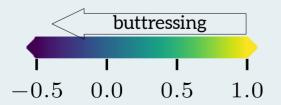


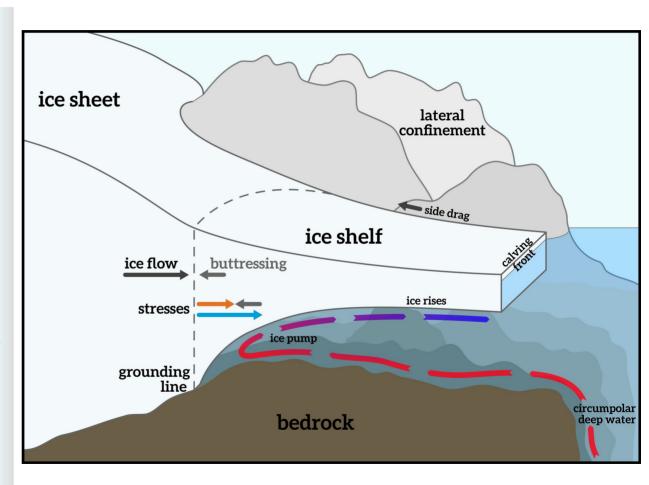
Why buttressing of ice shelves?

Weaker ice shelves lead to decreased longitudal back-stress and increased flux accross the grounding line

Buttressing ratio at GL

$$\Theta_N = \frac{\text{stresses with ice shelf}}{\text{stresses without ice shelf}}$$





Methods



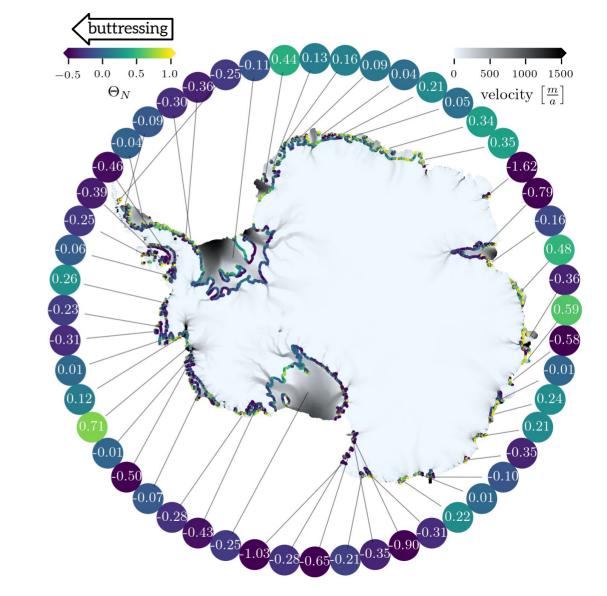
- transient simulations
- 8km resolution
- Spin-up: equilibrium close to present-day observations



- diagnostic mode
- > 200m resolution at GL
- Inversion to observed velocities

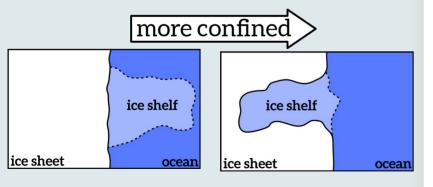
Present-day buttressing

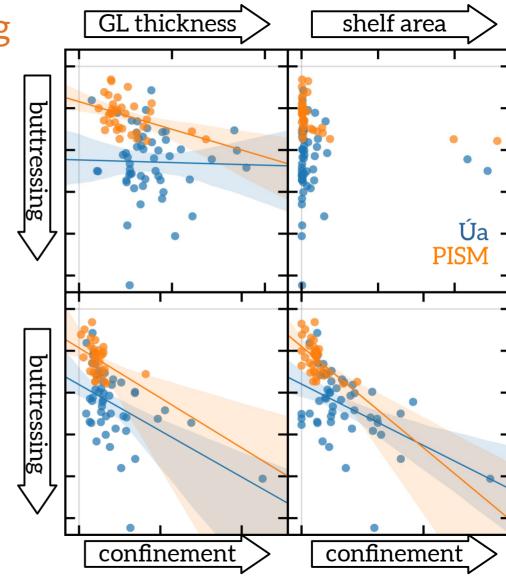
- aggregated shelf-wide buttressing values
- Strong buttressing in large ice shelves (e.g. Filchner-Ronne, Ross and Amery)
- Lowest buttressing in Thwaites glacier



Characteristics of Buttressing

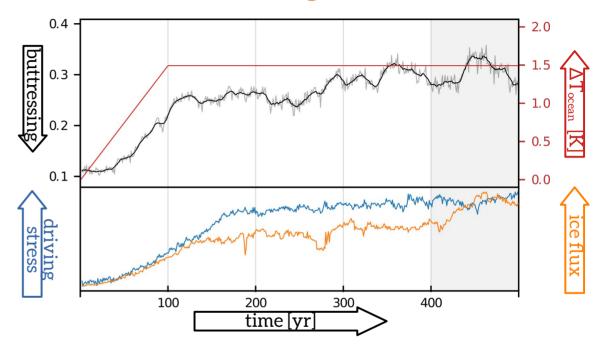
- Positive correlation for more confined ice shelves
- No correlation with shelf area, thickness, ice flux or driving stress





Transient buttressing under ocean warming

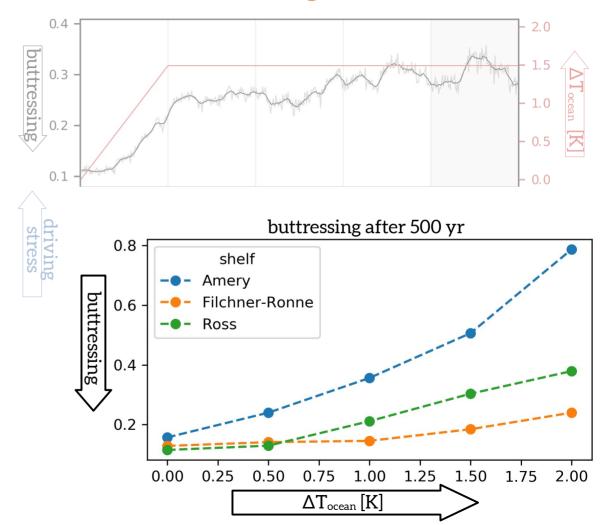
idealized ocean warming with melt rates from the boxmodel PICO leads to reduced buttressing and increased driving stress and ice flux



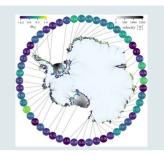
Transient buttressing under ocean warming

idealized ocean warming with melt rates from the boxmodel PICO leads to reduced buttressing and increased driving stress and ice flux

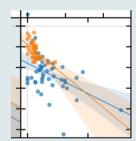
consistent decrease of buttressing after 500 years in different ice shelves across range of temperature offsets



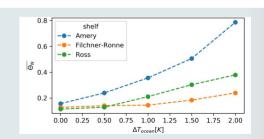
Shelf-wide buttressing measure allows to compare different ice shelves and transient evolution



Buttressing is higher in more confined ice shelves



Buttressing decreases under idealized warming accross several temperatures







Tipping Points in Antarctic Climate Components



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