Kwanza Basin

## **1.** Introduction

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•Rationale: Strike slip faults are a prominent tectonic feature in Earth to accommodate horizontal and/or oblique slip, commonly forming on plate boundary or salt-bearing slope settings.

•**Problem**: Little is still known about natural examples of three-dimensional geometry and growth of strike-slip faults.

•Our study document the two-and three-dimensional geometry to illustrate evolution of salt-detached strikeslip faults in the Outer Kwanza Basin, Offshore Angola.



Figure 3. Albian variance map illustrating spatial relationship between strike-Figure 2. (a) Base-salt depth and (b) Isopach map illustrating sub-and salt structural framework. Both slip and other supra-salt faults maps are overlaid (c) to illustrate spatial relationship between strike-slip faults and base-salt reliefs.





The salt isopach and Albian On 2D profiles, the variance map displays five left faults are character--lateral, NE striking, strike-slip ised by a planar, norfaults (Figs 2b and 3). The faults bound salt anticline, connects and cut the salt wall into several segments.

mal and/or reverse offset and, locally, a negative flower structure (Fig. 4).

Figure 5. Isopach thickness map of (a) Albian, (b) Eocene, (c) Late Miocene-Seabed illustrating changes of subsidence and accommodation that may relate to growth of the strike-slip fault The Overburden maps show that the F1 and F4 accumulated its near final length early, in Albian and, relatively late in Miocene respectively (Fig. 5). Albian and Eocene isopach maps illustrate lateral separation on and/or difference of depocenters along strike-slip faults (Fig. 5a, b). The Late Miocene-seabed isopach maps show that the strike-slip fault growth was influenced by growth of salt walls (Fig. 5c).

# Three-dimensional geometry and growth of salt-detached strike-slip faults, Outer Kwanza Basin, offshore Angola Aurio Erdi<sup>1,2</sup>, Christopher A-L. Jackson<sup>1</sup>

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fault are oriented par-





2. Methodology

a) Horizon displacement

•The vertical throw-distance are characterized by 3-to-10, now hard-linked sub segments that has polarity reversal (i.e. reverse and normal) along distance (Fig. 6b, d).

•The Tz profiles generally record throw maxima at the top of Albian and Eocene, and/ or locally Early Miocene, (Fig. 6e, f), indicating nucleation in Albian and reactivation in Eocene or Miocene age. The profile also record low throw gradient in Eocene interval, reflecting propagations in Eocene age.

 The EI show complex patterns of relationship with Tz profiles, suggesting the lateral separation and/or differences of depocenters along strike of strike-slip faults. Still, at the Miocene-Seabed, the EI>1 is consistent with decreasing of Tz, reflecting grabennfluenced by growth of the salt walls (Fig. 6e, f).



Figure 1. Loca-(after Erdi and Jackson, 2021)



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