Nonlinear forced change and nonergodicity: The case of ENSO-Indian monsoon and global precipitation teleconnections

#### <u>Gábor Drótos</u>, Tamás Bódai, Kyung-Ja Ha, June-Yi Lee, Eui-Seok Chung

Instituto de Física Interdisciplinar y Sistemas Complejos (CSIC-UIB), Palma de Mallorca, Spain MTA-ELTE Theoretical Physics Research Group, Budapest, Hungary

vEGU21 30th April 2020









- ENSO and Indian monsoon: important teleconnection
- In the MPI Grand Ensemble, Bódai et al. (2020) found
  - o increasing teleconnection strength on the long term,
  - a slight drop after 2000

between traditional ENSO indices and the all-Indian summer monsoon rainfall (AISMR)

• Are the findings robust? Spatial aspects?



Max Planck Institute Earth System Model, MPI Grand Ensemble: 63 members from 1880

(we discard 37 members out of 100 due to incomplete spin-up, and 30 years from 1850 to lose correlations between members  $\rightarrow$  natural probability measure of a snapshot/pullback attractor (Bódai and Tél, 2012; Drótos et al., 2015; 2017; Ghil et al., 2008; Chekroun et al., 2011; Romeiras et al., 1990))



Q: radiative forcing, historical+RCP8.5 (Meinshausen et al., 2011)



- ENSO: JJA mean sea surface temperature
- Indian monsoon: JJAS mean precipitation

In each year and each ensemble member



Pearson's correlation coefficient  $r(\Phi, \Psi)$  evaluated

- in each year
- with respect to the ensemble

between  $\Phi$  (ENSO-side quantity) and  $\Psi$  (IM-side quantity)



Φ	Ψ
Niño3	AISMR
Niño3.4	AISMR
PC of EOF1	AISMR
PC of EOF1	PC of EOF1
PC of MCA1	PC of MCA1
PC of CCA1	PC of CCA1
PC of EOF2	AISMR
PC of EOF2	PC of EOF2
PC of MCA2	PC of MCA2
PC of CCA2	PC of CCA2
PC of single CCA mode with AISMR	AISMR

where ...



#### ...where

- PC: principal component
- EOFn: nth empirical orthogonal function
- MCAn: nth mode of maximum covariance analysis
- CCAn: nth mode of canonical correlation analysis

Everything evaluated with respect to the ensemble! (With a little smoothing, such that a conservative estimate results for r.)



Beyond  $r(\Phi, \Psi)$ : TFSTVE: total fraction of spatio-temporal IM variability explained by the PCs of the Pacific-side MCA modes:

$$\text{TFSTVE} = \frac{\sum_{n=1}^{N} \sigma_n^2 r_n^2}{\sum_{n=1}^{N} \sigma_n^2},$$

where

- r<sub>n</sub> is evaluated between the PCs of MCAn
- $\sigma_n^2$  is the variance of the PC of MCA*n* on the IM side
- N is the number of MCA modes





- Teleconnection: mainly between the first modes
- Trends are difficult to see

Niño3-AISMR

Mann–Kendall test for each subinterval: H0: no monotonic trend  $\rightarrow$  colour: significant trend (5%) in the given subinterval



#### Niño3.4-AISMR



Mann–Kendall test for each subinterval: H0: no monotonic trend  $\rightarrow$  colour: significant trend (5%) in the given subinterval



#### EOF1-AISMR

EOF1-EOF1



Mann–Kendall test for each subinterval: H0: no monotonic trend  $\rightarrow$  colour: significant trend (5%) in the given subinterval



MCA1

CCA1





Mann–Kendall test for each subinterval: H0: no monotonic trend  $\rightarrow$  colour: significant trend (5%) in the given subinterval



TFSTVE

CCA-AISMR

- Long-term increase in strength confirmed
- Nontrivial further details, including a drop after 2000, but see Wilks (2016)

 $\rightarrow$ 



Linear regression model:

$$\Psi = a\Phi + \xi,$$

where  $\xi$  is random noise

 $r=rac{1}{\sqrt{1+rac{1}{\left((a/\sigma_{\xi})\sigma_{\Phi}
ight)^{2}}}},$ 

where  $\sigma$  denotes standard deviation





- Increase in  $a/\sigma_{\xi}$  and mostly in  $\sigma_{\Phi}$  (exception: late 21st c.)
- An increase in a outcompetes that in σ<sub>ξ</sub>



## Blue ( $\alpha\beta$ < 1): the increase in $a/\sigma_{\xi}$ is more important than that in $\sigma_{\Phi}$

Niño3–AISMR

Niño3.4-AISMR





EOF1-EOF1

#### Blue ( $\alpha\beta < 1$ ): the increase in $a/\sigma_{\xi}$ is more important than that in $\sigma_{\Phi}$



Summary: main driver: coupling strength a

EOF1-AISMR

# Nonlinearity, nonergodicity



Nonlinearity (as observed as a drop in the teleconnection strength) implies nonergodicity (a nonzero expected difference between the correlation coefficients evaluated with respect to time and the ensemble) (but not the other way around!).  $\rightarrow$  Can we explicitly detect nonergodicity gridpoint-wise over the globe?

Results for an ad hoc test statistic: contour lines: significant nonergodicity (5%)



#### Conclusions



ENSO-Indian monsoon teleconnection strength in the MPI-GE:

- long-term increase
- a drop after 2000
- nonlinearity

Main driver: coupling strength

Nonergodicity over the globe

Confidence thanks to the ensemble, but in a model

See more in Bódai et al., Front. Earth Sci. 8, 599785 (2021)

Cf. EGU21-3904



- Bódai and Tél, Chaos 22, 023110 (2012)
- Bódai et al., J. Clim. 33, 2163–2182 (2020)
- Chekroun et al., Physica D 240, 1685–1700 (2011)
- Drótos et al., J. Clim. 28, 3275–3288 (2015)
- Drótos et al., Eur. Phys. J. Spec. Top. 226, 2031–2038 (2017)
- Ghil et al., Physica D 237, 2111–2126 (2008)
- Meinshausen et al., Climatic Change 109, 213 (2011)
- Romeiras et al., Phys. Rev. A 41, 784 (1990)
- Wilks, Bull. Am. Meteorol. Soc. 97, 2263–2273 (2016)