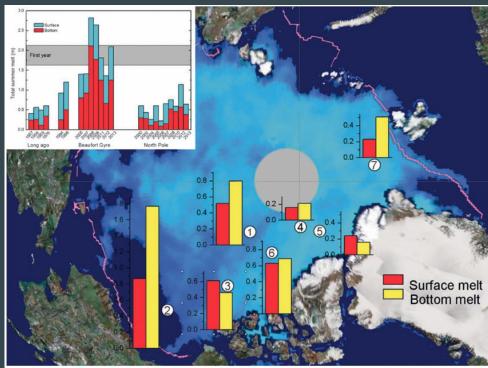
Quantifying the impact of submesoscale dynamics on the evolution of Arctic freshwater fronts

vEGU21 - Arctic changes – processes and feedbacks in climate, ocean and cryosphere (CL4.7)

Marion Alberty, Sonya Legg, Robert Hallberg, Jennifer MacKinnon, Janet Sprintall, Matthew Alford, John Mickett, and Elizabeth Fine

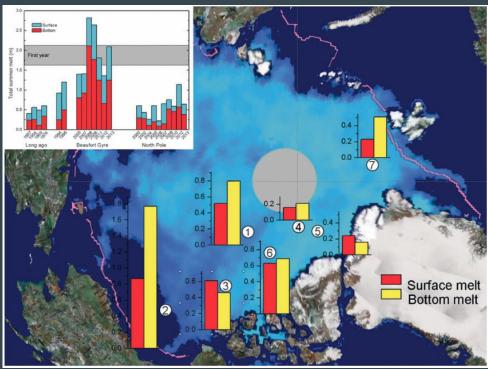
Seasonal cycle of freshwater flux due to sea ice melt.

Summer 2008 sea ice melt [m]



Seasonal cycle of freshwater flux due to sea ice melt.

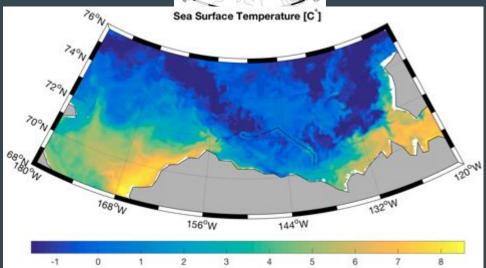
Over recent decades, seasonal ice melt has increased particularly in the Canada Basin. Summer 2008 sea ice melt [m]



Seasonal cycle of freshwater flux due to sea ice melt.

Over recent decades, seasonal ice melt has increased particularly in the Canada Basin.

This source of freshwater is inherently patchy.



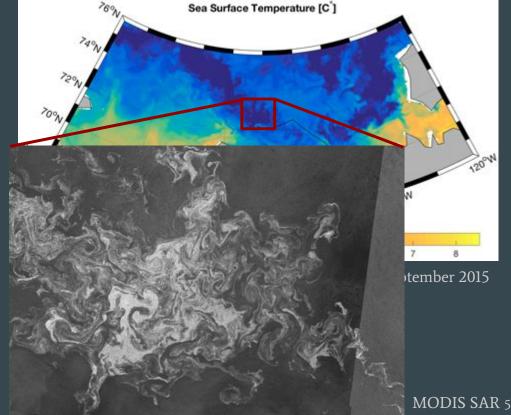
Hycom reanalysis, September 2015

Seasonal cycle of freshwater flux due to sea ice melt.

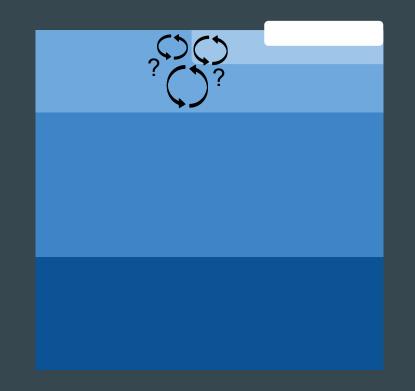
Over recent decades, seasonal ice melt has increased particularly in the Canada Basin.

This source of freshwater is inherently patchy.

The resulting fronts are stirred and strained by eddies and are observed to have Rossby numbers O(1).

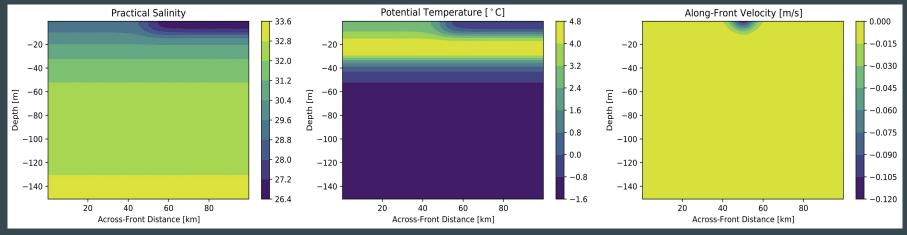


Questions



- What processes control the evolution of these freshwater fronts?
- What is the importance of submesoscale processes for the lateral and vertical exchange of heat and salt?

Model Setup



Vertical profiles and frontal length scales based on in-situ observations with a geostrophically balanced jet centered at the front. Salinity field seeded with white noise.

MITgcm, channel configuration, 25 vertical layers (100 km x 100 km x 150 m)

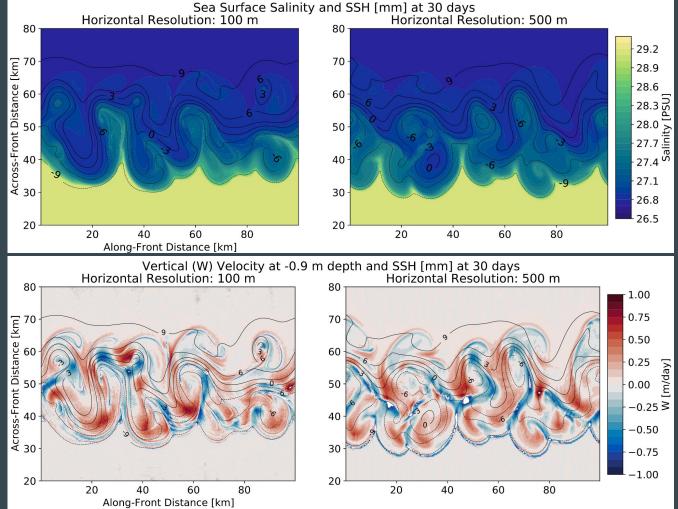
Mixed layer baroclinic Rossby radius is $\sim 2 \text{ km}$ for configuration. Employ a high (dx = 100 m) and low (dx = 500 m) lateral resolution simulations.

 $kappa_T = 1e-6 \text{ m}^2/\text{s}, nu = 1e-4 \text{ m}^2/\text{s}$

Front Evolution

Formation of baroclinic instabilities, jet meanders, and isopycnals slump and flatten.

Warm, salty intrusions into the cold, fresh side of the front oppose the tracer transport of the overturning.



Upper ocean temperature budget

 ∂T

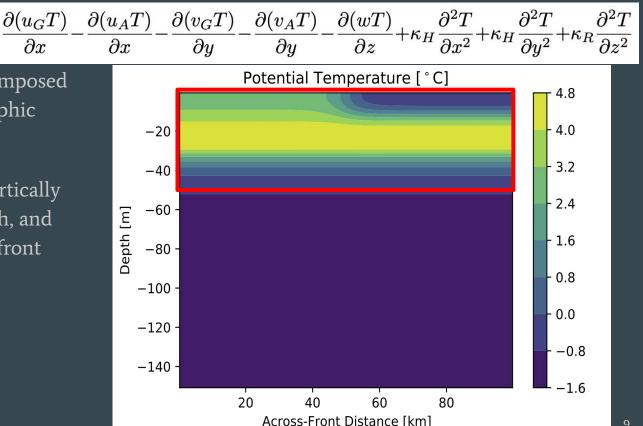
∂t.

Horizontal velocities are decomposed into geostrophic and ageostrophic components.

Integrate in time to day 60, vertically from the surface to 50 m depth, and over the domain in the along-front direction.

x - along-front direction

y - across-front direction



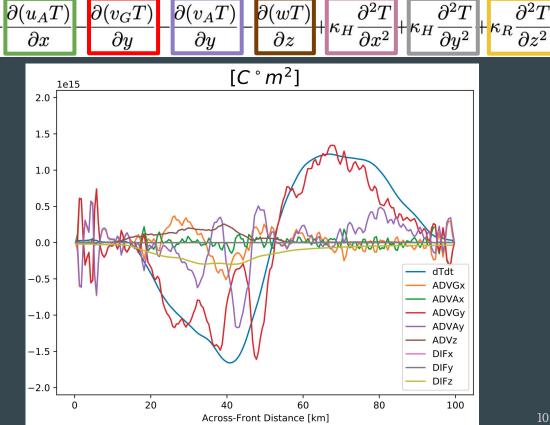
Upper ocean temperature budget ∂T $\partial(u_G T)$

 ∂x

 ∂t x - along-front direction

y - across-front direction

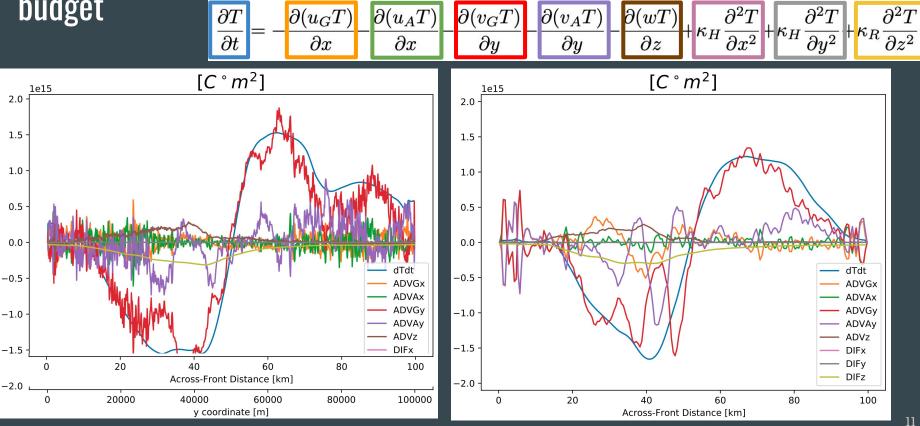
Temperature tendency is dominated by across-front geostrophic motions with across-front ageostrophic motions having greater influence on the salty side of the front.



 $\partial^2 T$

 $\partial^2 T$

Upper ocean temperature budget $\partial T = \partial (u_G T)$



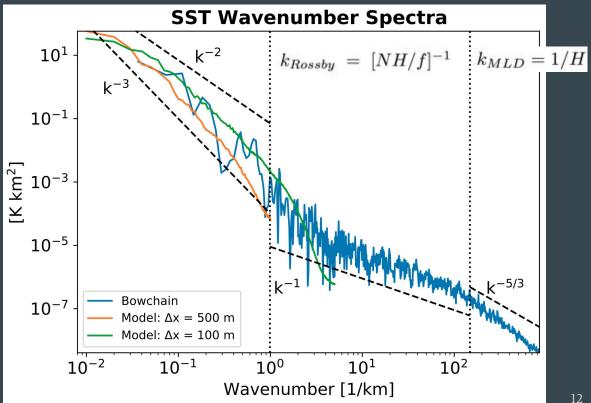
High resolution (100 m)

Low Resolution (500 m)

Comparison with observations

Simulations envelope the observations over the mesoscale range.

Still not quite resolving the first decade of the submesoscale range.



Takeaways

- 1. Across-front geostrophic motions enhance lateral transport of heat across Arctic surface mixed layer fronts.
- 2. Permitting a larger range of dynamics further increases the extent over which geostrophic motions drive lateral heat fluxes and increase the total temperature tendency.
- 3. Both simulations reproduce relatively steep spectral slopes observed over the mesoscale range but still can't capture full submesoscale range.

Thank you

Please feel free to send me your questions ahead of time at <u>malberty@princeton.edu</u>

I look forward to talking with you during the CL4.7 breakout sessions on Wednesday, 28 April 2021, 09:30-10:30 CEST.