

Ocean-only FAFMIP: Understanding Regional Patterns of Ocean Heat Content and Dynamic Sea Level Change

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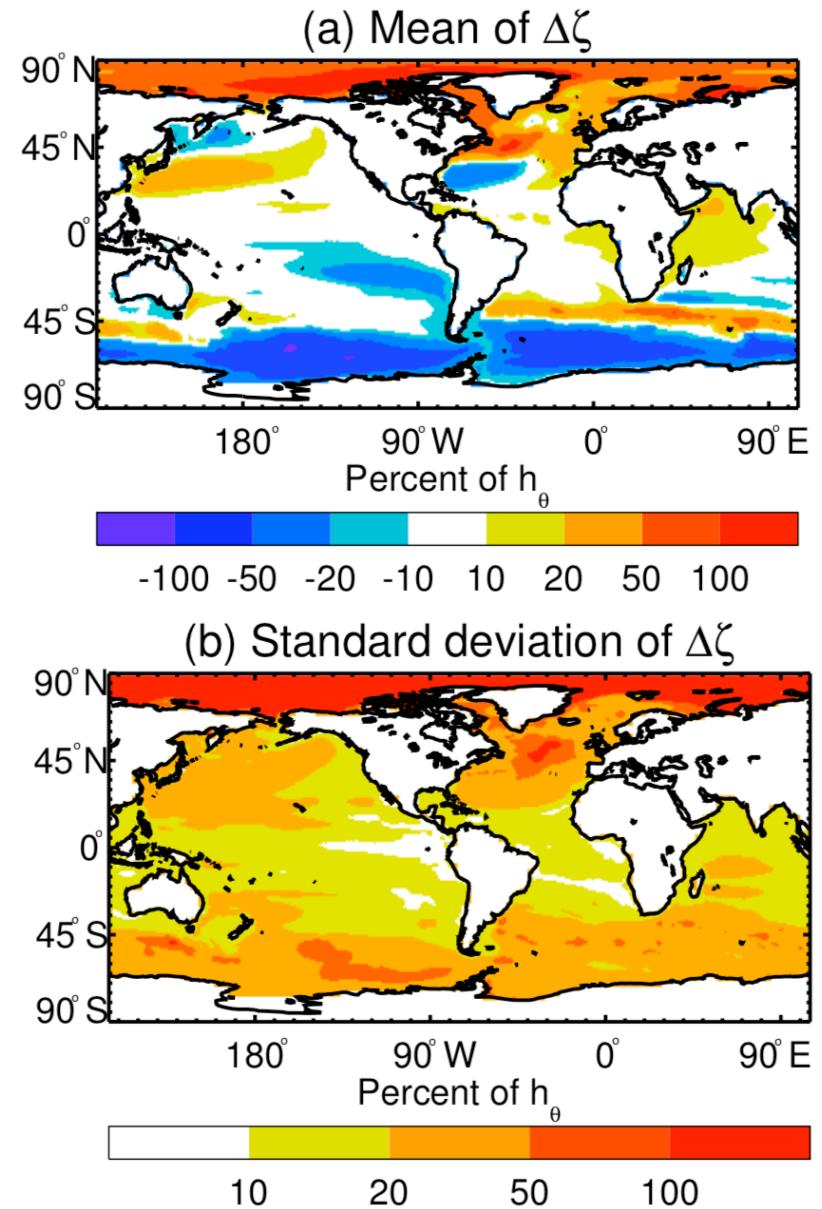
CL4.5 Live Chat - 4 May 2020 08:30-10:15 CEST: <https://meetingorganizer.copernicus.org/EGU2020/session/36739>

Introduction

Coupled (AOGCM) CMIP5 projections demonstrate a large uncertainty in dynamic sea level (DSL, ζ) and ocean heat content (OHC) change.

Uncertainty arises from the use of different ocean models and uncertainty in air-sea buoyancy and momentum flux changes.

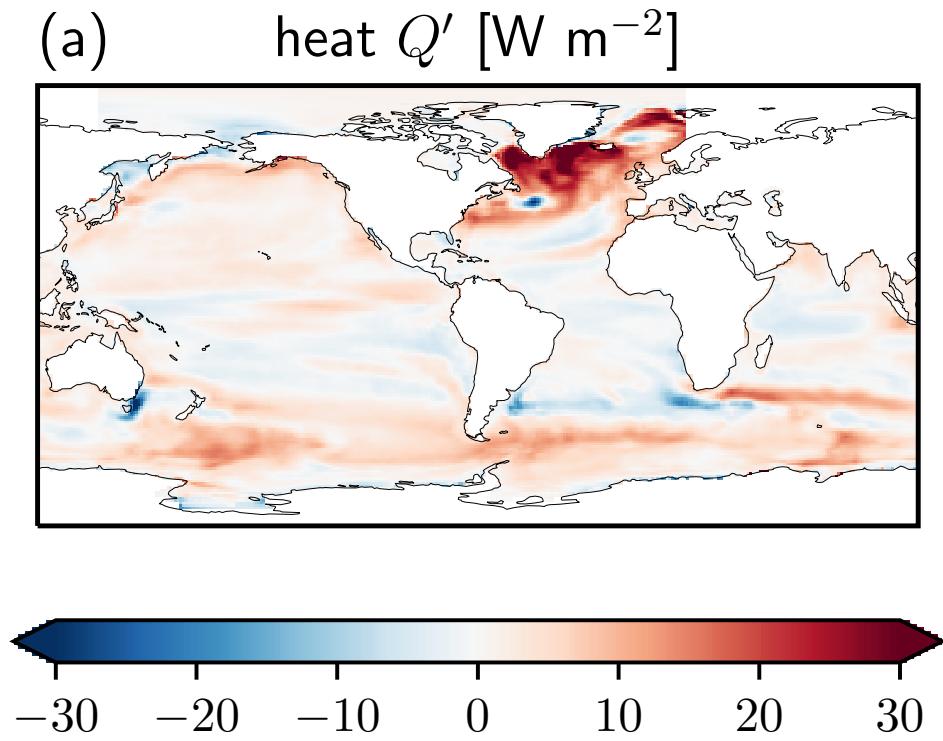
1. How much of the spread in DSL and OHC change is due to the use of different ocean models?
2. What is the effect of atmosphere feedbacks on ocean climate change?



(Gregory et al., 2016)

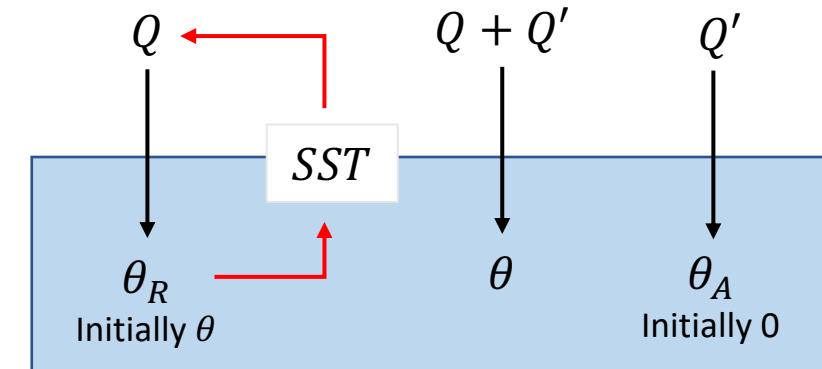
Surface Heat Flux Experiments

Perturbation: CMIP5 ensemble (13 member) mean 1pctCO₂ yr 61-80 mean minus piControl mean 12 month climatology.



1) faf-heat (AOGCM, Method B)

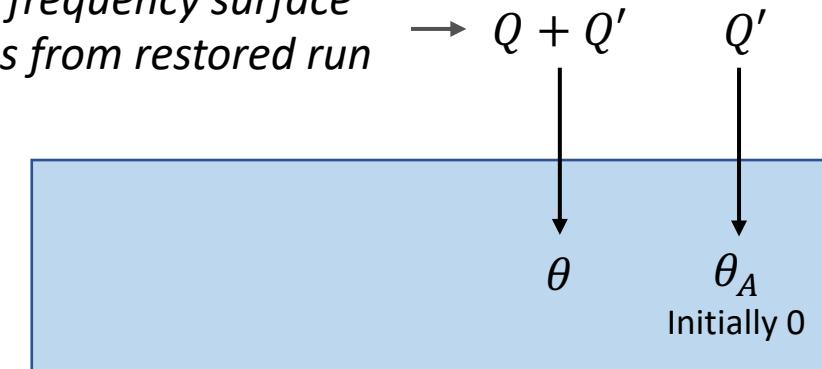
2*AOGCMs:
CanESM5
HadCM3



(Banks & Gregory, 2006;
Bouttes & Gregory, 2014;
Gregory et al., 2016)

2) faf-heat-o (OGCM, Blaker Method)

5*OGCMs:
NEMO3.4
HadOM3
GFDL-MOM5
ACCESS-OM2
MITgcm

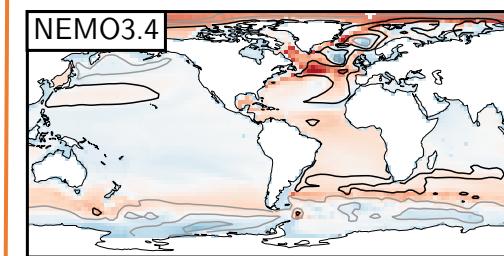


(Marshall et al., 2015;
Zika et al., 2018)

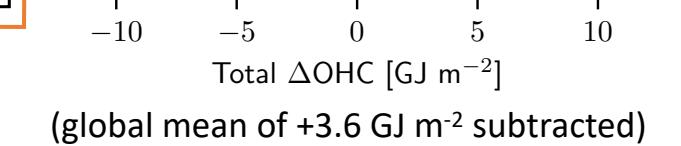
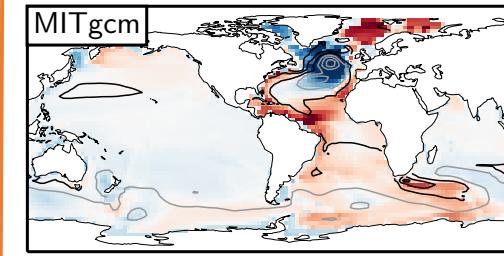
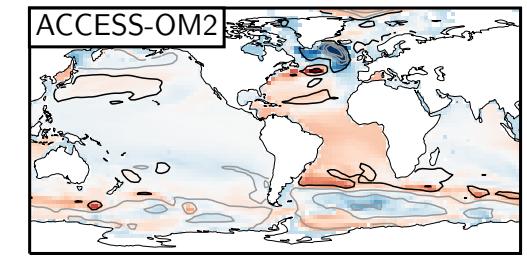
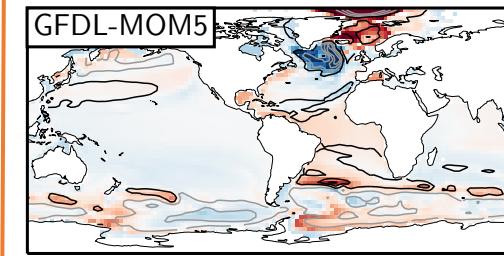
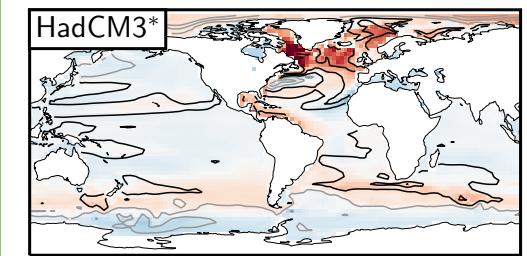
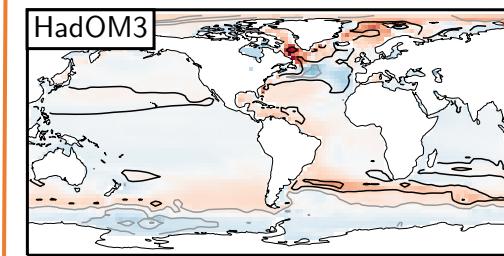
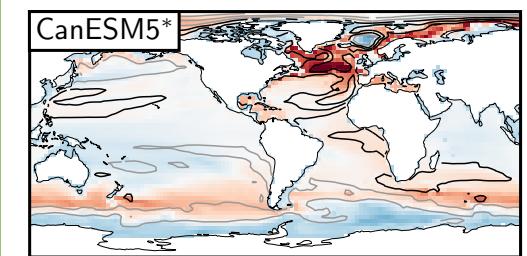
Ocean Heat Content Change

- Large disagreement over warming/cooling of mid to high latitude North Atlantic.
- Greater agreement for Southern Ocean meridional pattern.
- AOGCMs indicate more warming in North Atlantic relative to OGCMs.
 - AO-feedback amplifies warming.

faf-heat-o

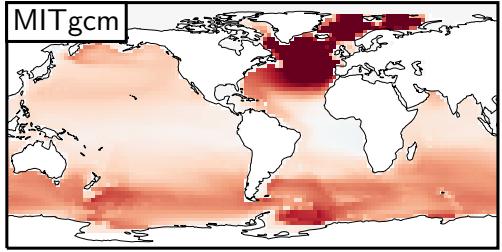
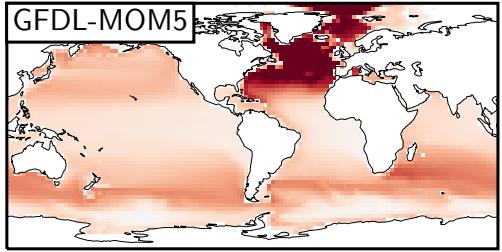
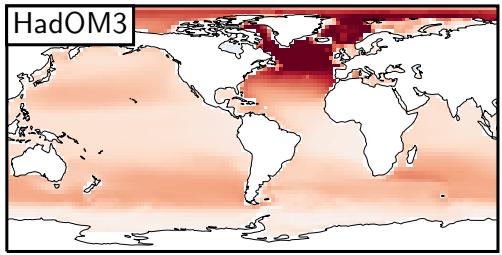
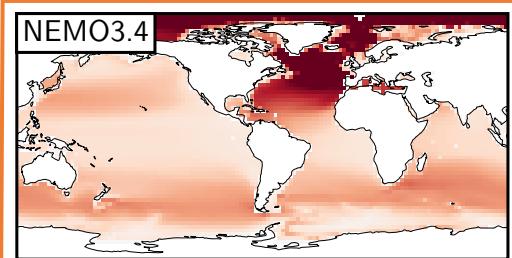


faf-heat

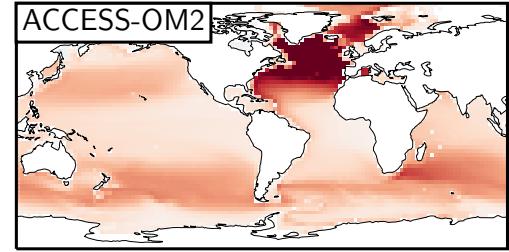
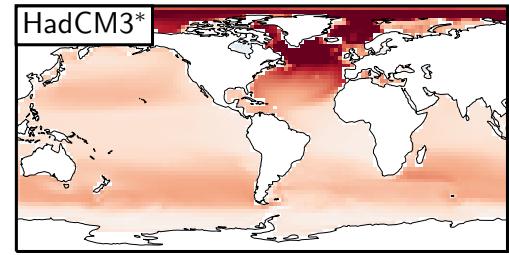
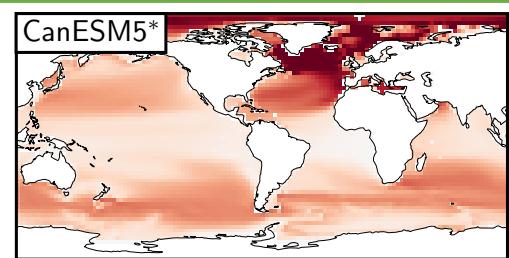


Added vs. Redisistributed Ocean Heat Content Change: $\theta = \theta_R + \theta_A$

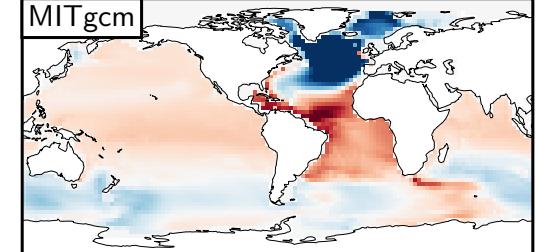
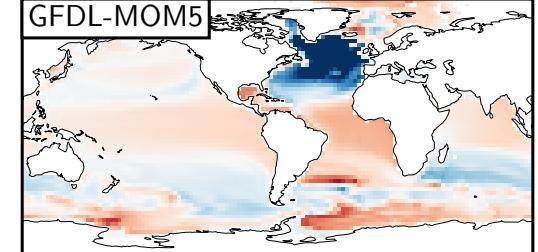
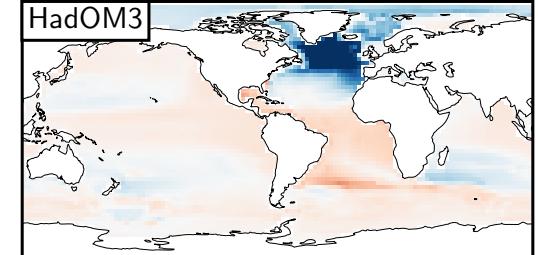
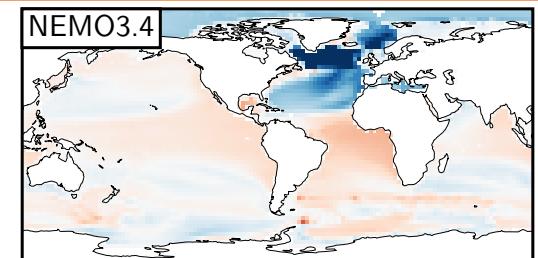
faf-heat-o



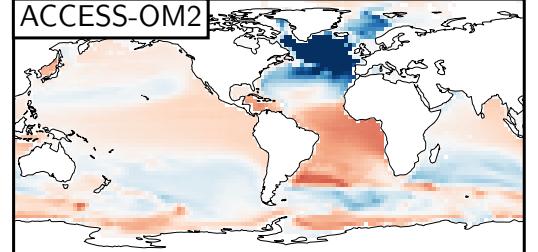
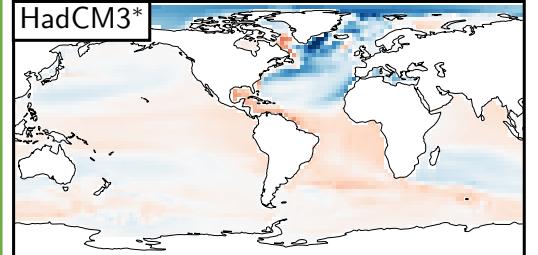
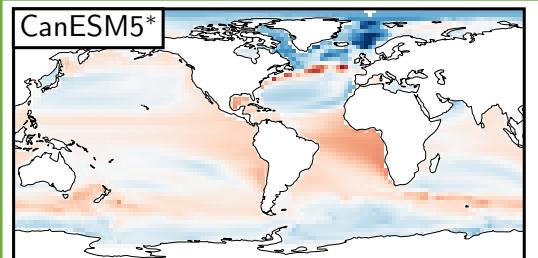
faf-heat



faf-heat-o



faf-heat



Added ΔOHC [GJ m $^{-2}$]

Redisistributed ΔOHC [GJ m $^{-2}$]

Thermosteric Sea Level Change

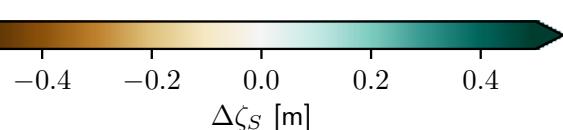
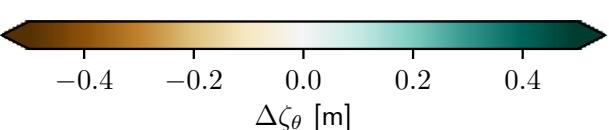
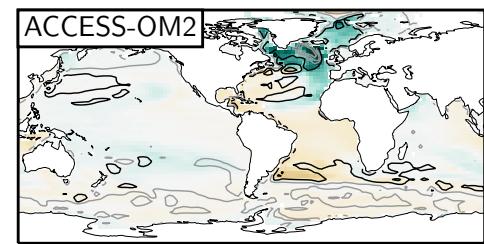
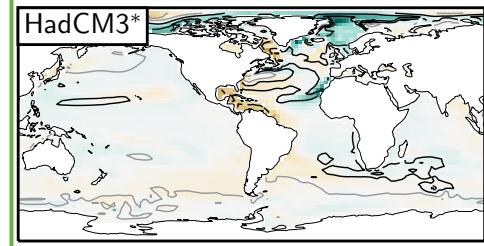
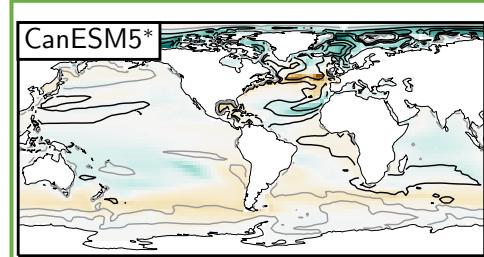
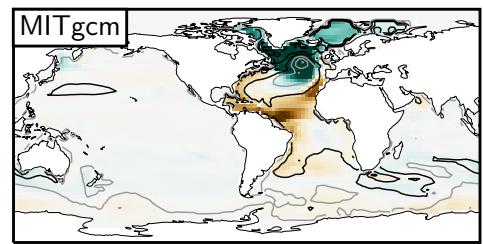
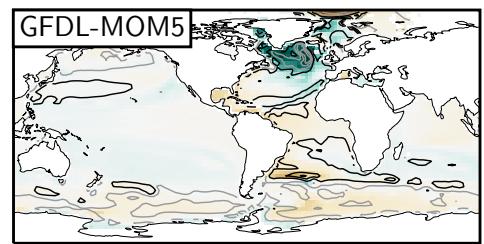
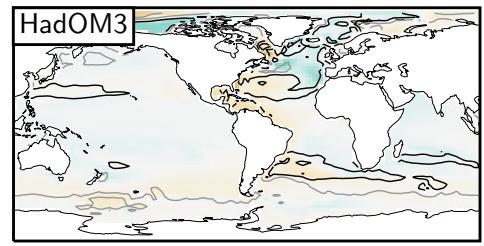
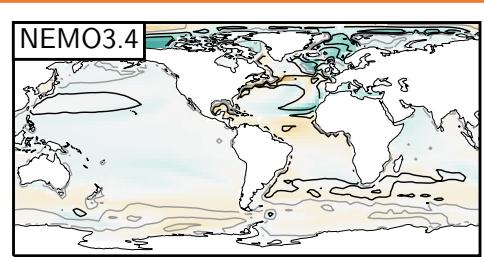
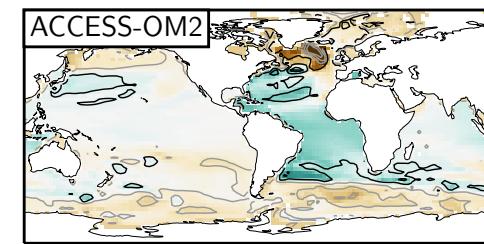
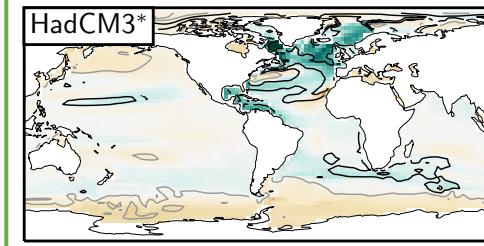
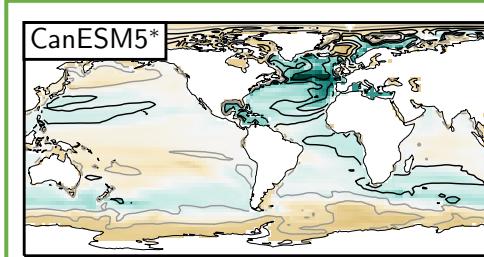
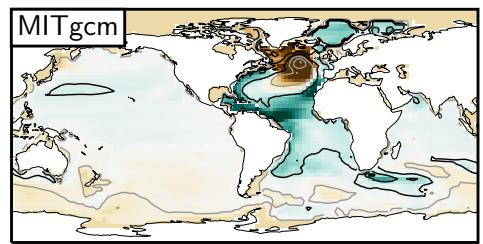
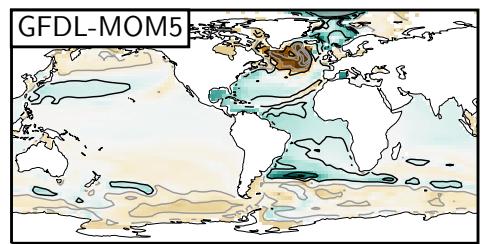
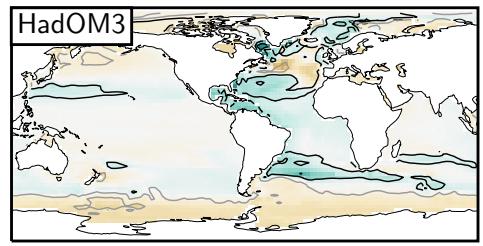
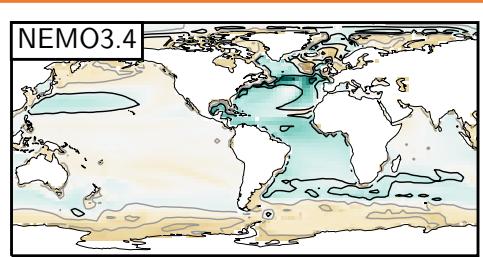
Halosteric Sea Level Change

faf-heat-o

faf-heat

faf-heat-o

faf-heat

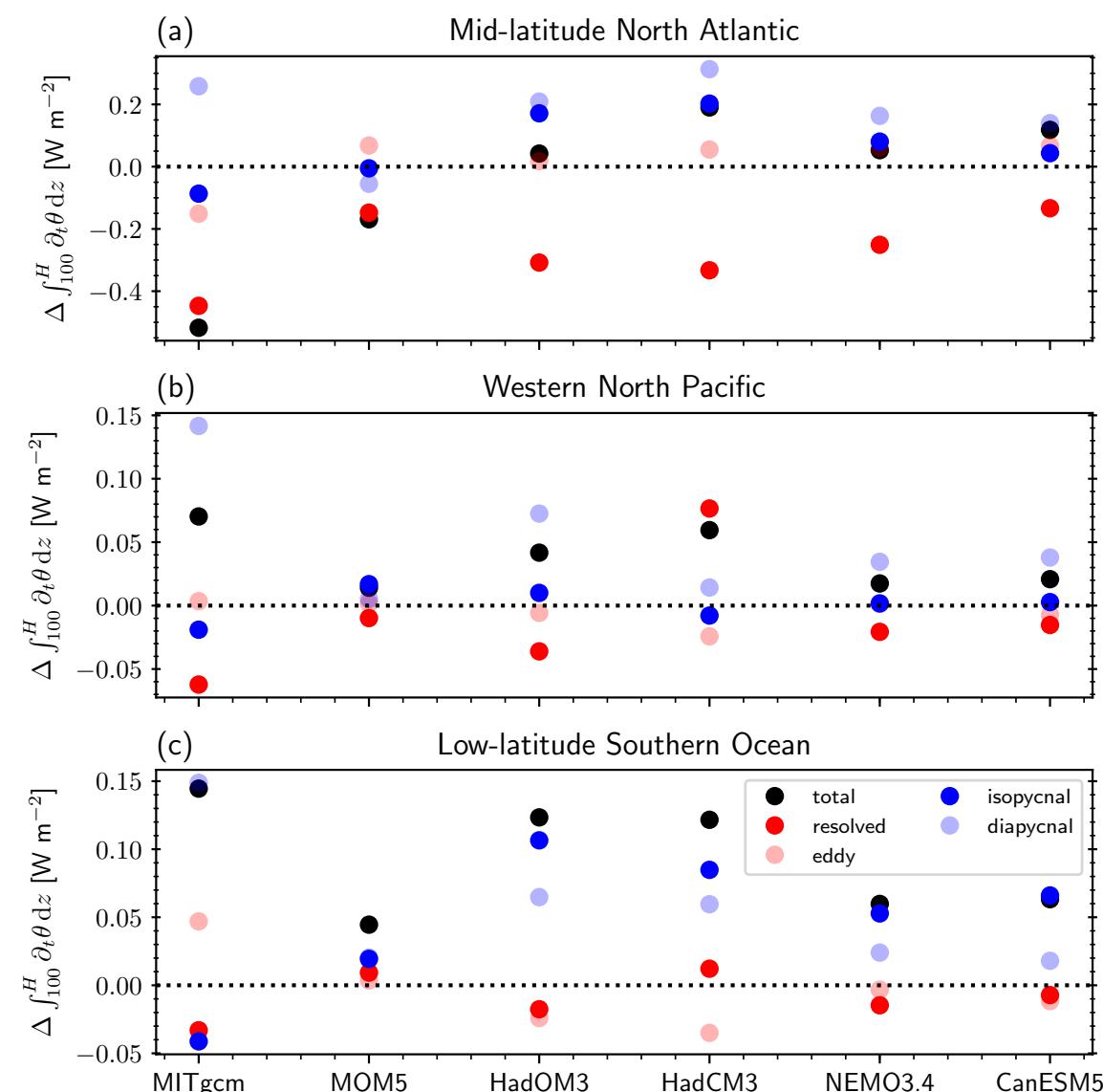
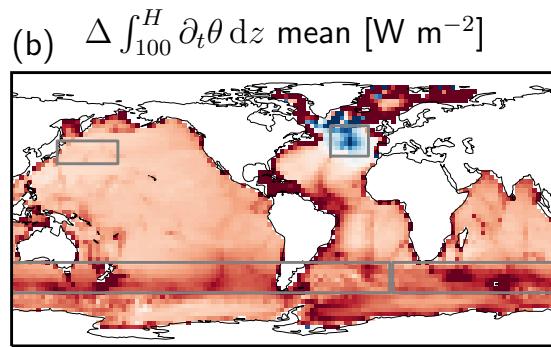


ΔOHC processes

Consider the heat budget:

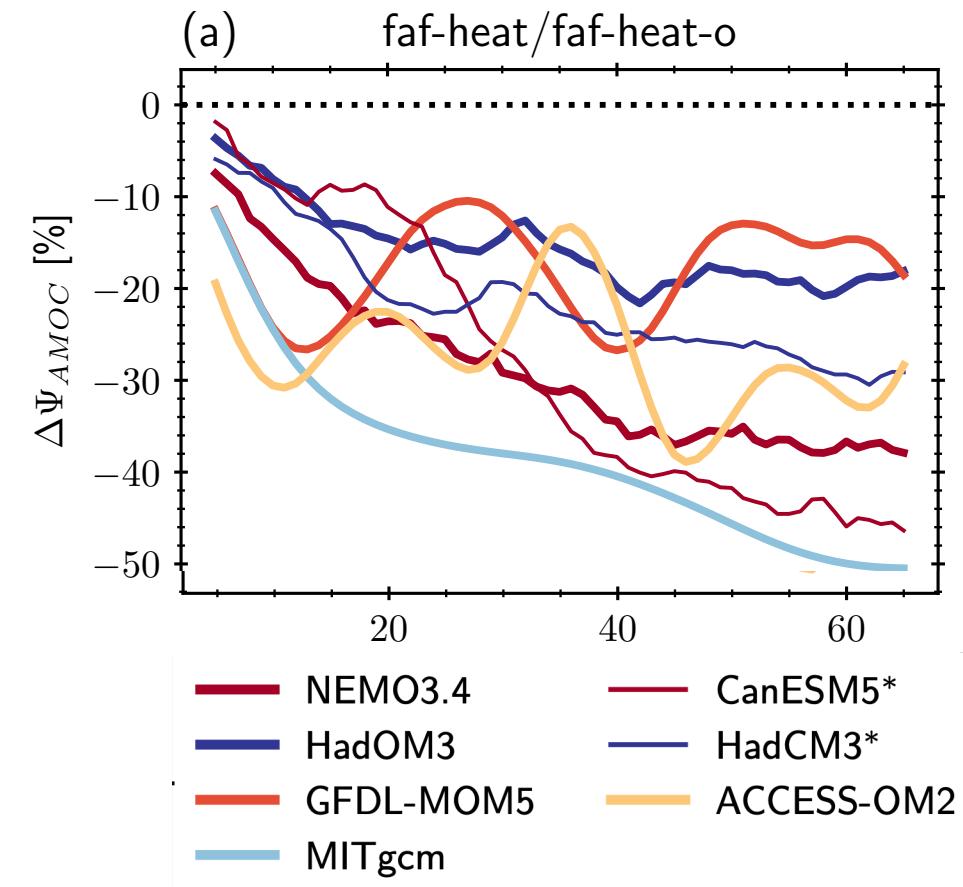
$$\partial_t \theta = \partial_t \theta_{resolved} + \partial_t \theta_{eddy} + \partial_t \theta_{isopycnal} + \partial_t \theta_{diapycnal}$$

- Mainly cooling due to residual circulation ($\partial_t \theta_{resolved} + \partial_t \theta_{eddy}$) change in North Atlantic
- Contrasting processes in HadOM3/HadCM3 in Pacific.
- Mainly diffusive processes in Southern Ocean.



AMOC Change

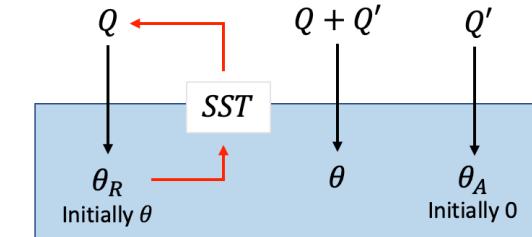
- Large spread (20-50% weakening) in relative AMOC change after 70 years.
- AMOC weakening amplified by ~10% in AOGCMs relative to OGCMs.
 - AO-feedback increases AMOC weakening
- MITgcm is an outlier: 21 Sv control and -11 Sv change.



North Atlantic Atmosphere-Ocean System

(Banks & Gregory, 2006;
Bouttes & Gregory, 2014;
Gregory et al., 2016)

1) faf-heat (AOGCM, Method B)



faf-heat (AOGCM)

vertical air-sea gradient
 $\Delta(T_{air} - \theta_R) > 0$

perturbation
 $Q' > 0$
 surface heat flux
 $Q \uparrow$
 passive heat uptake
 $\Delta\theta_A > 0$

High latitude
 North Atlantic
 $\Delta\theta < 0$

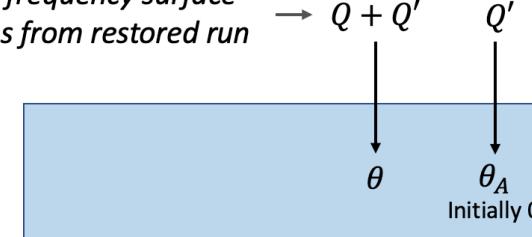
meridional
 $\Delta\rho \downarrow \rightarrow \Psi_{AMOC} \downarrow \rightarrow \Delta\theta_R < 0$
 $\Delta\rho \uparrow \rightarrow \Psi_{AMOC} \uparrow \rightarrow \Delta\theta_R > 0$

faf-heat-o (OGCM)

2) faf-heat-o (OGCM, Blaker Method)

(Marshall et al., 2015;
Zika et al., 2018)

High frequency surface
 fluxes from restored run



Summary

- 1) Large spread in ΔOHC and $\Delta \zeta$ from prescribing the same surface heat flux perturbation to different OGCMs.
 - Structural and background state differences alone contribute a large spread in ocean sensitivity to boundary condition changes.
 - Little agreement on which processes dominate ΔOHC on regional scales.
 - Large spread in both thermosteric and halosteric contributions to $\Delta \zeta$, consistent with Couldrey et al. (in rev.)
- 2) FAFMIP Method B introduces an AO-feedback which amplifies $\Delta \Psi_{AMOC}$ weakening by $\sim 10\%$
 - Mitigated by the negative AMOC- θ -surface heat flux feedback

