

'Water' content as a tool to estimate rheological differences in the lithosphere of young extensional basins

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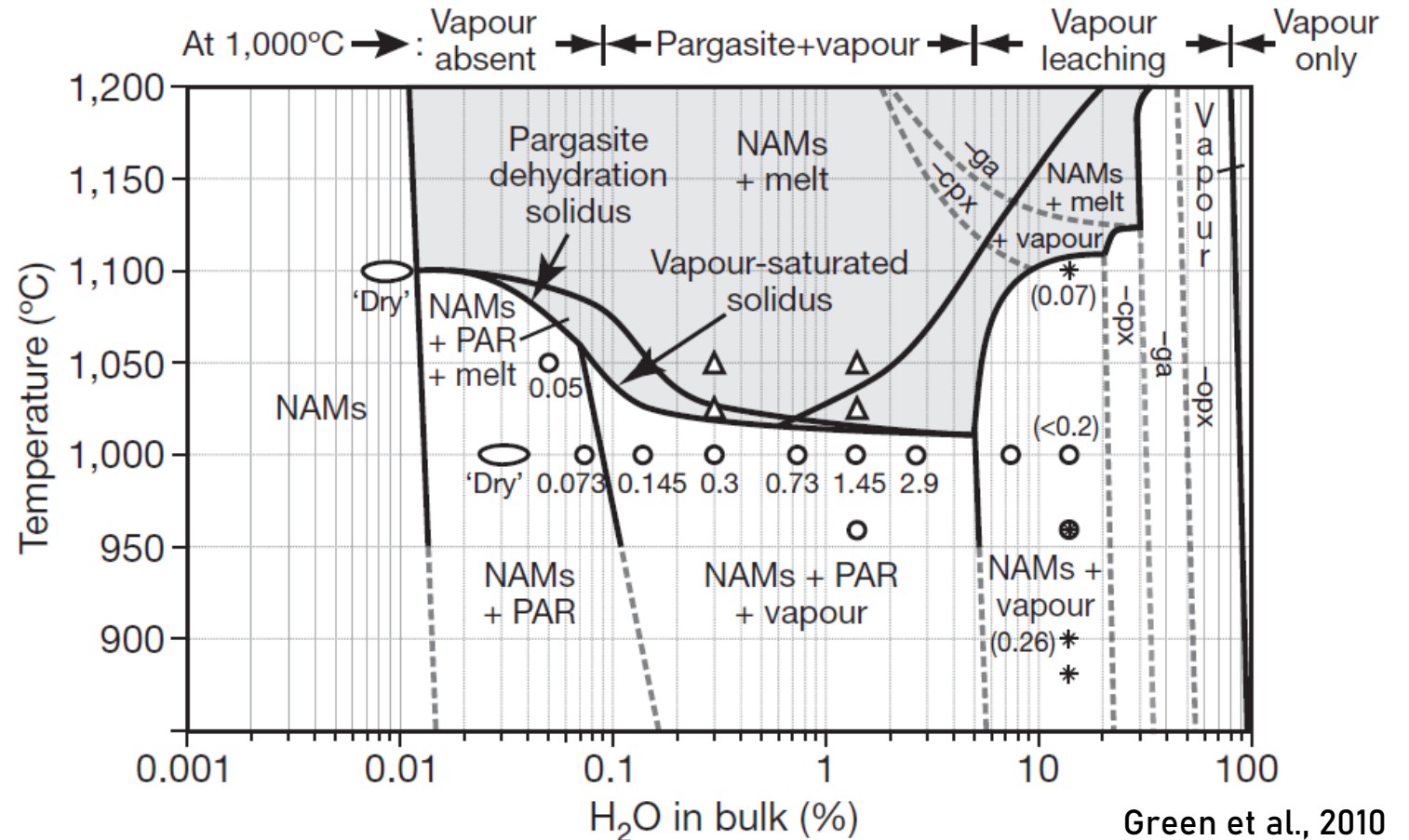
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Presence of 'water' in the upper mantle

- H₂O in fluid/melt inclusions
- Structurally bound hydroxyl in mineral structures
 - Volatile-bearing mantle minerals (e.g., pargasite, phlogopite) - ~2 wt.%
 - Nominally anhydrous mantle minerals (olivine, pyroxenes) - tens to hundreds of wt. ppm

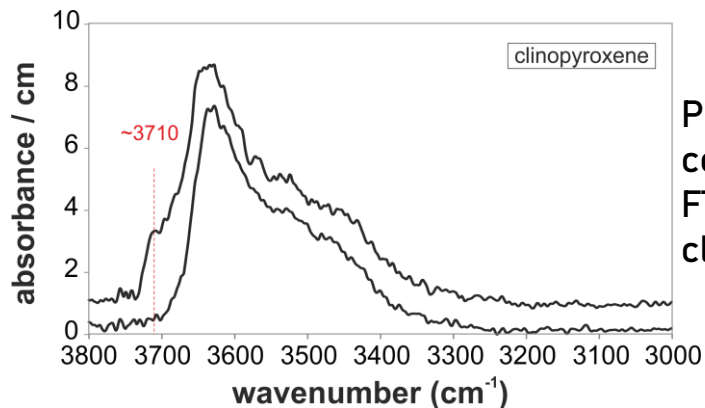
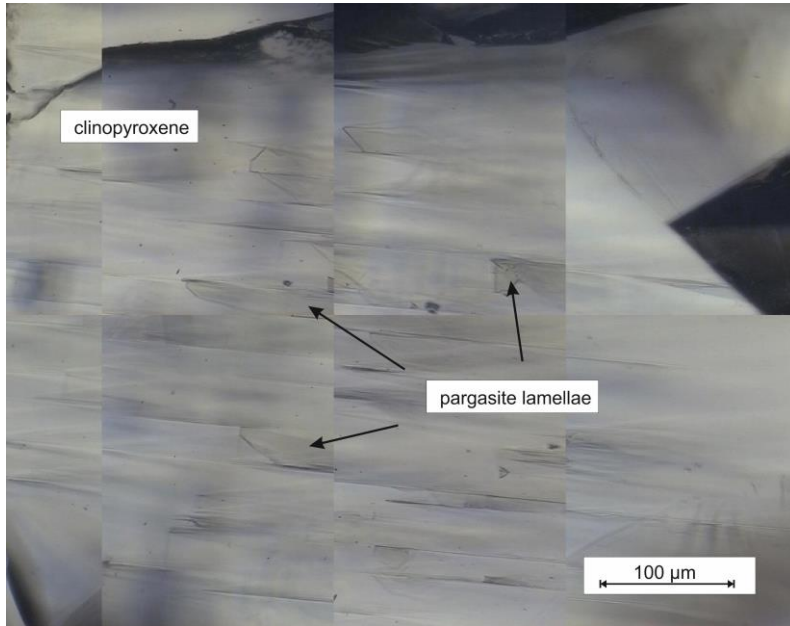


The upper limit of pargasite stability:

1050-1150°C or ~90-100 km depending on the lithospheric composition (fertility)

Pargasite stability

- Interstitial grains
- Lamella in pyroxenes (exsolving upon entering stability)



Pargasite contribution on FTIR spectrum of clinopyroxene

Outside pargasite stability

'Water' appears as:

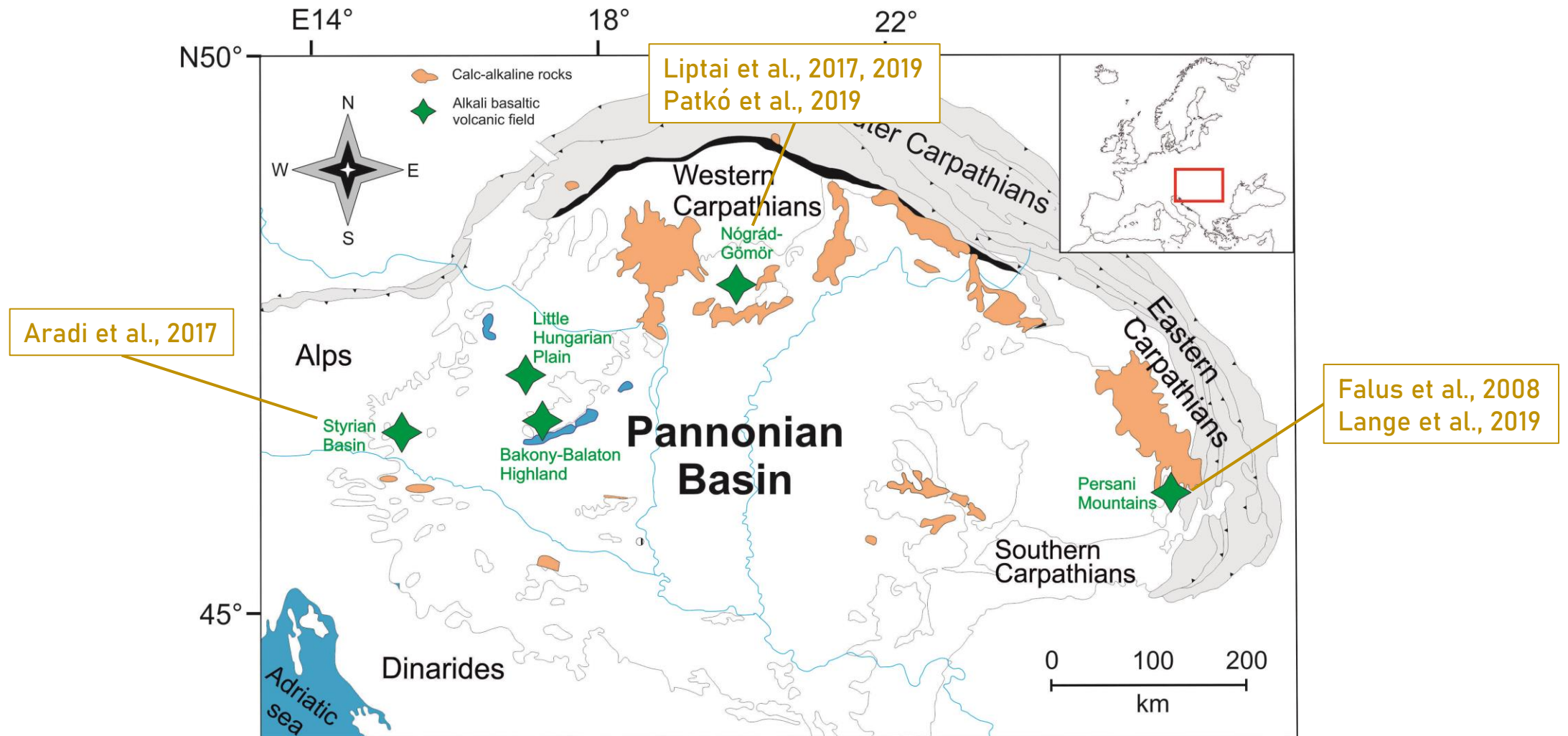
- Aqueous phase (e.g., in fluid inclusions)
- Dissolved in incipient melt
- Partitioning in nominally anhydrous minerals (NAMs) as H⁺ in cation vacancies



Influence on rheological properties (hydrolitic weakening):

- Partial melting temperature
- Effective viscosity
- Electrical conductivity / resistivity

Water in the xenoliths of the Carpathian-Pannonian region



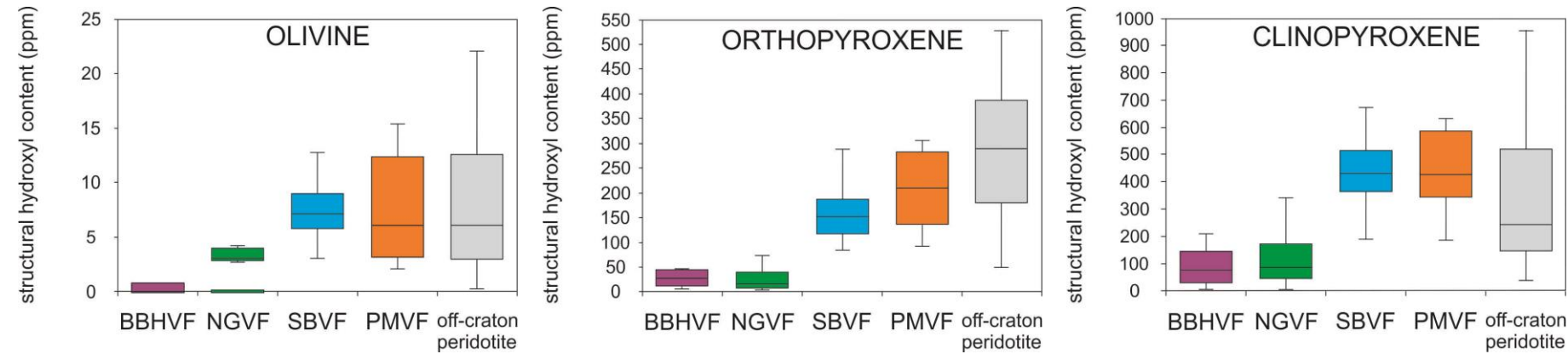
Late Miocene – Pleistocene xenolith-bearing alkali basalt localities:

'Marginal' locations: Styrian Basin, Perşani Mountains (in the vicinity of subduction zones)

'Central' locations: Bakony-Balaton Highland, Little Hungarian Plain, Nógrád-Gömör

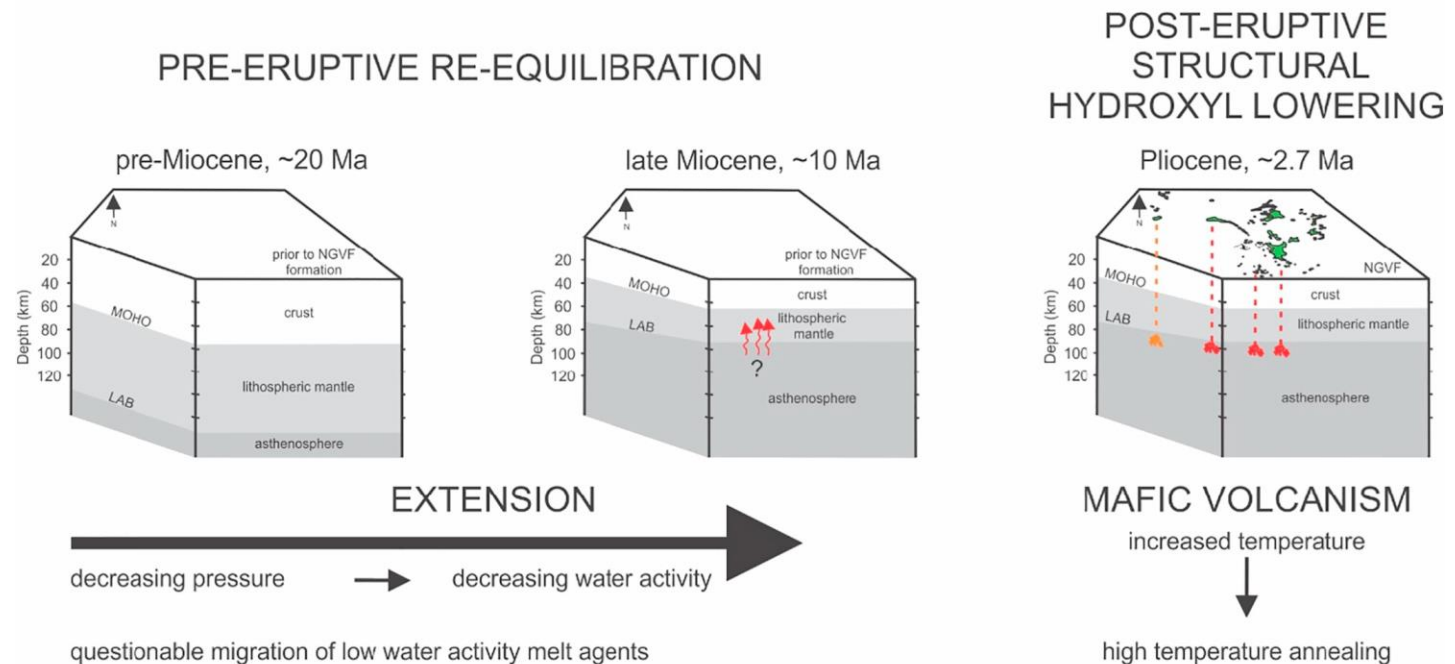
Water in the xenoliths of the Carpathian-Pannonian region

NAMs of xenoliths from Bakony-Balaton Highland (BBHVF) and Nógrád-Gömör (NGVF) have significantly lower water content than those from Styrian Basin (SBVF) and Perșani Mountains (PMVF)



Reasons for lower water content in xenoliths from areas more affected by lithospheric thinning:

- Pre-eruptive re-equilibration
- Post-eruptive hydrogen loss during cooling

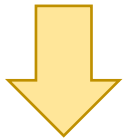


Effective viscosity

Expressed from stress and strain rate: $\eta_{\text{eff}} = \sigma / \dot{\epsilon}$

Strain rate is calculated with an Arrhenius equation containing a term related to water content (water fugacity or concentration)

$$\dot{\epsilon} = A_{cre} \sigma^{n_1} f_{H_2O}^r \exp\left(-\frac{Q + PV_{cre}}{RT}\right)$$



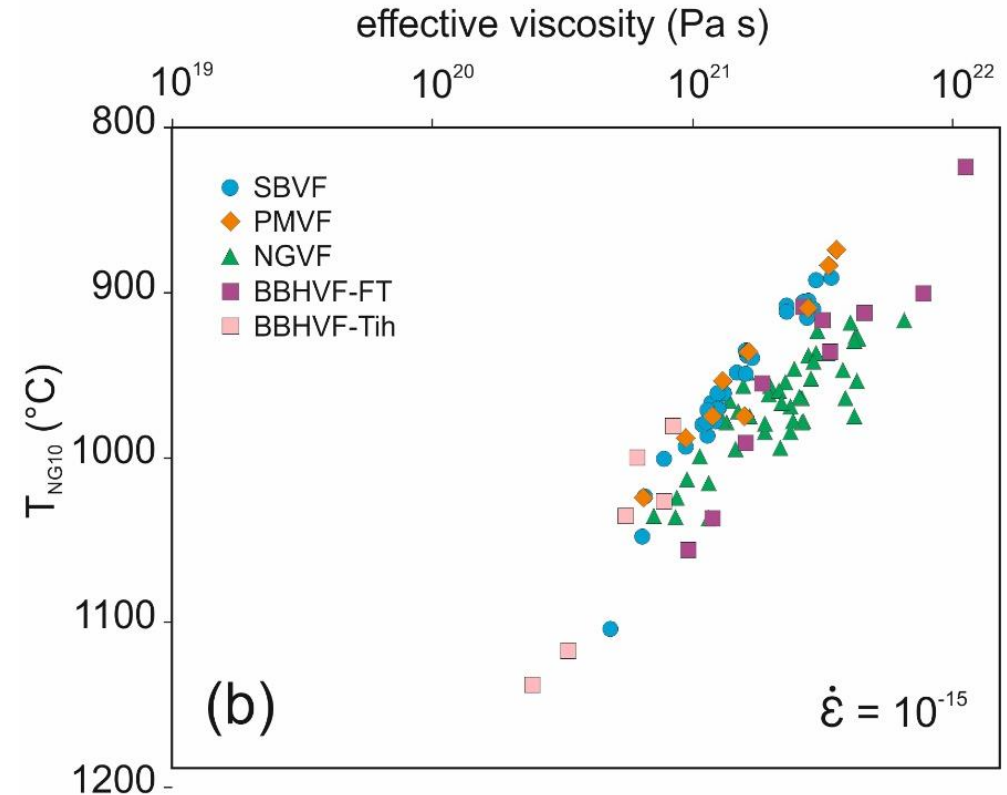
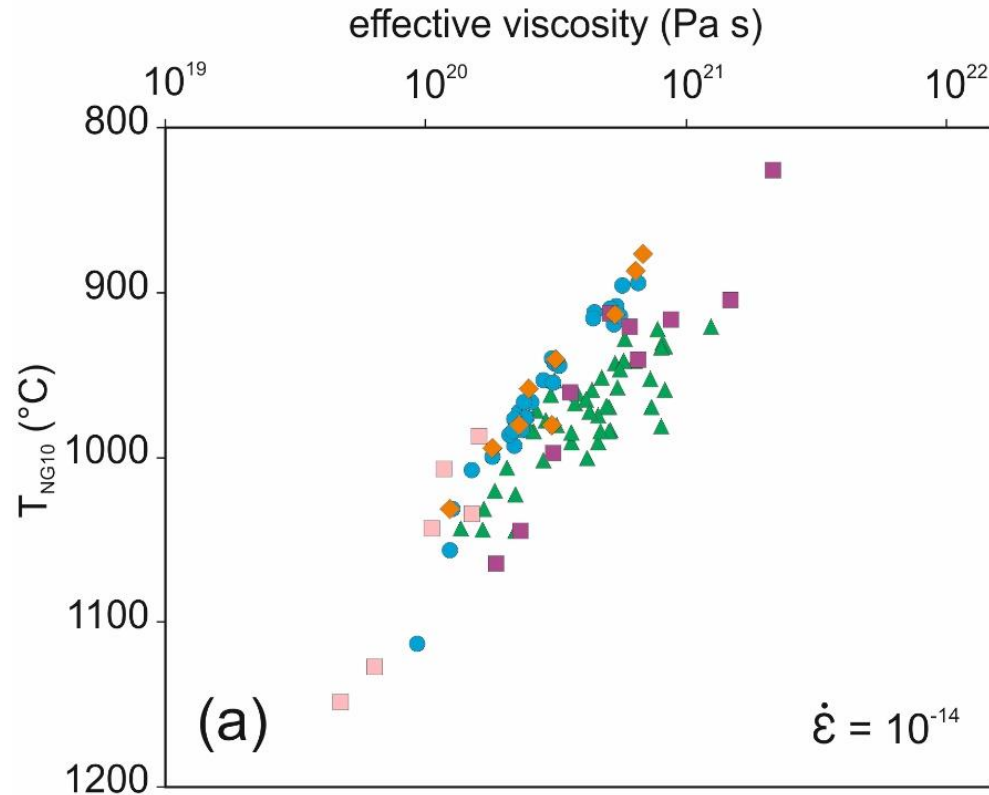
$$\eta_{\text{eff}} = \dot{\epsilon}^{(1-n)/n} f_{H_2O}^{-r/n} (A \exp^{-(H^*/RT)})^{-1/n}$$

Material constants		$H^* = Q^* + PV^*$	
	A	Q^* (J/mol)	V^* (m ³ /mol)
Dry dislocation	1.1×10^5 (MPa) ^{-n} /s	5.30×10^5 (± 0.04)	20×10^{-6}
Wet dislocation (constant f_{H_2O})	1600 (MPa) ^{-($n+r$)} /s	5.20×10^5 (± 0.4)	22×10^{-6}
Wet dislocation (constant C_{OH})	90 (MPa) ^{-($n+r$)} /s	4.80×10^5 (± 0.4)	11×10^{-6}

n (stress exponent)=3.5, r (fugacity exponent)=1.2 for wet dislocations [40].

Water fugacity (f_{H_2O}) can be replaced with water concentration (C_{OH})

Effective viscosity



Effective viscosity decreases with depth due to increasing T

- 'Marginal' localities have similar trends
- 'Central' localities show higher viscosity on average

Beside water content, strain rate also has an important effect

T_{NG10} : Ca-in-opx thermometer of Brey and Köhler (1990) modified by Nimis and Grütter (2010)

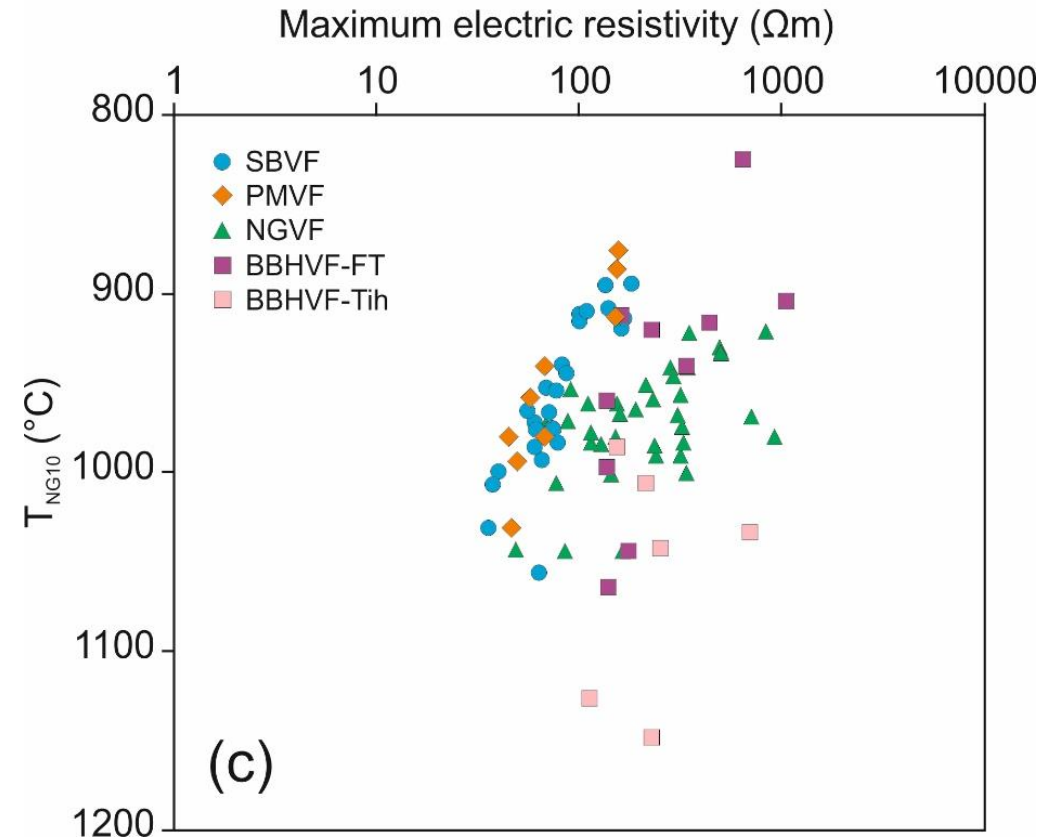
BBHVF-FT and -Tih are the youngest and oldest xenolith locality of the volcanic field, respectively

Electrical conductivity / resistivity

Model of Fullea, 2017

- Includes pyroxenes and their proportion alongside olivine
- Used parameters: modal ratio, composition (Fe-content), and structural hydroxyl concentration of olivine and pyroxenes

Higher resistivity of 'central' localities → lower water content (no significant difference in Fe-content)
Presence of melts/fluids can have a significant effect



T_{NG10} : Ca-in-opx thermometer of Brey and Köhler (1990) modified by Nimis and Grütter (2010)

BBHVF-FT and -Tih are the youngest and oldest xenolith locality of the volcanic field, respectively

Conclusions

- In the Carpathian-Pannonian region, xenoliths from the 'marginal' localities are more water-rich compared to those from 'central' localities
- Water content has an effect on rheological properties (effective viscosity, resistivity) → central areas are more rigid than marginal areas
- This may be similar in subduction – back-arc basin systems worldwide
- Additional effects (change in strain rate, presence of melts/fluids) need to be taken into account
- Comparable with geophysical observations (e.g., deep MT soundings)