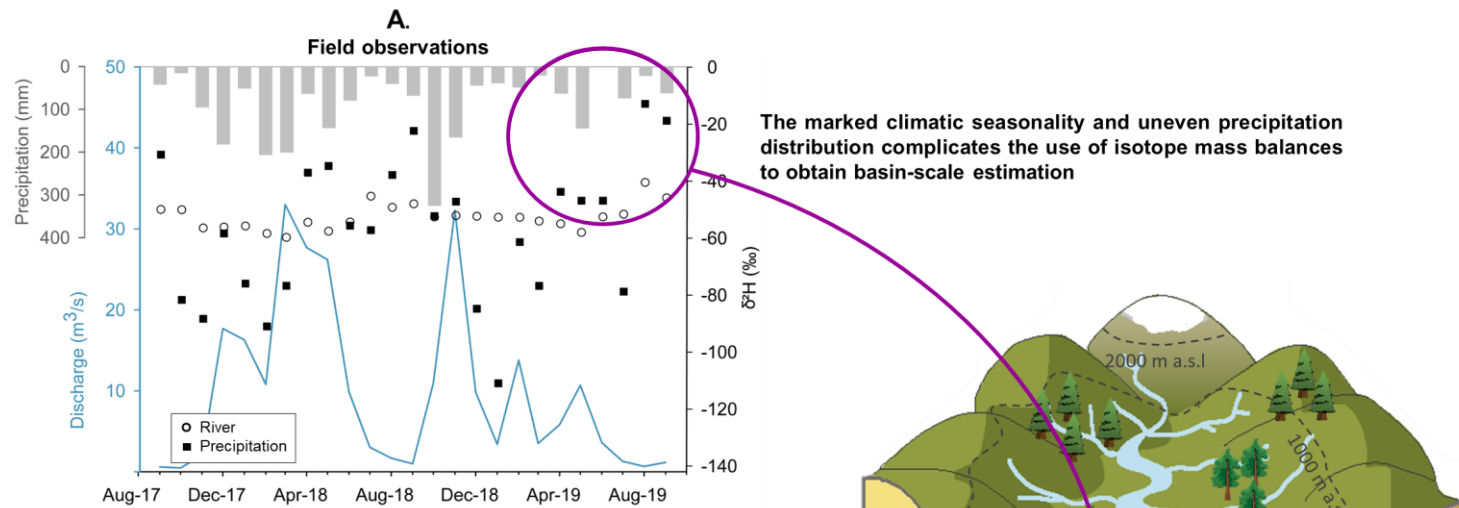


Basin-scale evaporation rate estimation based on isotopic approach: adaptation to the Mediterranean conditions

Alexandra Mattei, Frédéric Huneau, Emilie Garel, Sébastien Santoni, Yuliya Vystavna

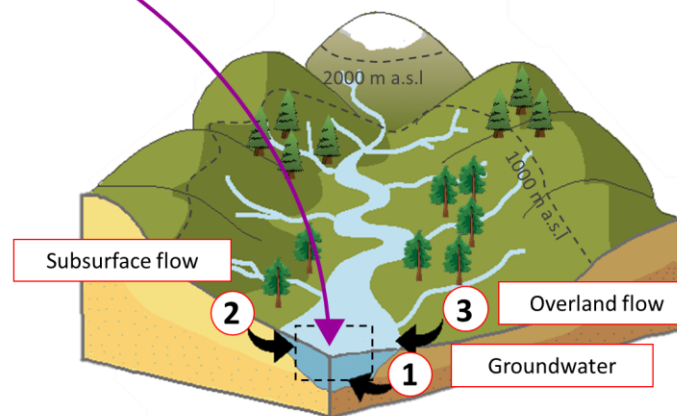
*

Navigate through the presentation by clicking on purple framed objects



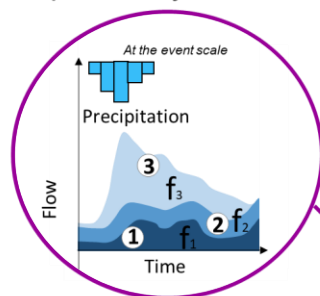
Key findings

- In regions with marked climatic seasonality, various flow contributions characterized by different flow paths and residence times have contrasting isotope signatures. The **mixing of these water sources** in various proportions and depending on the hydrological wetness of the watershed **controls the isotopic composition of the river water**.
- Applying **isotope mass balance** approach in regions with a **marked climatic seasonality** can lead, as highlighted here in Mediterranean context, to a **huge misinterpretation of evaporation rate**.
- Including **seasonal dynamics of river contribution sources** in the isotope mass balance approach, as proposed with the developed modified mass balance approach, is necessary to obtain **reliable evaporation rate estimate** in marked climatic seasonality regions.



Evaporation is not the only process altering the seasonal isotopic composition in the river

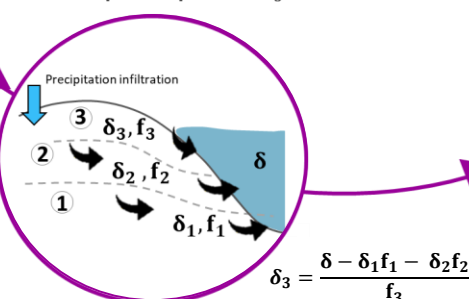
B. Quantifying reservoir mixing processes by EMMA



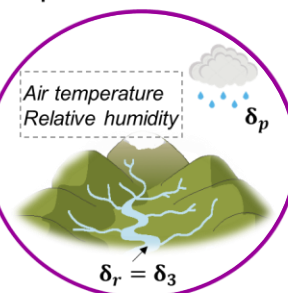
Contribution of the different water reservoirs f_1 , f_2 and f_3

- 1: Groundwater
- 2: Subsurface water (subsurface stormflow)
- 3: Surface water (Hortonian and saturated overland flow)

C. Determining overland flow isotope composition δ_3



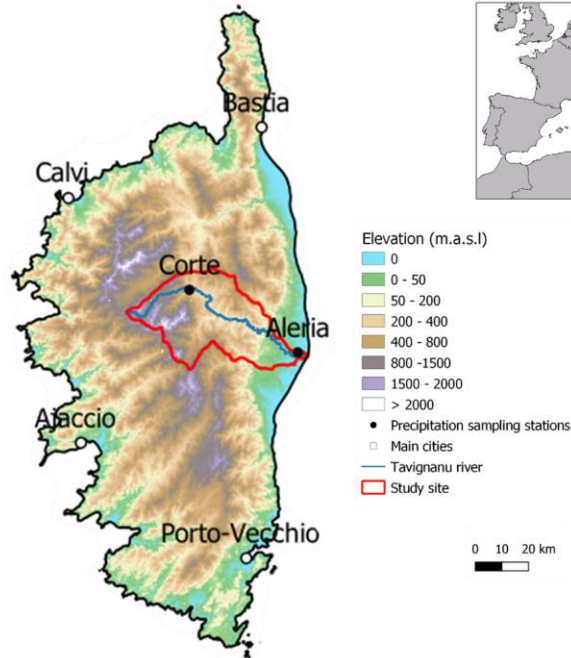
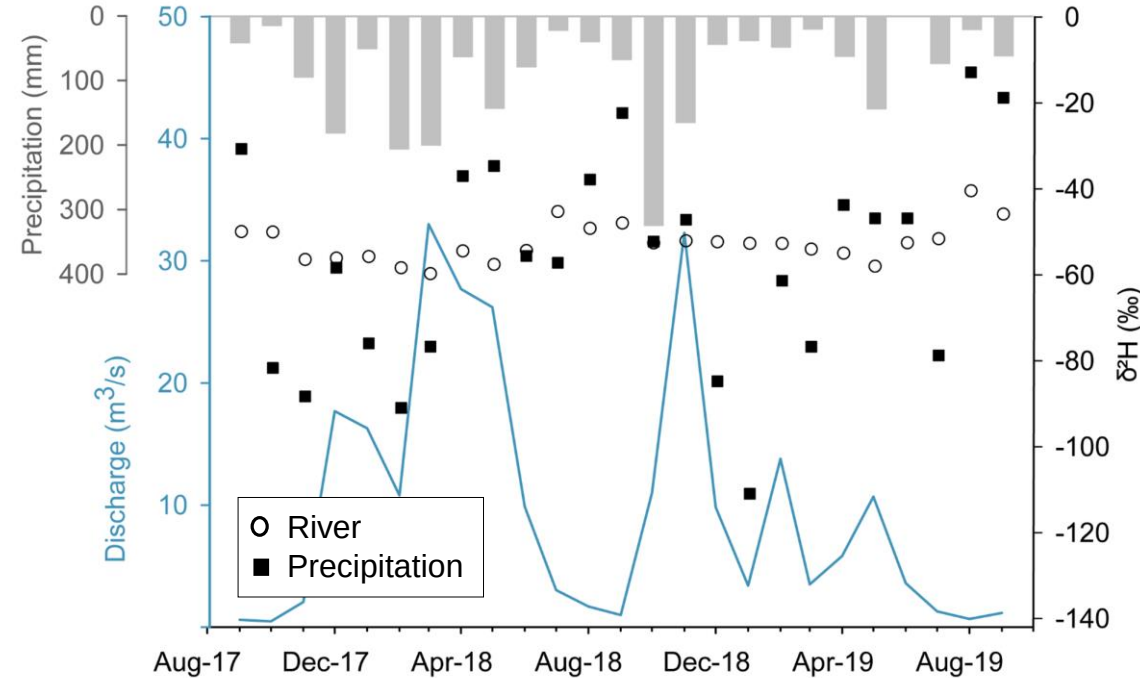
D. Estimating evaporation rate using isotope mass balance method



For more information read the published, full-length, article:

Mattei, A, et al. (2021). "Evaporation in Mediterranean conditions: estimations based on isotopic approaches at the watershed scale." Hydrological Processes. <https://doi.org/10.1002/hyp.14085>

Study site localisation

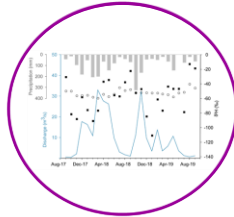
Monthly variations in precipitation and stream water $\delta^2\text{H}$, discharge and rainfall

- Precipitation pattern shows marked seasonal variations, following the thermal seasonality.
 - River water shows seasonal differences that are, in comparison to precipitation, lagged and highly damped.
- ⇒ **Important contribution of water reservoirs with longer residence times and more stable $\delta^2\text{H}$, integrating seasonally-variable precipitation inputs.**
- Spring and summer months values are ^2H -depleted compared to precipitation and relatively stable
- ⇒ **Important contribution of water reservoirs with even longer residence times, which integrate precipitation events over years**

Classical isotope mass balance approach is not reliable for estimating evaporation rates as overland flow is not the main year-round contributor of river discharge.

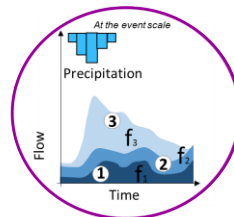
A.

Field observations

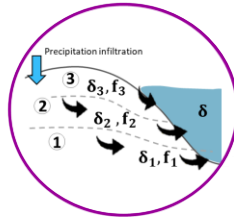


B.

Quantifying reservoir mixing processes by EMMA

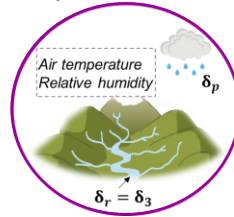


C.

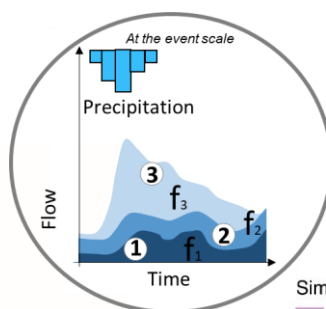
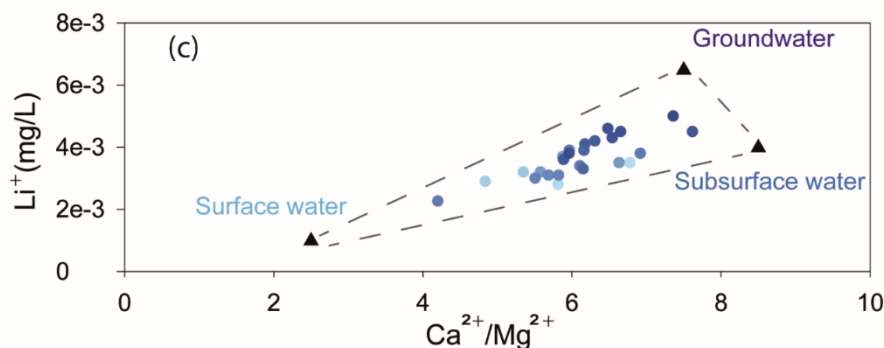
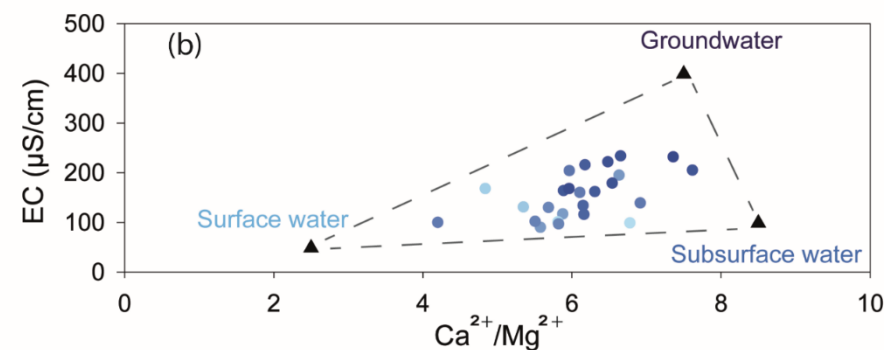
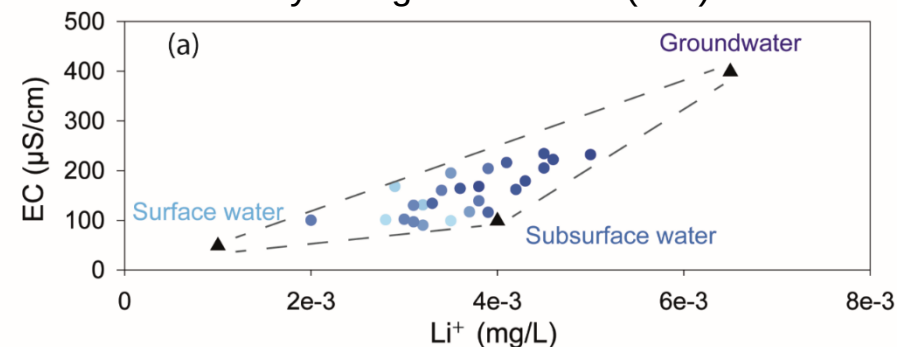
Determining overland flow isotope composition δ_3 

D.

Estimating evaporation rate using isotope mass balance method



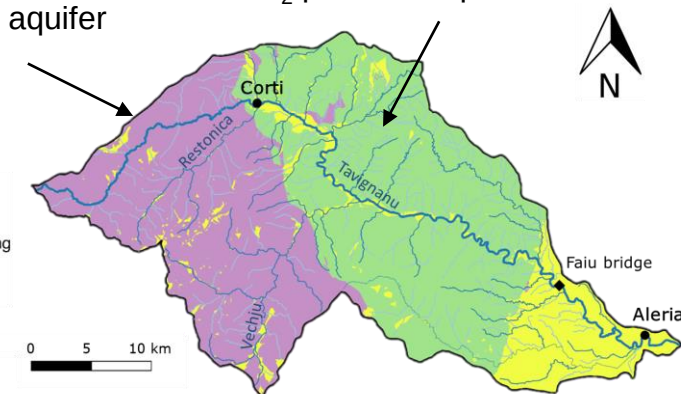
Plots of the conductivity and concentrations of Li^+ , Ca^{2+} , Mg^{2+} . The shade of blue illustrates the hydrological wetness (API).



Simplified geology

- Igneous rocks
- Metamorphic rocks carbonate bearing
- Clastic sedimentary rocks
- River sampling point

Simplified geology of the Tavignanu watershed

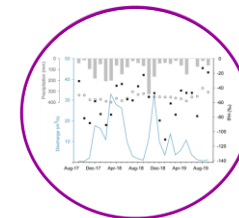
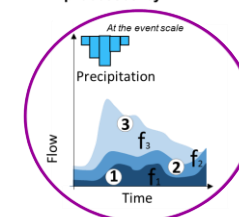
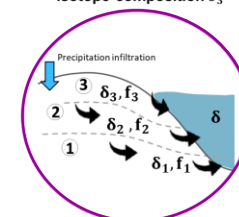
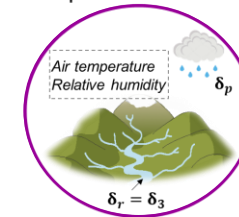
f₁ potential aquiferf₂ potential aquifer

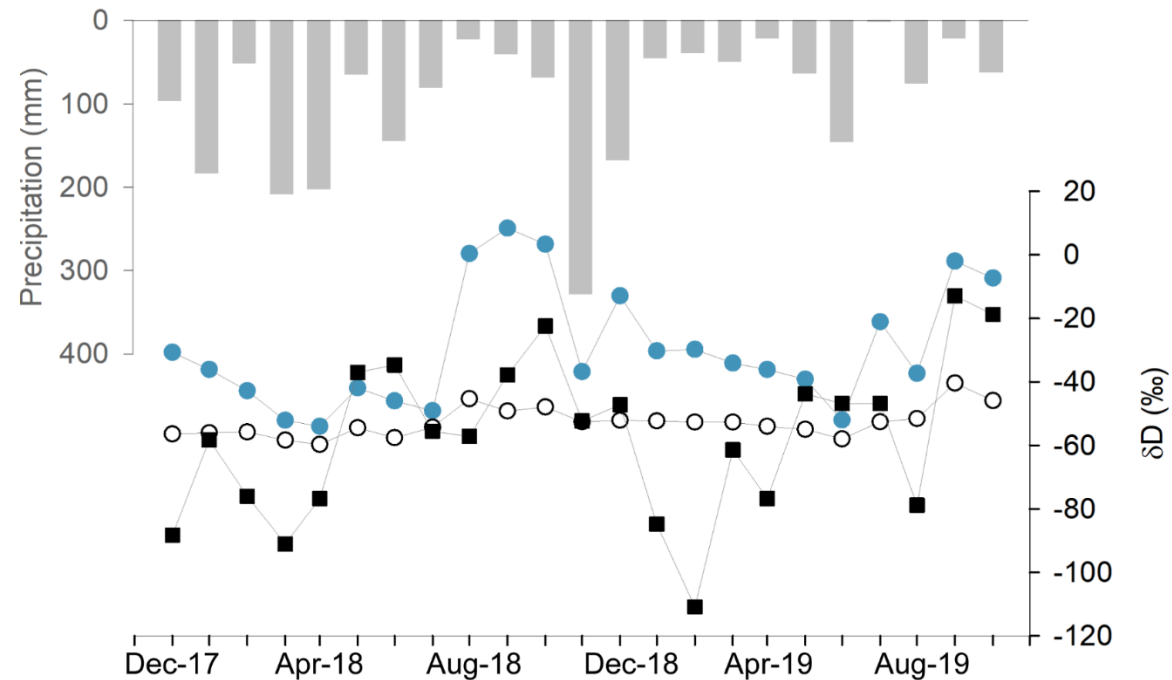
Mean annual contributions* (%)

	Surface (f ₃)	Subsurface (f ₂)	Groundwater (f ₁)
EC - Li ⁺	32	45	23
Ca ²⁺ /Mg ²⁺ - Li ⁺	35	37	28
Ca ²⁺ /Mg ²⁺ - EC	36	40	24
Mean	34	41	25
σ	2	4	3

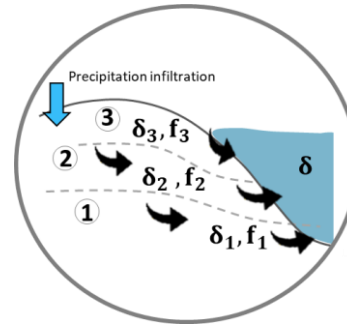
* Values obtained using End Member Mixing Analysis with three end members

The quantified water reservoir contributions obtained are in agreement with the first hypothesis about the hydrological functioning of the river discussed earlier

A.
Field observationsB.
Quantifying reservoir mixing
processes by EMMAC.
Determining overland flow
isotope composition δ₃D.
Estimating evaporation rate using
isotope mass balance method

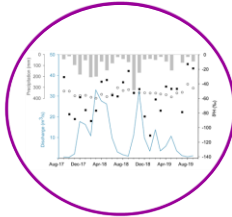
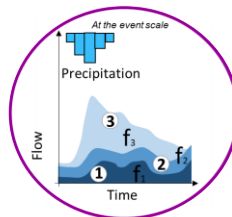
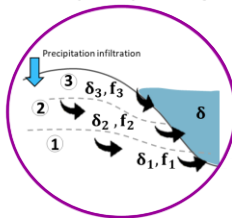
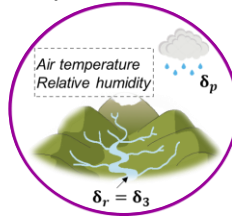
Monthly variations in precipitation, stream water, and surface water reservoir $\delta^2\text{H}$ 

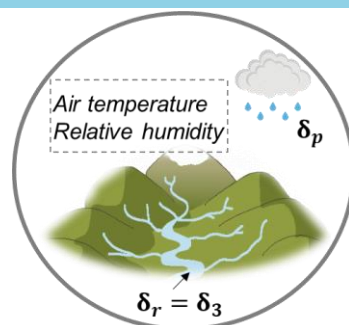
○ River
■ Precipitation
● Surface water reservoir

 $\delta^2\text{H}$ of the river (‰)

	Non modified approach	Modified approach
2017 - 2018	-56.5 ± 0.5	-45.6 ± 9
2018 - 2019	-53.0 ± 0.5	-31.1 ± 9

- In comparison to stream water, the surface water reservoir has a strong seasonal pattern, following the one of precipitation.
- The influence of evaporation during spring and summer months is evident in the surface water reservoir, this is not the case for stream water.

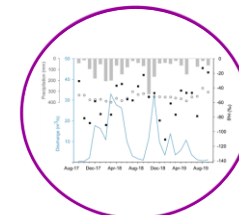
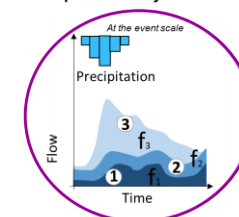
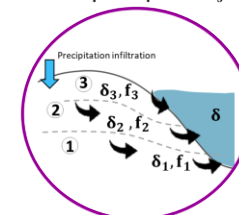
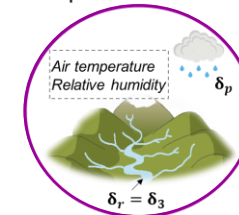
A.
Field observationsB.
Quantifying reservoir mixing processes by EMMAC.
Determining overland flow isotope composition δ_3 D.
Estimating evaporation rate using isotope mass balance method



Evaporation rate (%)

	T (°C)	h (%)	$\delta^2\text{H}_p$ (‰)	Non modified approach	Modified approach	Hydro-meteorological modelling approach
2017 - 2018	12.8 ± 1	74 ± 5	-63.7 ± 1	10 ± 14	40 ± 15	40
2018 - 2019	13.2 ± 1	68 ± 5	-53.8 ± 1	1 ± 14	46 ± 15	50

- The non-modified and modified approaches lead to very different estimates in evaporation rates
- Evaporation rates estimated using a non-modified isotope mass balance approach lead to inconsistent results
- Evaporation rates estimated using a modified isotope mass balance approach lead to reliable results
- Uncertainty analysis demonstrates that evaporation rates obtained with the modified approach are not altered compared to those obtained with the non-modified approach, as long as tracers chosen to determine water volume contributions are appropriate, i.e. with analytical uncertainty lower than the estimated value.

A.
Field observationsB.
Quantifying reservoir mixing processes by EMMAC.
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² CNRS, UMR 6134, SPE, BP 52, F-20250 Corte, France.



³ International Atomic Energy Agency, Isotope Hydrology Section, Vienna International Centre B.P. 100, 1400 Vienna, Austria.



⁴ Biology Centre CAS, Institute of Hydrobiology, Na Sádkách 7, České Budějovice, 37005, Czech Republic.