

# BOUNDING GLOBAL AEROSOL RADIATIVE FORCING OF CLIMATE CHANGE



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# A new assessment approach

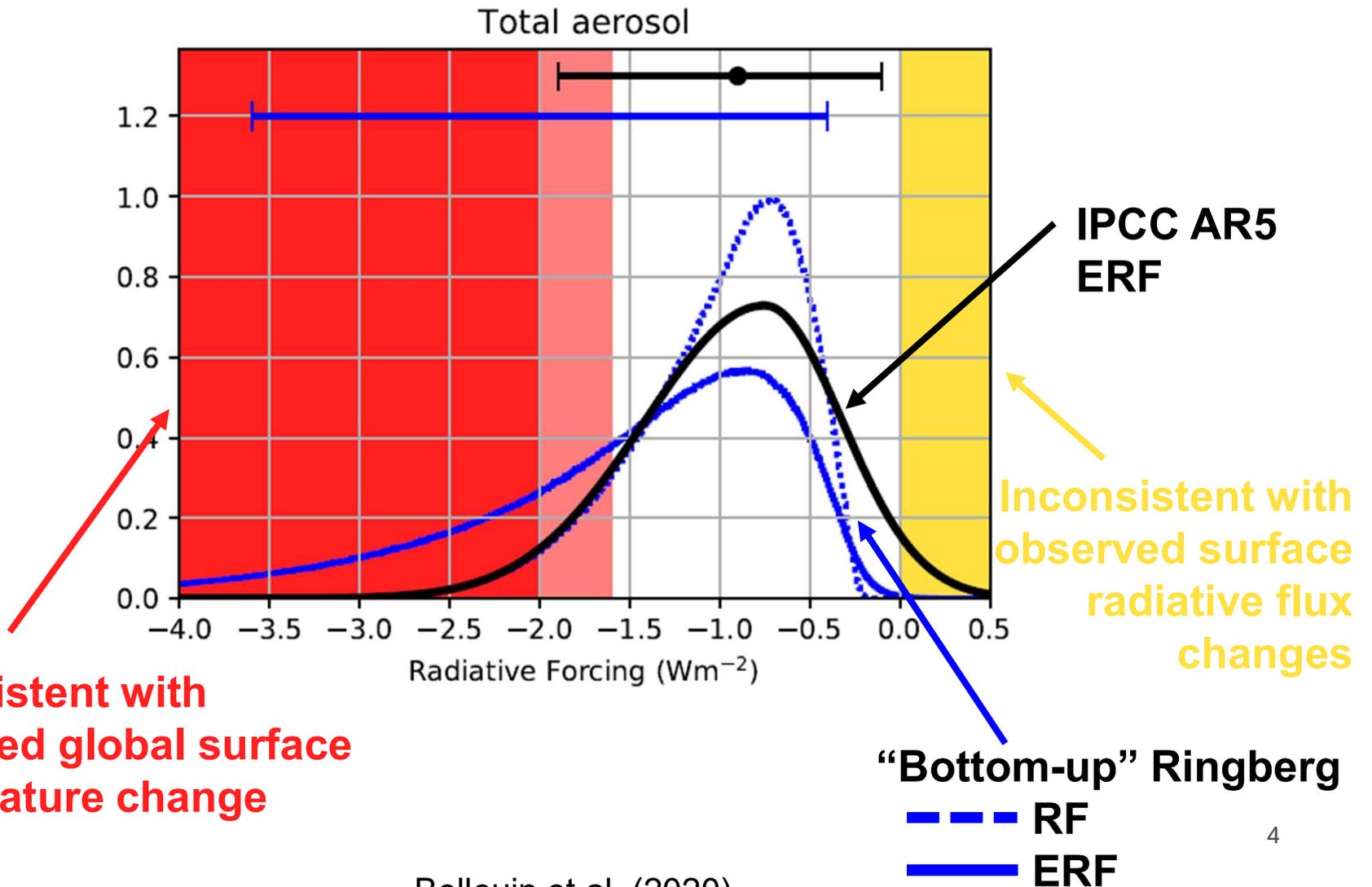
- Base the assessment on multiple, traceable and arguable lines of evidence.
  - In-situ observations, remote sensing from the ground, aircraft or satellite.
  - Cloud resolving models, large eddy simulation.
  - Climate models, including multi-model and perturbed parameter ensembles.
  - Inferences based on observed changes (top-down approaches).
- Clearly identify the questions that are settled and those that remain open.
  - Which aerosol radiative forcing mechanisms are based on more speculative, untested hypotheses?
- Keep the task manageable.
  - Radiative forcing between “average present-day” 2005-2015 and 1850.
  - Global averages only.
  - 68% confidence intervals (IPCC “*likely*”).

# Conceptual model

Split the aerosol radiative forcing problem into three sets of questions:

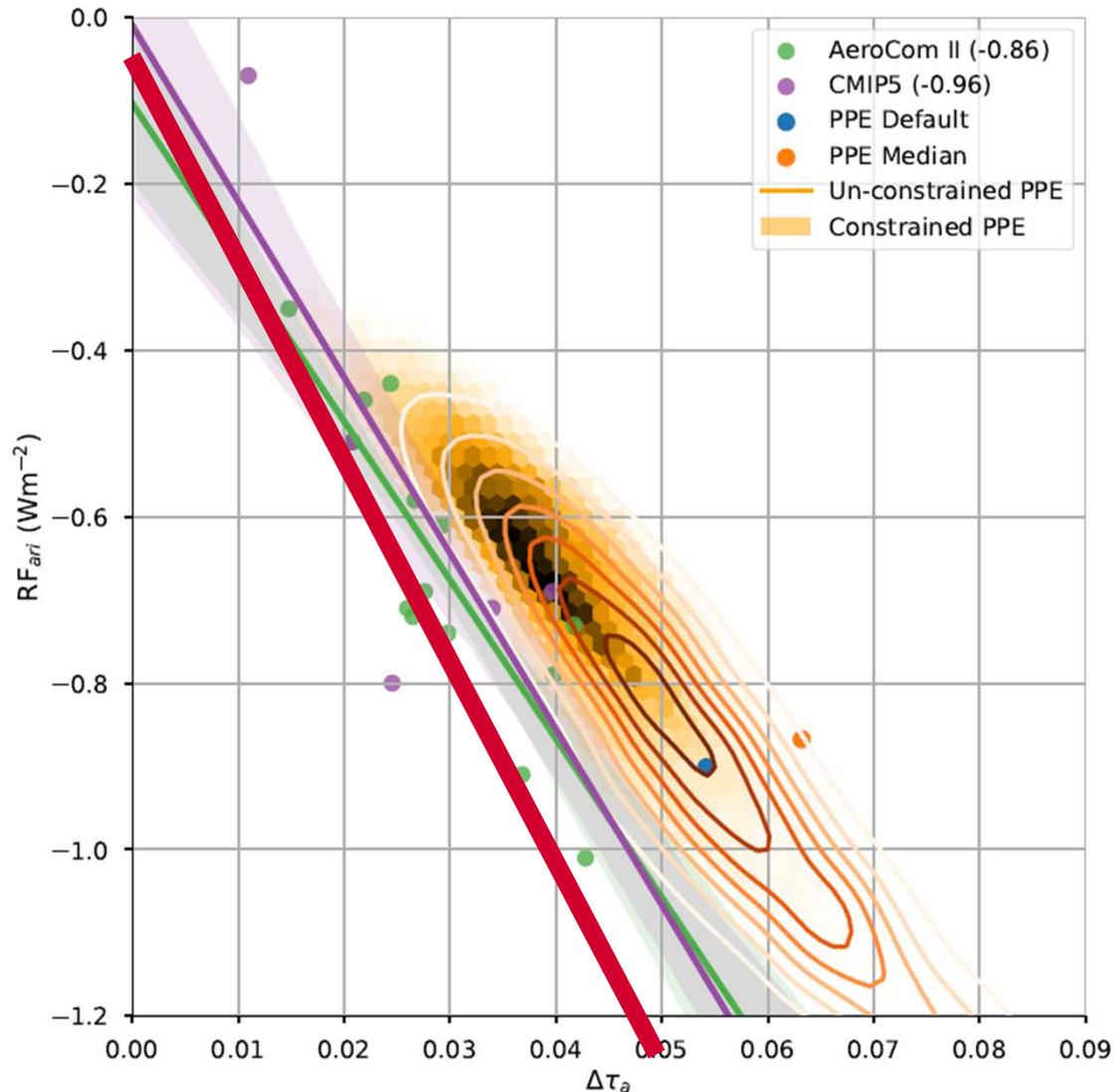
1. What are the large-scale changes in aerosol optical depth ( $\tau_{\text{aerosol}}$ ) and cloud droplet number ( $N_{\text{cloud}}$ ) over the industrial era?
2. What is the sensitivity of top-of-atmosphere radiation to those changes?  
What are the sensitivities of absorption and clouds to those changes?  
What is the sensitivity of top-of-atmosphere radiation to those rapid adjustments?
3. Over what fractions of the globe are the different radiative forcing mechanisms exerted?

# Bounding aerosol ERF



Bellouin et al. (2020)

# The role of absorption



High level of agreement among large-scale models about sensitivities of top-of-atmosphere to changes in aerosol optical depth.

Assuming purely scattering aerosols (thick red line) is only slightly wrong on the global scale.

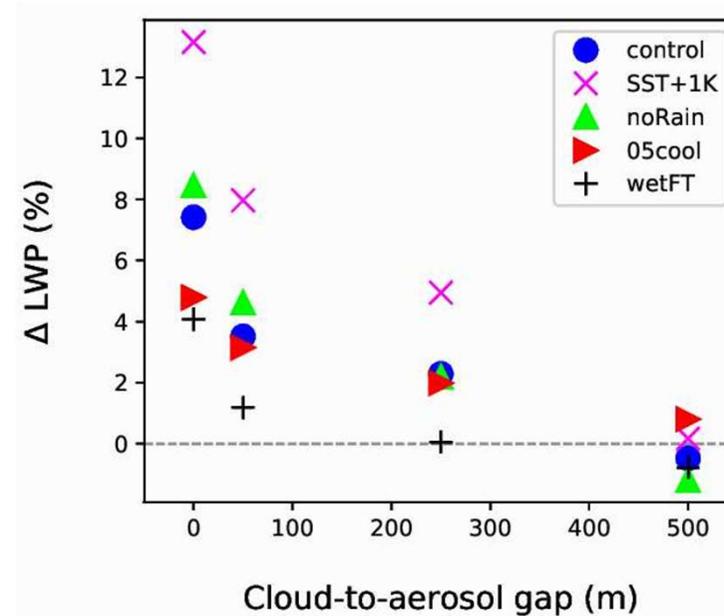
# The role of adjustments to ari

- For black carbon, PDRMIP models estimate a negative rapid adjustment to ari (semi-direct effects): the forcing and adjustment for BC may oppose each other.

Radiative forcing ( $\text{W m}^{-2}$ )	BCx10	SO <sub>2</sub> x5
Instantaneous	+2.42	-3.21
Rapid adjustments	-1.25	-0.32
Effective	+1.17	-3.52

Myhre et al. (2018)

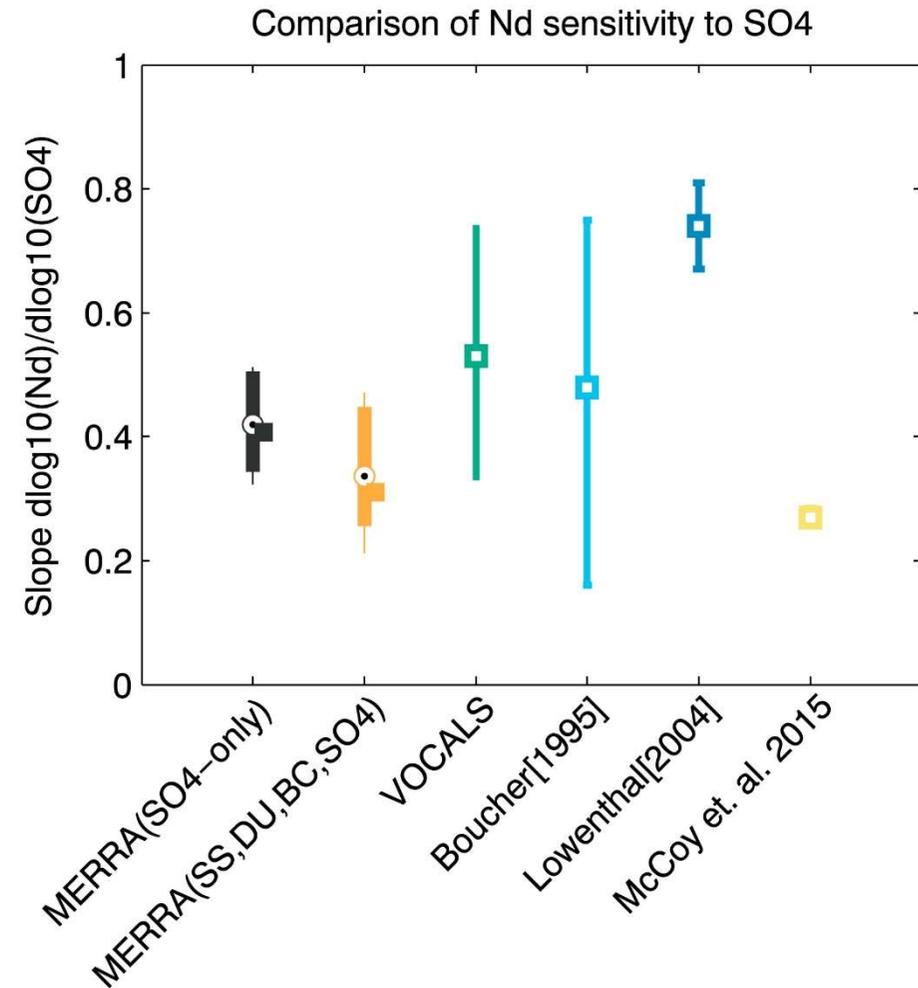
- In addition, it is probably more difficult than previously thought for an overlying absorbing aerosol layer to affect stratocumulus cloud water content.



based on Herbert et al. (2019)

# The role of the Twomey effect

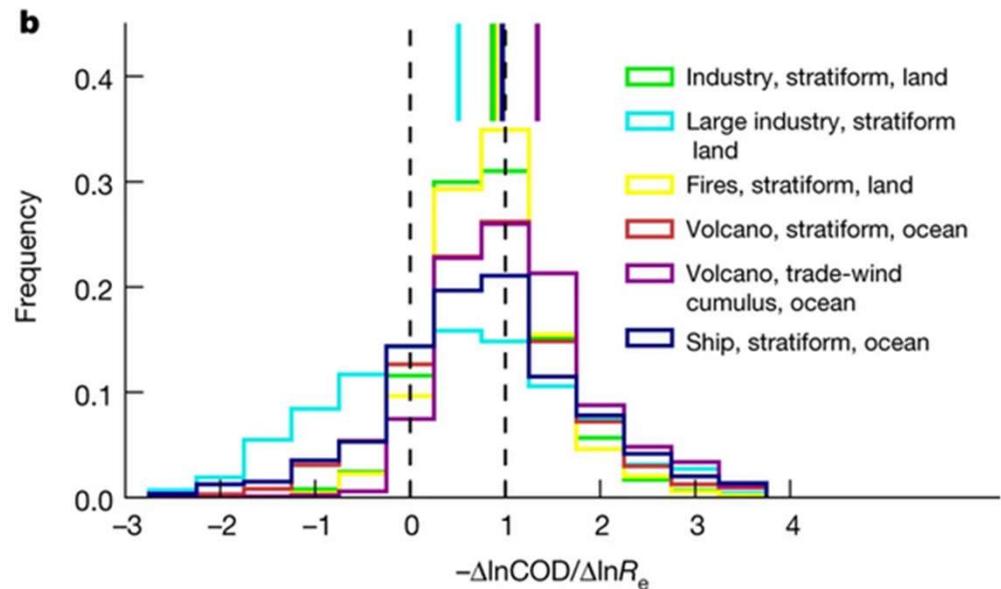
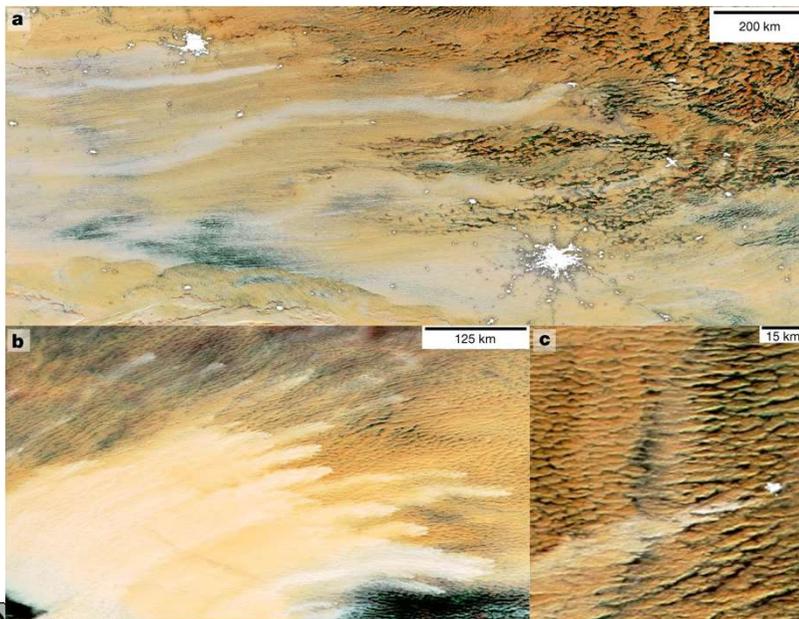
- The theory behind the Twomey effect is supported by many observations.
- But quantifying sensitivities is more challenging, as they depend on cloud regime.
  - Cloud droplet number concentrations are generally much more sensitive to changes in aerosol optical depth in aircraft observations than in satellite studies.



McCoy et al. (2017)

# The role of adjustments to aci (1/2)

- The challenge is to separate the impact of aerosols from natural variability in cloud properties.
- A **large-scale response in cloud liquid water content** is probably weak, based on satellite analyses
  - Chen et al. (2014), Malavelle et al. (2017), Toll et al. (2017, 2019), Rosenfeld et al. (erratum, 2019), Gryspeerd et al. (2019)



Toll et al. (2019)

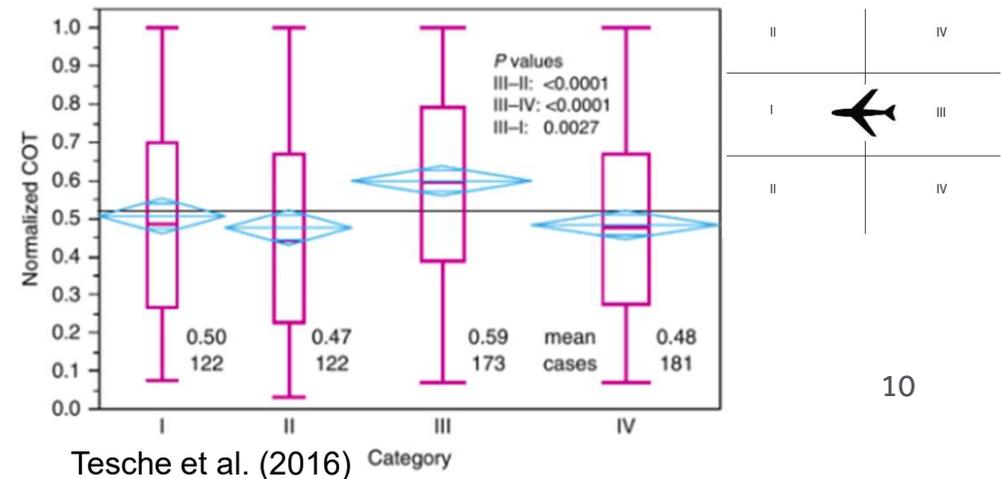
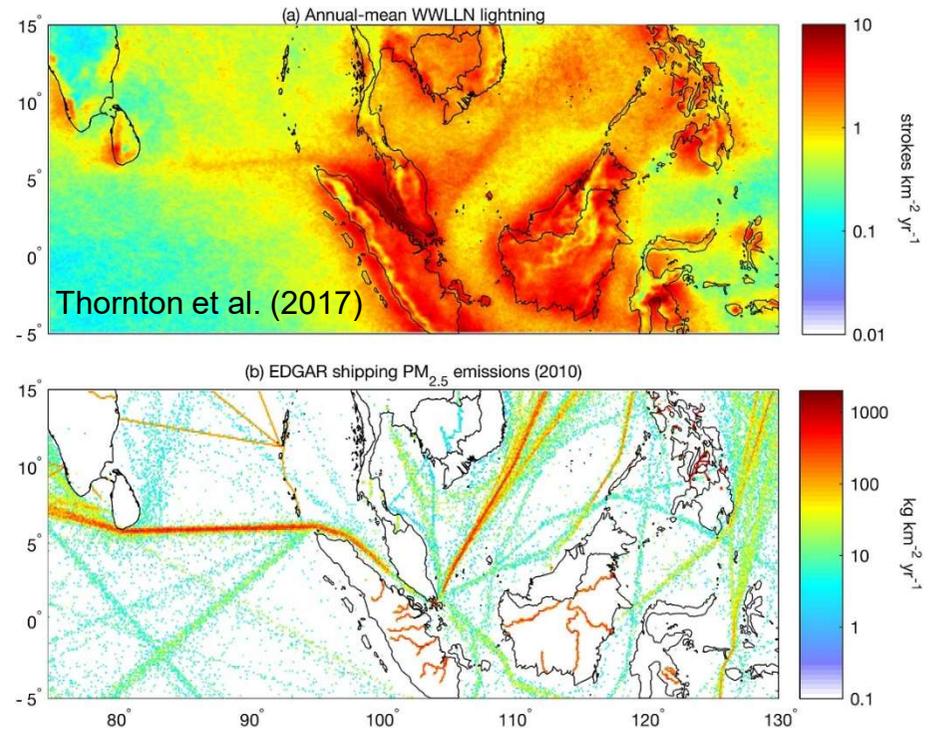
# The role of adjustments to aci (2/2)



- A **large scale response in cloud fraction** is still possible, and could double or triple R<sub>Faci</sub>
  - Gryspeerdt et al. (2016), Christensen et al. (2017), Andersen et al. (2017)
  - But the time dependence of that response is unclear.

# What about ice clouds?

- Many potential mechanisms for aerosols affecting ice clouds, but most are untested.
- RFaci[ice] depends on the balance between homogeneous and heterogeneous nucleation in ice clouds, which is poorly known.
- It now seems likely that black carbon is not a good ice nucleating particle.
- Some evidence to support aerosol impacts on cirrus and convective clouds, but the associated sensitivities are unknown.

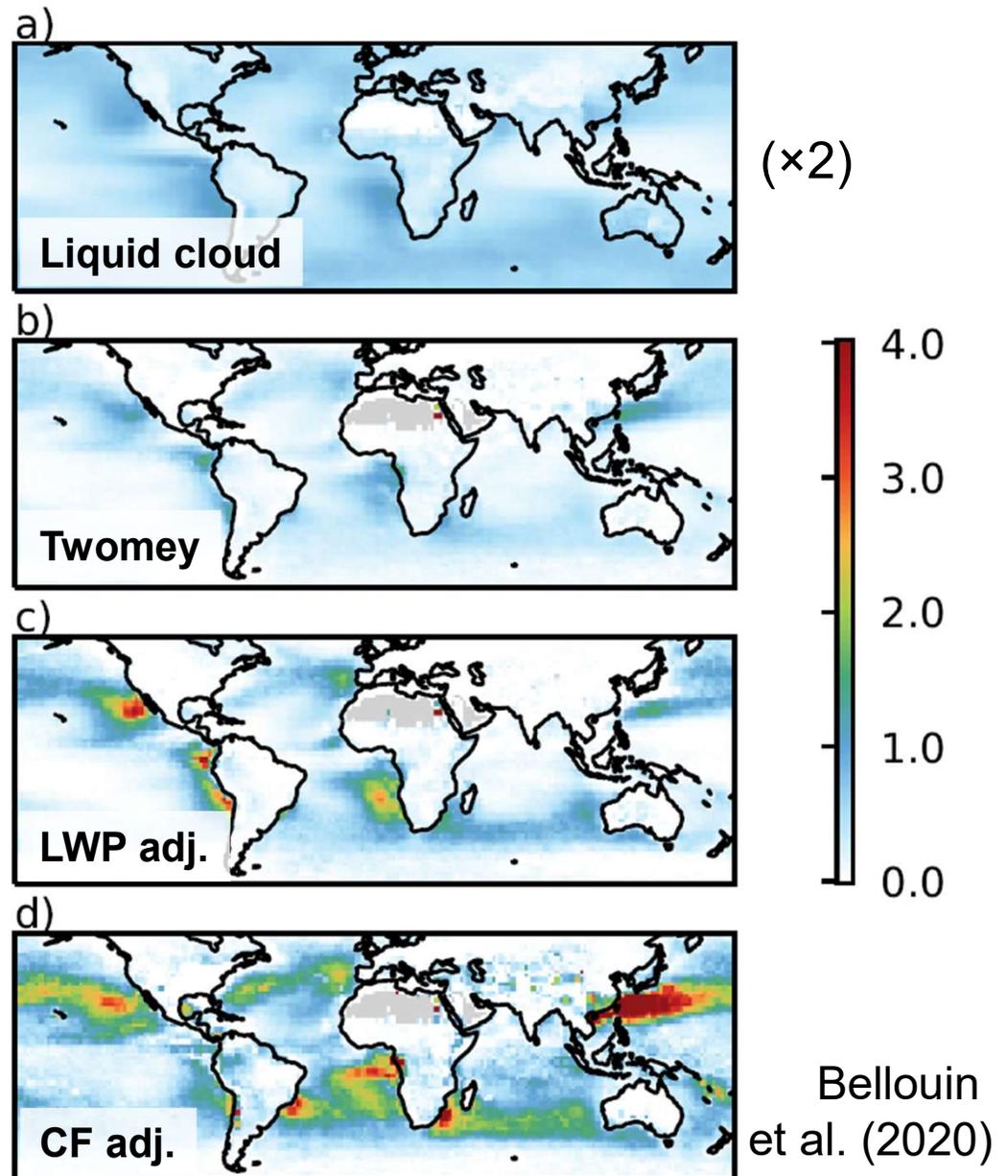


# Effective cloud fractions

A forcing is exerted by:

- A perturbation;
- Sensitivity of radiation to the perturbation;
- Where cloud/moisture can respond, sensitivity of cloud/moisture to the perturbation.

Those requirements can be expressed as fractional areas analogous to cloud fractions.



# Lines of evidence

**Table 4**  
Ranges Obtained by This Review in the 16–84% Confidence Interval for the Variables of Equations (8), (15), and (24)

Section	Variable	Lower bound	Upper bound	Line of evidence
4	$\tau_a^{PD}$	0.13	0.17	Satellite retrievals
4	$\Delta \tau_a$	0.02	0.04	Global modeling
4	$\Delta \ln \tau_a = \Delta \tau_a / \tau_a^{PD}$	0.14	0.29	Modeling/satellite
6	$\Delta \ln N_d = \Delta N_d / N_d$	0.05	0.17	Modeling/satellite
Aerosol-radiation interactions				
5	$S_r^{clear} [W m^{-2} \tau_a^{-1}]$	-27 (0.08)	-20 (0.06)	Global modeling
5	$c_r$	0.59	0.71	Global modeling
5	$S_r^{cloudy} c_r [W m^{-2}]$	-0.1	+0.1	Global modeling
5	<i>RF of ari</i> [ $W m^{-2}$ ]	-0.37	-0.12	
7	$dR/dR_{atm}$	-0.3	-0.1	Global modeling
7	$dR_{atm}/d\tau_a [W m^{-2} \tau_a^{-1}]$	17	35	Global modeling
7	<i>RA of ari</i> [ $W m^{-2}$ ]	-0.25	-0.06	
7	<i>ERF of ari</i> [ $W m^{-2}$ ]	-0.58	-0.23	
Aerosol-cloud interactions				
6	$\beta_{\ln N - \ln \tau}$	0.3	0.8	Modeling/satellite
6	$S_N [W m^{-2}]$	-27 (0.079)	-26 (0.076)	Satellite retrievals
6	$c_N$	0.19	0.29	Modeling/satellite
6	<i>RF of aci</i> [ $W m^{-2}$ ]	-1.10	-0.33	
8	$\beta_{\ln c - \ln N}$	-0.36	-0.011	Satellite analyses
8	$S_{C,N} [W m^{-2}]$	-54	-56	Mixed
8	$c_C$	0.21	0.29	Mixed
8	<i>RA of aci (liquid water path)</i> [ $W m^{-2}$ ]	0.01	+0.56	
8	$\beta_{C - \ln N}$	0	0.1	Global modeling, LES
8	$S_{C,N} [W m^{-2}]$	-91	-153	Satellite analysis
8	$c_C$	0.59	1.07	Mixed
8	<i>RA of aci (cloud fraction)</i> [ $W m^{-2}$ ]	-1.14	0.0	
8	<i>ERF of aci</i> [ $W m^{-2}$ ]	-1.73	-0.27	
11	<b>Total aerosol ERF</b> [ $W m^{-2}$ ]	<b>-2.19</b>	<b>-0.61</b>	
11	(constrained by observational inferences)	-1.60	-0.61	

Bellouin et al. (2020) relies on global modelling for

- industrial-era aerosol changes
- ari

and global satellite studies for

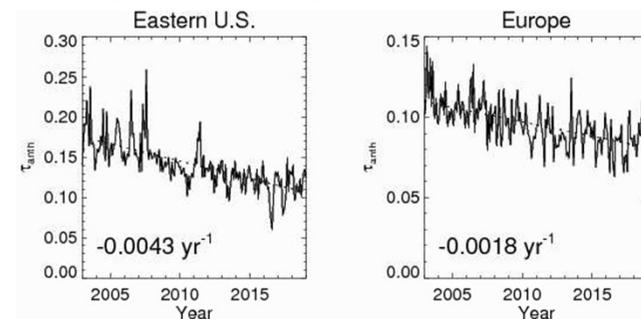
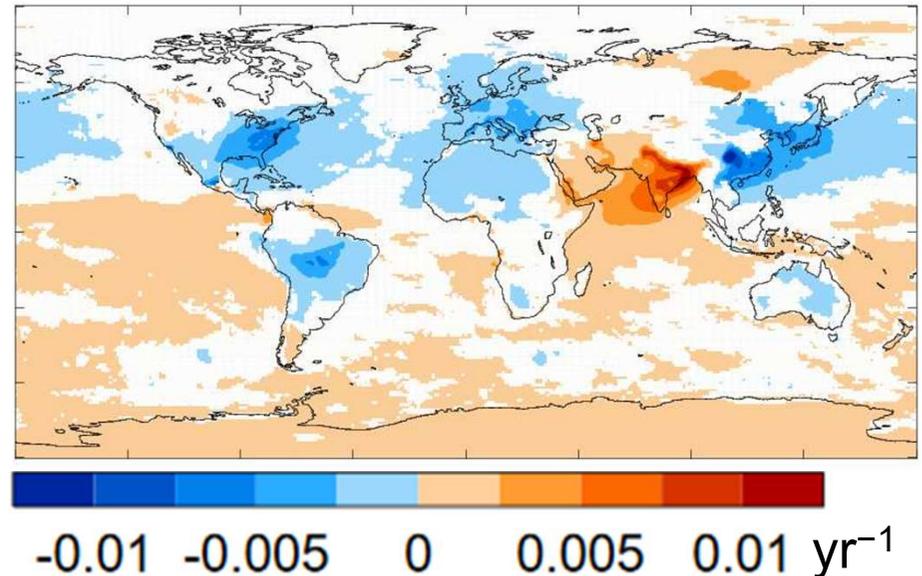
- aci.

Stronger lines of evidence like in-situ measurements and process modelling ended up not being used because nobody knows how to properly weigh their results into a global average.

# Promising avenues

- Scale effects are increasingly being considered in model development. Global large eddy simulation is now becoming possible.
- Cloud responses to regional aerosol trends, and volcanic eruptions and ship tracks may provide insights into cloud regime shifts and ice cloud responses.
- Observational inferences are promising but their uncertainties need to be better understood. Large regional trends (right) may provide strong constraints.
- Models of all scales involve a large number of poorly known parameters, and statistical methods to explore model uncertainties are being adopted.

Deseasonalised linear trends in anthropogenic aerosol optical depth 2003—2019, based on CAMS Reanalysis



# Thank you for your attention!

Bellouin, N., Quaas, J., Gryspeerdt, E., Kinne, S., Stier, P., Watson-Parris, D., et al. (2020). Bounding global aerosol radiative forcing of climate change. *Reviews of Geophysics*, 58, e2019RG000660. <https://doi.org/10.1029/2019RG000660>