ITS3.1/NP0.1/CL4.23 Tipping Points in the Earth System (TiPES) EGU21-2462



vEGU21, 27 April 2021

## **Tipping Points Due to Parameter Drift in a Simple Model** of the Wind-driven Ocean Circulation

### **Stefano Pierini and Michael Ghil**







Pls. see further relevant material at

http://dist.altervista.org/PIERINI/ and

https://dept.atmos.ucla.edu/tcd/people/michael-ghil & https://www.researchgate.net/profile/Michael\_Ghil

## **Motivation**

- The self-consistent description, simulation and prediction of climate variability subject to climate change requires an understanding of nonautonomous dynamical systems (NDSs) and their pullback attractors (PBAs).
- A key feature of nonlinearity in classical, autonomous dynamical systems, is the presence of bifurcations that give rise to more complex behavior: multiple equilibria, limit cycles (LCs) and strange attractors.
- The generalization of bifurcations to NDSs is given by tipping points (TPs) that likewise lead to more complex behavior, including the coexistence of multiple PBAs with distinct stability properties.
- A particularly interesting case is that of excitable systems, in which a fixed point or small-amplitude, smooth solution coexists with a large-amplitude, more irregular one, like a relaxation oscillation (RO).
- > The paradigmatic example of such a system is the Van der Pol oscillator.
- Interesting climatic examples are given by paleoclimatic oscillators on various time scales (glaciation cycles, Heinrich and Dansgaard-Oeschger events), as well as interannual variability of the wind-driven circulation.
- > It is the latter case that is studied here in some detail.

#### The Simple Wind-driven Ocean Circulation Model – I

- The quasi-geostrophic (QG) model (Pierini, 2011) is projected onto 4 basis functions that include westward intensification.
- This low-order, spectrally truncated QG model is quadratically nonlinear and excitable: for timeindependent wind stress a bifurcation leads from a small limit cycle to a large amplitude self-sustained RO.
- > We will study its TPs under the action of a smooth drift in the external forcing + periodic perturbation.

$$\frac{\mathrm{d}\Psi}{\mathrm{d}t} + \Psi \mathbf{J}\Psi + \mathbf{L}\Psi = G(t)\mathbf{w}$$
$$\Psi(t) = (\Psi_1, \Psi_2, \Psi_3, \Psi_4)$$
$$\psi(\mathbf{x}, t) = \sum_{i=1}^4 \Psi_i(t) E_i(x, y)$$

Four successive snapshots of the streamfuntion showing a typical relaxation oscillation:



#### The Simple Wind-driven Ocean Circulation Model – II Time

- examples  $\succ$ Two autonomous ot behavior are identified by the thin, vertical blue arrows.
- The thick orange line shows that our model possesses the fundamental property of **excitability** when subjected to noisy forcing. This property is common to excitable systems relevant to the climate sciences, ranging from paleoclimatic multidecadal and down to to interannual time scales.



forcing

of  $\Psi_3$ 



#### Effects of a Ramp and of a Periodic Perturbation – I



### Effects of a Ramp and of a Periodic Perturbation – II

#### Numerical experiments

In Exp1 – Exp6,  $\tau$  takes on 80 different values that range from 32.5 yr to 1316.25 yr.



For each value of  $\tau$ , an **ensemble simulation (ES)** of *N* = 12x14 = 168 members differing only by their initial points (dots in the  $\Gamma$ -plane on the right) is carried out in order to simulate the irreducible uncertainty associated with the system's internal variability

This allows us to estimate the model's pullback attractor (PBA)

Numerical					
experiment	γ	α	$\tau$ (yr)	β	<i>T</i> (yr)
Exp1	0.9	0.3	32.5 - 1316.25	0	_
Exp2	0.8	0.4	"	"	_
Exp3	0.9	0.3	"	0.025	5
Exp4	"	"	"	0.050	"
Exp5	0.8	0.4	"	0.025	"
Exp6	"	"	"	0.050	"
Exp7	0.9	0.3	800	0.050	1 - 100
Exp8	"	"	"	0.100	"



#### Effects of a Ramp and of a Periodic Perturbation – III





#### *Effects of a Ramp and of a Periodic Perturbation – IV*

#### Impact of ramp steepness $\delta$

The **ramp steepness**  $\delta$  of the forcing G(t) can also be thought of as the **drift rate**:

$$\delta = \alpha R_{\tau}' \Big|_{t = (t_1 + t_2)/2}$$

- Forcing level  $G_{tp}$  at the TPs decreases with ramp length  $\tau$  and **increases** overall with ramp steepness  $\delta$ .
- ➢ A periodic perturbation lowers the forcing level required to reach a TP:  $G_{tp}^* ≤ G_{tp}$ .
- > Surprises occur where the PBA splits.



#### *Effects of a Ramp and of a Periodic Perturbation – V*



### *Effects of a Ramp and of a Periodic Perturbation – VI*



**EXP2**: Multiple PBAs appear in (a,b), total independence on initial data (TIID) is found in (c,d) and no TP arises in (c). **EXP5**: With a periodic perturbation ( $\beta \neq 0$ ) these features are lost, except the TIID, with the TP shifting away as  $\tau$  increases. **EXP6**: When the periodic perturbation is larger (larger  $\beta$ ), the TPs stay almost the same but the ROs become irregular.

#### *Effects of a Ramp and of a Periodic Perturbation – VII*

#### Nonlinear resonance

Exp7 & Exp8 differ from EXP1 in that a periodic forcing is now present, with *T* that varies from 1 to 100 yr, while  $\tau$  = 800 yr is constant.

The abrupt reduction of  $G_{tp}^*$  occurs for periods T that are comparable to the typical time scale of the ROs; thus, **nonlinearly resonant**–like behavior seems to occur.



#### **Clustering of Trajectories and Phase Coherence – I**

To study these aspects of the model's ROs, Pierini (2014) proposed to use the **parameter**  $\overline{C}$ , given by

$$\bar{C}(r,\tau) = \frac{1}{T_0} \int_{t_{\rm tp}(\tau)}^{t_{\rm tp}(\tau)+T_0} C(r,t,\tau) dt \quad \text{with} \quad C(r,t,\tau) = \frac{1}{N^2} \sum_{i,j}^{N} H\left[ r - \left| \Psi_{\tau}^{(i)}(t) - \Psi_{\tau}^{(j)}(t) \right| \right]$$

to obtain information about the **clustering** and **phase dependence** of ensemble members. Here H(r - x)



is a Heaviside function that counts proximity of orbit pairs.

#### **Clustering of Trajectories and Phase Coherence – II**







#### Clustering of Trajectories and Phase Coherence – III

Effect of the periodic perturbation on the clustering



Comparison with the previous two slides shows that the degree of clustering is basically preserved under periodic perturbation. Thus, TIID appears to be a fairly robust feature.

# **Concluding Remarks**

- ➤The simple, low-order QG model of the wind-driven circulation studied herein provides an excellent illustration of an excitable nonlinear model with relaxation oscillations (ROs).
- The ramp R(t) in the forcing represents anthropogenic warming or, rather, cooling, while the periodic forcing represents a seasonal effect.
- The time-dependent forcing leads to a remarkable plethora of phenomena that distinguish the model's NDS from the autonomous case.
- $\succ$  Overall, both the timing  $t_{tp}$  and the forcing value,  $G_{tp}$  or  $G_{tp}^*$ , at which the NDS tipping points (TPs) occur are distinct from the bifurcations of the autonomous model.
- > The duration  $\tau$  and the steepness  $\delta$  of the forcing ramp play a major role in determining the changes in timing and in the forcing threshold of the TPs.
- The fine structure of the model's global PBA changes, with multiple local attractors equipped with distinct stability properties appearing and disappearing as parameters change.
- Periodic perturbations of small amplitude have a major effect on the TPs, as well as on the PBAs fine structure.
- > There's much to do to fully understand this rich phenomenology.

## Some references

- Arnold, L., 1998: Random Dynamical Systems, Springer Monographs in Math., Springer, 625 pp.
- Caraballo T, Han X (2017) Applied Nonautonomous and Random Dynamical Systems: Applied Dynamical Systems. Springer Science + Business Media, x + 108 pp.
- Chekroun, M. D., E. Simonnet, and M. Ghil, 2011: Stochastic climate dynamics: Random attractors and time-dependent invariant measures, *Physica D*, **240**, 1685–1700, doi :<u>10.1016/j.physd.2011.06.005</u>.
- Dijkstra, H.A., and M. Ghil, 2005: Low-frequency variability of the large-scale ocean circulation: A dynamical systems approach, *Rev. Geophys.*, **43**, RG3002, doi:<u>10.1029/2002RG000122</u>.
- Ghil, M., 2017: The wind-driven ocean circulation: Applying dynamical systems theory to a climate problem, *Discr. Cont. Dyn. Syst. A*, **37**(1), 189–228, doi:10.3934/dcds.2017008.
- Ghil, M., 2019: A century of nonlinearity in the geosciences, Earth & Space Science, 6, 1007–1042, doi: 10.1029/2019EA000599.
- Ghil, M., and V. Lucarini, 2020: The physics of climate variability and climate change, Rev. Mod. Phys., 92(3), 035002, doi: 10.1103/RevModPhys.92.035002.
- Ghil, M., M.D. Chekroun, and E. Simonnet, 2008: Climate dynamics and fluid mechanics: Natural variability and related uncertainties, *Physica D*, **237**, 2111–2126, doi:<u>10.1016/j.physd.2008.03.036</u>.
- Pierini, S., 2011: Low-frequency variability, coherence resonance and phase selection in a low-order model of the wind-driven ocean circulation, *J. Phys. Oceanogr.*, **41**, 1585-1604, doi: <u>https://doi.org/10.1175/JPO-D-10-05018.1</u>.
- Pierini, S., 2014: Ensemble simulations and pullback attractors of a periodically forced double-gyre system, *J. Phys. Oceanogr.*, **44**, 3245–3254, doi: <a href="https://doi.org/10.1175/JPO-D-14-0117.1">https://doi.org/10.1175/JPO-D-14-0117.1</a>.
- Pierini, S., M. Ghil, and M.D. Chekroun, 2016: Exploring the pullback attractors of a low-order quasi-geostrophic model: The deterministic case, *J. Climate*, **29**, 4185–4202, doi: <u>https://doi.org/10.1175/JCLI-D-15-0848.1</u>.
- Pierini, S., M.D. Chekroun, and M. Ghil, and 2018: The onset of chaos in nonautonomous dissipative dynamical systems: A low-order ocean-model case study, Nonlin. Processes Geophys., 25, 671–692, doi https://www.nonlin-processes-geophys.net/25/671/2018/.

Pierini, S., and M. Ghil, 2021: Tipping points induced by parameter drift in an excitable ocean model, Scientific Reports, in press.

Sushama, L., M. Ghil, and K. Ide, 2007: Spatio-temporal variability in a mid-latitude ocean basin subject to periodic wind forcing. *Atmosphere-Ocean*, **45**, 227–250, doi: <u>10.3137/ao.450404</u>.