

# A Regionally Adaptable Ground-Motion Model from European Strong Motion Dataset

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## 1. Datasets

- European Strong Motion dataset (Lanzano et al. 2018, Bindi et al. 2018)
- TSUMAPS-NEAM (Basili et al. 2019)
- TECTO ESHM2020 (Danciu et al.)

## 2. Ground-motion model (Kotha et al. 2020)

- Robust linear mixed-effects regressions
- Random-effect and residual variances
- Spatial variability of GMM coefficients

## 3. Application and testing

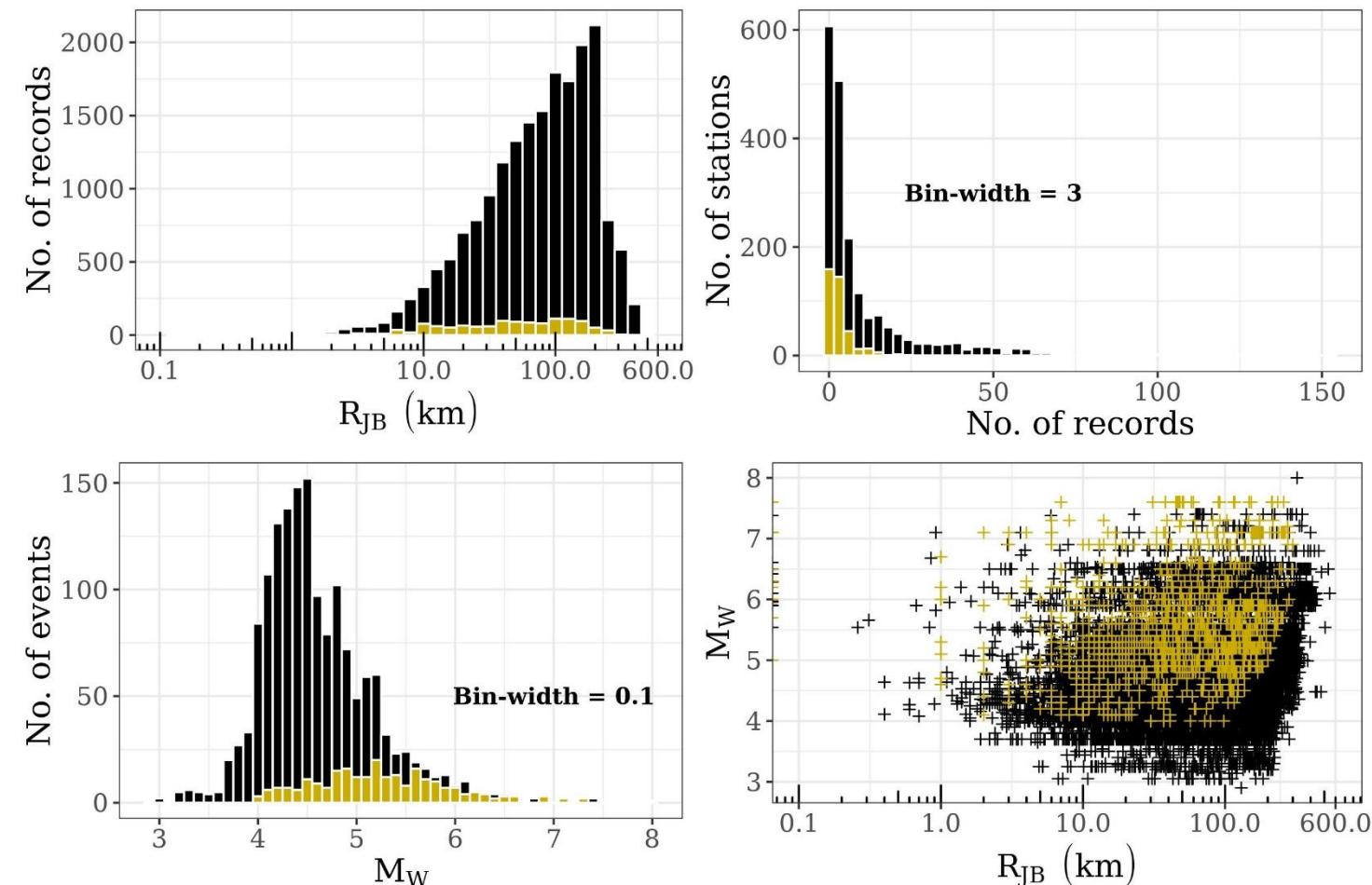
- Comparison with observed response spectra
- 10-fold cross-validation

# European Strong Motion dataset

ESM dataset: Lanzano et al. (2018) & Bindi et al. (2018)

RESORCE dataset: Akkar et al. (2014)

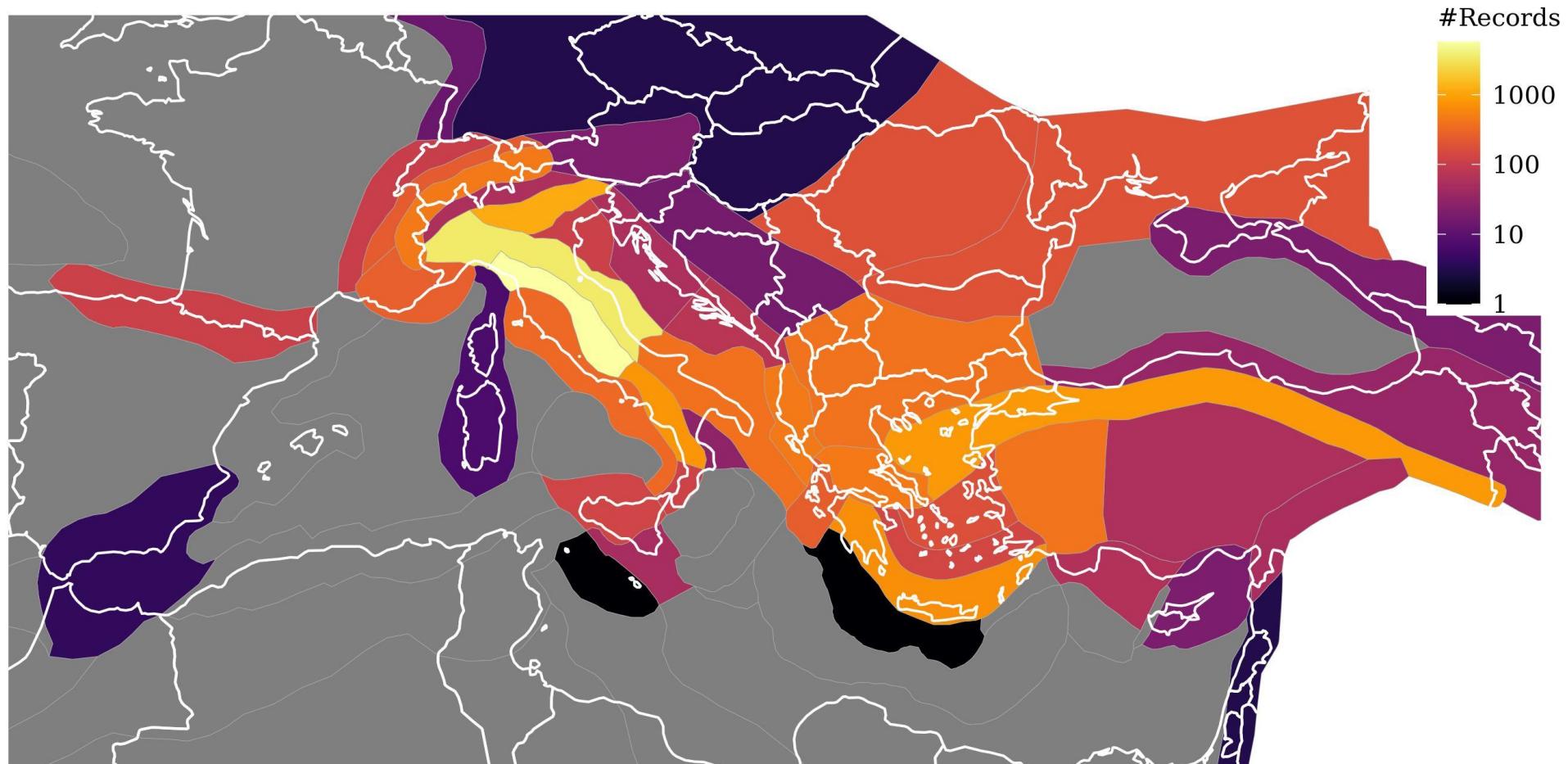
	<b>ESM</b>	<b>RESORCE</b>
$M_W$	3.0 – 7.4	4.0 – 7.6
$R_{JB}$	0 – 471km	0 – 300km
Sites	1829	384
Events	927	246
Tectonic localities	56	-
Attenuation regions	46	2
Records	18222	1251



# Regionalization of paths

Regionalization of apparent anelastic attenuation (TSUMAPs-NEAM - Basili et al. 2019)

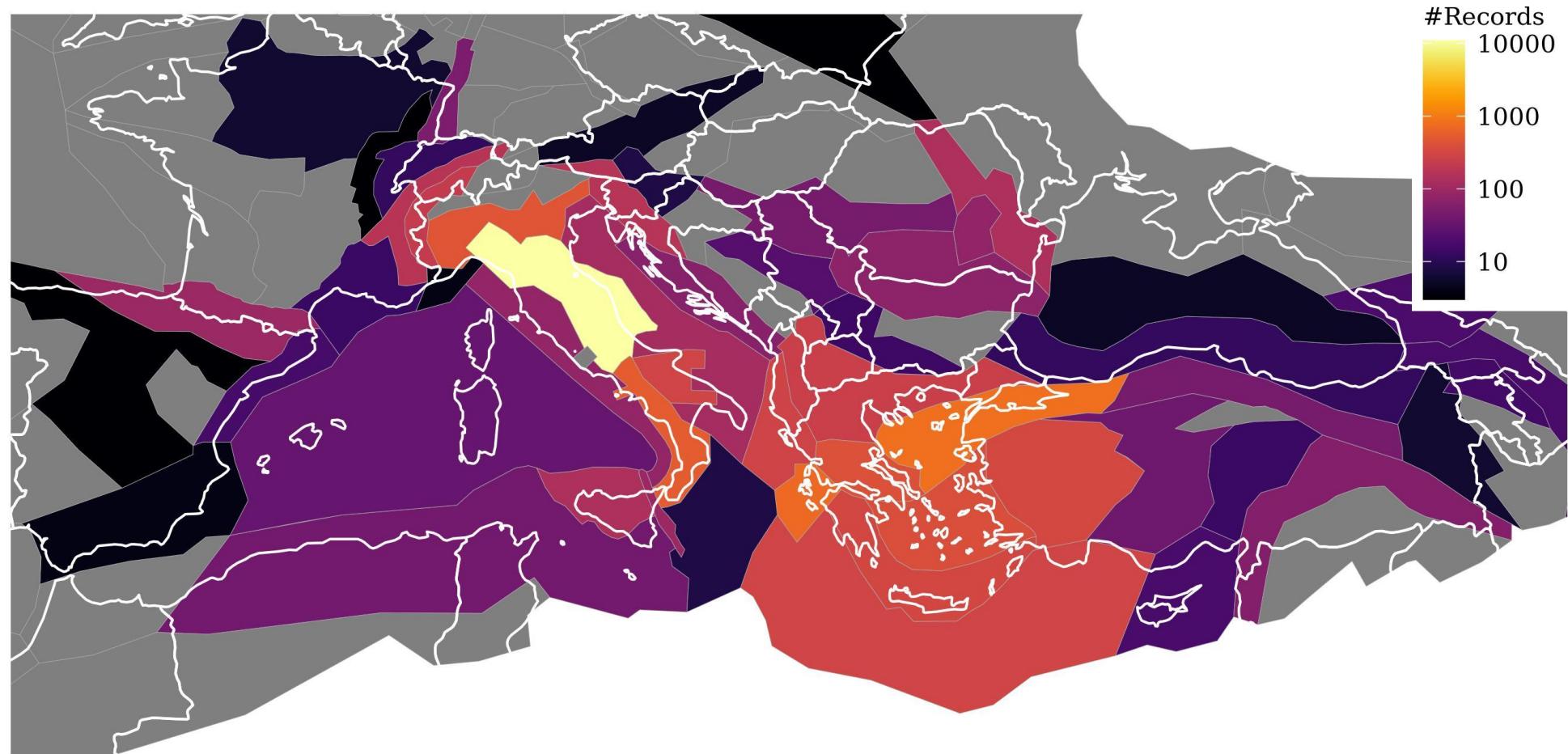
18222 records at 1829 sites into 46 regions (Kotha et al. 2020)



# Regionalization of sources

Localization of shallow crustal events into tectonic localities (ESHM2020 - Danciu et al.)

18222 records from 927 events into 56 tectonic localities (Kotha et al. 2020)



## Ground-motion model for PGA, PGV, SA for T = 0.01-8s - Kotha et al. (2020)

$$\ln(SA_{obs}) = F_M(M_W) + F_G(M_W, R_{JB}) + F_{A,r}(R_{JB}) + \delta L2L_l + \delta B_{e,l}^0 + \delta S2S_s + \varepsilon$$

1. Robust linear mixed-effects regression to flag outliers
2. Regionalization of anelastic attenuation (TSUMAPs-NEAM - Basili et al. 2019)
  - $F_{A,r}(R_{JB})$
3. Localization of earthquakes into tectonic localities (ESHM2020 - Danciu et al.)
  - $\delta L2L_l$
4. Site-terms for regional risk assessment (Crowley et al. 2020, Weatherill et al.)
  - $\delta S2S_s$

## ESM (solid) v/s RESORCE (dashed) based Kotha et al. (2016) GMM

$$\sigma = \sqrt{\Phi_{S2S}^2 + \Phi^2 + \tau^2}$$

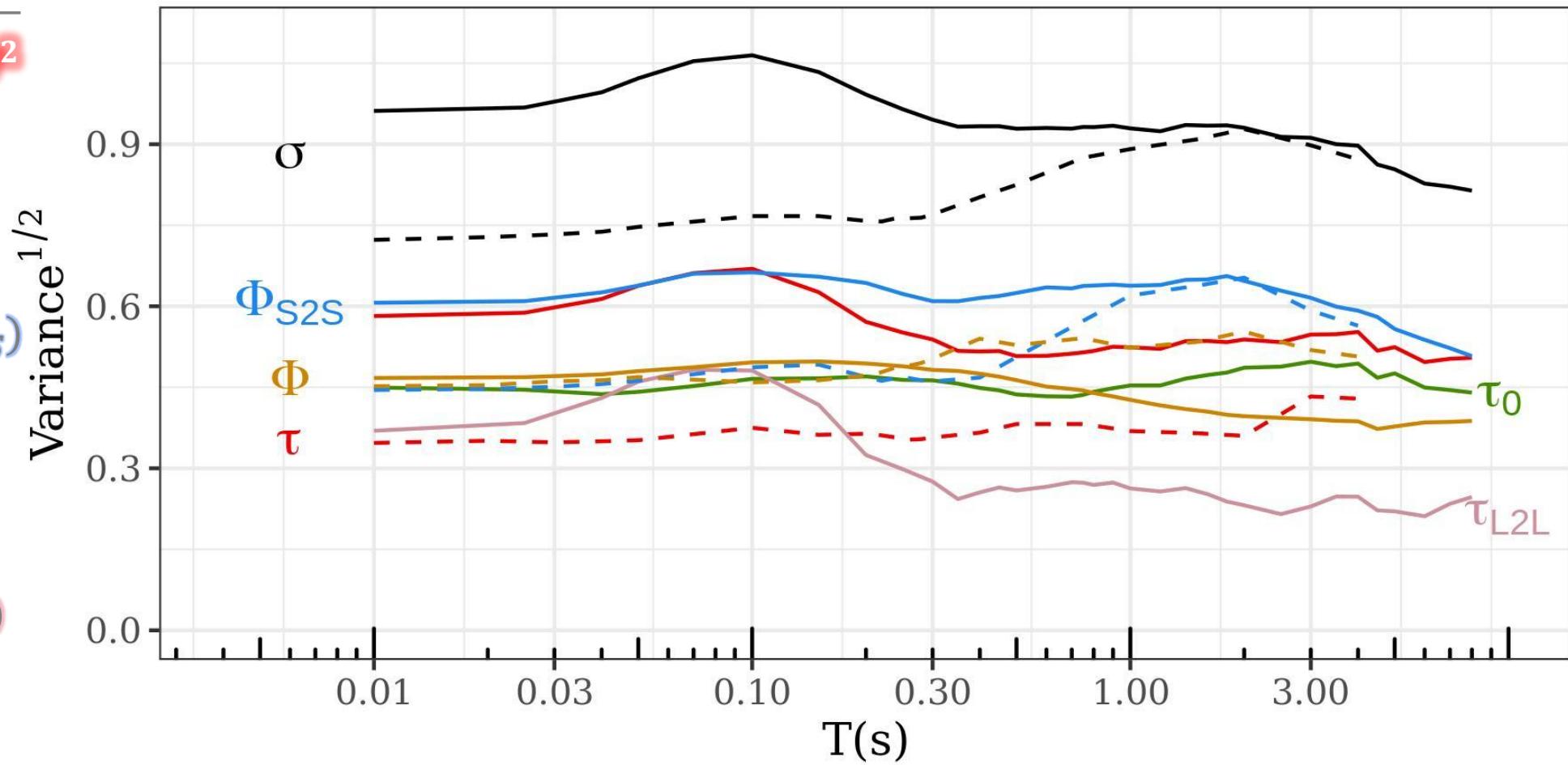
$$\tau = \sqrt{\tau_0^2 + \tau_{L2L}^2}$$

- $\Delta S2S_s = N(0, \Phi_{S2S})$

- $\varepsilon = N(0, \Phi)$

- $\Delta B_{e,l}^0 = N(0, \tau_0)$

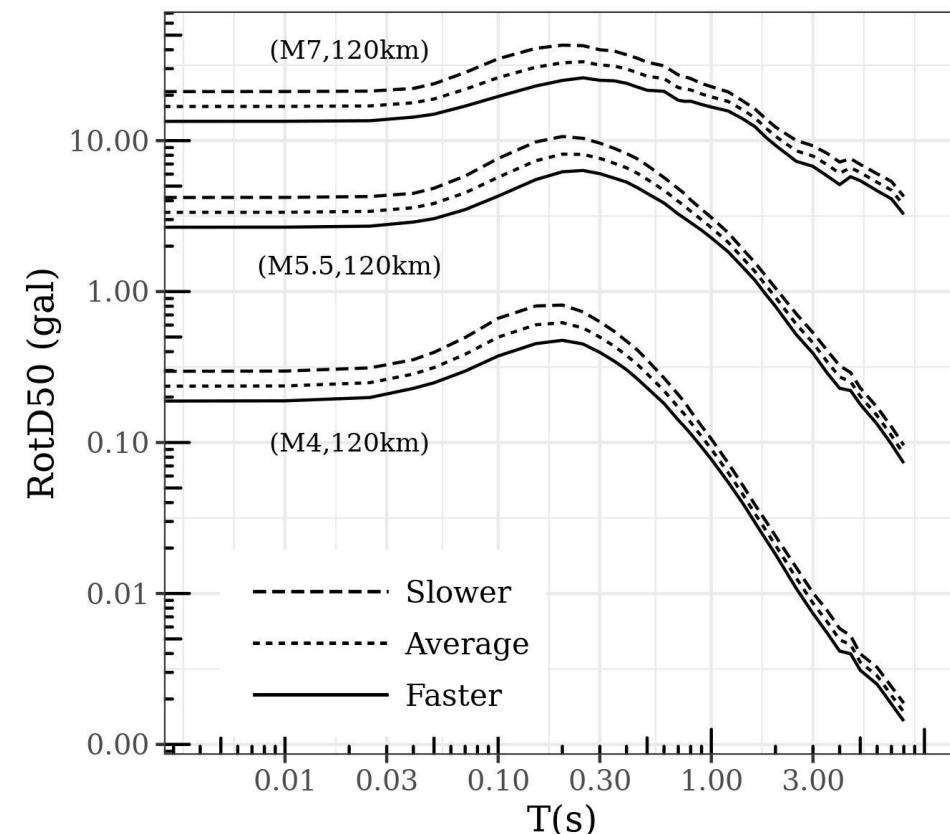
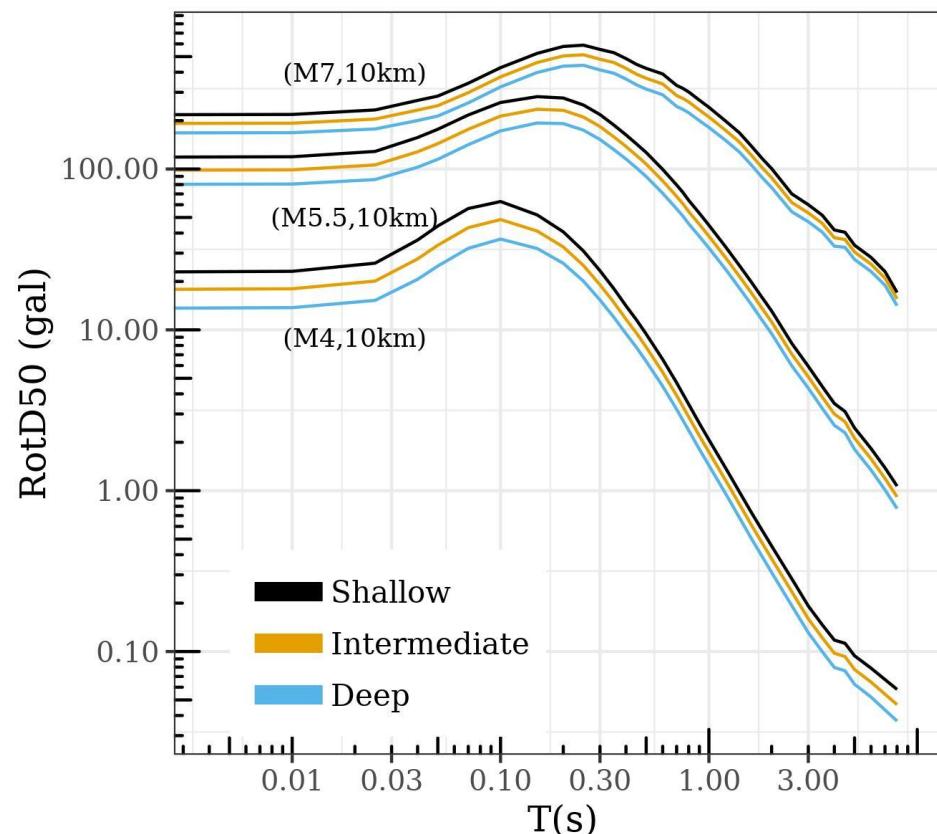
- $\Delta L2L_l = N(0, \tau_{L2L})$



# Response spectra

## Regionalized attenuation and event hypo. depth dependence – 9 combinations

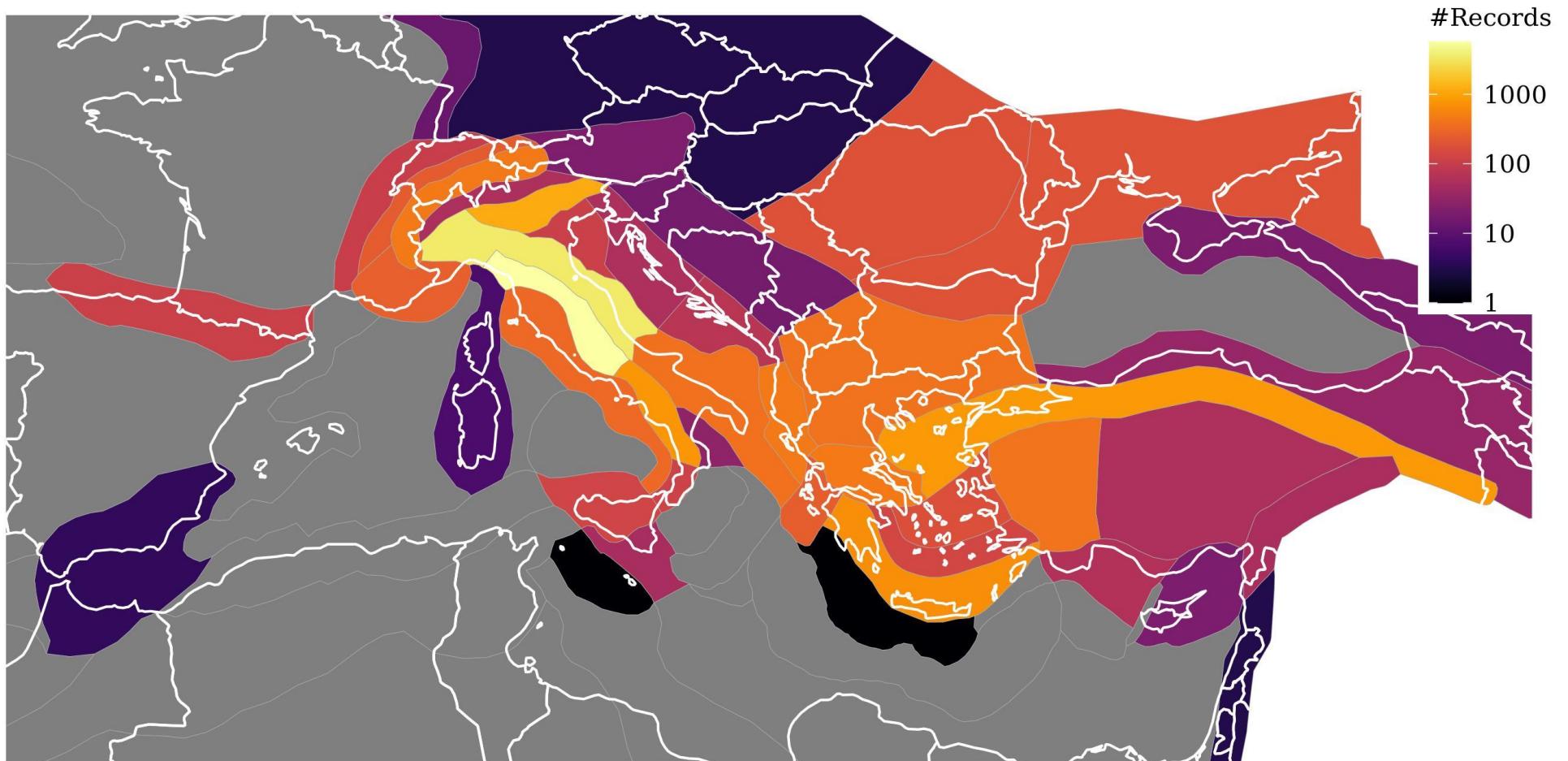
- Shallow events -  $D_{\text{hypo}} < 10\text{km}$  ; Inter. -  $10 \leq D_{\text{hypo}} < 20\text{km}$  ; Deep -  $D_{\text{hypo}} \geq 20\text{km}$
- Slower attenuation -  $\delta c_{3,r} = \tau_{c3}$  ; Average -  $\delta c_{3,r} = 0$  ; Faster -  $\delta c_{3,r} = -\tau_{c3}$



# Regionalized Anelastic Attenuation

$\Delta c_{3,r} = N(0, \tau_{c3})$  - High seismicity regions attenuate faster than average

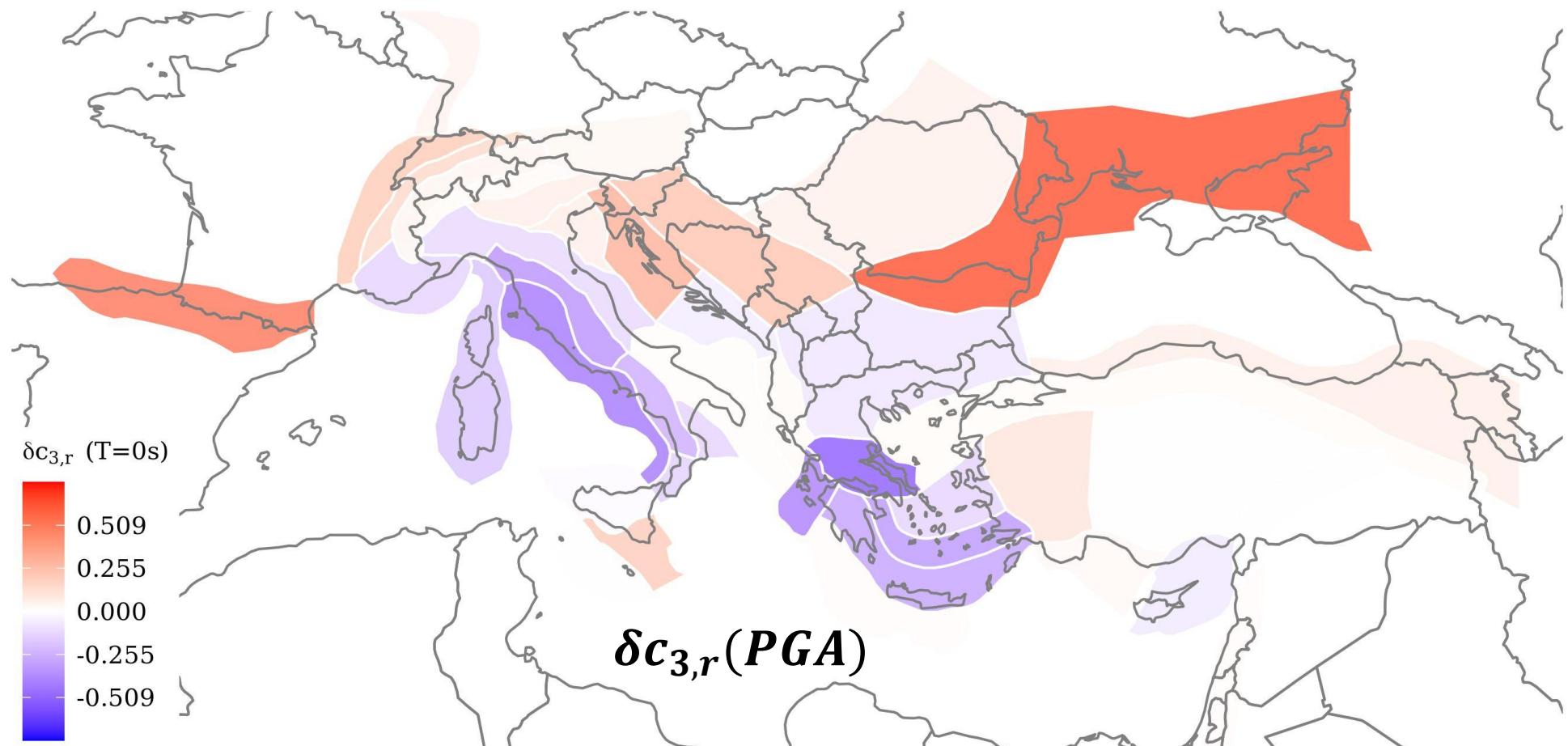
Volcanic regions attenuate even faster, e.g. Gulf of Corinth



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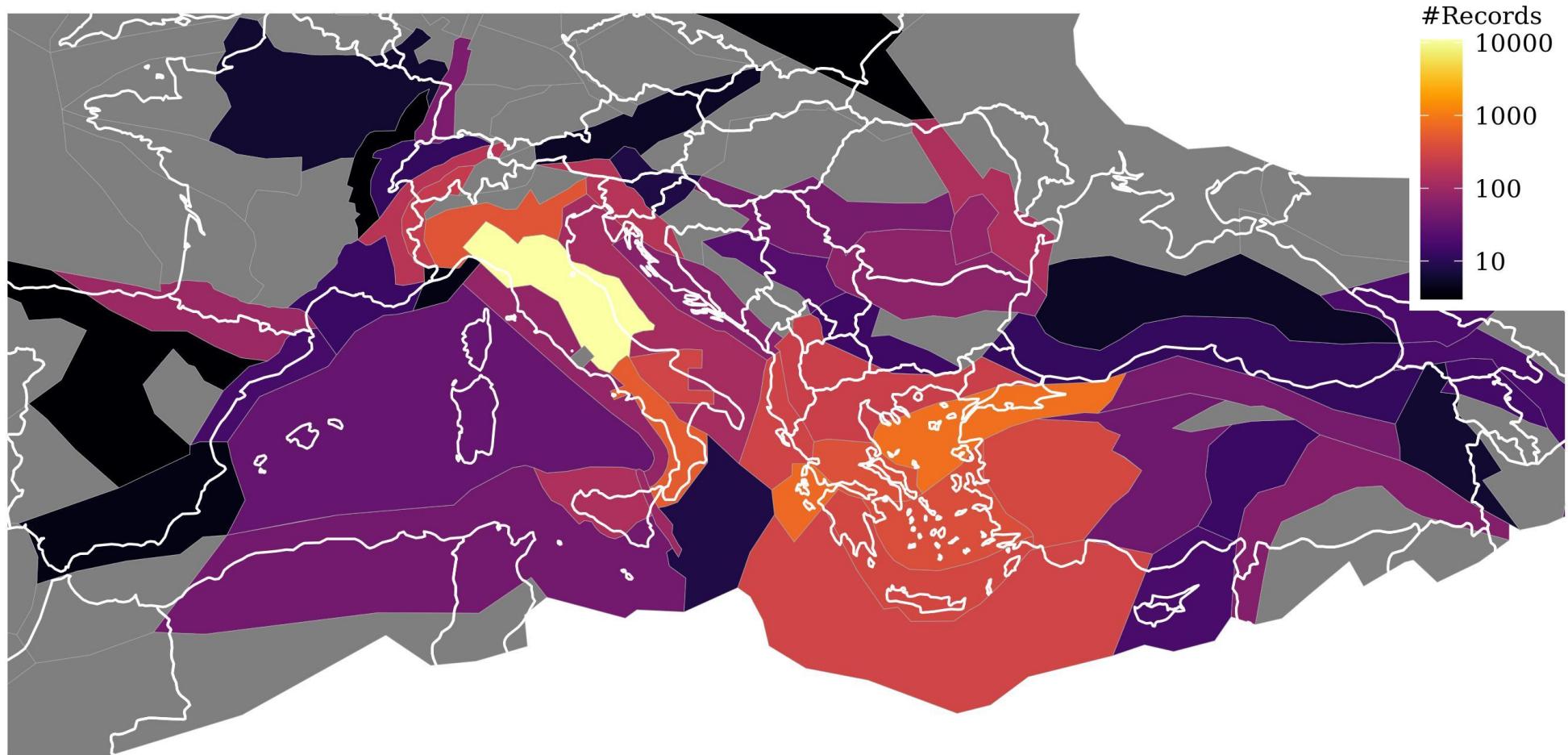
Volcanic regions attenuate even faster, e.g. **Gulf of Corinth**



# Regionalized Sources

$\Delta B_{e,l}^0 = \mathcal{N}(0, \tau_0)$  – No discernible spatial pattern, after location correction

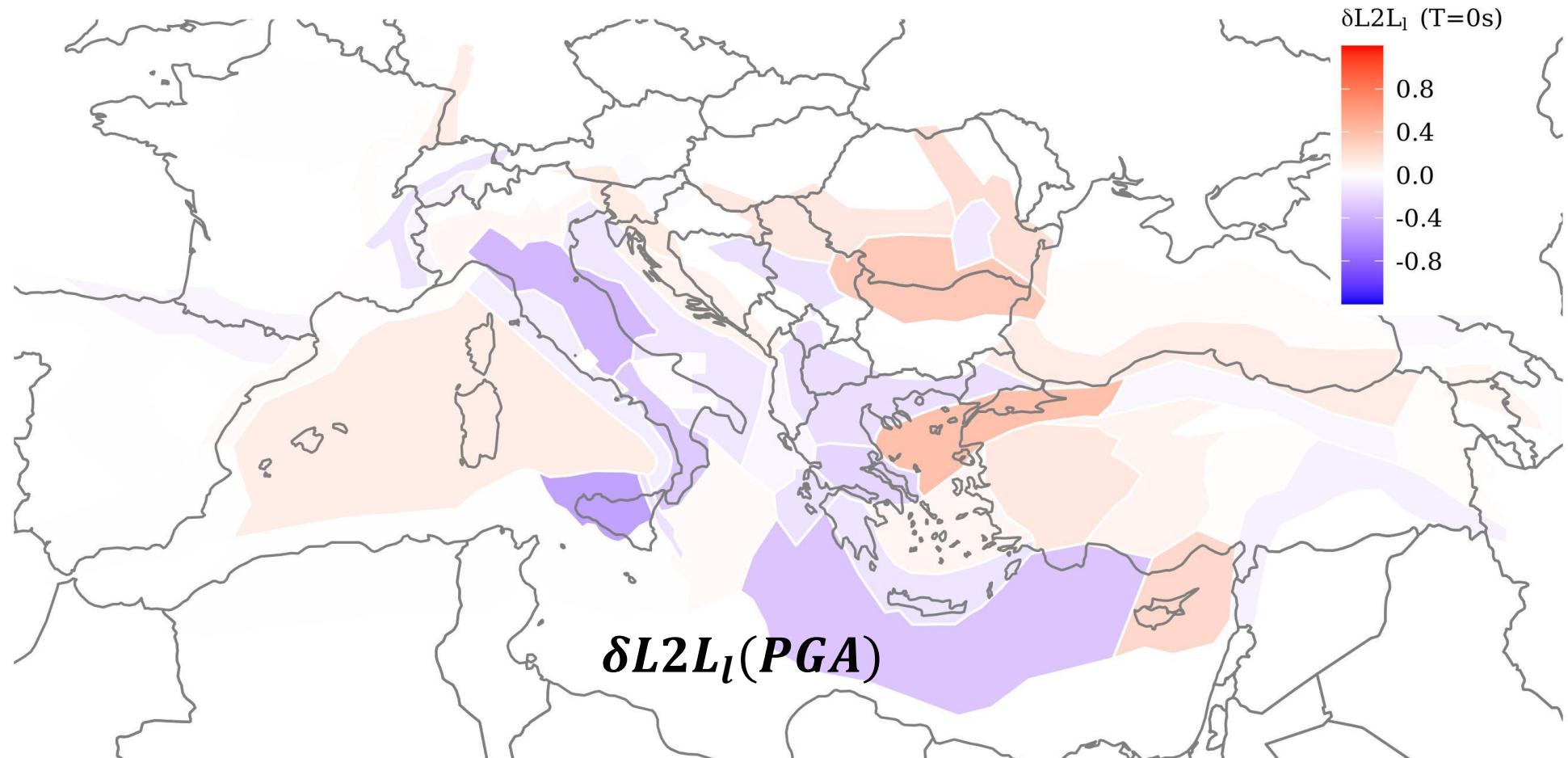
$\Delta L2L_l = \mathcal{N}(0, \tau_{L2L})$  – Stronger events in North-West Anatolia than in Central Apennines



# Regionalized Sources

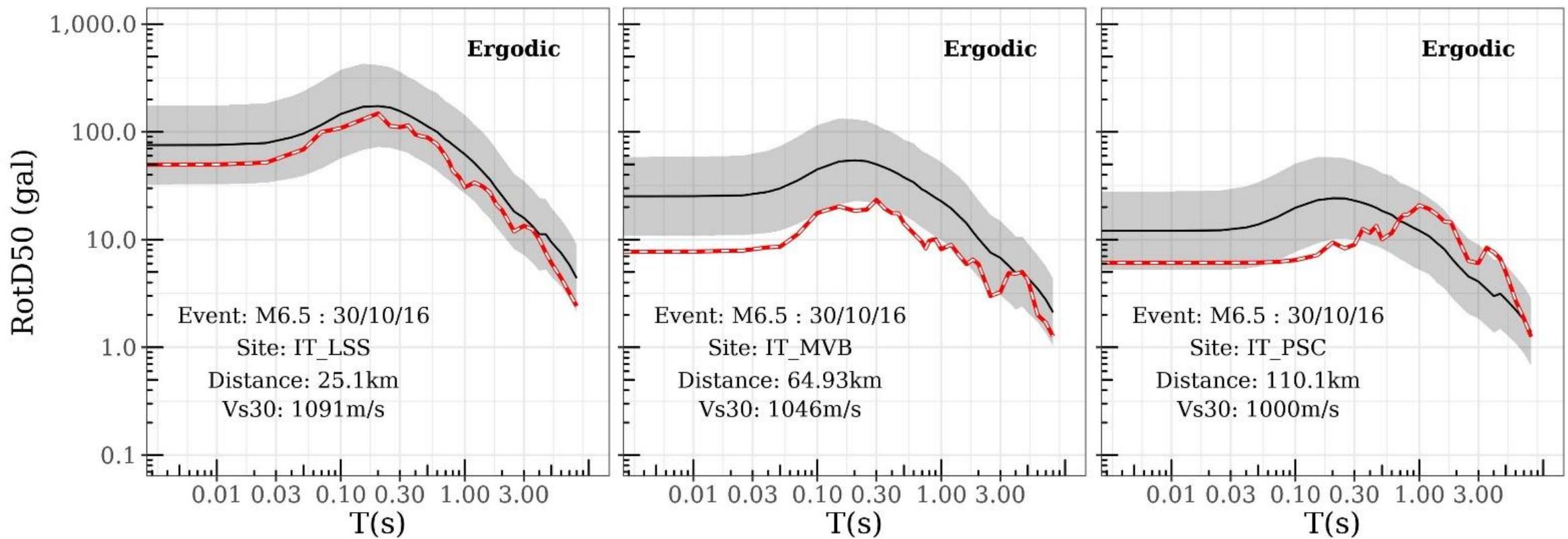
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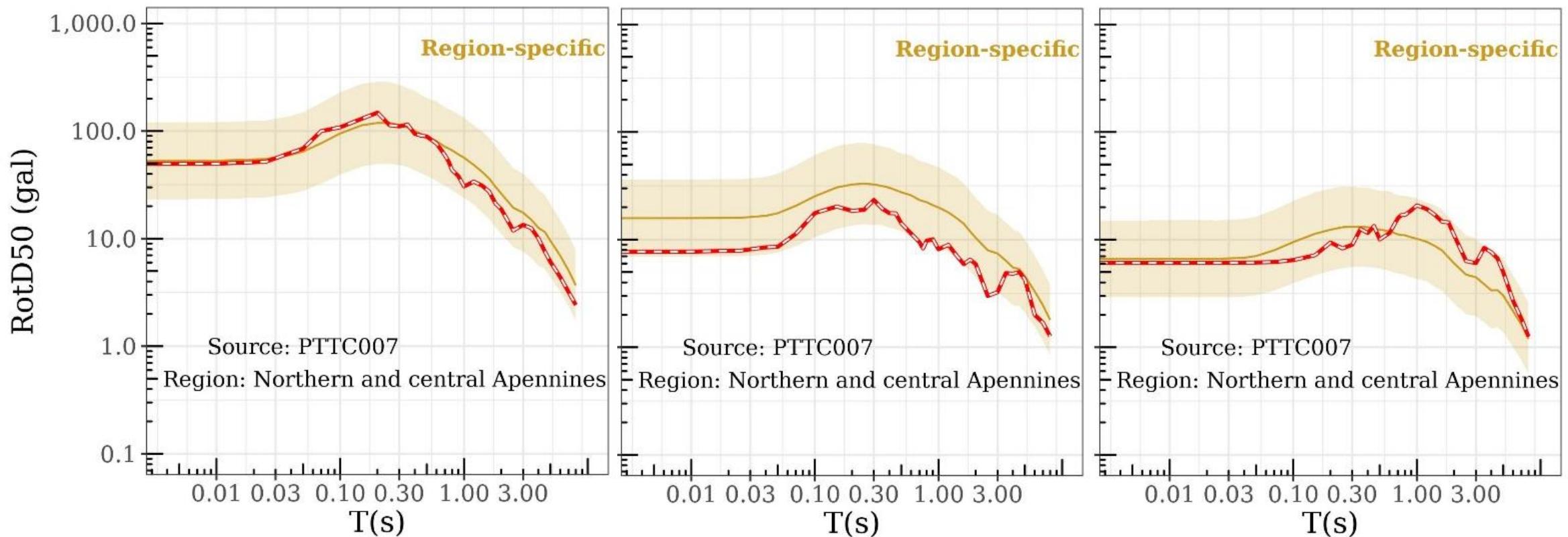
# Comparison with M6.5 Norcia (2016) spectra

*Ergodic prediction* –  $N(\mu, \sqrt{\tau_0^2 + \tau_{L2L}^2 + \Phi_{S2S}^2 + \Phi^2})$



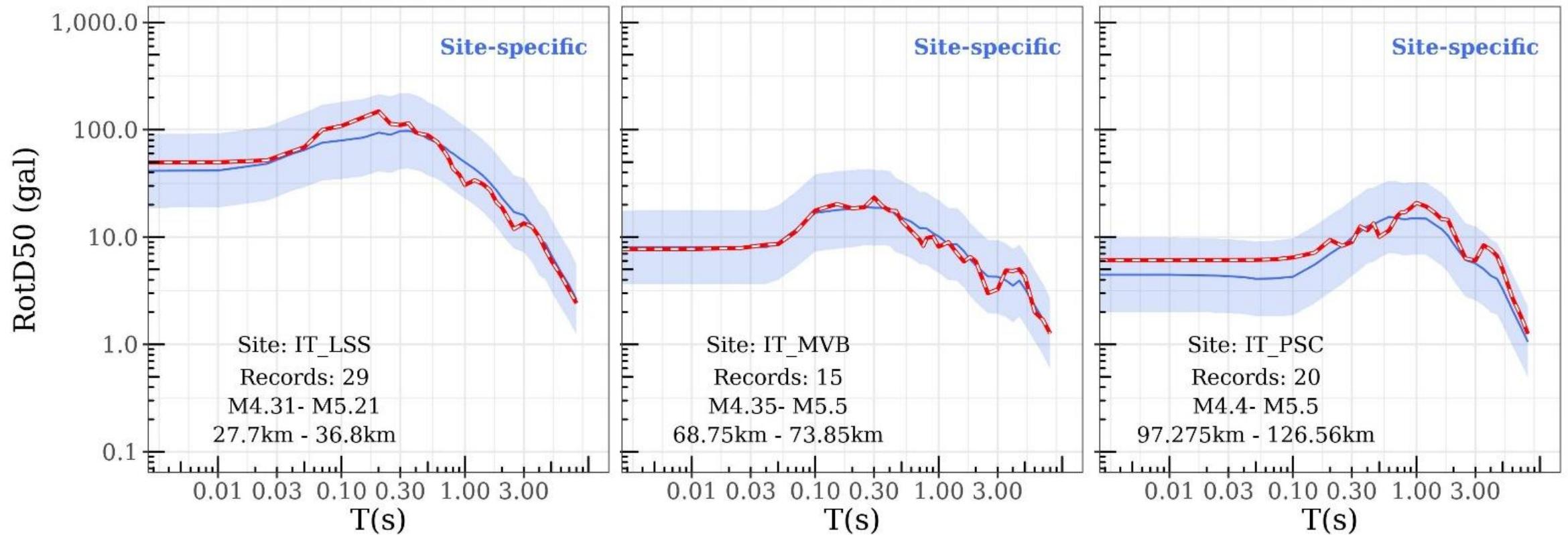
# Comparison with M6.5 Norcia (2016) spectra

*Regionalized prediction* –  $N(\mu(\delta c_{3,r}) + \delta L2L_l, \sqrt{\tau_0^2 + \varphi_{S2S}^2 + \varphi^2})$



# Comparison with M6.5 Norcia (2016) spectra

*Site – specific prediction –  $N(\mu(\delta c_{3,r}) + \delta L2L_l + \delta S2S_s, \sqrt{\tau_0^2 + \varphi^2})$*

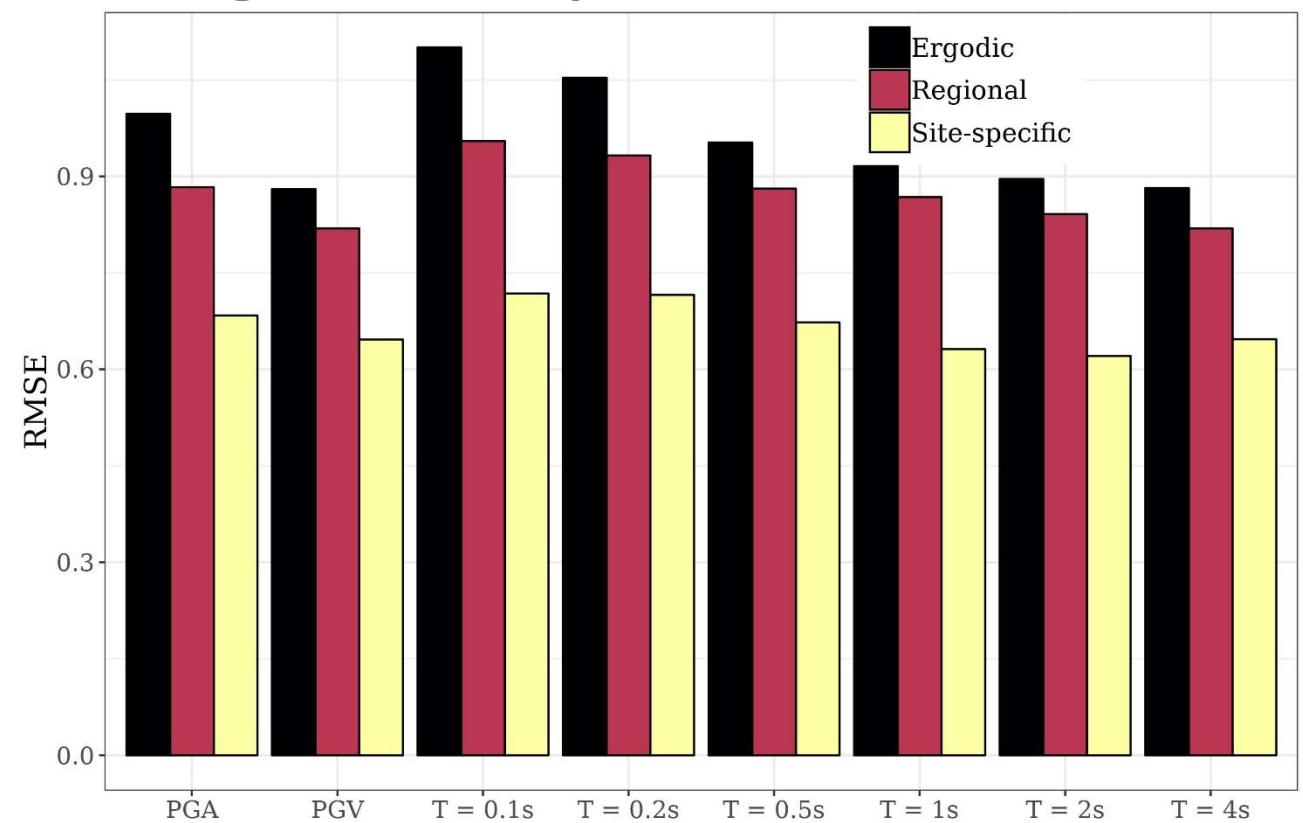


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Recognizable improvement from ergodic to region to site-specific GMMs

- 10 subsets with exclusive events
- Train on 9 subsets
- Test on 1 subset
- Repeat 10 times for each period
  - Ergodic
  - Regional
  - Region and site-specific



# Thank you!!!

## Bibliography

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