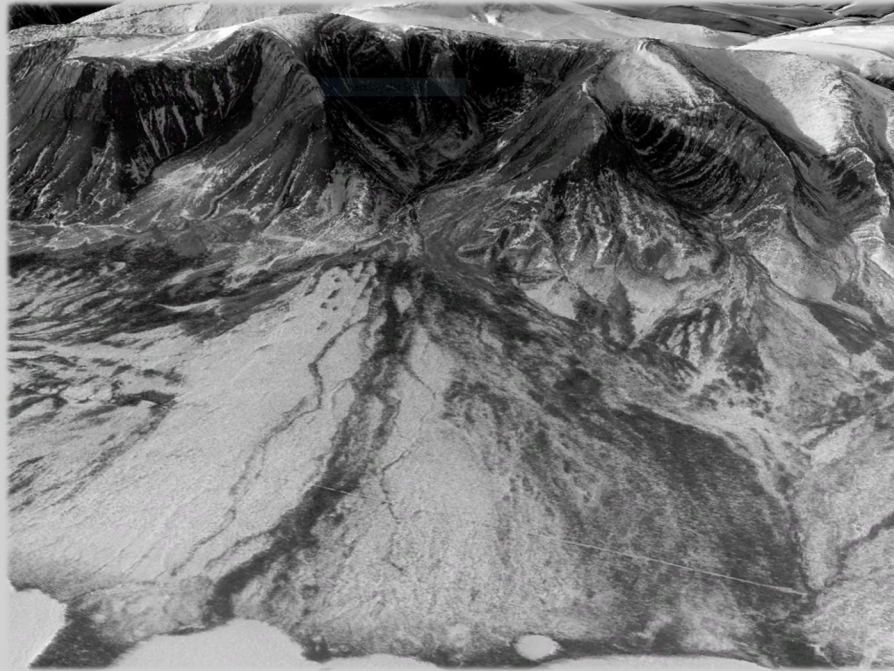
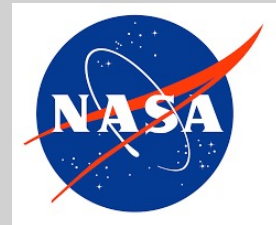


# Sediment transport processes in an arctic watershed undergoing climate change



**Marisa C. Palucis, Justin V. Strauss, and Jill Marshall**



# Motivation



Rowland et al., 2010

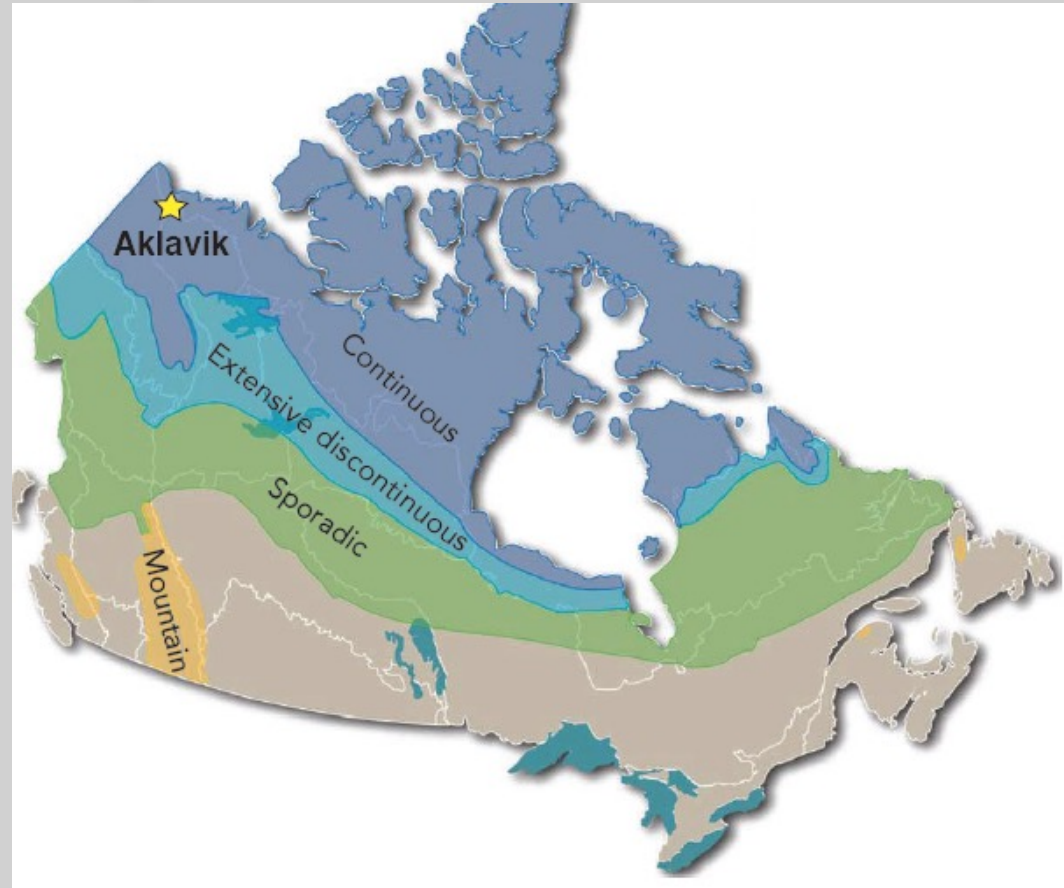
- Arctic warming  $\sim 2\times$  faster than temperate landscapes
- Degradation of permafrost will change water and sediment fluxes (and processes)
- Carbon, nutrients, and trace elements will follow similar pathways

# Want to quantify water and sediment fluxes to an alluvial fan (closed system) in an arctic catchment

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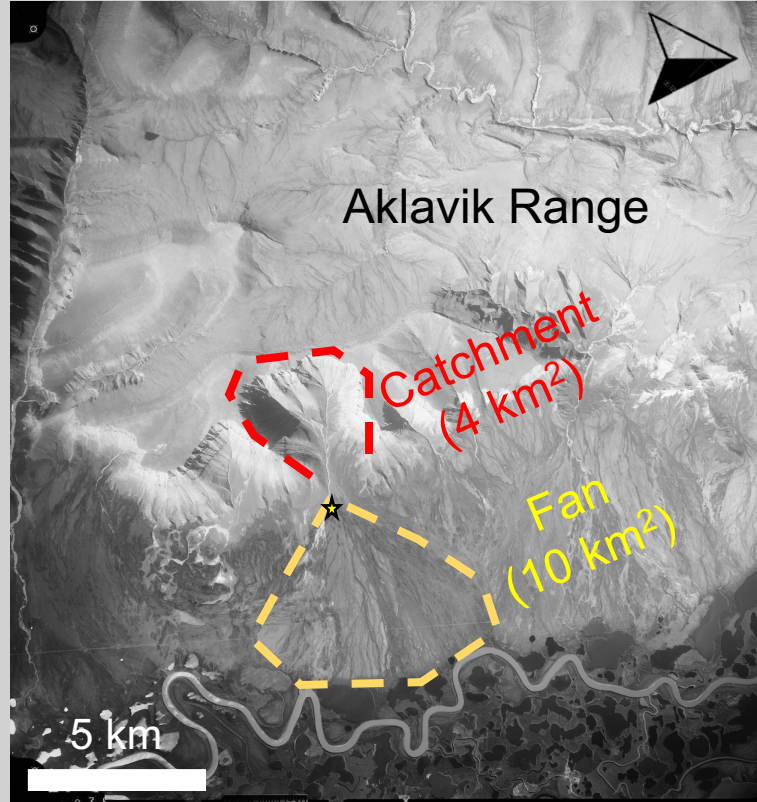
- What are the dominant processes that build periglacial fans?
  - Low rock-to-water fluvial events
  - High rock-to-water debris flow events
- What are typical flow discharges, grain sizes, and sediment fluxes?
- How does this fan compare to temperate systems?

# Field site: Richardson (Aklavik) Mountains (NWT, Canada)





# Field site: Black Mountain Fan



# Field Site: Black Mountain Fan

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- **Climate:** semi-arid with cold and dry winters and short summers (12°C mean temp in June and -35°C mean temp in January), mean annual precipitation is ~200-250 mm (>50% snow)

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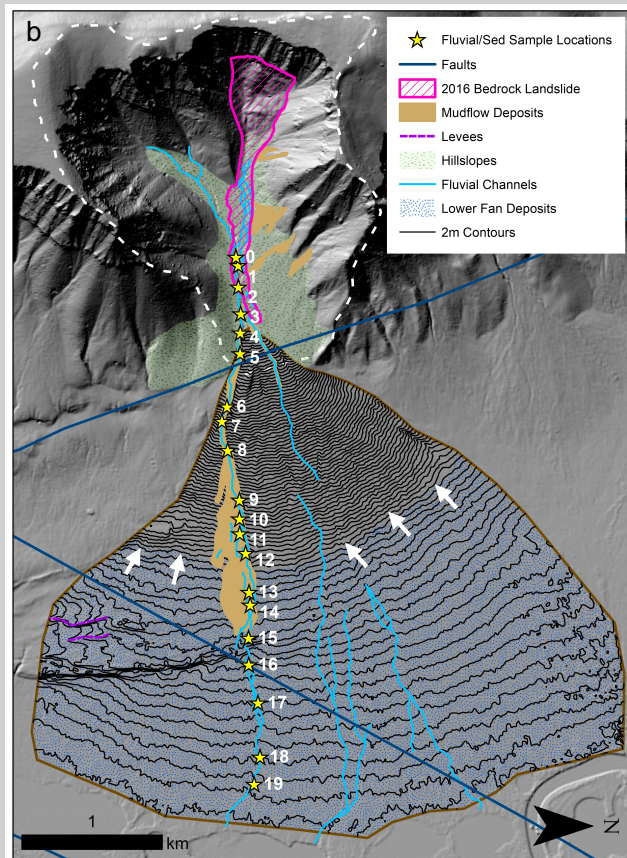
# Field Site: Black Mountain Fan

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- **Geology:** source material for fan is Cretaceous (~113-133 Ma) shale, siltstone, and sandstone; small amounts of sediment are glaciogenic (westward advance of Laurentide ice sheet)
- **Geomorphology:** fan mapped in the 1990s to be a debris flow dominated fan with periglacial modification



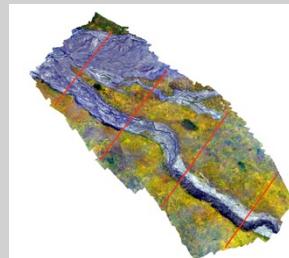
# Summer 2019 Field Campaign



Fan sedimentology and geomorphology



Monitor active flow event



Fan topography

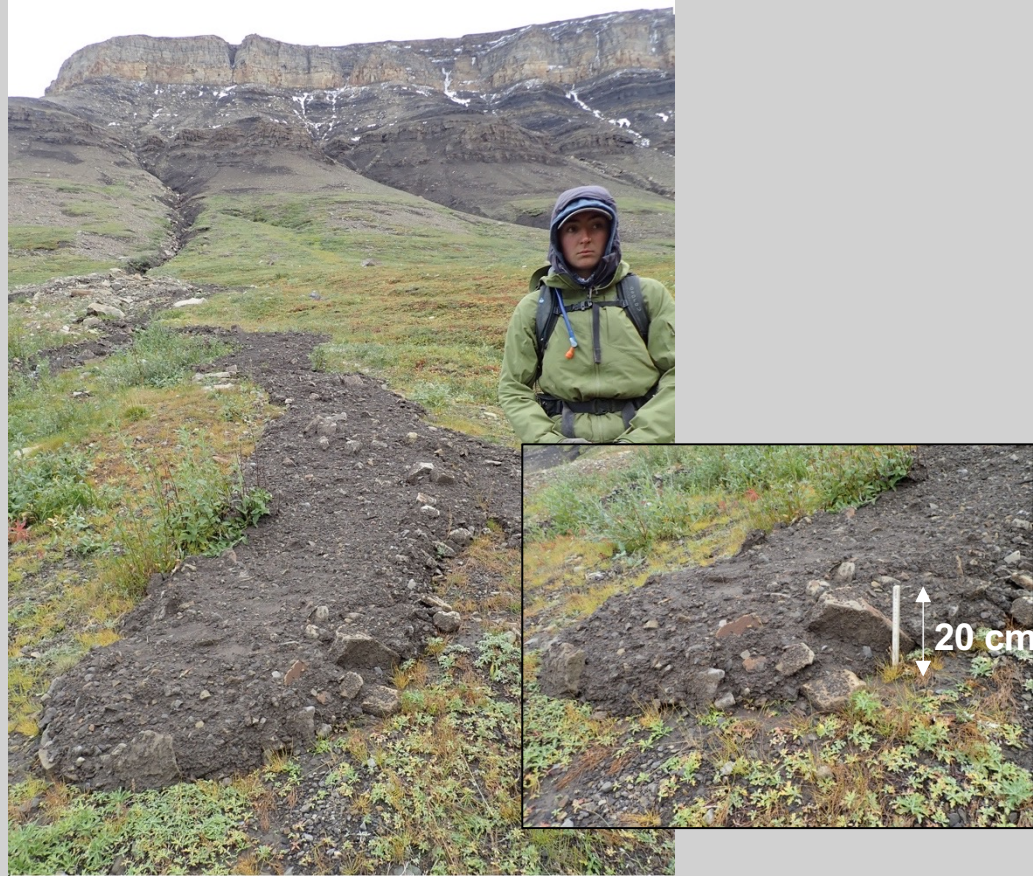
# Results

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- What are the dominant processes that build periglacial fans?
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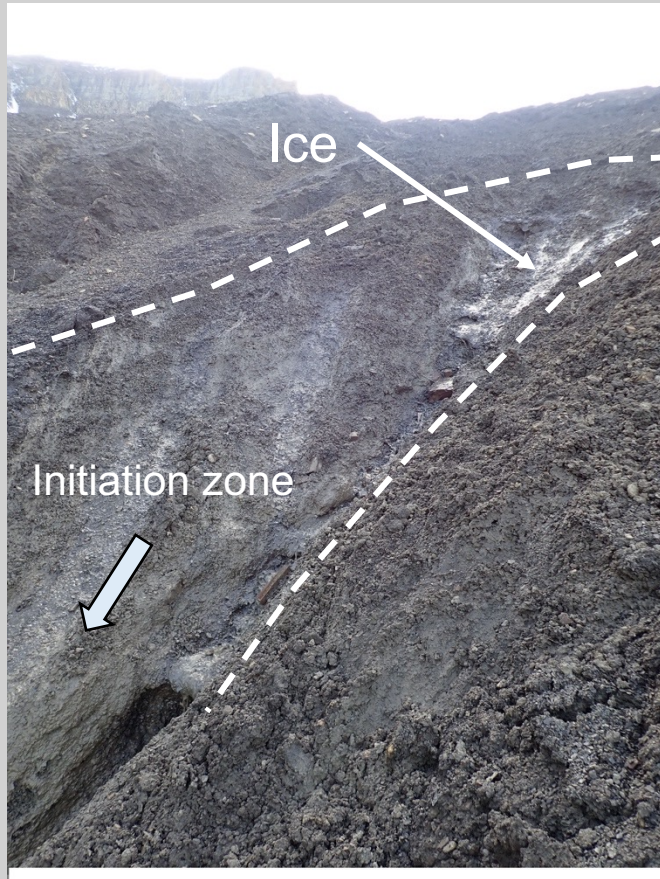


# Results: debris flow event from icy gully





# Results: debris flow event from ice-melt-induced landslide





Results: Lower fan had active fluvial channels across it



# Results: mainly debris flow deposits within incised channel walls of upper fan



- Where we have good exposures, clasts tend to be sub-angular to sub-rounded, poorly sorted, no imbrication or sedimentary structures, occasionally reverse grading

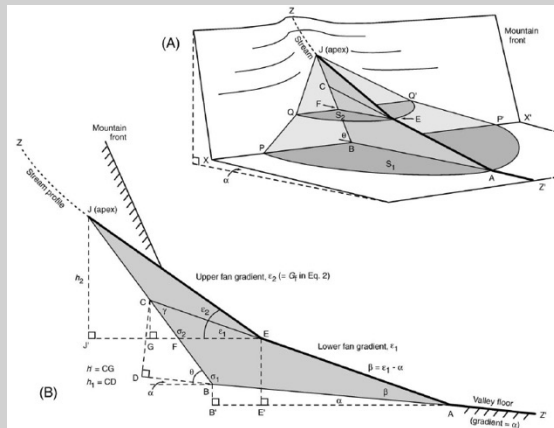


# Results: fluvial processes dominate lower fan

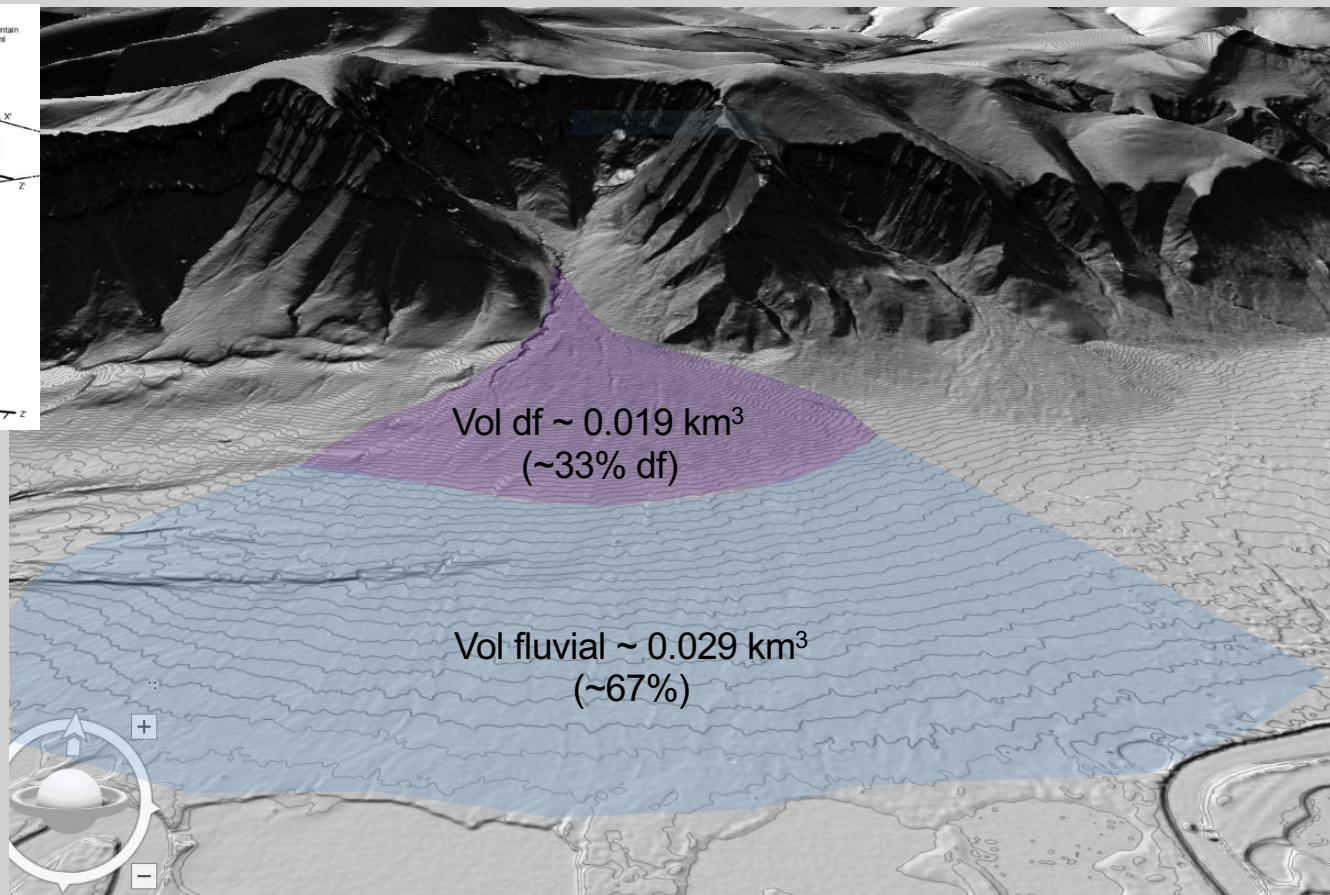


- Alternating layers of silt and clay
- Layering is ~5 cm thick, with silt often in lenses with imbricated pebbles
- Lower fan deposits tend to be overlain with ~0.5 to 0.75 m of soil/tundra

# Results: % debris flow versus fluvial

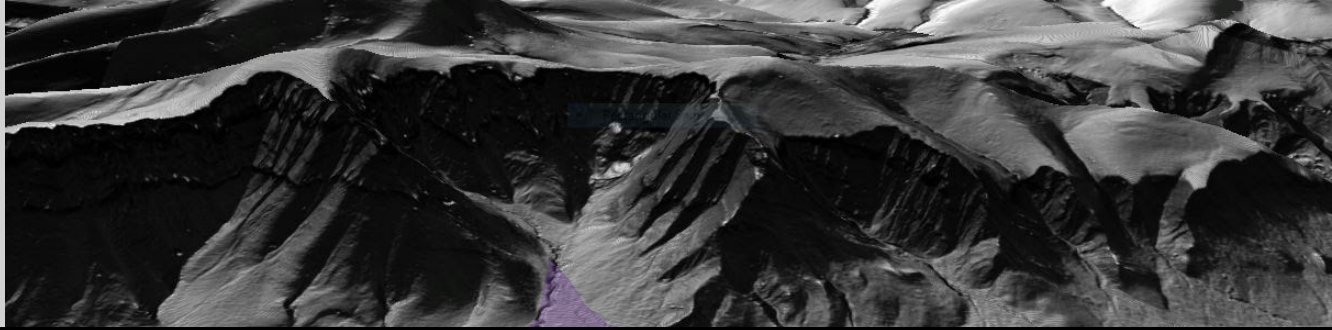


Giles et al. (2010)

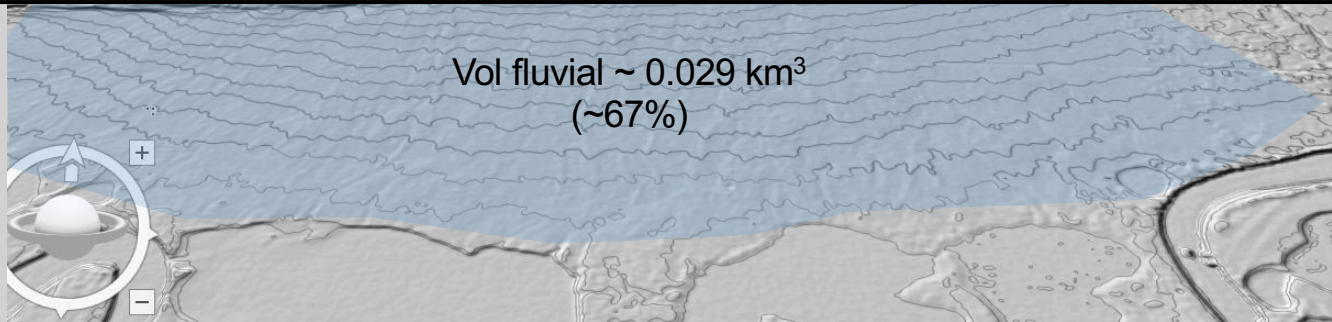




## Results: % debris flow versus fluvial



Compared to earlier work, surface processes and sedimentology suggest that much of the fan is build by fluvial processes, not debris flows, and hence we might be able to use standard sediment transport equations to estimate fluxes...

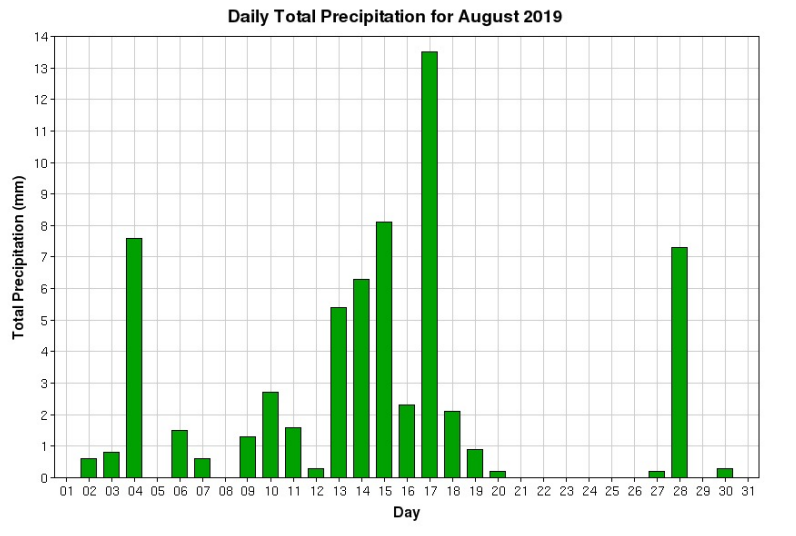


# Results

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- What are the dominant processes that build periglacial fans?
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- What are typical flow discharges, grain sizes, and sediment fluxes?
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# Results: Observed a rain/snow event on Aug 17 2019



- ~14 mm of rain and snow occurred over several hours on the evening of the 17<sup>th</sup>
- Max rate ~ 1 - 2 mm/hr
- Snow filled the gullies within the catchment; rain occurred lower on the fan (at camp)

# Results: Discharge (Q)

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$$(1) \quad U = \sqrt{gHS} \frac{6.5(\frac{H}{D_{84}})}{[(\frac{H}{D_{84}})^{5/3} + 6.76]^{1/2}}$$

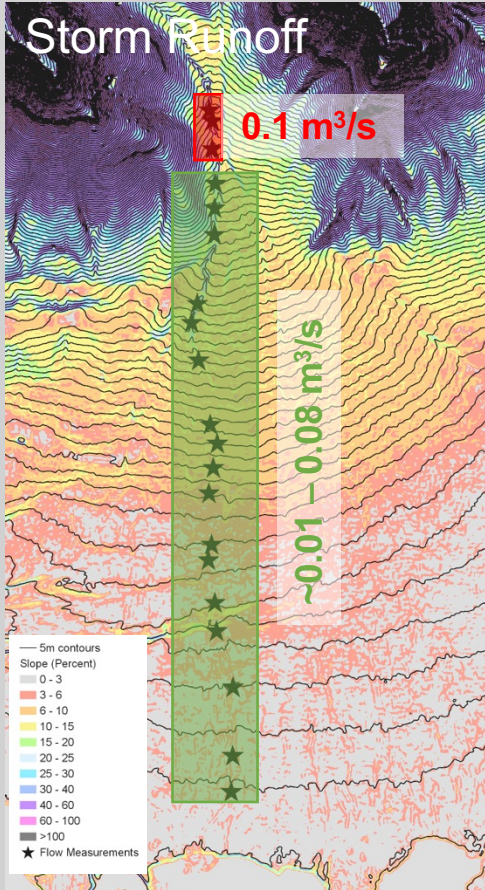
$$(2) \quad Q = \text{discharge} = UWH$$

Variable power flow resistance equation (VPE) (Ferguson, 2007 and 2012)

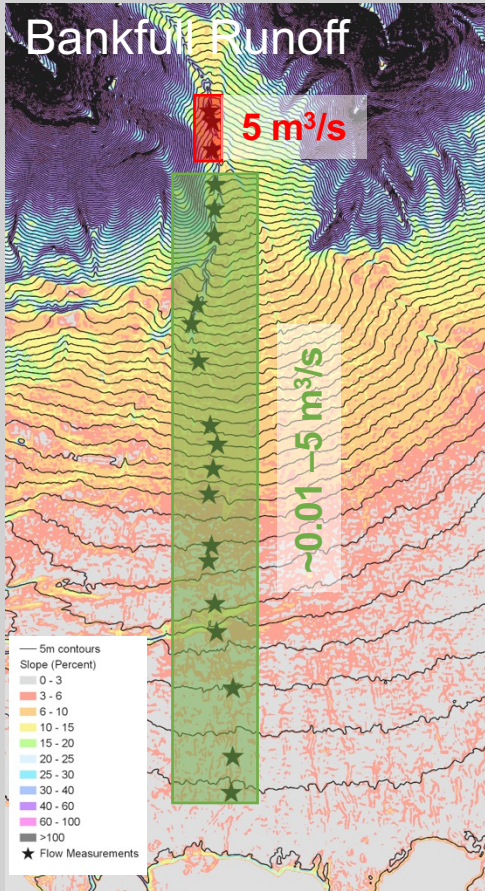
- H = flow depth (field)
- $D_{84}$  = grain size (field, 84<sup>th</sup> percentile)
- S = slope (field, DEM)
- U = average flow velocity



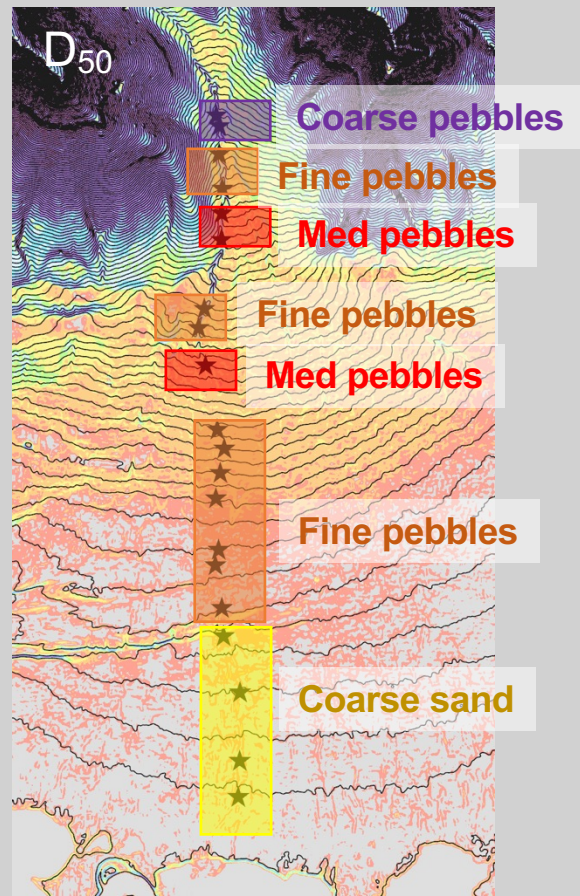
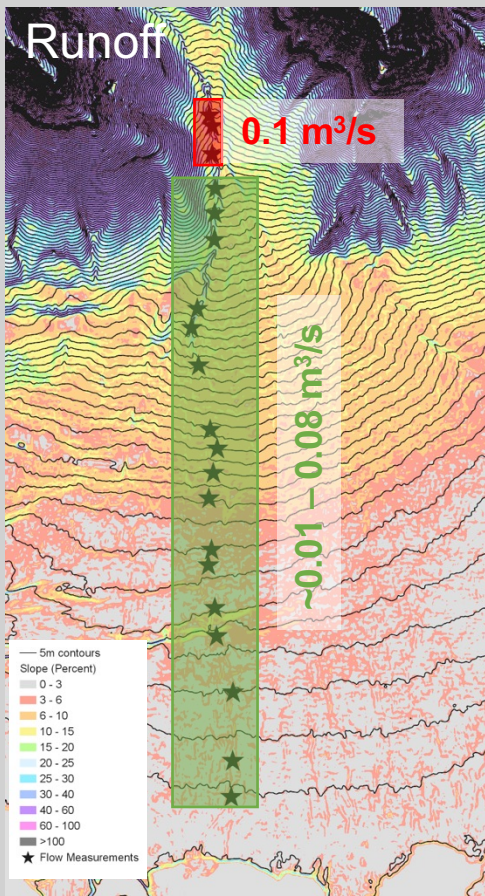
# Results: Storm discharge



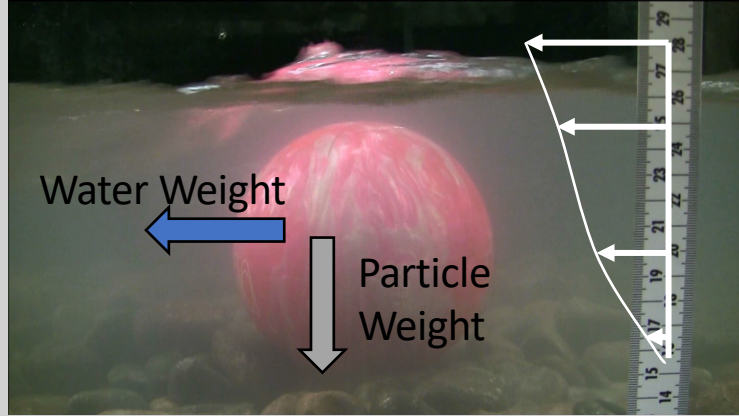
# Results: Bankfull discharge



# Results: Grain size



# Results: Fluvial sediment flux



Force of flow  
Particle weight



$$\tau_* = \frac{\rho g h S}{(\rho_s - \rho) g D}$$

Sediment flux  $\propto \tau^* / \tau_c^*$ , where  $\tau_c^*$  is the threshold over which sediment starts to move ( $\tau^* / \tau_c^* > 1$  means sediment moving)



$$\tau_c^* = 0.15 S^{0.25}$$

Lamb et al. (2008)



$$\tau_c^* = 0.021 + 0.015 \exp[-20 F_s]$$

Wilcock and Crowe (2003)

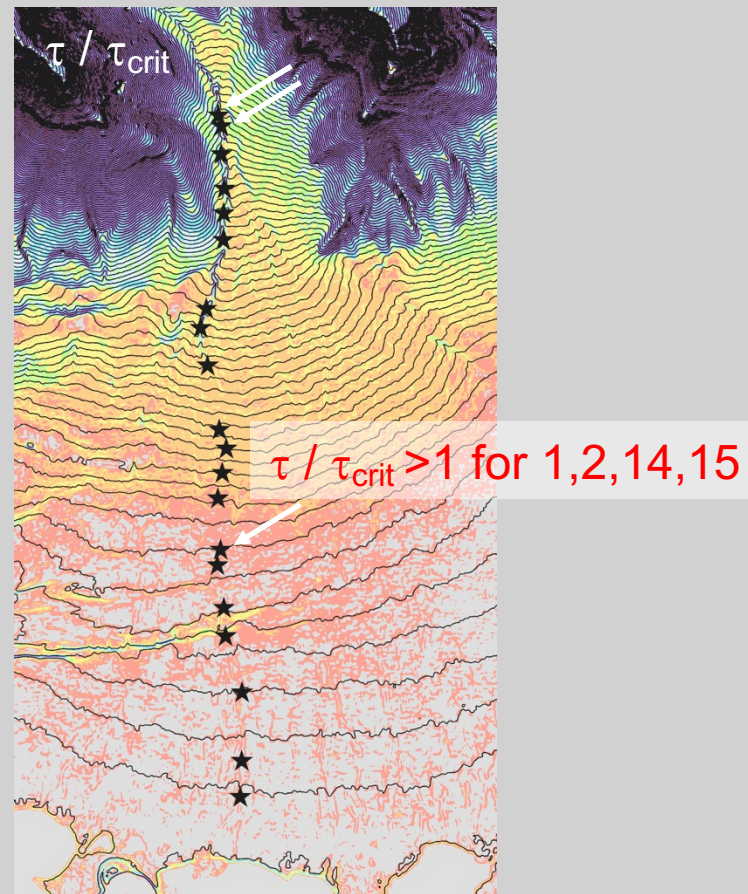
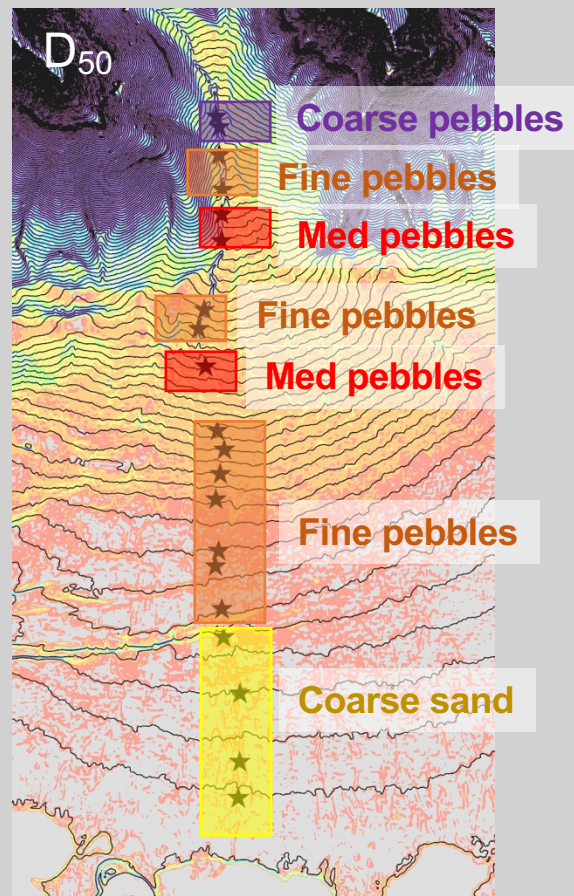
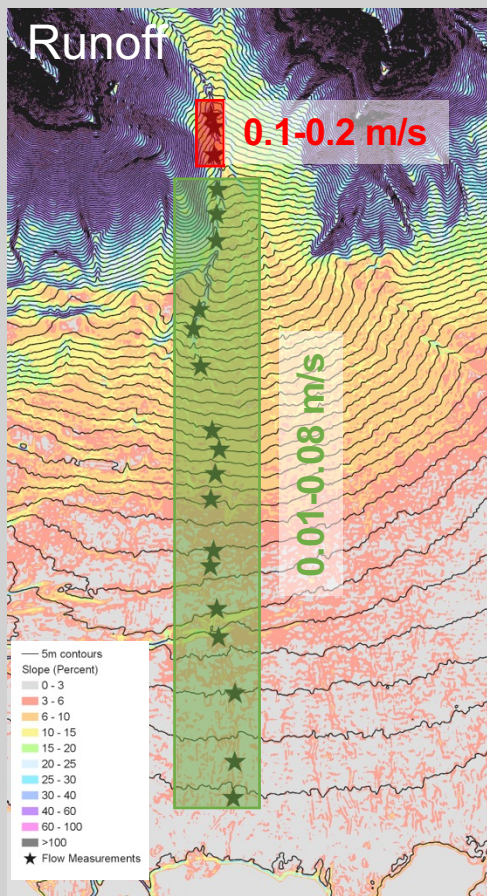


$$\tau_c^* = 0.56 S^{0.5}$$

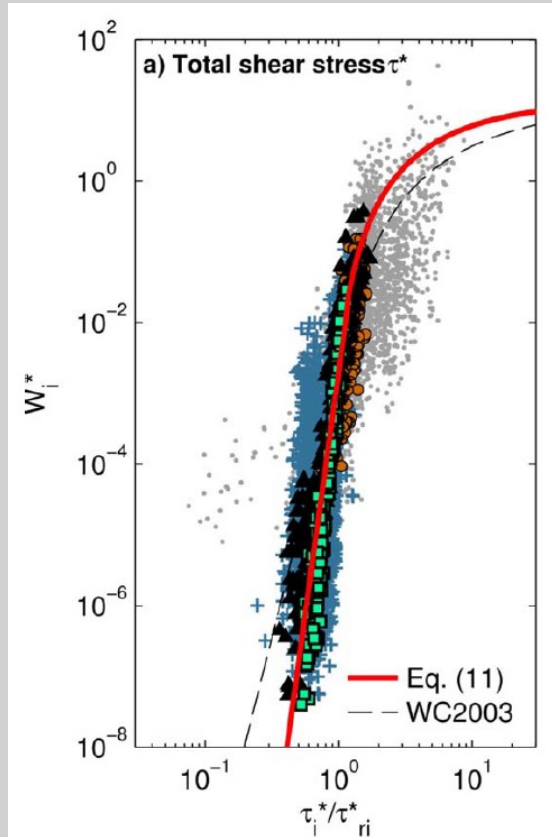
Schneider et al. (2015)



# Results: Fluvial sediment flux



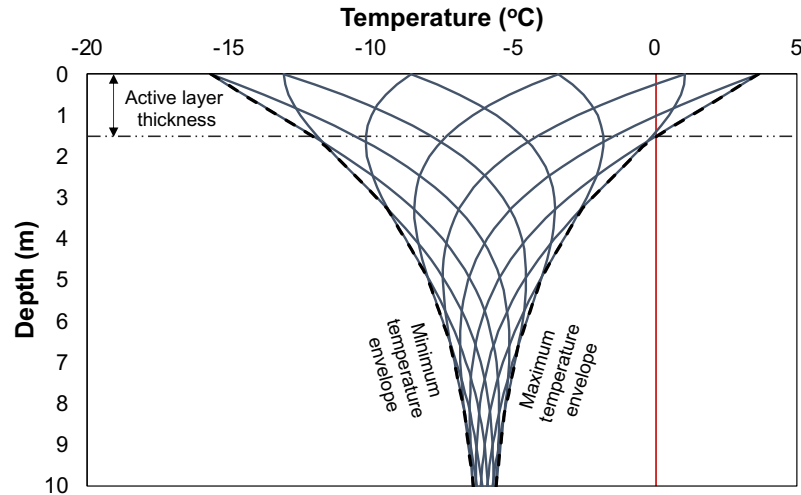
# Results: Fluvial sediment flux



- $W^* = \frac{q_s}{D_{50} \sqrt{RgD_{50}}} = \text{Einstein number}$
- For range of  $\tau^* / \tau_c^*$  estimated, dimensionless sediment flux per unit width ( $W$ ) ranges from 0 to 0.4 (for storm) and  $10^{-6}$  to 5 (for bankfull)
- Converted to dimensional flux, we estimate fluxes of 0 to 0.04 m<sup>3</sup>/hr during storm and 0.3 to 440 m<sup>3</sup>/hr (**70 m<sup>3</sup>/hr average**) during bankfull

# Results: Debris flow sediment flux

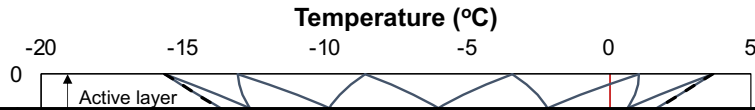
$$T(z,t) = \bar{T} + \Delta T e^{-z/z^*} \sin \left[ \frac{2\pi t}{P} - \frac{z}{z^*} \right]$$



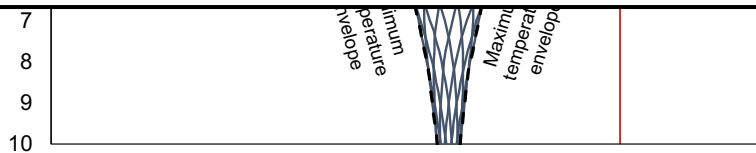
- Using simple diffusion model with periodic BC, estimate active layer depth is  $\sim 1.5$  m
- Most debris flow deposits were 1 – 2 m thick
- Possible that size of mass failures controlled by melting and failure along permafrost table – but intermittency unknown...

# Results: Debris flow sediment flux

$$T(z,t) = \bar{T} + \Delta T e^{-z/z^*} \sin \left[ \frac{2\pi t}{P} - \frac{z}{z^*} \right]$$



Appreciable amounts of sediment can be moved by fluvial processes during ~yearly melt events. Amount of sediment delivered by debris flows may be controlled by active layer thickness on hillslope/talus.



- Using simple diffusion model with periodic BC, estimate active layer depth is ~ 1.5 m

- Possible that size of mass failures controlled by melting and failure along permafrost table – but intermittency unknown...

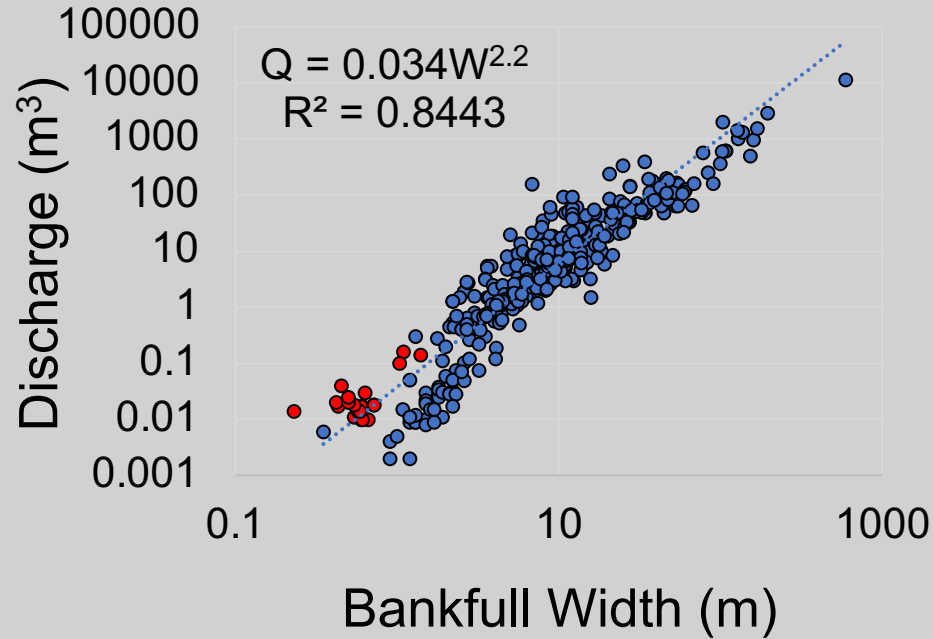


# Results

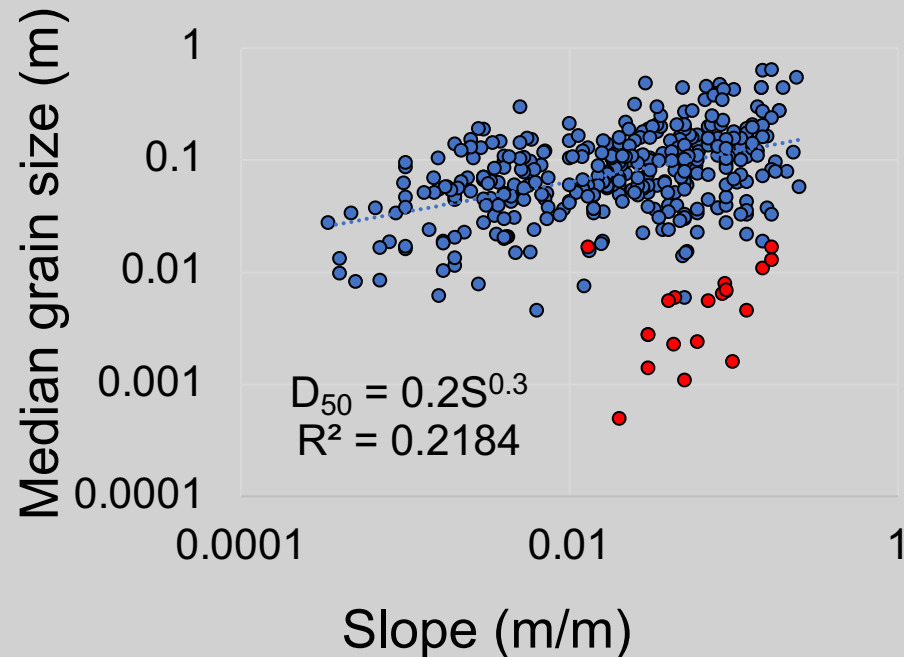
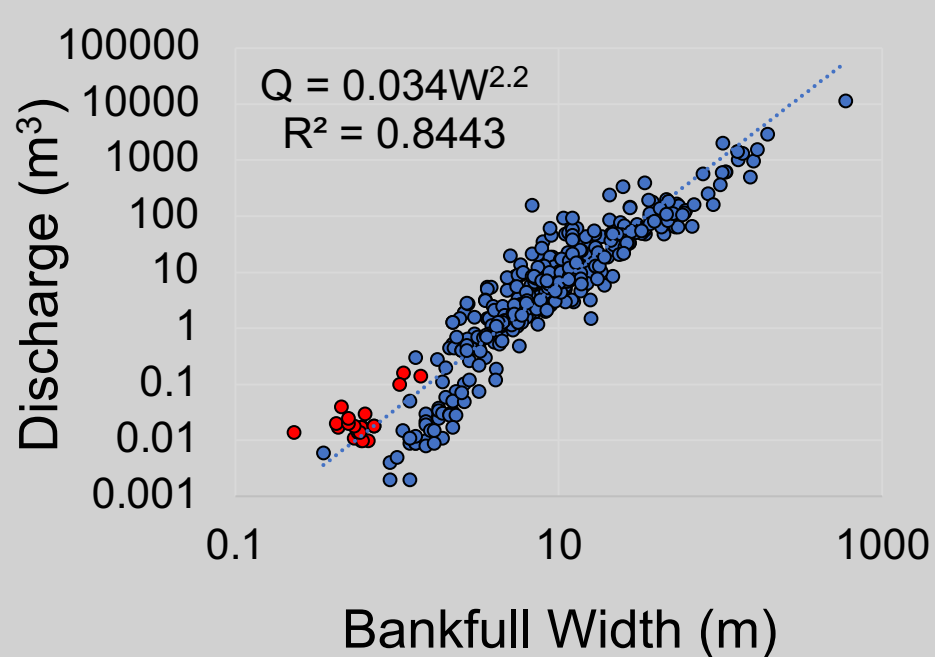
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# Results: Arctic fan compared to terrestrial alluvial rivers



# Results: Arctic fan compared to terrestrial alluvial rivers



# Results: Shale fracturing





# Results: Shale fracturing



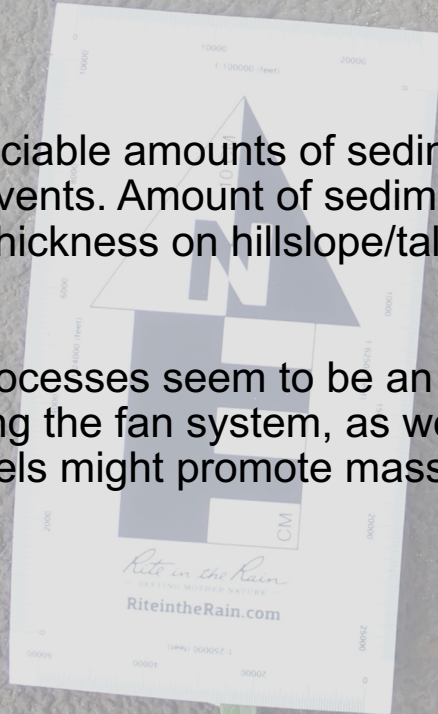
Discharge-width relation still holds for this arctic fan channel, but grain sizes within channels are finer relative to other (more temperate) steep streams



Shale fracturing in-place

# Conclusions

- Compared to earlier work and other periglacial fans, surface processes and sedimentology suggest that much of the fan is build by fluvial processes, not debris flows
- Appreciable amounts of sediment can be moved by fluvial processes during ~yearly melt events. Amount of sediment delivered by debris flows may be controlled by active layer thickness on hillslope/talus
- Icy processes seem to be an important control on sediment amounts and grain size entering the fan system, as well as sorting along fan → fine-grained sediment in steep channels might promote mass flow events / in-channel debris flows





# Thanks!

- NASA SSW grant 80NSSC19K0539
- Dartmouth College, Smithsonian Institution, University of Arkansas and the Center for Advanced Spatial Technologies
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- Yukon Geological Survey (YGS), Craig Nicholson (YGS), and Galen Halverson (McGill University)
- NASA SUPPR Program for funding Diane Wagner (summer intern, Franklin and Marshall)
- and Sonny McDonald

