Temporal and spatial earthquake clustering near Athens, Greece, revealed through comparison of millennial strain-rates measured with $^{36}$Cl cosmogenic exposure dating and decadal GPS strain-rate.

Francesco Iezzi (francesco.iezzi.15@ucl.ac.uk), G. Roberts, J. Faure Walker, I. Papanikolaou, A. Ganas, G. Deligiannakis, J. Beck, S. Wolfers, D. Gheorghiu
Introduction

• Clustered earthquake activity causes problems in the identification of the locations of active faults and associated strain-rates. Moreover, earthquake activity has been suggested to swap across strike onto neighboring faults to maintain regional strain-rates (e.g. Cowie et al., 2012). It becomes therefore important to gain observations able to resolve the precise locations of active structures over different time scales.

• GPS observations are commonly used to map regional strain-rates using decadal observations. However, debate is centered on whether GPS results (a) apply over multiple seismic cycles, that is, hundreds to thousands of years, and (b) can resolve the location of active faults if the region between GPS stations contains multiple active faults, but the seismic activity is clustered on specific faults within the GPS stations.

• To answer these questions, we performed $^{36}$Cl cosmogenic dating results on three parallel active faults arranged across strike in the direction of the principal extensional strain in the region of Athens, central Greece.

• The material shown in this presentation is currently under review in Iezzi et al., 2020, in Geology.
Cosmogenic dating of fault planes

Cosmic particles from supernovae are constantly bombarding the Earth’s surface.

Mt. Vettore fault scarp, white free face is the coseismic surface rupture following 2016 Central Italy earthquakes (modified after Iezzi et al., 2018)
The impact of neutrons and muons with Calcium atoms causes production of $^{36}\text{Cl}$ atoms.
Cosmogenic dating of fault planes

The concentrations of $^{36}\text{Cl}$ can be measured with a particle accelerator.
Location of study area

Attica, central Greece

- Extension is accommodated by multiple parallel low slip-rate faults (0.2-0.5 mm/yr; Deligiannakis et al. 2018) with little or no evidence of historical surface-rupturing earthquakes, although historical and instrumental earthquakes of $M \geq 6$ occurred in the area.

- The area presents a set of seven active faults arranged across-strike, crossed by the transect across GPS stations 007A-025A.

- We sampled three out of the seven faults across the GPS transect (other faults did not present sites suitable for cosmogenic dating).
Attica, central Greece

- The three faults present low slip-rates, but they are all capable of releasing earthquakes with $M \geq 6$.

- The faults are in the proximity of Athens (~4 million population), with the southernmost of them (Fili fault) having its SE tip probably underlying the city.

- The definition of the fault slip histories is fundamental for the evaluation of the seismic hazard of Athens.
Sampling of fault planes

Milesi fault

Map view

Samples are collected on a planar fault plane with no post-15 ka erosion or sedimentation, with undisturbed upper and lower slopes, parallel hangingwall and footwall cutoffs.

Section view

Fault scarp profile

Dip upper slope 28°
Dip fault 40°
Dip lower slope 27°
Slip on-fault 653 cm
Colluvial wedge

Metres
\( v = h \)

Sample ladder up the fault plane

Right-step in order to sample preserved higher portion of fault plane

Sample size 4.5 cm x 5 cm
2.5 cm thickness

Section view of trench

Contact between Holocene and LGM deposits

Sample 0, base of Holocene deposits

Planar fault plane, with constant dip along the outcrop (distance between 1-2 is ~1.5 m)
Sampling fault planes

**Malakasa fault**

Samples are collected on a planar fault plane with no post-15 ka erosion or sedimentation, with undisturbed upper and lower slopes, parallel hangingwall and footwall cutoffs.

**Map view**
- Sample site

**Section view**
- Undisturbed upper slope
- Undisturbed lower slope
- Horizontal hangingwall and footwall cut-offs

**Malakasa fault scarp**
- Preserved upper slope
- Eroded upper slope
- Dip upper slope 29°
- Dip fault 43°
- Dip lower slope 22°
- Height of slip (red line) 612 cm
- Throw = 4.15 m
- Colluvial wedge

**Samples collected on planar fault plane with preserved striae indicating minimal erosion**
- Left step to sample upper part of fault plane
- Contact Holocene deposits on LGM sediments
- Sample 0

**Detail of samples and the trench**
- Sample size 15 cm x 5 cm
- 2.5 cm thickness

**Section view of the fault plane**
- Fault plane is planar, with constant dip along the outcrop (clipboard as scale, ~30 cm long).
Fili fault

Samples are collected on a planar fault plane with no post-15 ka erosion or sedimentation, with undisturbed upper and lower slopes, parallel hangingwall and footwall cutoffs.
Modelling of cosmogenic $^{36}$Cl measured on fault planes

Modelling is performed using a Bayesian reversible-jump MCMC approach (Beck et al., 2018). It iterates the slip history many thousands of times, forward modelling expected $^{36}$Cl concentrations each time, to search for the best-fit to the measured $^{36}$Cl.

The model has been allowed to explore slip histories for a large time interval to cope with possible early preservation of fault scarps.
Results

Combined slip histories of the three studied faults

- Episodic behaviour, with non-systematic alternance of periods of rapid slip accumulation and periods of quiescence.

- We interpret the slip pulses as periods during which the faults experienced repeated surface-rupturing earthquakes in a short time interval (i.e. earthquake clusters), alternating with periods of quiescence (i.e. earthquake anticlusters).
Results

Combined slip histories of the three studied faults during the last 15 ka

- Earthquake clusters are alternating on the three parallel faults.
- Cluster on one fault corresponds to periods of quiescence on other faults.
- The clusters on different faults do not overlap in time, with activity migrating rapidly across strike as a cluster terminates.
• The horizontal strain-rates implied by slip during clusters are a large percentage (50-95%) of the regional strain-rate if the GPS rate applies over longer time scales.

• These high percentage values imply that only a small number of the active faults on an across-strike transect contribute to the regional strain-rate at any given time.
Results

Correlation between the age of preserved fault scarps and elevation

The retrieved slip histories highlight a difference in the ages since the three fault scarps started to be preserved: older fault scarps are located at lower elevations, younger fault scarps are located at higher elevations.

Our interpretation is that, during the last glaciation, erosion related to frost-shattering increased with elevation. This explains why the lower elevation sites preserve a longer portion of the scarp history.
• Earthquake clusters migrate both into the hangingwall and footwall of faults spaced at least 5-20 km across strike.

• Clusters begin on faults within a few hundred years of the time when activity ceases on faults across strike.

• Clusters last several millennia.

• Clusters involve 2.0-3.5 meters of slip which is 78-100% of the slip measured over 15 millennia.

• Intense clustering implies a factor of x3.75-4.23 difference between slip-rates and earthquake recurrence intervals within clusters and compared to that calculated since 15 ka.
• The implied strain-rates during measured earthquake clusters do not exceed that implied by decadal measurements with GPS, implying that the latter remains a candidate for the long-term regional deformation rate.

• If several active faults exist between GPS stations, it will not be possible to resolve whether strain is concentrated on one fault or shared between all faults or how strain localization changed through the Holocene.
Conclusions

• Faults near Athens, although characterized by low slip-rates, are active and capable of surface-rupture during earthquakes, and given the uncertainty of whether clusters and anti-clusters are ongoing or about to end, the identification of clustered fault activity is vital for probabilistic seismic hazard assessments for the city.

• The combination of palaeoseismological data, such as $^{36}$Cl dating, and geodetic data is a powerful tool to study continental deformation and seismic hazard because it combines the long and short-term views of the deformation.

• We advocate dense GPS networks with stations on every fault block combined with InSAR observations that provide continuous spatial coverage of strain accumulation if the precise location and width of actively deforming zones is required, alongside palaeoseismology covering many millennia, for example $^{36}$Cl studies, where possible.
Thank you!

Please contact francesco.iezzi.15@ucl.ac.uk or @Francelezzi on Twitter if you want to know more!