

# Observational evidence of solar activity interaction with chlorine chemistry curbing Antarctic ozone loss

Emily Gordon\*, Annika Seppälä Department of Physics, University of Otago, NZ

Bernd Funke Instituto de Astrofísica de Andalucía, Granada, Spain

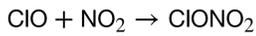
Johanna Tamminen Finnish Meteorological Institute, Finland

Kaley A. Walker Department of Physics, University of Toronto, Toronto, Canada

\*Now at Colorado State University, United States

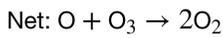
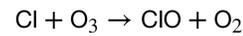
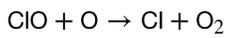
## Background

- Several mechanisms for solar influence on climate, some link to variability on **regional** rather than global **scale**.
- Solar wind brings electrons and protons from the Sun to the Earth. Earth's magnetic field causes some to "rain" into the atmosphere in the **polar areas** → **Energetic Particle Precipitation (EPP)**.
- EPP** increases ionisation in the atmosphere, influencing **chemical balance of the stratosphere and mesosphere**.
  - Produces  $\text{NO}_x$  gases (e.g.  $\text{NO} + \text{NO}_2$ , so called **EPP- $\text{NO}_x$** ), known to catalytically destroy ozone.
- $\text{NO}_x$  is long-lived in polar winter and descend into the stratosphere. If still present when spring arrives,  $\text{NO}_x$  increases could potentially interact with chlorine chemistry:



converting active chlorine to inactive reservoir

- Removes some ClO from the active chlorine catalytic cycles such as:



**Hypothesis:  $\text{NO}_x$  can react with ClO, acting as a limiter for Cl<sub>x</sub> driven ozone loss.** Test with existing satellite datasets!

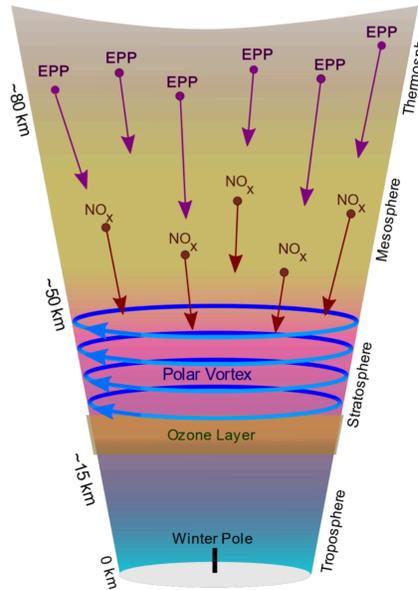


Figure 1. EPP- $\text{NO}_x$  formation over the winter pole. EPP penetrates the mesosphere-lower thermosphere resulting in  $\text{NO}_x$  increases. The  $\text{NO}_x$  descends downward into the stratosphere in the polar vortex.

## Datasets

### OMI/Aura

- $\text{NO}_2$  stratospheric column,  $\text{O}_3$  total column

### MLS/Aura

- $\text{O}_3$ , ClO stratospheric profiles

### MIPAS/Envisat

- ClONO<sub>2</sub> stratospheric profile

### ACE-FTS/SCISAT

- ClONO<sub>2</sub> profile (not shown here)

Overall, covering years 2005-2017.

**These results have been published. Please see further details of data and results in:**

Gordon, Seppälä, and Tamminen (2020) Evidence for energetic particle precipitation and quasi-biennial oscillation modulations of the Antarctic  $\text{NO}_2$  springtime stratospheric column from OMI observations, *Atmos. Chem. Phys.*, doi: 10.5194/acp-20-6259-2020

Gordon, Seppälä, Funke, Tamminen, and Walker (2021) Observational evidence of energetic particle precipitation  $\text{NO}_x$  (EPP- $\text{NO}_x$ ) interaction with chlorine curbing Antarctic ozone loss, *Atmos. Chem. Phys.*, doi:10.5194/acp-2020-847, 2021

## 1. Solar activity and EPP levels vary

Geomagnetic activity index  $A_p$  is a well known proxy of EPP variability. We use correlations with  $A_p$  to estimate the contribution of EPP to stratospheric  $\text{NO}_x$ .

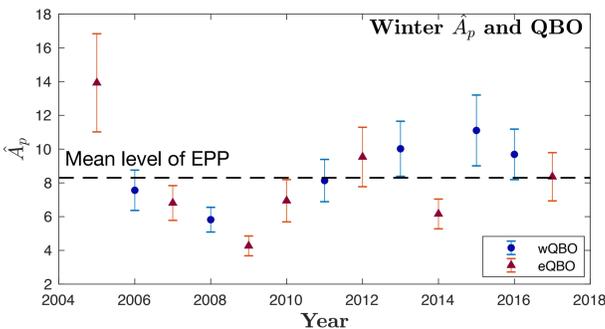


Figure 2. The SH polar winter average  $A_p$  (level of EPP), denoted  $\hat{A}_p$  and the Quasi-Biennial Oscillation (QBO) phase (easterly=eQBO; westerly=wQBO) during the polar winter. We account for the QBO as it is known to modulate 1) transport of air from lower latitudes to the polar atmosphere affecting non-EPP source of  $\text{NO}_x$ , and 2) polar temperatures (wQBO: cold, effective removal of  $\text{NO}_x$  from stratosphere; eQBO: opposite)

## 3. November year-to-year variability

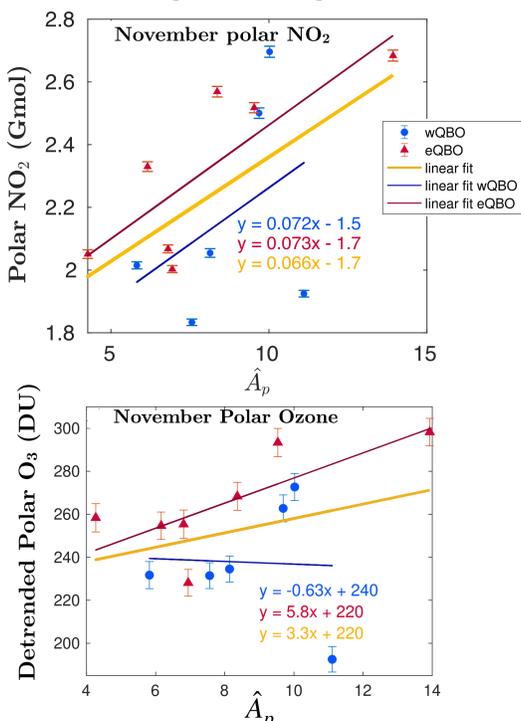


Figure 5. Top: OMI mean November polar (60S-90S)  $\text{NO}_2$  column in Giga moles vs wintertime EPP level. Bottom: OMI mean November polar (60S-90S)  $\text{O}_3$  column in DU vs wintertime EPP level.

## 2. Solar activity and stratospheric $\text{NO}_x$ and $\text{O}_3$

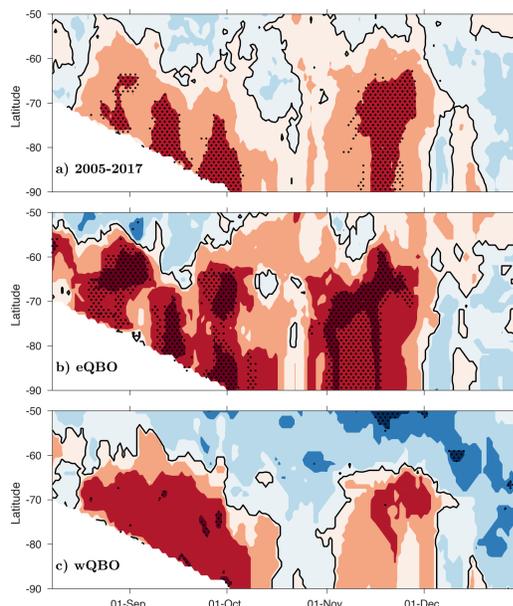


Figure 3. Correlation of  $A_p$  and OMI stratospheric  $\text{NO}_2$  column for a) all years, b) eQBO years and c) wQBO years. **Winter EPP is highly correlated with stratospheric  $\text{NO}_x$  until December during eQBO.** EPP- $\text{NO}_x$  significantly contributes to stratospheric  $\text{NO}_x$  levels. Stippling =  $\geq 95\%$  significance level.

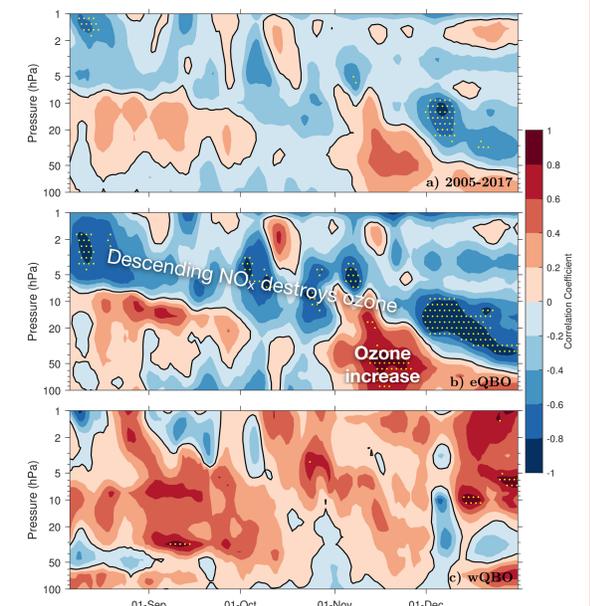


Figure 4. MLS SH polar (60S-82S)  $\text{O}_3$  profile correlation with EPP levels. Descending  $\text{NO}_x$  drives ozone loss from upper stratosphere (in August) down to 50hPa (December). **Significant positive correlation takes place below 20hPa in November indicating increased ozone.**

## 4. What about stratospheric chlorine?

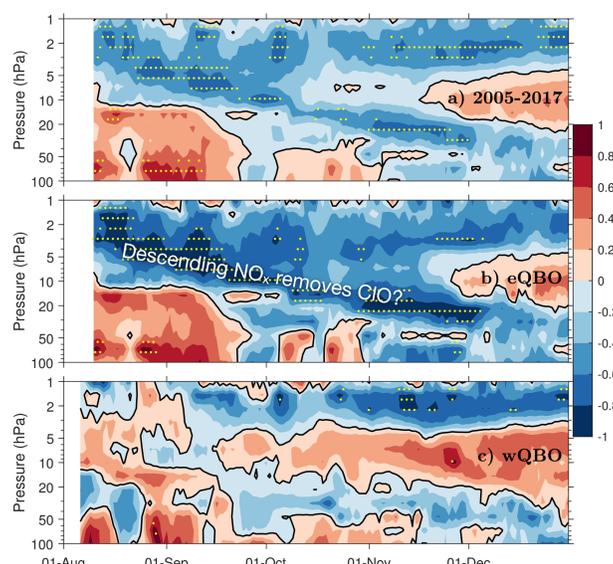


Figure 7. Polar (60S-82S) MLS ClO profile correlation with EPP levels for a) all years, b) eQBO years and c) wQBO years. The descending  $\text{NO}_x$  feature also present in ozone (Figure 4) is also present in ClO. Reduction in active chlorine would explain increase in Nov ozone.

The reaction to remove the active ClO is  
 $\text{ClO} + \text{NO}_2 \rightarrow \text{ClONO}_2$   
 If this is taking place we should observe an increase in the inactive  $\text{ClONO}_2$ !

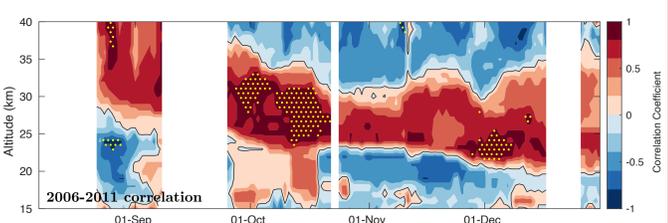


Figure 8. MIPAS polar (60S-90S) ClONO<sub>2</sub> profile correlation with EPP levels. High EPP levels are correlated with increased ClONO<sub>2</sub>, indicating **conversion to chlorine reservoir is taking place during high solar activity years.**

## Summary

More EPP = More  $\text{NO}_x$  → Reacts with ClO producing inactive  $\text{ClONO}_2$  →  $\text{O}_3$  recovery.  
**First observational evidence of this unaccounted for source of ozone variability!**