

# Examining saturated and unsaturated hydraulic parameter changes as a result of geochemical reactions in tailings

R. Akeseh<sup>1</sup>, M. Edraki<sup>2</sup> and T. Baumgartl<sup>3,4</sup>

<sup>1</sup>PhD student, Sustainable Minerals Institute, The University of Queensland, St Lucia QLD 4069, r.akeseh@uq.edu.au

<sup>2</sup>Principal Research Fellow, Sustainable Minerals Institute, The University of Queensland, St Lucia, QLD 4069. [m.edraki@cmlr.uq.edu.au](mailto:m.edraki@cmlr.uq.edu.au)

<sup>3</sup>Director, Geotechnical and Hydrogeological Engineering Research Group (GHERG), Federation University Australia, Gippsland VIC 3841, t.baumgartl@federation.edu.au

<sup>4</sup>Honary Professor, Sustainable Minerals Institute, The University of Queensland, St Lucia QLD 4069. t.baumgartl@uq.edu.au



## 1. INTRODUCTION

Water flow and solute transport in the unsaturated zone is governed by the hydraulic properties and the hydraulic conductivity function. However, geochemical reactions induce changes in the hydraulic properties as oxidation occurs in reactive materials. Transient flow experiment can be designed to indirectly estimates these changes. This is an easy and practical method which relies on easily measurable parameters. Our aim was to test the hypothesis of induced hydraulic properties changes due to geochemical reaction using inverse numerical modelling.

## 2. HYDRAULIC PROPERTIES ESTIMATION

- One dimensional vertical flow in rigid porous media is governed by Richards equation,

$$C(h) \frac{\partial h}{\partial t} = \frac{\partial}{\partial z} \left[ k(h) \left( \frac{\partial h}{\partial z} - 1 \right) \right]$$

- Initial and Boundary condition

$$h(z, t) = h_0(z) \quad t = 0, 0 \leq z \leq L$$

$$k(h) \left( \frac{\partial h}{\partial z} - 1 \right) = E(t)$$

- van Genuchten-Mualem hydraulic function

$$S_e(h) = \frac{(\theta(h) - \theta_r)}{\theta_s - \theta_r} = [1 + (\alpha h)^n]^{-m}$$

$$K(S_e) = K_s S_e^{0.5} \left[ 1 - (1 - S_e^{1/m})^m \right]^2$$

### INVERSE ANALYSIS

- Hydraulic parameters  $\alpha$  and  $n$  determined by minimisation of objective function

$$O(b) = \sum_{i=1}^2 \sum_{j=1}^M [h(z_i - t_j) - h^*(z_i, t_j, b)]^2$$

- Solution by Levenberg-Marquardt algorithm
- Solution was implemented in HYDRUS-1D

## 3. MATERIALS AND METHODS

- Material Composition
  - Dolomite (25%)
  - Pyrite (25%)
  - Quartz (35%)
  - Chlorite (15%)
- Two categories of particle size range
  - <75  $\mu\text{m}$
  - 75-200  $\mu\text{m}$
- Column: 4 cm, bulk density: 1.4 g/cm<sup>3</sup>
- Nine drying and wetting cycles over nine months
- Initial hydraulic parameters determined by desiccation
- Curve fitting in RETC
- Transient water flow experiment by evaporation
- Mass of water loss was measured by a balance
- Hydraulic conditions in the tailings measured by tensiometer

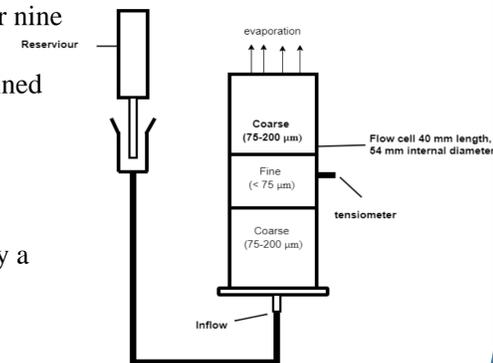


Fig 1: Evaporation Experiment

## 4. RESULTS

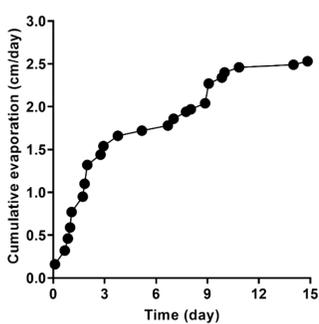


Fig 2: Cumulative water loss vs time

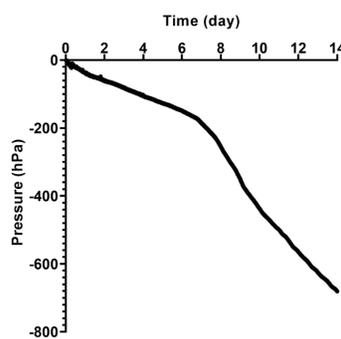


Fig 3: Time series of hydraulic condition in the column

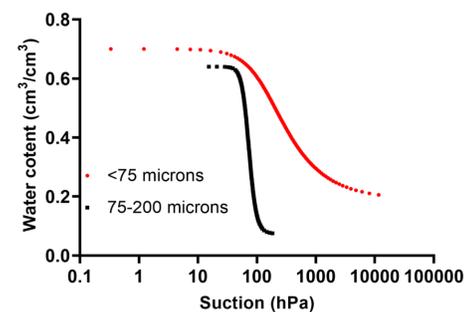


Fig 4: Initial water retention curve for different particle sizes

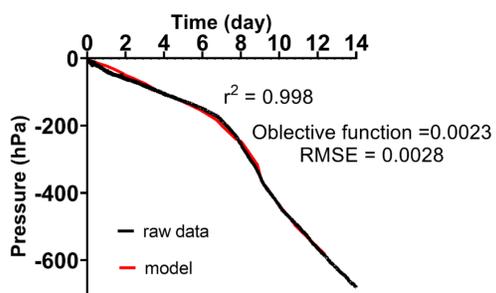


Fig 5: Plot of inversely modelled hydraulic conditions in the column

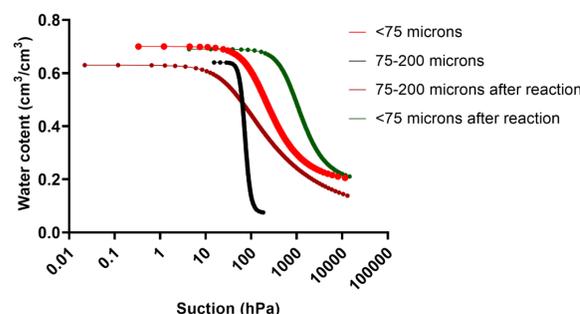


Fig 6: Comparison of initially measured hydraulic parameters and inversely measured parameters

Table 1: Initial parameter and optimised values after changes

| Parameter                              | Initial value |        | Optimised value |        |
|--|---------------|--------|-----------------|--------|
|  | Fine          | Coarse | Fine            | Coarse |
| $K_s$ (cm/day)                         | 1.06          | 128.76 | 1.05            | 127.47 |
| $O_s$ ( $\text{cm}^3 \text{cm}^{-3}$ ) | 0.7           | 0.64   | 0.69            | 0.63   |
| $O_r$ ( $\text{cm}^3 \text{cm}^{-3}$ ) | 0.19          | 0.07   | 0.19            | 0.07   |
| $n$                                    | 1.78          | 7.36   | 1.36            | 2.09   |
| $\alpha$ ( $\text{cm}^{-1}$ )          | 0.007         | 0.01   | 0.0013          | 0.026  |

## 5. CONCLUSION

- The model results show that hydraulic parameters of the studied material changed over time
- The inverse model was effective in establishing these changes
- On the assumption of the saturated volumetric and saturated conductivity decreased by 1% during this period the  $\alpha$  and  $n$  parameters of the coarse and fine material changed with time

## REFERENCES

- NAKHAEI, M., ŠIMŮNEK, J. J. O. H. & HYDROMECHANICS 2014. Parameter estimation of soil hydraulic and thermal property functions for unsaturated porous media using the HYDRUS-2D code. 62, 7-15.
- ROMANO, N. & SANTINI, A. 1999. Determining soil hydraulic functions from evaporation experiments by a parameter estimation approach: Experimental verifications and numerical studies. *Water resources research*, 35, 3343-3359.
- ŠIMŮNEK, J., VAN GENUCHTEN, M. T. & WENDROTH, O. 1998. Parameter estimation analysis of the evaporation method for determining soil hydraulic properties. *Soil Science Society of America Journal*, 62, 894-905.
- TOORMAN, A., WIERENGA, P. & HILLS, R. 1992. Parameter estimation of hydraulic properties from one-step outflow data. *J Water Resources Research*, 28, 3021-3028.