The continental record of the Eocene-Oligocene Transition in the Eastern Ebro Basin. Decoding the paleoclimatic signature from sediments and clay mineralogy.

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The context of the EOT.

The **Eocene-Oligocene Transition (EOT)** represents one of greatest shifts of the global climate as recorded in deep sea isotope records.

The impact of the EOT in the terrestrial realm has been associated to a dramatic faunal turnover ("Grande Coupure"), and cooling and aridification in some, but not all nonmarine records. It appears that the response to global cooling on continents presents a significant ———regional variability.

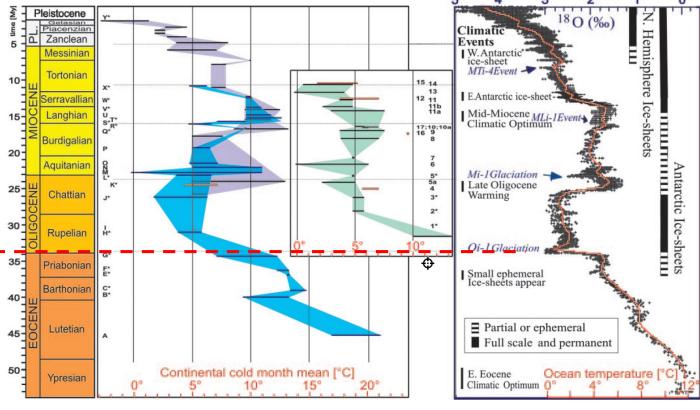


Fig. 3. Continental temperature curves (CMM) for Central Europe during the last 45 My in comparison with the global marine oxygen isotope record of Zachos et al. (1) adapted to the International Commission on Stratigraphy 2004 time scale (12). For code numbers and bar colors, see Fig. 2.



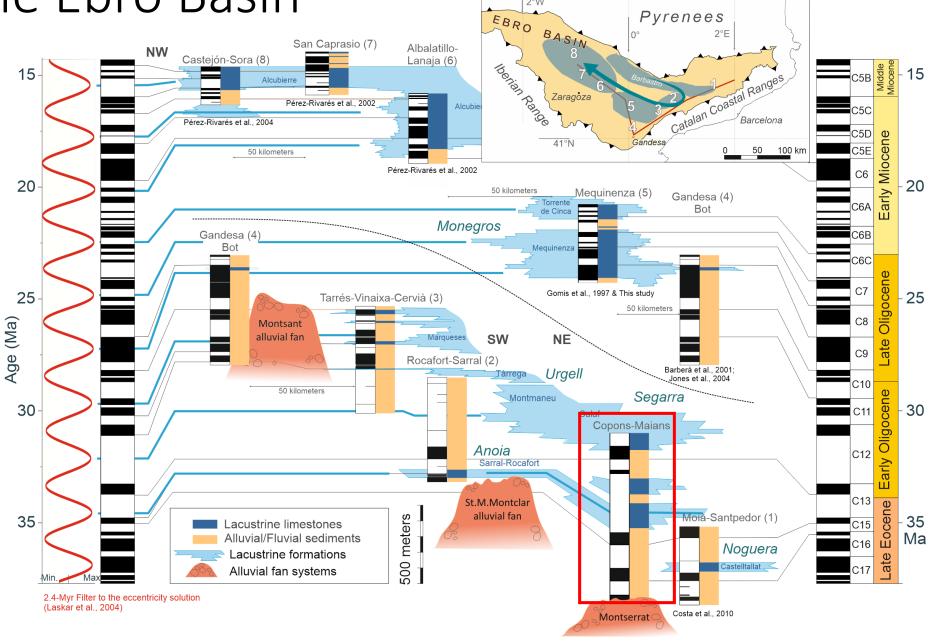
The EOT in the Ebro Basin

The Ebro basin underwent an internally drained basin since the late Priabonian. A rim of alluvial fans along its margins fed a central saline to fresh-water lake.

Cycles of lake expansion were correlated with times of 2.45myr long eccentricity màxima (Valero et al, 2014).

Study section



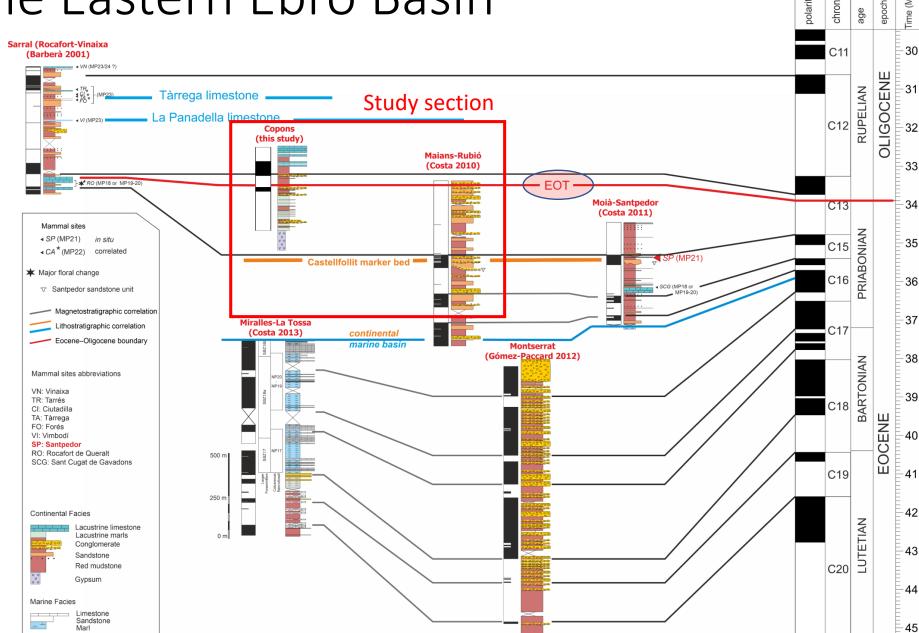


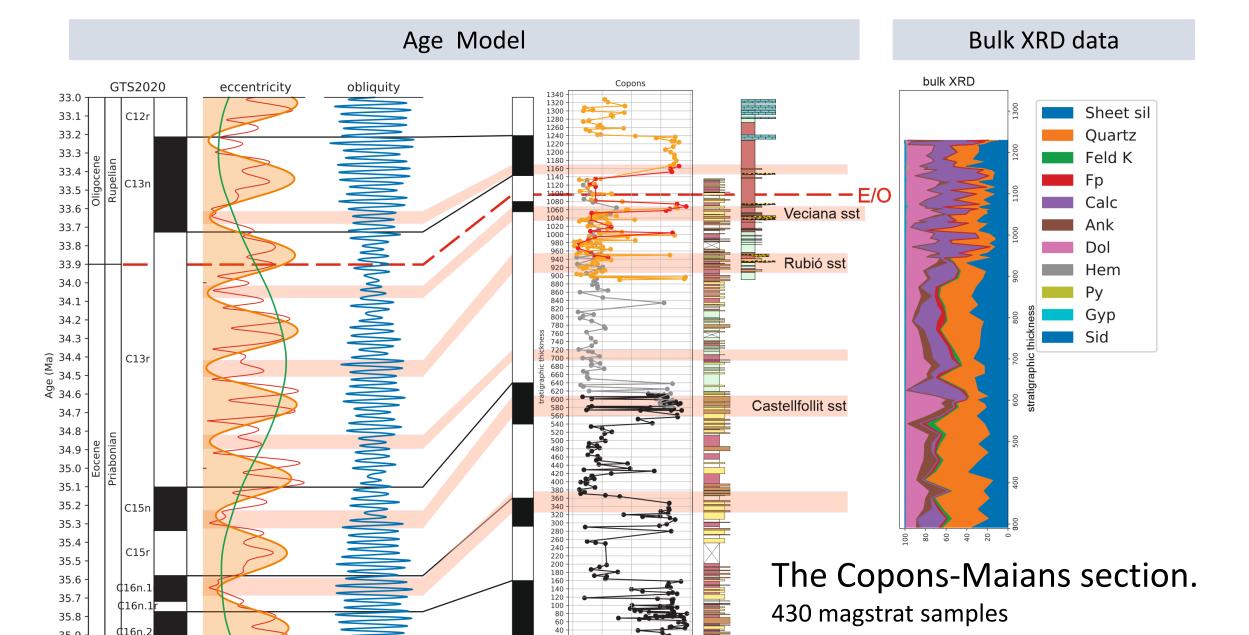
The EOT in the Eastern Ebro Basin

The **EOT** in the Eastern Ebro basin is now **magnetostratigraphically** constrained, located in a distal alluvial/shallow lacustrine succession.

This chronology challenges biostratigraphic studies (Vianey-Liaud et al. 2019) that suggest occurrence of "Oligocene faunas" in a lower stratigraphic position (→ see GarcesEGU2021, for further details)







-90

-45

VGP latitude

45

35.3 35.4

35.5

35.6

35.7

35.8

35.9

C15r

¢16n.1

¢16n.1r

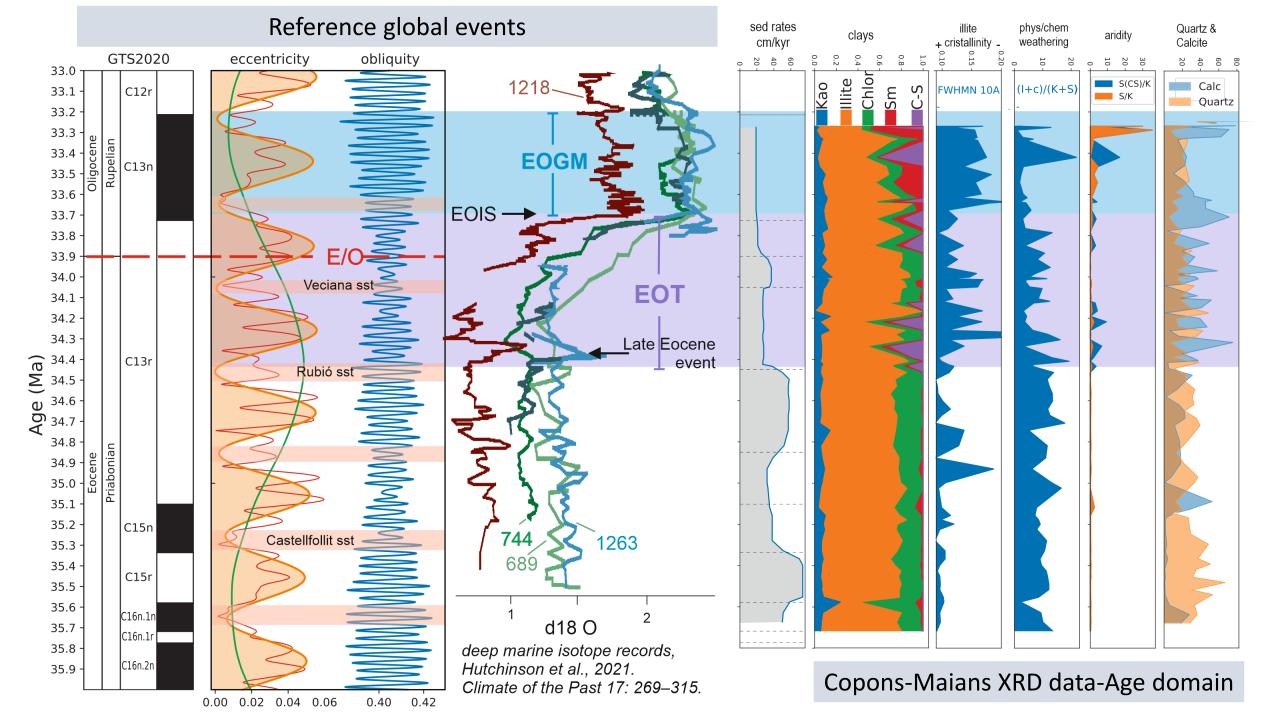
¢16n.2

0.00 0.02 0.04 0.06

0.40

0.42

The Copons-Maians section. 430 magstrat samples 75x3 XRD samples



Highlights on the EOT in the Copons-Maians section:

- Bulk composition shows:
 - Relative increase of detrital Calcite against Quartz → Increase of physical alteration, decrease dissolution of carbonate rocks. → aridity
- Clay composition
 - Independent from sedimentary facies. → Composition inherited from the source (detrital signal)
 - Increase of Smectite → formed in soils in the source under alternating wetting/drying conditions → increased seasonality.
- Two-step increase in Illite alteration (decreasing crystallinity):
 - starting at 34.4Ma (late Eocene event),
 - peaking during the EOGM (Glacial Maximum).
- Sediment accumulation trending to lowering rates (multiple causes to be discussed)



Conclusions.

Observations compatible with increasing cooling and aridity:

- Relative increase of detrital Calcite against Quartz
- Increase of Smectite against Kaolinite.
- Increase of aridity indexes (S/K), (S+CS)/K

Unanticipated observations (← seasonality?):

- Decreasing illite crystallinity (FWHMN 10A)
- Decreasing Physical against Chemical weathering index (I+C)/(S+K)

Tricky observations:



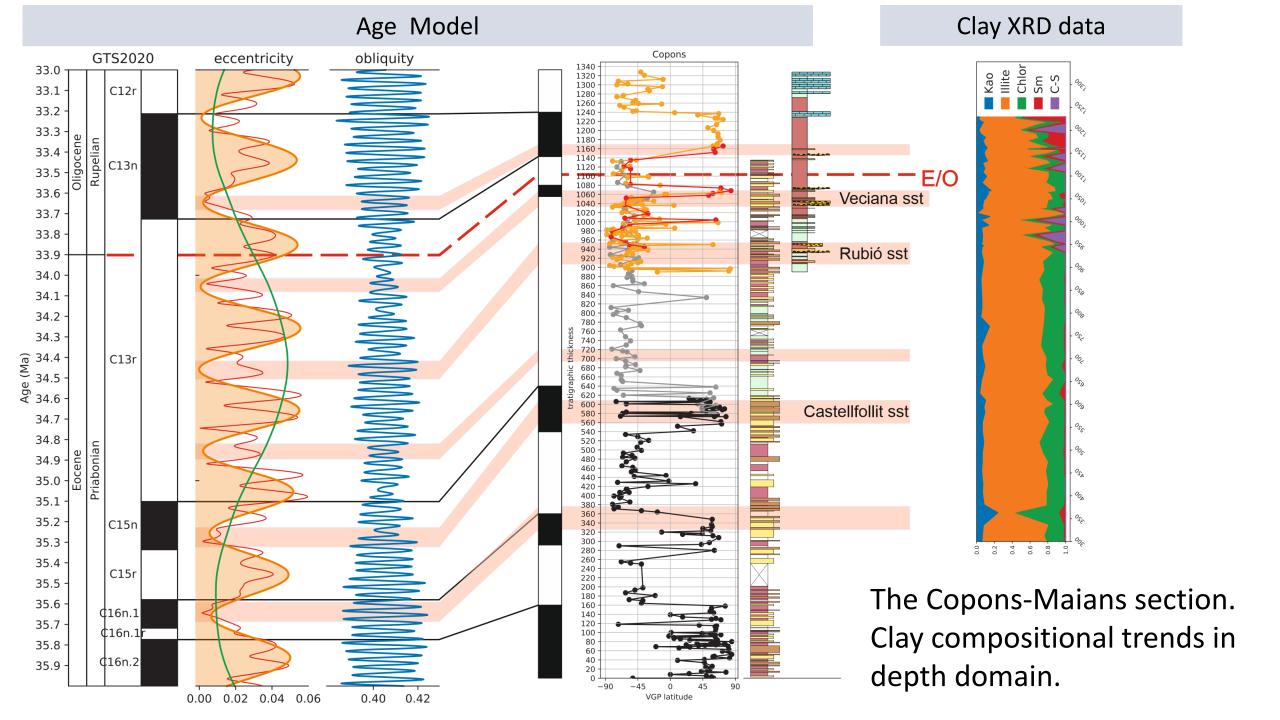
 Sediment accumulation trending to lowering rates (←lowering average Qw?)

Supplementary material

See next slides for additional material and figures from reference papers

- ☐ Clay compositional trends in depth domain
- Tiepoints for age model
- ☐ Clay composition correlation matrix
- ☐ Bulk composition correlation matrix
- ☐ Reference papers:
- Rennes Basin (Tramoy et al., 2016)
- ☐ The deep marine stable isotope record (Hutchinson et al., 2021)
- The EOT at Saint Stephens Quarry, Alabama (Katz et al., 2008)
- The EOT at Massignano section (Jovane et al., 2007)





Age constrains

- Tie points used for depth-to-age transformation.
- (ecc1, ecc2, ecc3: consecutive eccentricity mínima within C13r correlated with alluvial proggradion events)

	chron	age	level (m)
0	C16n.1r	35.774	161
1	C16n.1n	35.718	291
2	C15r	35.58	360
3	C15n	35.336	540
4	C13r	35.102	630
5	ecc1	34.85	710
6	ecc2	34.45	940
7	ecc3	34.05	1050
8	E/O	33.9	1070
9	C13n	33.726	1142
10	C12r	33.214	1240



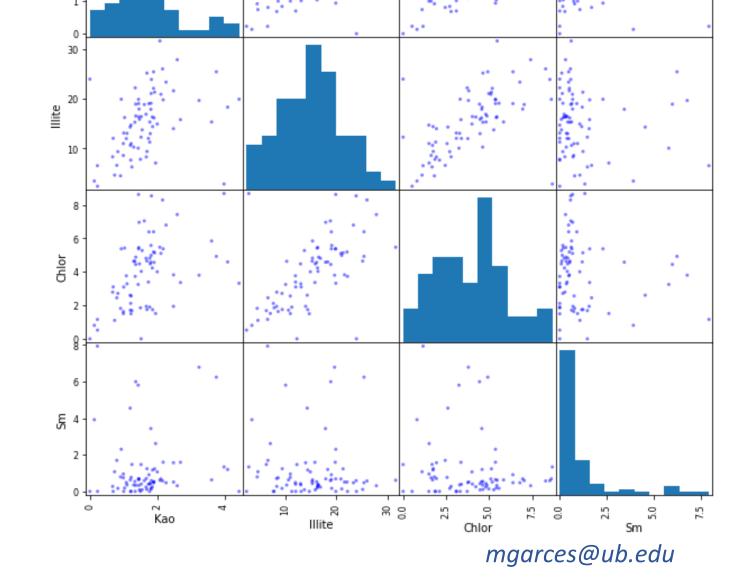
Clay composition Correlation matrix

Positive correlations:

Illite vs Chlorite
Illite vs Kaolinite (poor)

No correlation:

Smectite vs any other component





Bulk composition Correlation matrix

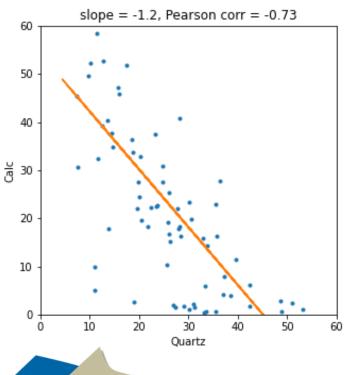
Only Calcite and Quartz reveal a negative correlation → Dissolution effect

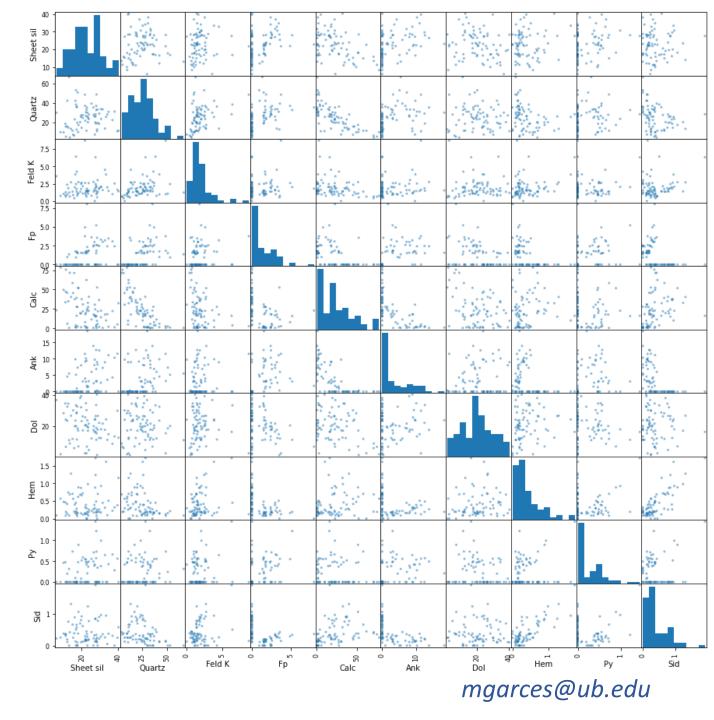
Geomodels

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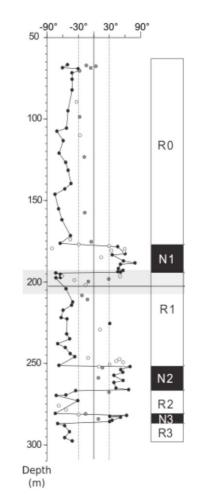


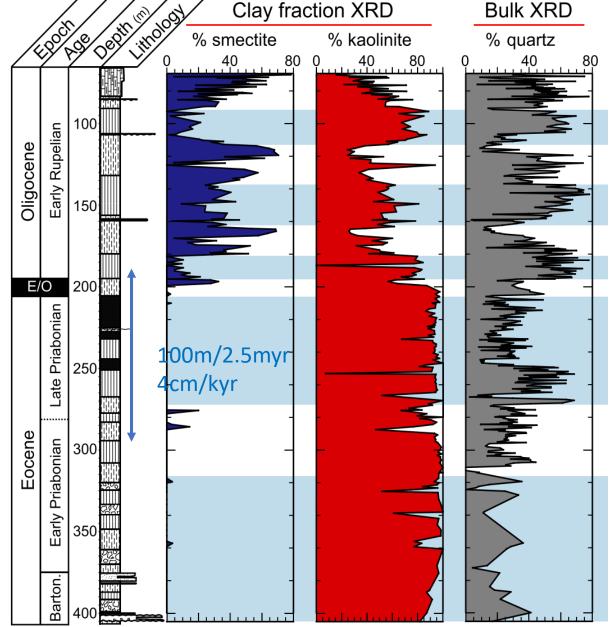


Rennes basin

CDB1 borehole

Tramoy, R., Salpin, M., Schnyder, J., Person, A., Sebilo, M., Yans, J., Vaury, V., Fozzani, J., and Bauer, H., 2016, Stepwise palaeoclimate change across the Eocene-Oligocene transition recorded in continental NW Europe by mineralogical assemblages and δ15Norg (Rennes Basin, France): Terra Nova 28: 212–220.

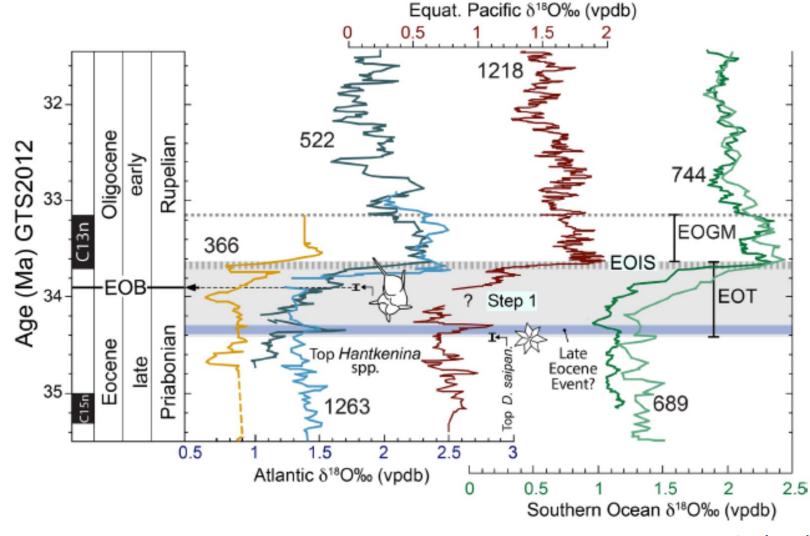






The deep marine stable isotope record

Hutchinson, D.K., Coxall, H.K., Lunt, D.J.,
Steinthorsdottir, M., De Boer, A.M., Baatsen, M.,
Von Der Heydt, A., Huber, M., Kennedy-Asser,
A.T., Kunzmann, L., Ladant, J.B., Lear, C.H.,
Moraweck, K., Pearson, P.N., et al., 2021, The
Eocene-Oligocene transition: A review of marine
and terrestrial proxy data, models and modeldata comparisons: Climate of the Past, v. 17, p.
269–315, doi: 10.5194/cp-17-269-2021.





Saint Stephens Quarry, Alabama

Position of "ghost subchron" C13r.1n between EOT-1 and the E/O boundary

C13r.1n

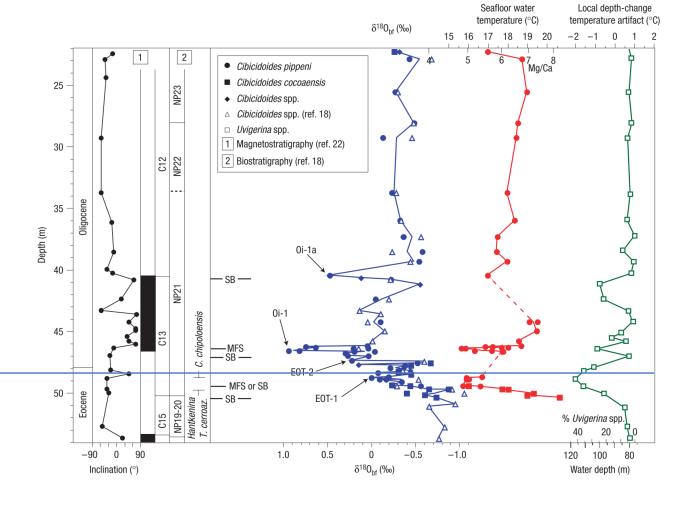


Figure 1 SSQ proxy data for ice volume, temperature and sea level. δ^{18} 0 and Mg/Ca were measured on benthic foraminifera (see Supplementary Information, Table S2). Multiple analyses of single samples are shown, connected with a line through average values. *Uvigerina* abundances were used to estimate palaeodepth and the local temperature change driven by water-depth changes (see the text). MFS = maximum flooding surface (deepest palaeodepth within a sequence). SB = sequence boundary



Katz, M.E., Miller, K.G., Wright, J.D., Wade, B.S., Browning, J. V., Cramer, B.S., and Rosenthal, Y., 2008, Stepwise transition from the Eocene greenhouse to the Oligocene icehouse: Nature Geoscience, v. 1, p. 329–334, doi: 10.1038/ngeo179.

The EOT at Massignano section

B11101

 Position of E/O boundary relative to "ghost subchron" C13r.1n

Jovane, L., Sprovieri, M., Florindo, F., Acton, G., Coccioni, R., Dall'Antonia, B., and Dinarès-Turell, J., 2007, Eocene-Oligocene paleoceanographic changes in the stratotype section, Massignano, Italy: Clues from rock magnetism and stable isotopes: Journal of Geophysical Research, v. 112, p. B11101



early Oligocene late Eocene GSSP marker 2 P1 LO hantkeninids [Coccioni et al., 1988] LO Cribrohantkenina inflata [Coccioni et al., 1988] P3 LO Globigerinatheka index [Coccioni et al., 1988] P4 FO Turborotalia cunialensis [Spezzaferri et al., 2002] LO Globigerinatheka semiinvoluta [Spezzaferri et al., 2002] P5 N1 LO Reticulofenestra reticulata [Monechi et al., 2000] N2 FO Cribrocentrum oamaroensis [Monechi et al., 2000] N3 FO Isthmolithus recurvus [Coccioni et al., 1988] K-Ar [Odin et al., 1988] Rb-Sr [Montanari et al., 1988] -87Sr-86Sr [Montanari et al., 1993] Rb-Sr and Ar-Ar [Premoli Silva and Jenkins, 1993] Sedimentary rate mean 0.73 cm/k.y. (7.3 m/Ma) N₂

JOVANE ET AL.: PALEOMAGNETISM AT EOCENE-OLIGOCENE GSSP

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