

Reactive transport of dichloromethane in porous media under dynamic hydrogeological conditions: from experiments to modelling

M. Prieto-Espinoza^{1*}, S. Weill¹, R. Di Chiara¹, B. Belfort¹, F. Lehmann¹, J. Masbou¹, E. Muller², S. Vuilleumier², G. Imfeld^{1*}

¹ University of Strasbourg, CNRS/EOST, LHyGeS UMR 7517, Laboratory of Hydrology and Geochemistry of Strasbourg, Strasbourg, France

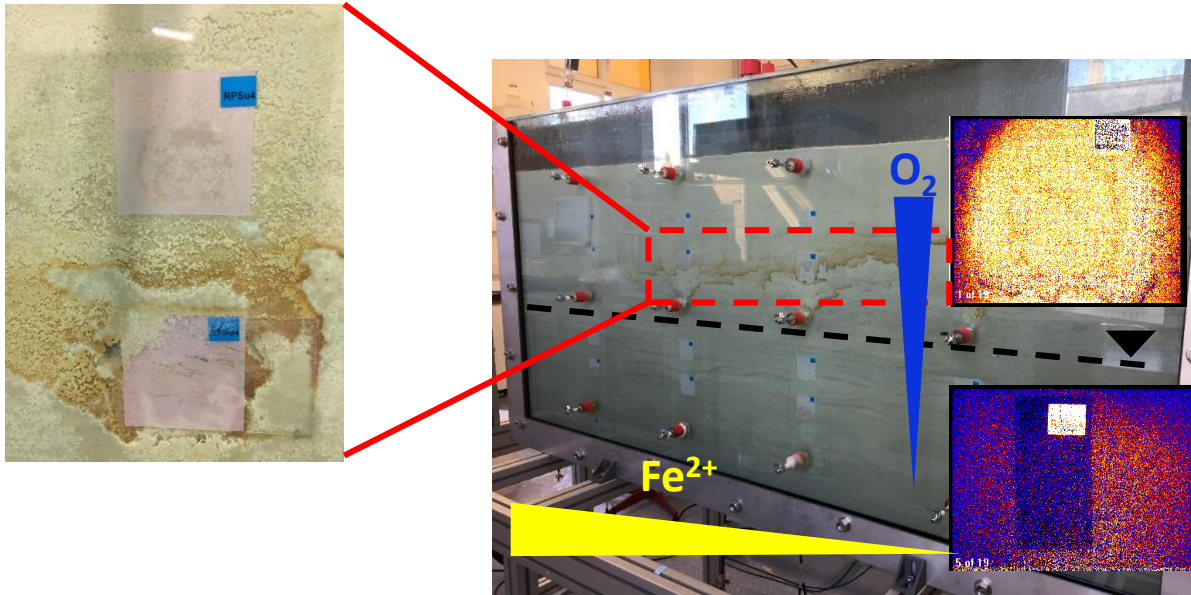
² University of Strasbourg, CNRS, GMGM UMR 7156, Génétique Moléculaire, Génomique, Microbiologie, Strasbourg, France

Motivation

- **Dichloromethane** (DMC) is one of the most common chlorinated methanes often detected in groundwater as a result of extensive use, inappropriate disposal and accidental spills.^{1,2}
- In surface-groundwater interfaces, **water table variations** can strongly impact the transfer of gases (e.g. oxygen), the evolution of redox conditions and the evolution/adaptation of bacterial/microbial populations that control biodegradation pathways of contaminants.
- In this study, **numerical modelling** is used to investigate DCM reactive transport under steady-state and fluctuating conditions based on our experimental study.
- Compound-stable isotope analysis (**CSIA**) has been used to derive the extent of DCM degradation and reaction mechanisms using a carbon-chlorine isotope approach.⁵
- This study **aims** to i) integrate a dual-element isotope modelling approach, ii) to validate the model with our experimental data and iii) to prove its feasibility under dynamic hydrogeological conditions.

Methods: Experimental study

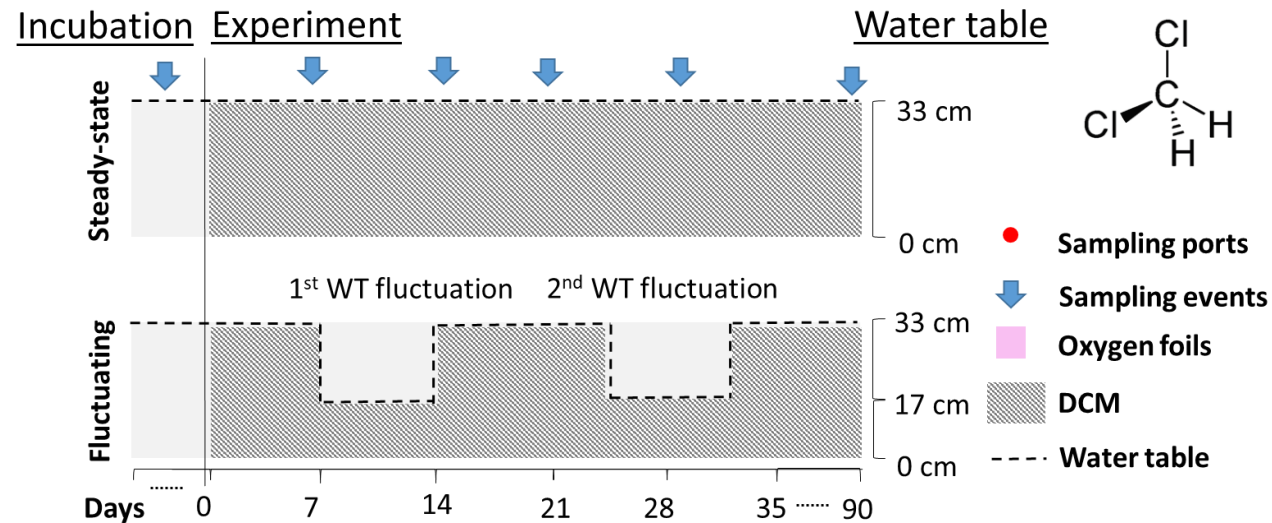
Two laboratory aquifers under near-natural setting (injection of field groundwater).



Dimensions: 160 x 80 x 7 cm³
Flow rate: 0.33 mL/min
DCM injection: 0.45 mM/L
Oxygen content: < 0.3 mg/L
Residence time: 31 days

Highlights:

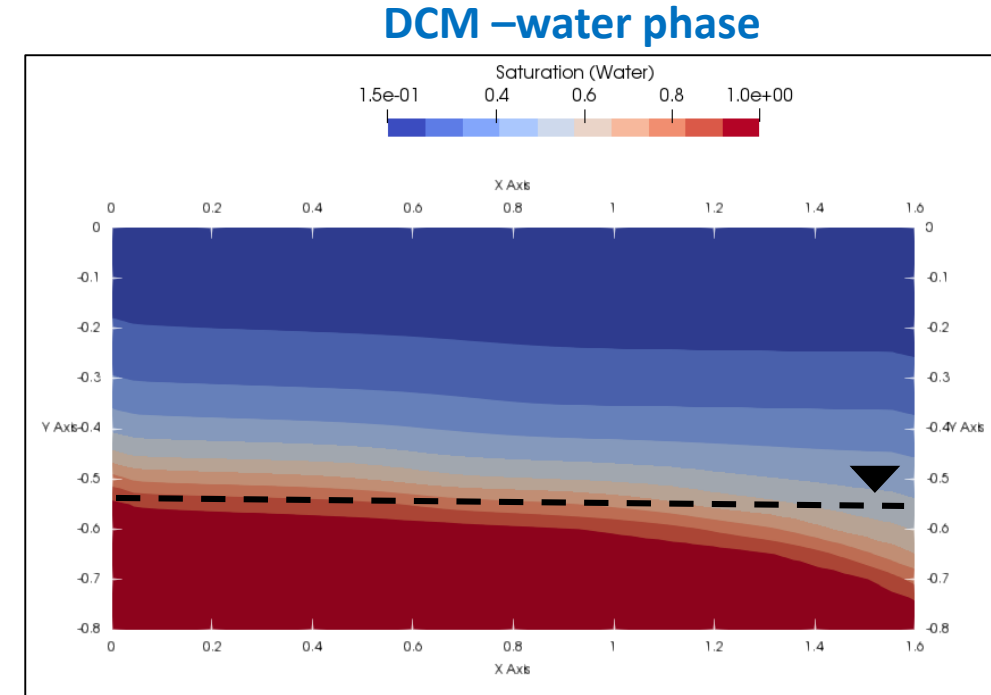
- Iron-reducing conditions prevailed in the systems.
- Water table fluctuations enhanced DCM biodegradation and influenced DCM degradation pathways compared to steady-state conditions.



Scheme of the operation of two laboratory aquifers.

Methods: Numerical simulations - 2D CubicM model⁸

- The presented numerical simulations account for:
 - i) transport processes (e.g., advection, dispersion) and
 - ii) reactive transformations (e.g., microbial transformation).
- Microbial transformation is derived from changes in isotopic composition³ of two elements: ^{13}C and ^{37}Cl . We also account for isotope fractionation by molecular diffusion.⁴
- Kinetics of biodegradation are described by Monod kinetics³.
- Simulation of the 4 most abundant DCM isotopologues. The abundance of each isotopologue is computed accounting for the occurrence of C and Cl isotopes.⁵

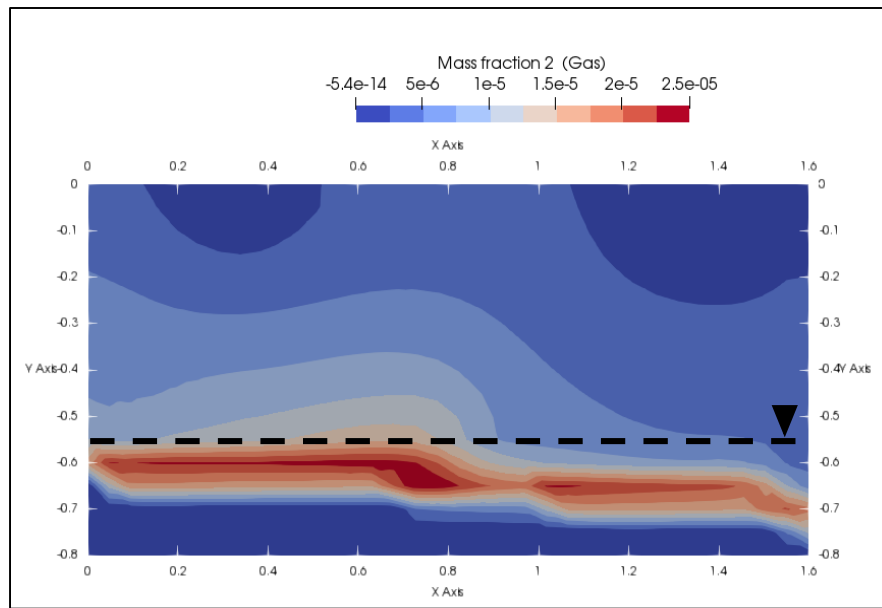


DCM plume in the saturated zone after 12 days of steady-state injection.

Two-phase flow: accounting for transient conditions

Current model simulation (12 days) under steady-state conditions have proven to grasp two-phase flow of water and gas > DCM volatilization and oxygen intrusion from the surface of the aquifer models.

DCM – gas phase

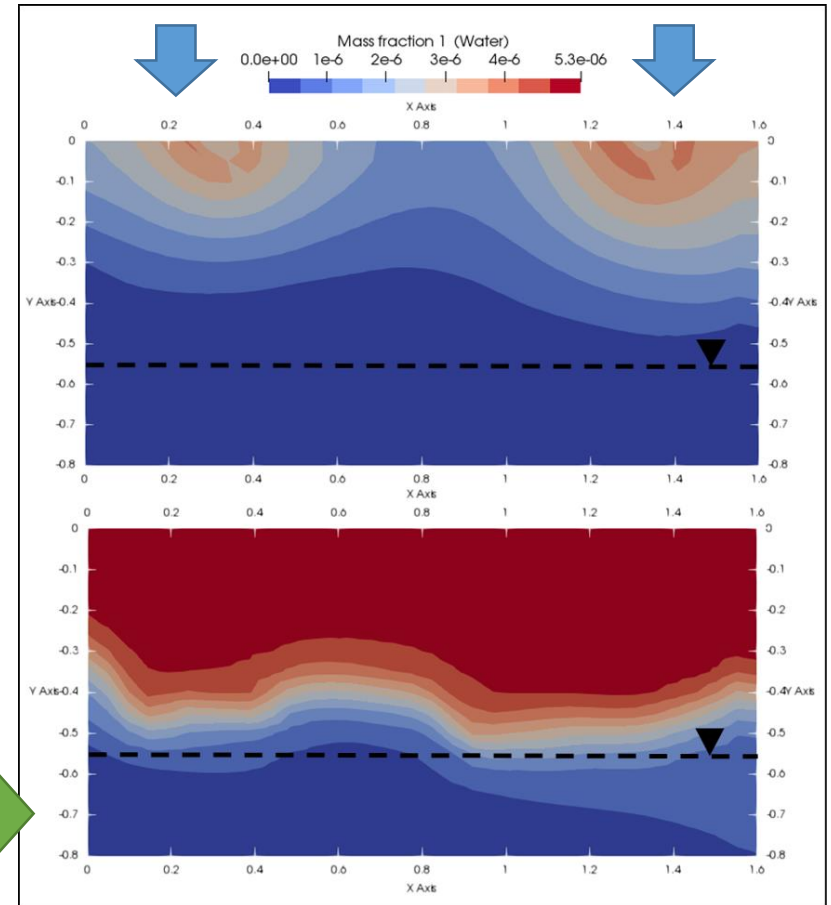


DCM transport in the gas phase – DCM volatilization.

Anoxic groundwater



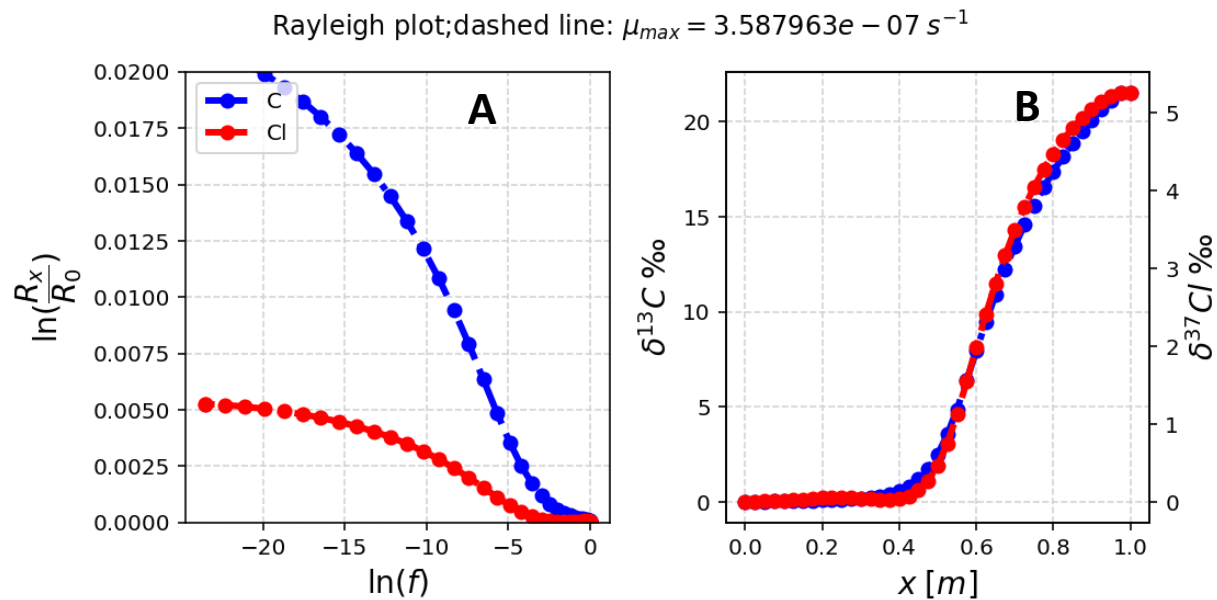
Ambient air



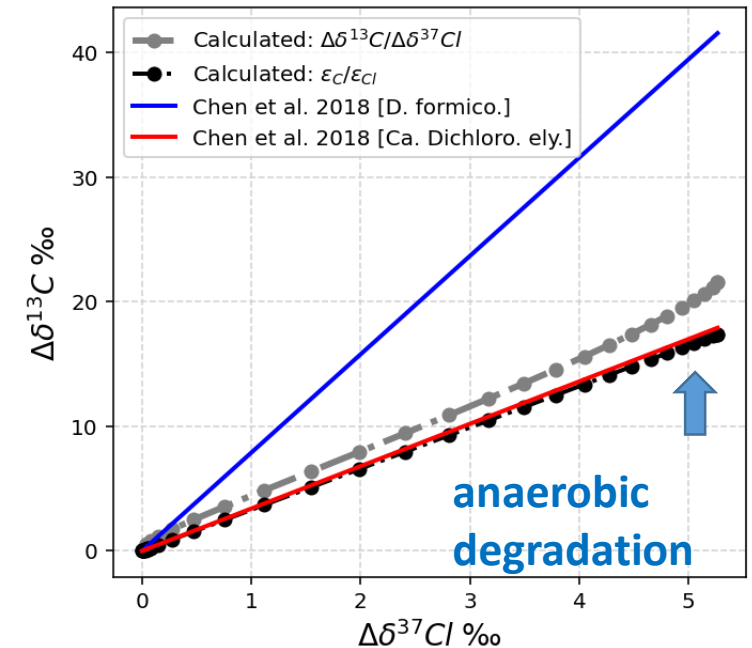
Oxygen intrusion from two ports (top).
Gradient of oxic to anoxic conditions from top to bottom.

Dual-element CSIA and DCM degradation pathway

- A simple 2D simulation of the 4 DCM isotopologues to validate the model isotope approach⁵:
 - ✓ DCM degradation pathways (under oxic and iron-reducing conditions).
- Preliminary results are in agreement with our experimental study where DCM degradation pathway followed the one described by the DCM degrader *Ca. Dichloromethamonas elyunquensis*^{6,7}.



- A) Distinct ^{13}C and ^{37}Cl isotope fractionation.
- B) ^{13}C and ^{37}Cl isotope fractionation over distance of domain.



Simulation of dual C-Cl isotope plot of DCM degradation.

Outlook

- The dual-element isotope modelling approach will be tested on the 2D-CubicM model (accounting for isotope evolution under transient conditions).
- Longer simulation times are needed to grasp differences between the steady-state and fluctuating system:
 - ⇒ hydrochemistry (i.e., oxygen and reduced iron)
 - ⇒ DCM mass balance
 - ⇒ isotope signatures
 - ⇒ Difference in microbial distribution and biomass evolution
- Coupling isotope data (dual-element) and bacterial kinetics may help to understand the highlighted results of our experimental study:
 - ⇒ distinct DCM degradation pathways under steady-state and fluctuating conditions?
 - ⇒ mechanistically distinct C-Cl bond cleavage reactions subjected to microbial adaptations during dynamic hydrogeological conditions?
- By considering dynamic hydrogeological conditions, our model may prove useful for improved assessment of natural attenuation of DCM at contaminated sites.

References

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THANK YOU!

¹LHyGES:

Gwenaël Imfeld
Sylvain Weill
Benjamin Belfort
François Lehmann
Raphaël Di Chiara
Benoit Guyot
Jérémy Masbou
Colin Fournet
Thierry Perrone

²GMGM:

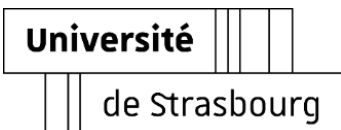
Stéphane Vuilleumier
Emilie Muller
Carmen Lazaro

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*E-mail corresponding authors:

M. Prieto-Espinoza: prietoespinoza@unistra.fr
G. Imfeld: imfeld@unistra.fr

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Governing equations

We solve the following system of partially ordinary equations:

$$^{12}C^{35}Cl_2 : \frac{\partial ^{12}C_s}{\partial t} + \mathbf{u} \nabla ^{12}C_s = \nabla \cdot (^{12}D_s \nabla ^{12}C_s) - r_{deg} \frac{^{12}C_s}{C_s^{tot}}$$

$$^{13}C^{35}Cl_2 : \frac{\partial ^{13}C_s}{\partial t} + \mathbf{u} \nabla ^{13}C_s = \nabla \cdot (^{13}D_s \nabla ^{13}C_s) - r_{deg} \frac{^{13}C_s}{C_s^{tot}} \alpha_C$$

$$^{12}C^{37}Cl^{35}Cl : \frac{\partial ^{12}C_s^*}{\partial t} + \mathbf{u} \nabla ^{12}C_s^* = \nabla \cdot (^{12}D_s^* \nabla ^{12}C_s^*) - r_{deg} \frac{^{12}C_s^*}{C_s^{tot}} \alpha_{Cl}$$

$$^{13}C^{37}Cl^{35}Cl : \frac{\partial ^{13}C_s^*}{\partial t} + \mathbf{u} \nabla ^{13}C_s^* = \nabla \cdot (^{13}D_s^* \nabla ^{13}C_s^*) - r_{deg} \frac{^{13}C_s^*}{C_s^{tot}} \alpha_C \alpha_{Cl} \quad \text{in which } C_s^{tot} = ^{12}C_s + ^{13}C_s + ^{12}C_s^* + ^{13}C_s^* [M/L^{-3}]$$

$$\frac{\partial C_{O_2}}{\partial t} + \mathbf{u} \nabla C_{O_2} = \nabla \cdot (D_{O_2} \nabla C_{O_2}) - r_{deg}^{O_2} f_{deg}^{O_2} \left(\frac{^{12}C_s}{C_s^{tot}} + \frac{^{13}C_s}{C_s^{tot}} \alpha_C + \frac{^{12}C_s^*}{C_s^{tot}} \alpha_{Cl} + \frac{^{13}C_s^*}{C_s^{tot}} \alpha_C \alpha_{Cl} \right) - k_{dec} f_{dec}^{O_2} X$$

$$\frac{\partial C_{Fe^{3+}}}{\partial t} + \mathbf{u} \nabla C_{Fe^{3+}} = \nabla \cdot (D_{Fe^{3+}} \nabla C_{Fe^{3+}}) - r_{deg}^{Fe} f_{deg}^{Fe} \left(\frac{^{12}C_s}{C_s^{tot}} + \frac{^{13}C_s}{C_s^{tot}} \alpha_C + \frac{^{12}C_s^*}{C_s^{tot}} \alpha_{Cl} + \frac{^{13}C_s^*}{C_s^{tot}} \alpha_C \alpha_{Cl} \right) - k_{dec} f_{dec}^{Fe} X$$

$$* \frac{\partial X}{\partial t} + \mathbf{u} \nabla C_X = \nabla \cdot (D_X \nabla C_X) + r_{deg} Y_s \left(\frac{^{12}C_s}{C_s^{tot}} + \frac{^{13}C_s}{C_s^{tot}} \alpha_C + \frac{^{12}C_s^*}{C_s^{tot}} \alpha_{Cl} + \frac{^{13}C_s^*}{C_s^{tot}} \alpha_C \alpha_{Cl} \right) \left(1 - \frac{X}{X_{max}} \right) - k_{dec} X$$

* All kinetic parameters are currently under revision.