



# Changing our ideas about the evolution of magmatic systems with improved temporal resolution: do we get it right?

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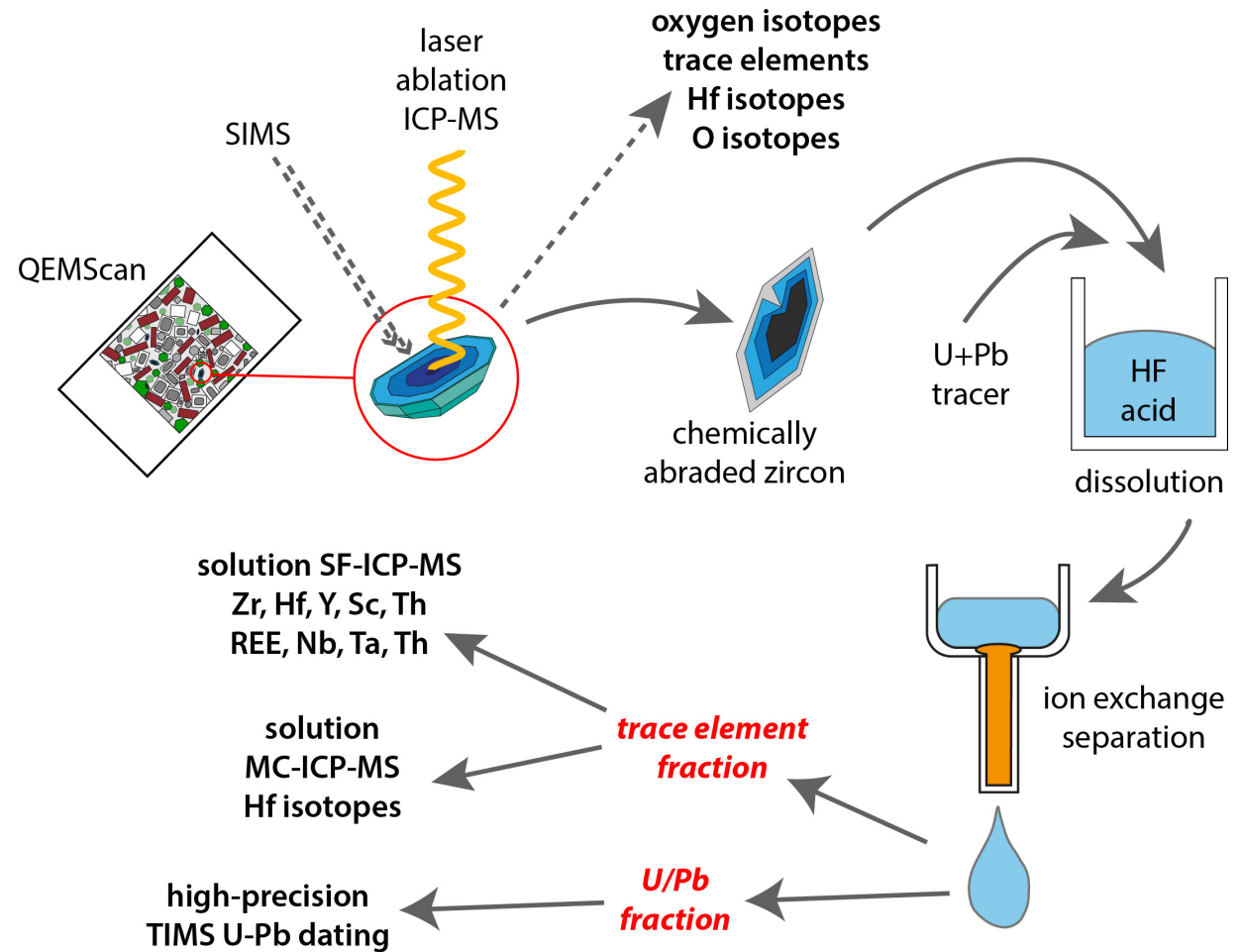


# Zircon petrochronology

## CA-ID-TIMS U/Pb dating

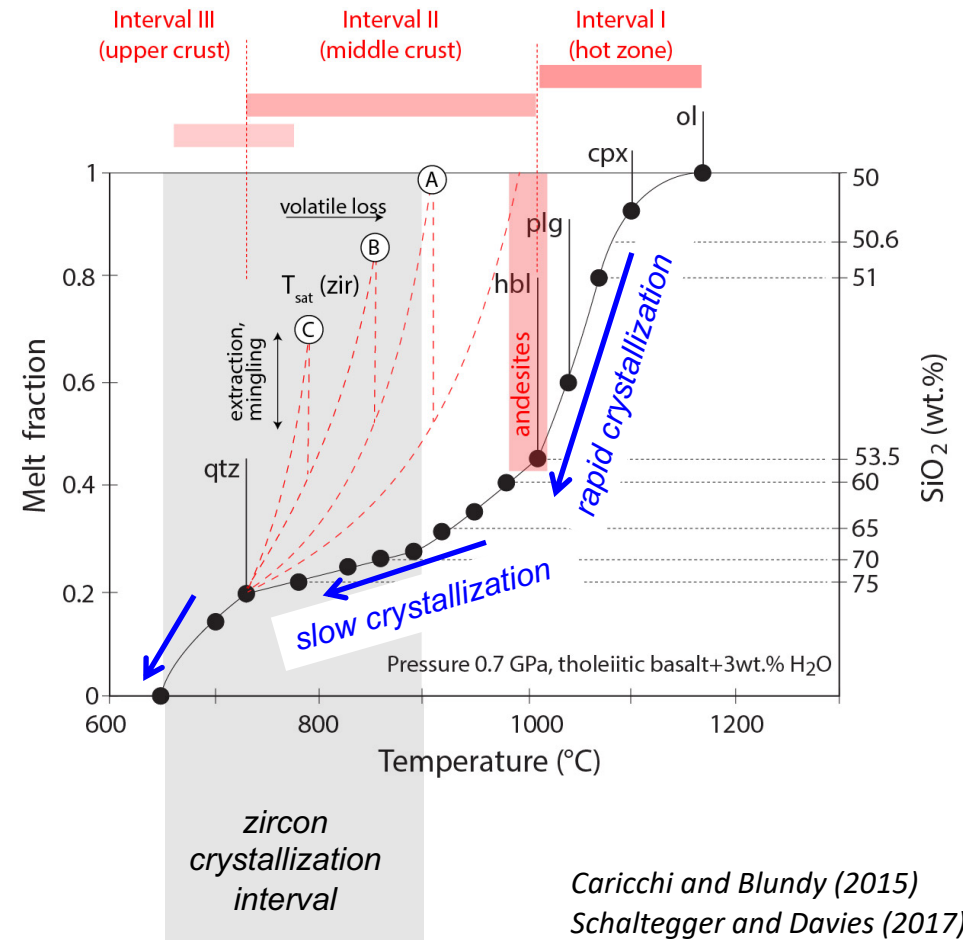
### Workflow:

- Imaging
- In-situ chemical and isotopic analysis
- Chemical abrasion
- High-precision dating
- Solution chemical and isotopic analysis



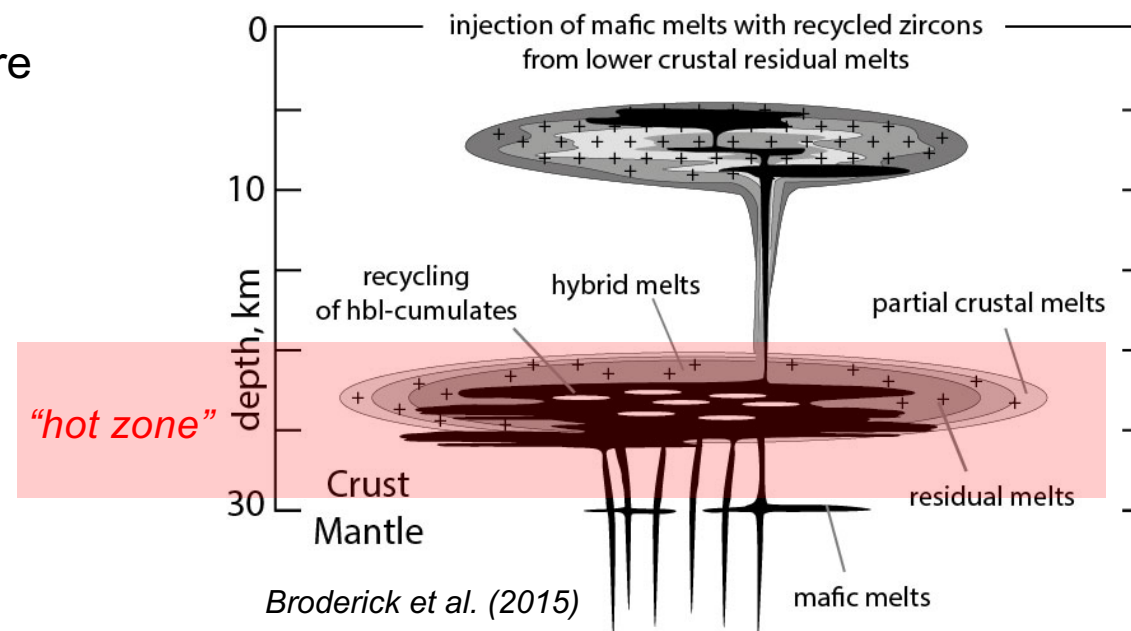
## Residence in magma and recycling processes in magmatic systems

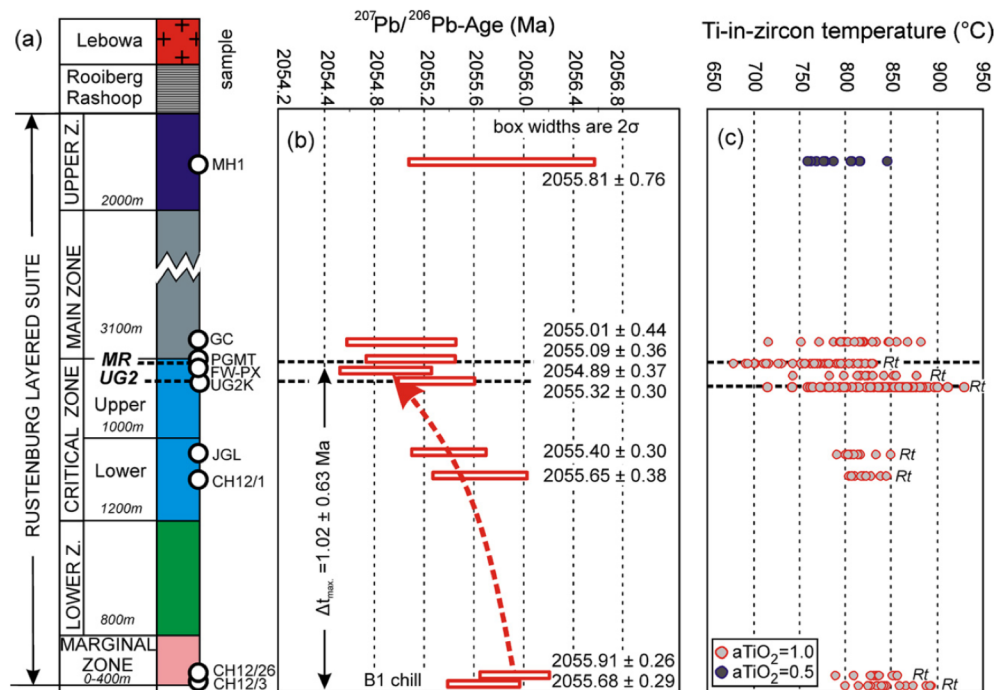
- Arrival at  $T_{\text{sat}}[\text{zirc}]$  as  $f_{(\text{Na, K, Ca, Al, Si, Zr; T})}$
- Rapid crystallization at 50-65%  $\text{SiO}_2$  (few zircons in marginal portions of a pluton)
- Slow crystallization at 65-75%  $\text{SiO}_2$  (many zircons forming between  $T_{\text{sat}}$  and solidus)
- Rapid crystallization at the solidus (few zircons)
- Sigmoidal age distribution – which grains to select?
- Rejuvenation (partial remelting) of crystal mushes (“cannibalization”) during recharge events
- Open system: mixing of liquids, mixing of mushes, transfer of crystals



## A conceptual model for arc systems

- Much (most?) of the zircon grows at a lower to middle crustal level
- Magma storage is “cold”: crystal mushes, which episodically get rejuvenated during recharge events
- Mixing of melt batches from different origin and their crystals
- Incremental accretion in the upper crust
- relative temporal relationships in the field are younger than zircon crystallization



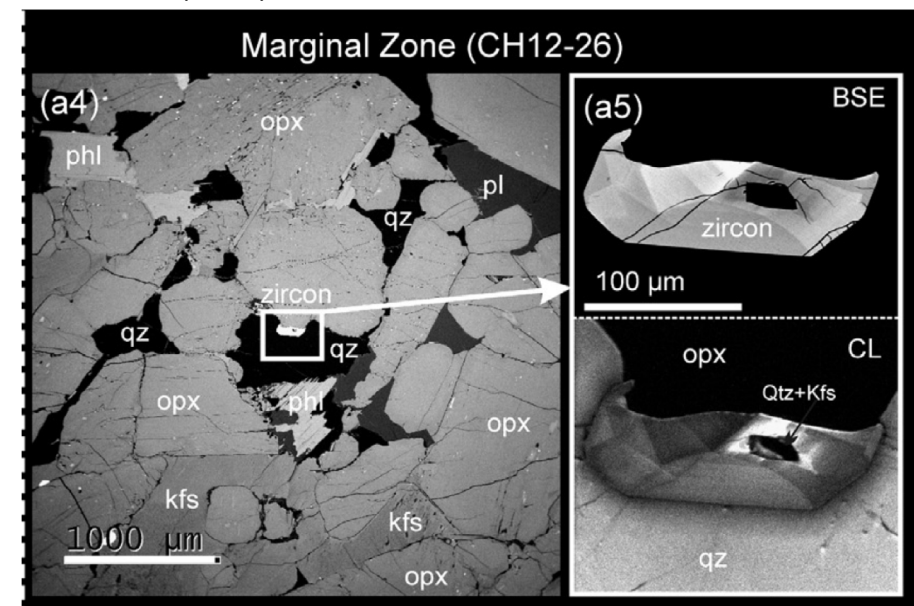


## Zircon growth in mafic magmatic systems

Zircon is surprisingly common in tholeiitic MOR gabbros and plagiogranites, sills and dykes of Large Igneous Provinces

Example: interstitial zircon in Bushveld magmas

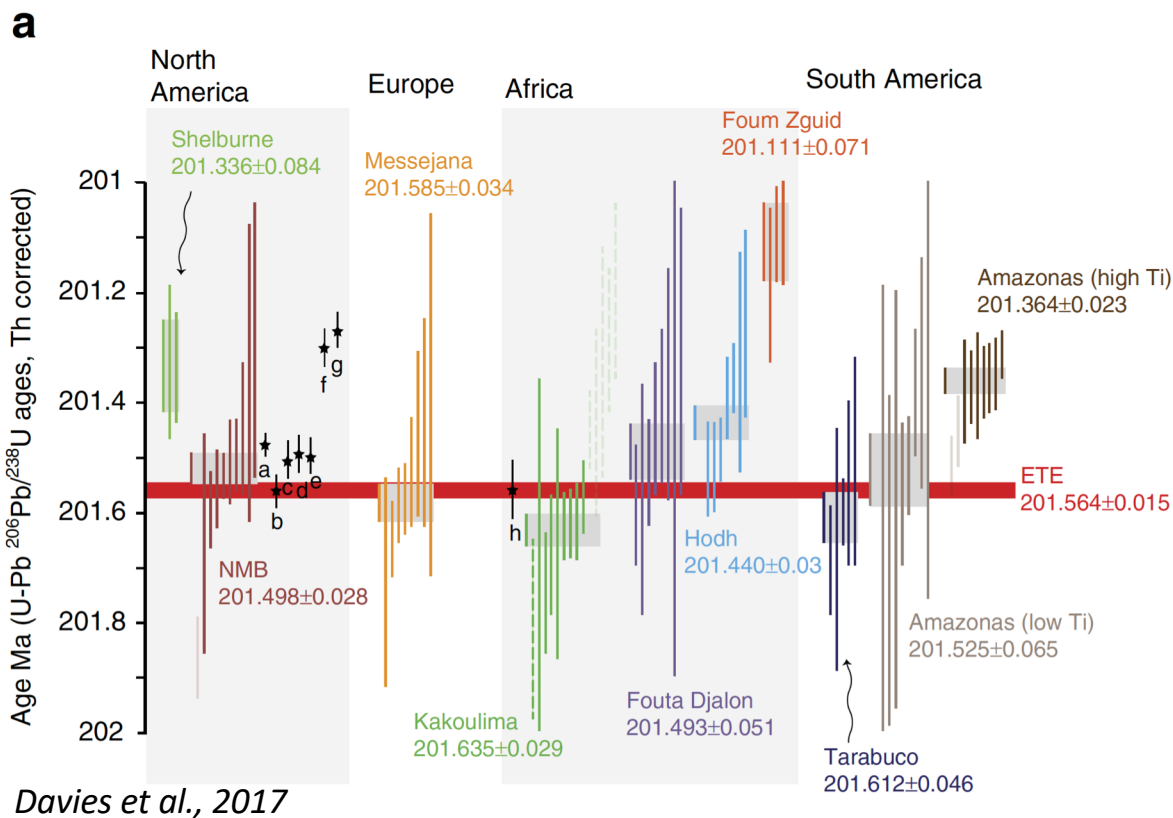
Zeh et al. (2015)



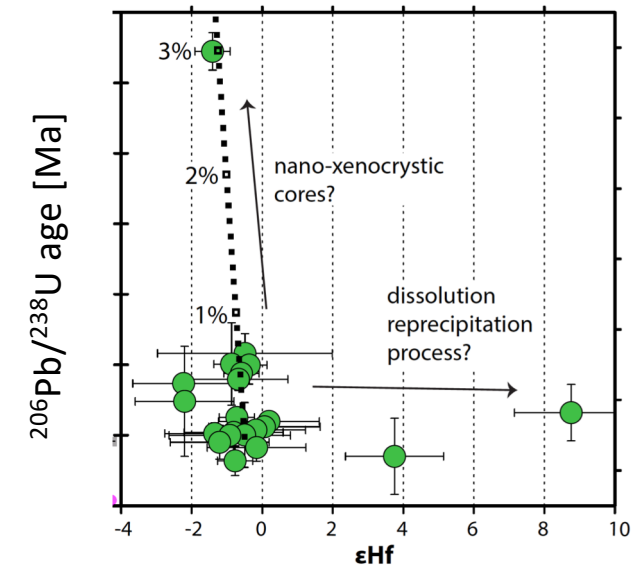
- Zircon is not crystallizing directly from tholeiitic melt but from evolved residual melt
- T estimates for zircon crystallization are at  $\sim 900\text{-}700^{\circ}\text{C}$
- zircon is thus forming after  $>90$  vol% crystallization of major minerals

## Zircon growth in mafic magmatic systems

Zircon can date emplacement of the dyke/sill complex of a LIP at highest temporal resolution  
Example: North Atlantic Magmatic Province



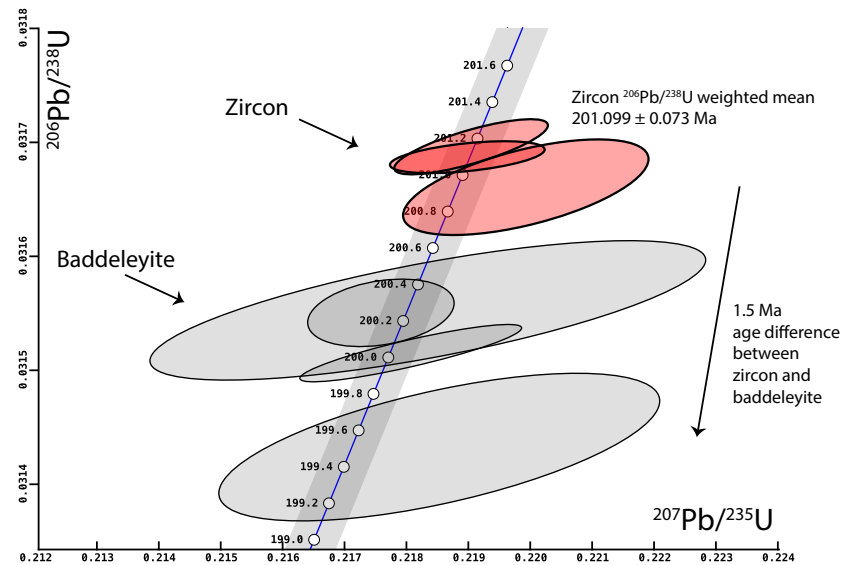
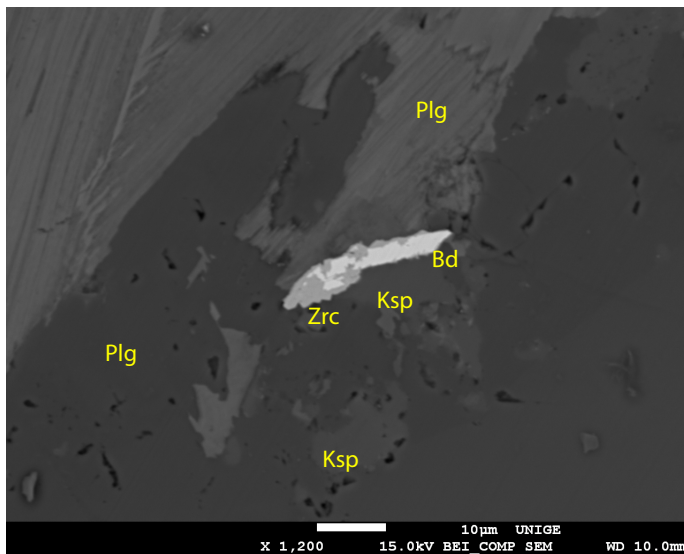
Their age variation, however, indicates participation of an inherited component



Davies et al., in prep.









## Alternative if there is no zircon in the rock: Baddeleyite ( $\text{ZrO}_2$ )



Example from a CAMP basaltic dyke; *Schaltegger and Davies (2017); Davies et al. (in prep.)*

- baddeleyite U-Pb ages are younger than chemically abraded zircon ages
- Main problem with baddeleyite petrochronology is that reliable age information can not be obtained at present due to unresolved Pb loss

## Conclusions

-  Arc plutons grow through incremental addition of small ( - to bigger) melt batches
-  The melts were saturated in zircon at intermediate crustal levels, zircon was transported in the melt to upper crustal level
-  Zircon crystals were also recycled from previous crystal mushes that were remobilized through incoming hot magma
-  Therefore few zircons are in chemical and isotopic equilibrium with the present-day host rock
-  Zircon in residual melts of mafic (tholeiitic) systems only saturates after >90% fractional crystallization
-  Excess age variation concurrent with Hf isotopes indicates that zircon in residual melts of mafic systems may at least partly nucleate around tiny relics of inherited zircon