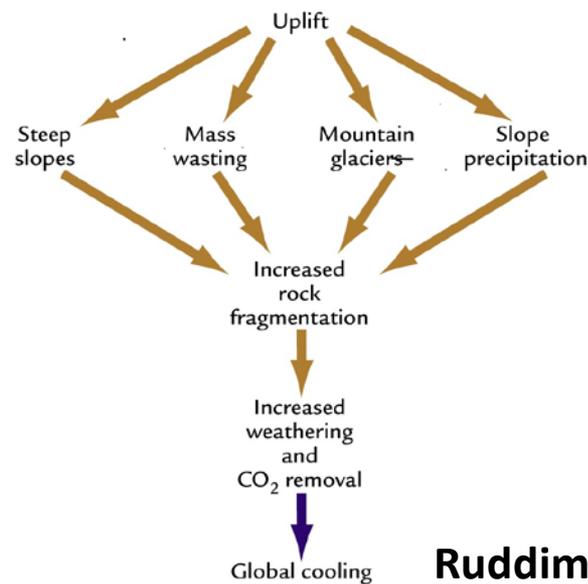


Zachos et al., 2008



Uplift-Weathering Hypothesis

Maureen Raymo



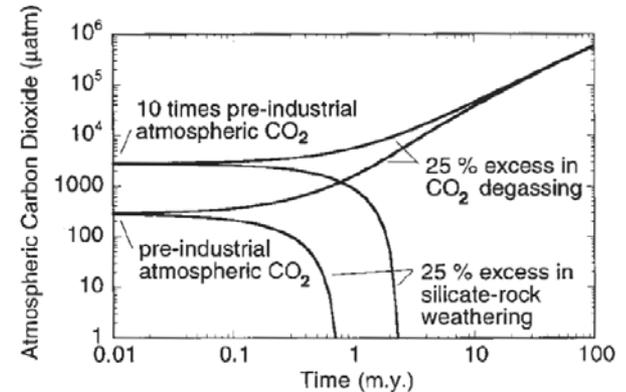
Ruddiman, 2008



Negative Feedbacks to maintain geological carbon cycle

The imbalance resulting from accelerated CO_2 consumption and a relatively stable CO_2 input from volcanic degassing during the Cenozoic should have depleted atmospheric CO_2 within a few million years.

Therefore, a negative feedback mechanism must have stabilised the carbon cycle.



Bener and Caldeira, 1997

- 1 Reduce organic carbon burial (Raymo and Ruddiman, 1992)
- 2 Enhance reversal weathering and CO_2 release (Raymo and Ruddiman, 1992)
- 3 Enhance metamorphism degassing (Bickle, 1996)
- 4 Enhance sulfide weathering and CO_2 release (Torres et al., 2014)
- 5 Reduce Temperature-regulated weathering (Kump and Arthur, 1997)

5.1 Ocean crust basalt (Coogan and Dosso, 2015)

5.2 Ocean island basalt (Li et al., 2013)

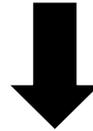
5.3 Continental arc (Lee et al., 2015)

5.4 Weatherability (Caves et al., 2016)

5.5 Weathering in other regions (Kump and Arthur, 1997)

Working hypothesis

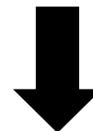
CO₂ consumption increase **in tectonically active region** (e.g., Himalayas, where erosion rate increases)



CO₂ decline reduced the degree of silicate alteration **in tectonically less active region** (where erosion rate is stable)

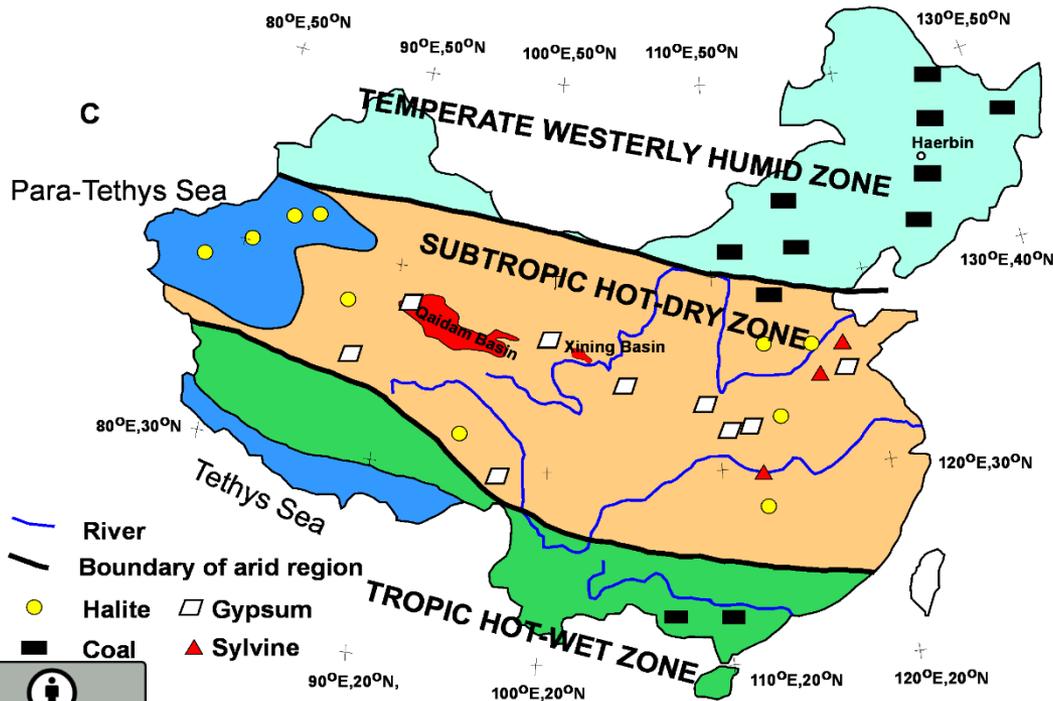
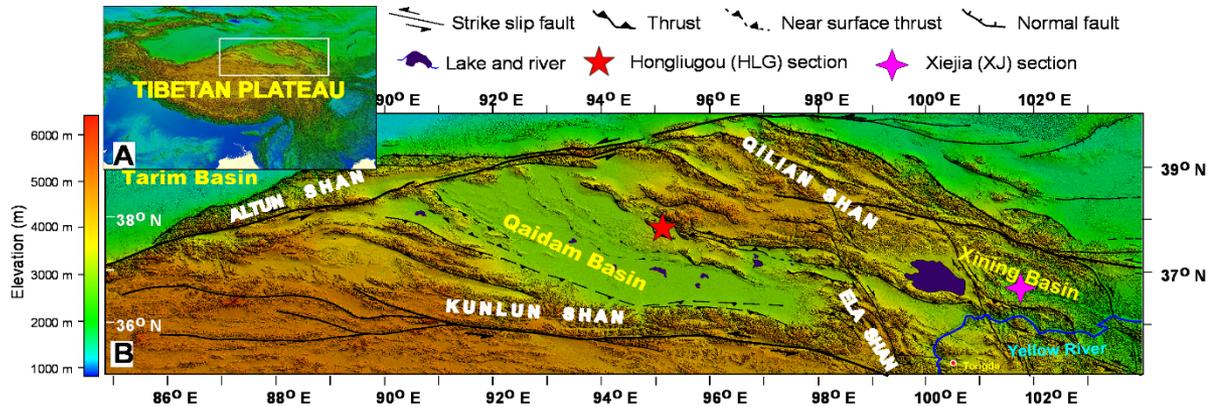


Decrease in silicate weathering flux and CO₂ consumption **in tectonically less active region**



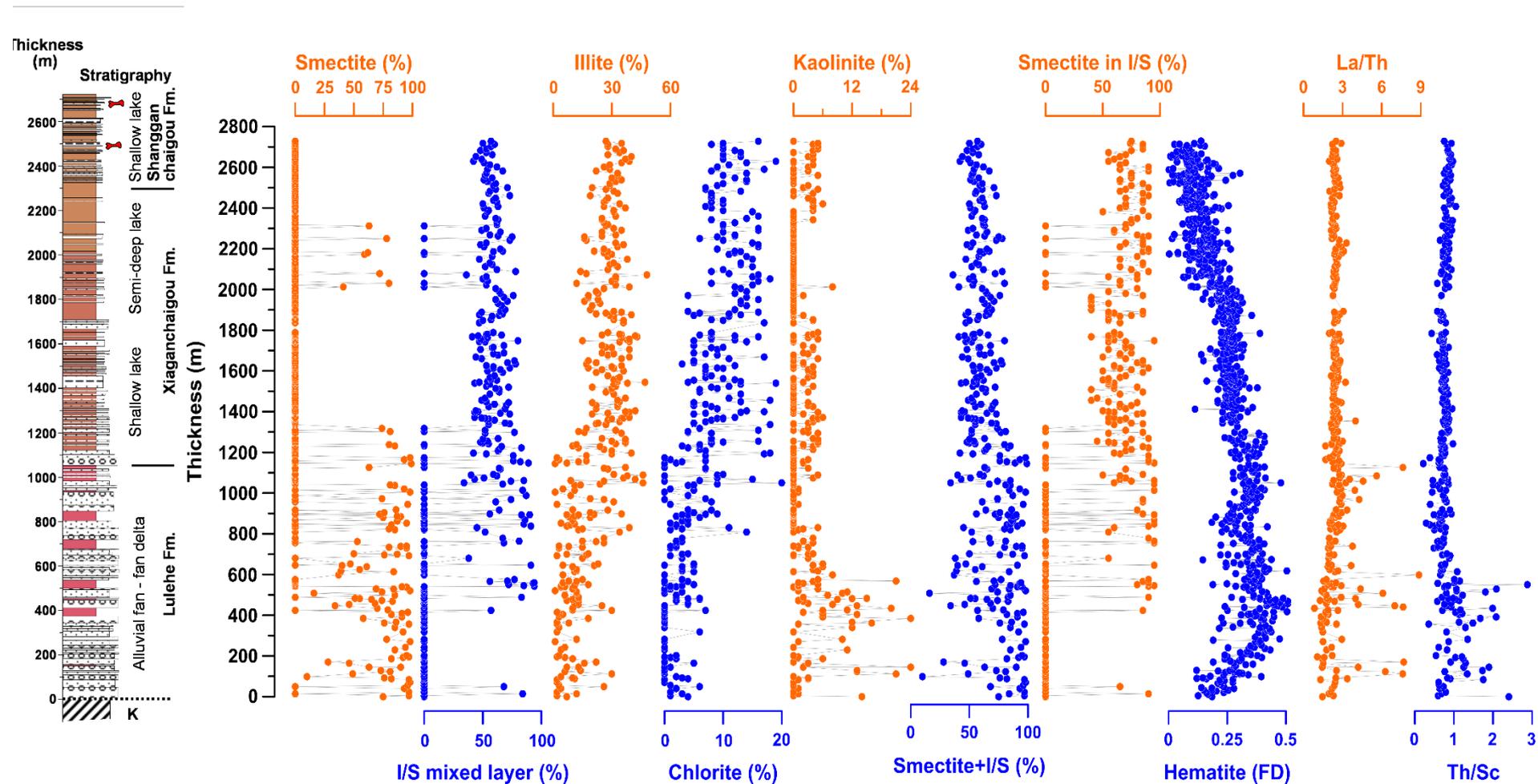
Carbon cycle balanced!!

Study region

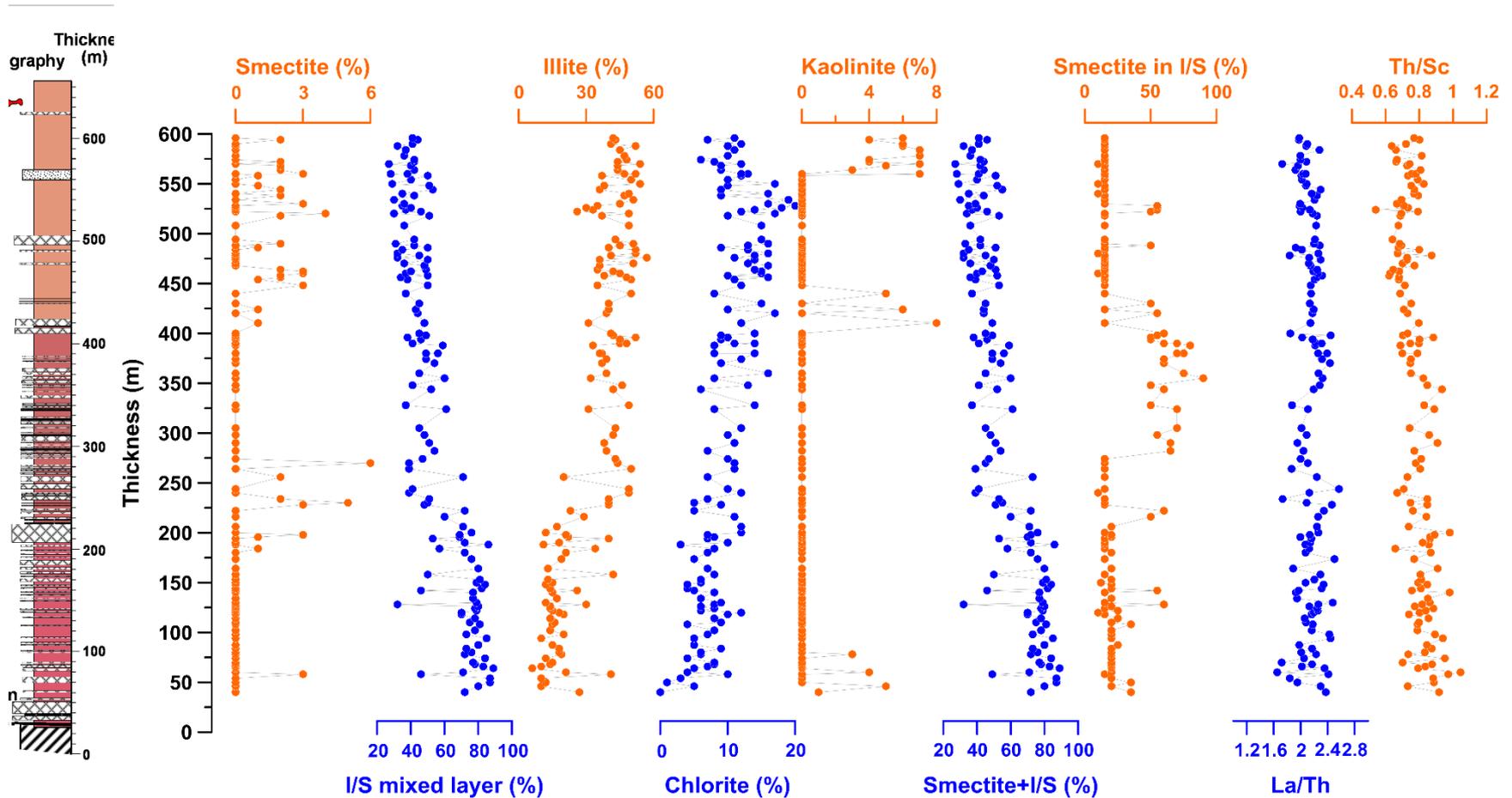


- 1) During the Paleogene, the northern Tibetan Plateau was a tectonically less active area;
- 2) The Asian summer monsoon did not reach this region until the beginning of the Neogene

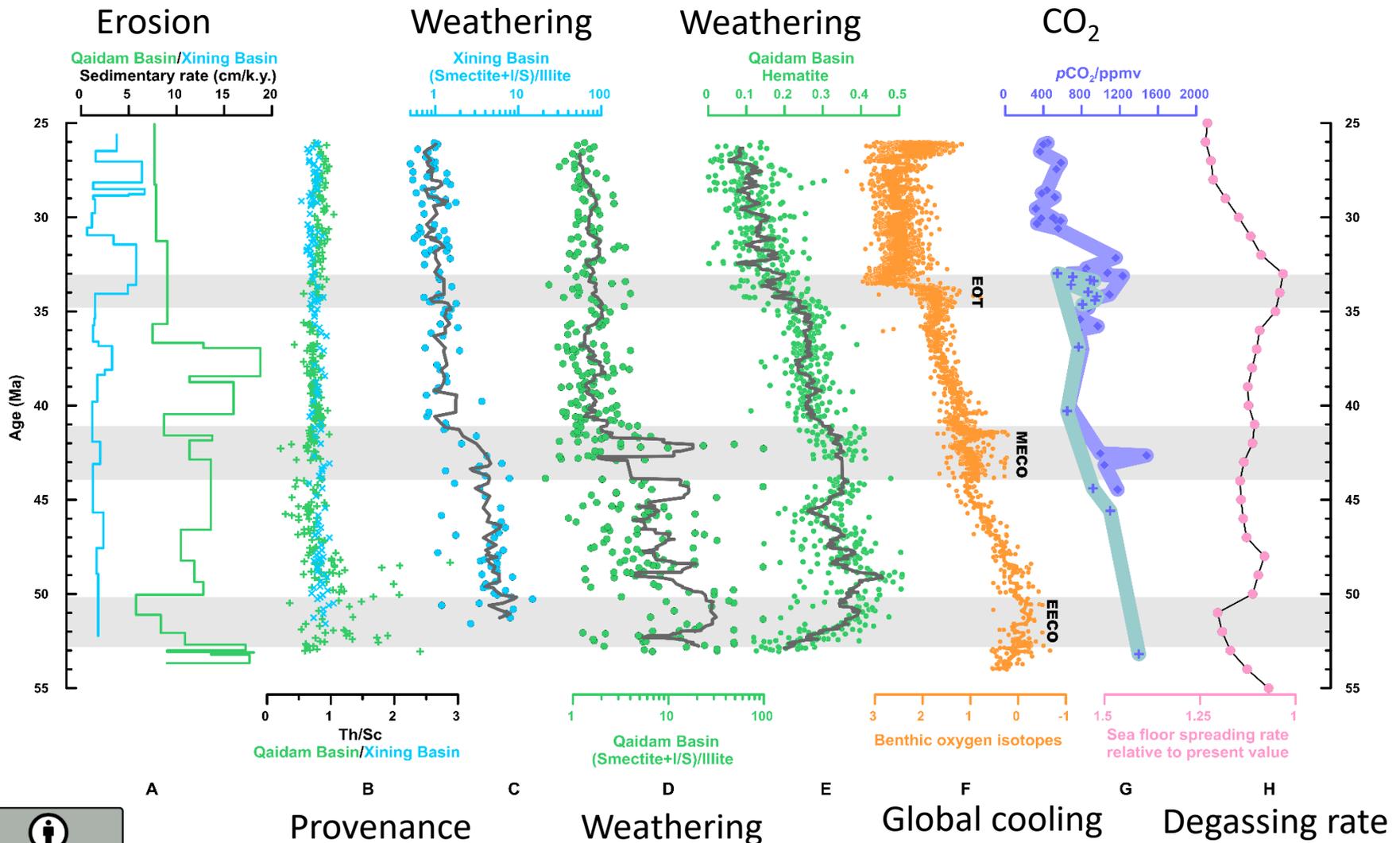
Clay mineral records in Qaidam Basin



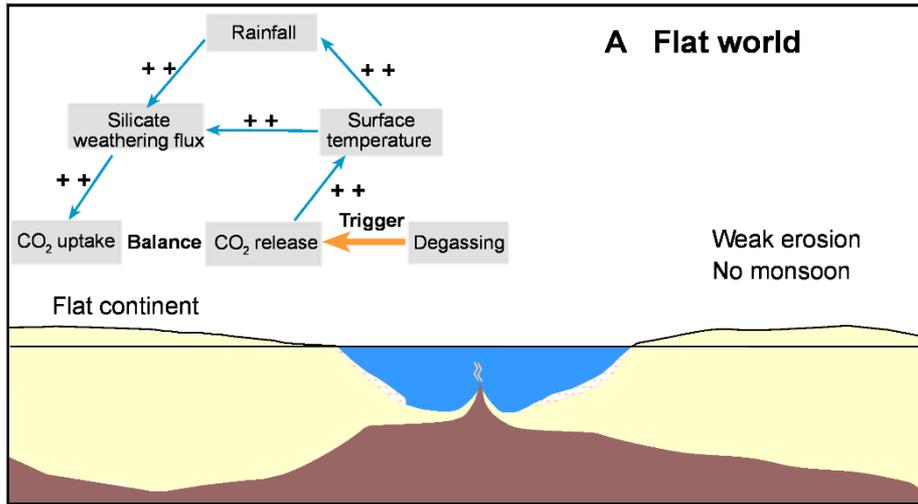
Clay mineral records in Xining Basin



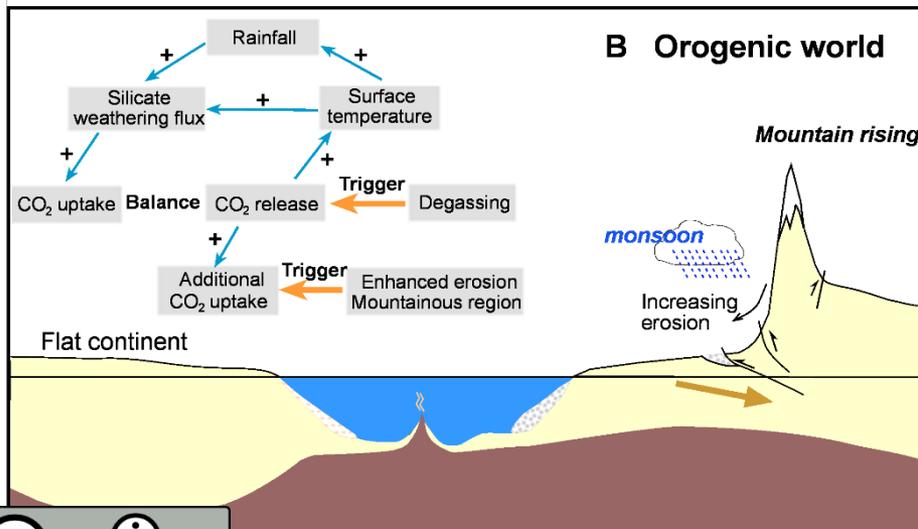
Global cooling regulate silicate weathering intensity in the northern Tibetan Plateau



Conceptual model



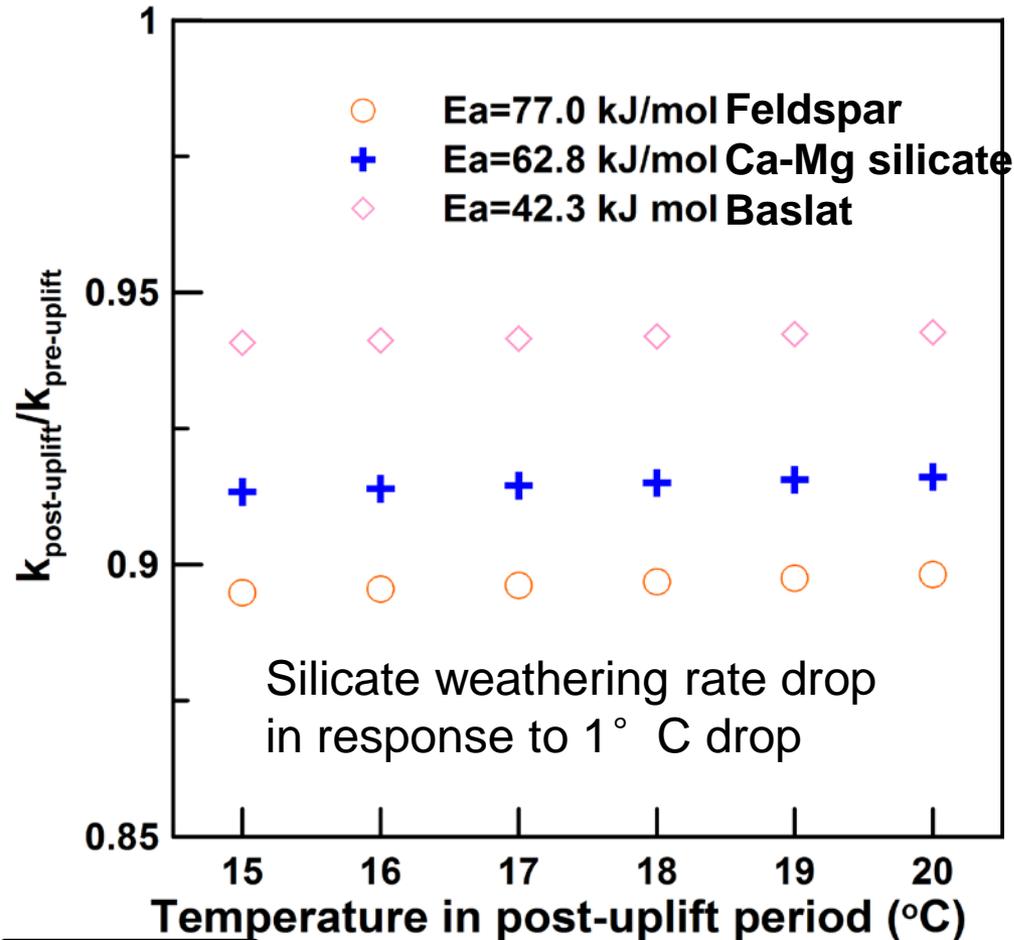
In scenario (A), any increase in degassing flux will lead to an increase in temperature and ultimately will be balanced by nearly the same amount of increase in CO₂ uptake flux through increases in the silicate weathering intensity and thus in the silicate weathering flux.



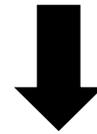
In scenario (B), given a constant degassing flux, any increase in erosion in orogenic belts will result in an overall increase in the CO₂ uptake flux in orogenic belts. This will be quasi balanced by nearly the same amount of decrease in CO₂ uptake flux in flat tectonically inactive regions through decreases in in the silicate weathering intensity and the silicate weathering flux

Model test

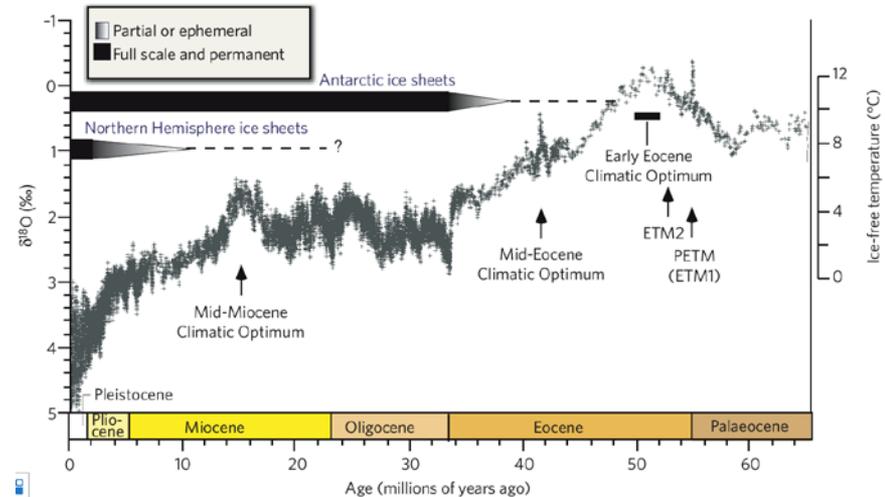
$$\ln(k_{\text{post-uplift}}/k_{\text{pre-uplift}}) = -Ea/R(1/T_{\text{post-uplift}} - 1/T_{\text{pre-uplift}})$$



52-36 Ma Temperature drop 7-8 ° C,
 ~0.44 ° C/Myr

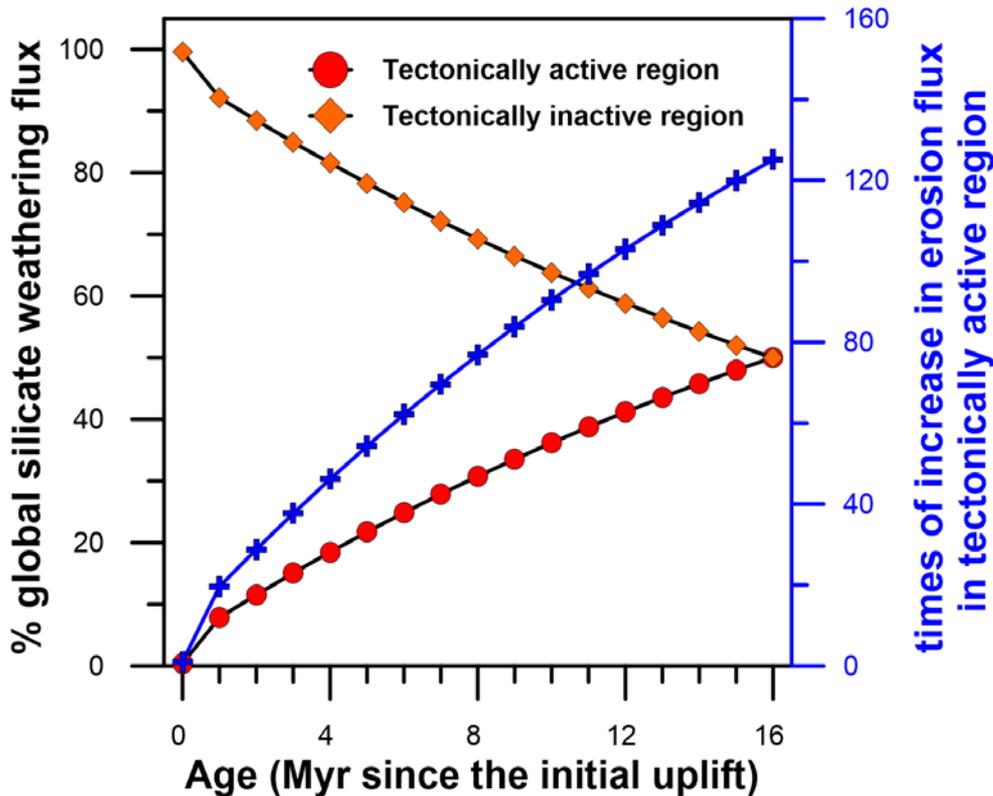


Silicate weathering rate drop
 4%/Myr



Zachos et al., 2008

Model test



Pre-uplift:

Area of tectonically inactive region : 99.6%

Area of tectonically active region : 0.4%

Uniform erosion flux everywhere,
Weathering flux in proportion to area

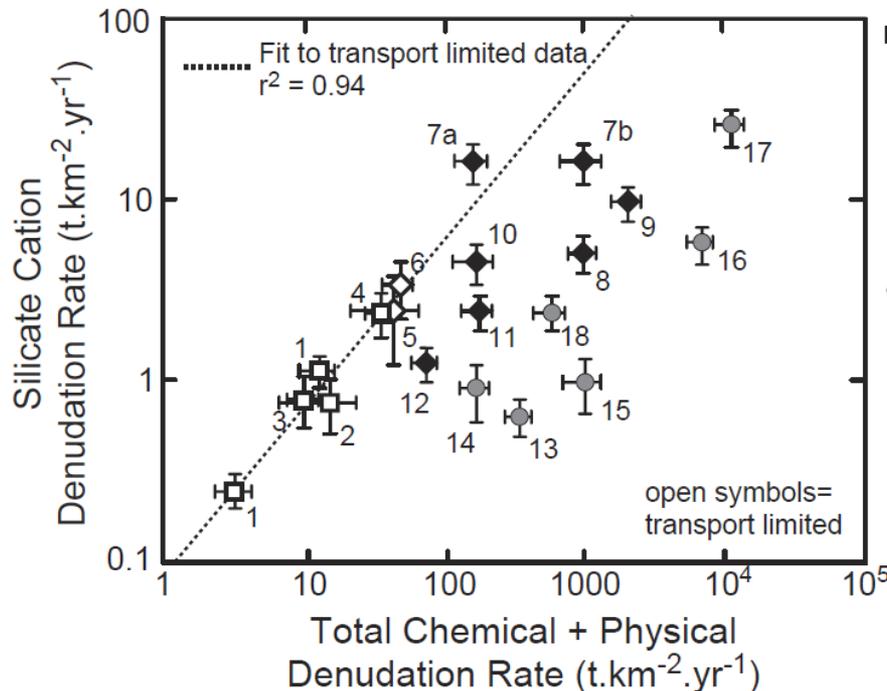
Post-uplift:

Weathering flux drop in tectonically inactive area would be balanced by erosion-induced weathering flux increase in tectonically active region

Result:

Erosion increase in tectonically active region should be >100-fold

Modern observations



- **Continental Cratons**
 - 1. Canadian Shield
 - 2. Siberian Shield
 - 3. African Shield
 - 4. Guyana Shield
- **Alpine Catchments**
 - 13. Colorado Rockies
 - 14. Sierra Nevada
 - 15. Svalbard
 - 16. High Himalaya
 - 17. West Southern Alps
 - 18. Swiss Alps
- ◆ **Submontane Catchments**
 - 5. British Columbia
 - 6. Sabah Malaysia
 - 7. Puerto Rico
 - a. Long Term Erosion
 - b. Modern Day Erosion
 - 8. East Southern Alps
 - 9. Lesser Himalaya
 - 10. Cote d'Ivoire
 - 11. Idaho Batholith
 - 12. Appalachians

Main display content can be found at

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Paleogene global cooling–induced temperature feedback on chemical weathering, as recorded in the northern Tibetan Plateau

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