Relationship between surface thermodynamic contrasts and precipitation intensity in idealised monsoon simulations: control runs and metrics

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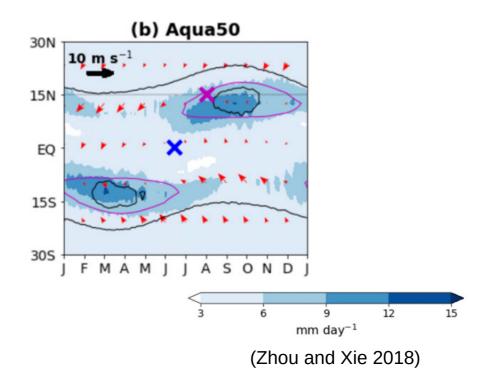




Does the surface temperature contrast matter at all for monsoon precipitation intensity?

Traditional view of monsoon: large-scale sea breezes...

Aquaplanet

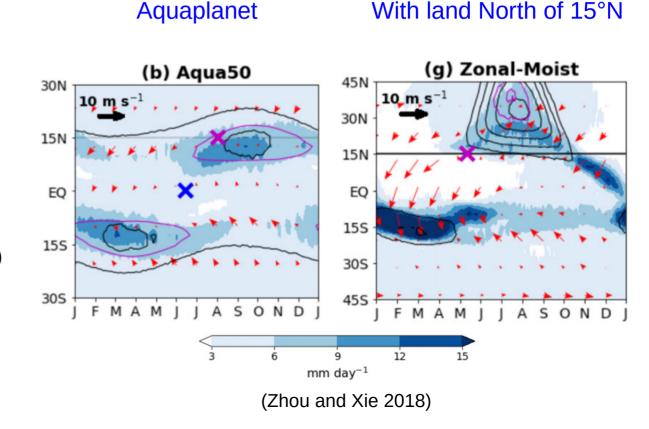


Does the surface temperature contrast matter at all for monsoon precipitation intensity?

Traditional view of monsoon: large-scale sea breezes...

That doesn't work so well...

(Gadgil 2018, Hill 2019, Geen et al 2020)



Questions and strategy

QUESTIONS:

- Does the surface temperature contrast matter at all for monsoon precipitation intensity?
- Do monsoon non-linearities come from convective processes, dynamical feedbacks, or non-linearities in the forcing?

STRATEGY:

- Very few idealised simulation studies of monsoons at relatively high resolution (and with resolved convection).
- So we devise a modular framework to simulate idealised monsoons at higher resolutions.

Idealised simulations with imposed SST

- WRF idealised simulations over an aquapatch
- Imposed SST with peak at 12°N (or 24°N, or Eq)
- Grid spacing: 10 km, 20 km, 30 km (for now)
- Domain: 21°S-21°N to 63°S-63°N
- Periodic BC on x-axis
- Symmetric BC on y-axis
- Either no convection scheme, or Kain-Fritsch

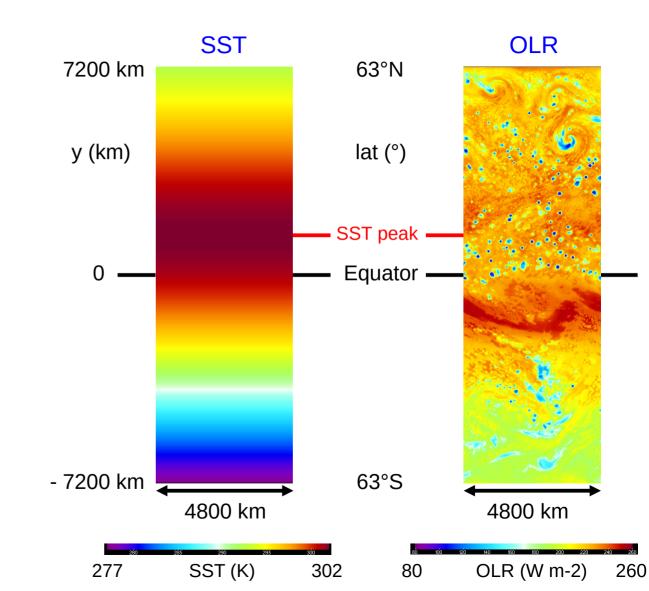
$SST_{LH} = T_0 -$	$\Delta T(\sin\!\phi - \sin\!\phi_0)^2$
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(Boos and Kuang 2010)

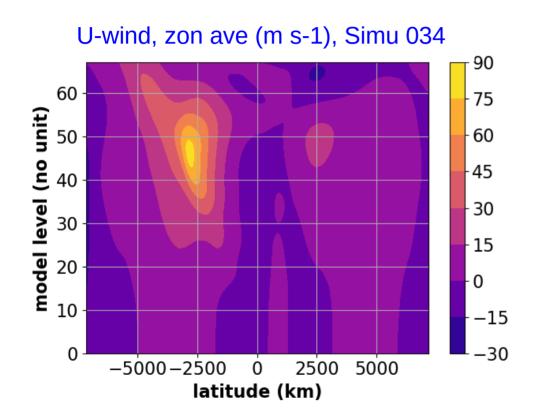
	Lat SST peak	Coriolis	Grid spacing	Domain	Convection	Notes
Simu 022	12°N	Yes	10 km	21°S-21°N	resolved	TCs, Pb at North boundary
Simu 029	12°N	Yes	20 km	42°S-42°N	resolved	TCs
Simu 034	12°N	Yes	30 km	63°S-63°N	resolved	TCs, better steady state
Simu 035	12°N	Yes	30 km	63°S-63°N	KF scheme	Very different spatial org
Simu 021	Eq	Yes	10 km	21°S-21°N	resolved	Double ITCZ

Spatial organisation

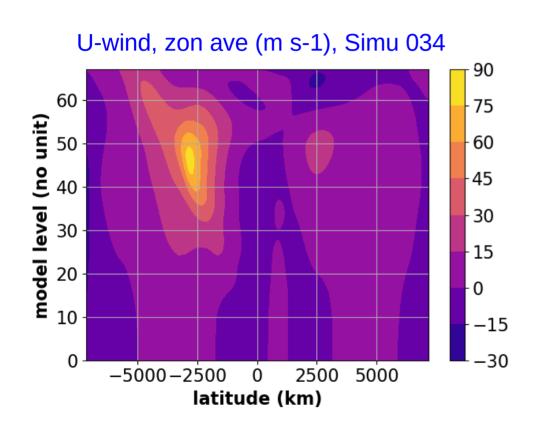
- Many Tropical Cyclones
- Mid-latitudinal features
- More convection in NH, slightly away from the Equator
- Drier subtropics in the winter hemisphere
- If SST peaks at Equator, double ITCZ

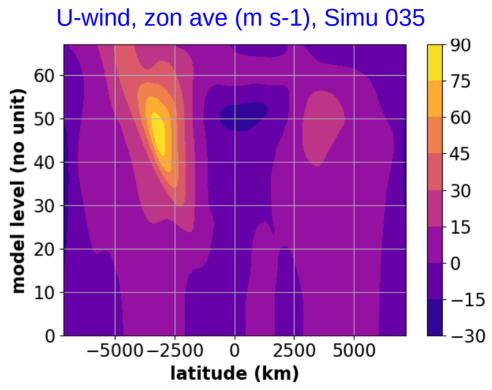


Circulation with subtropical jets and Trade winds



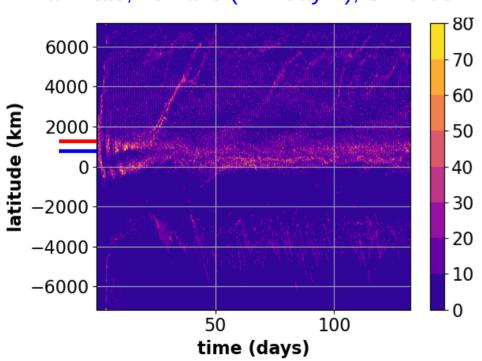
Circulation with subtropical jets and Trade winds





An ITCZ shifted to the North, but not as far as the SST peak

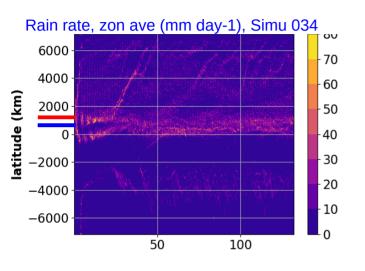


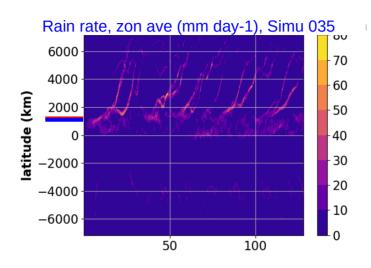


Reasonable summer ITCZ features ("monsoon-like")

How to define a good monsoon metrics?

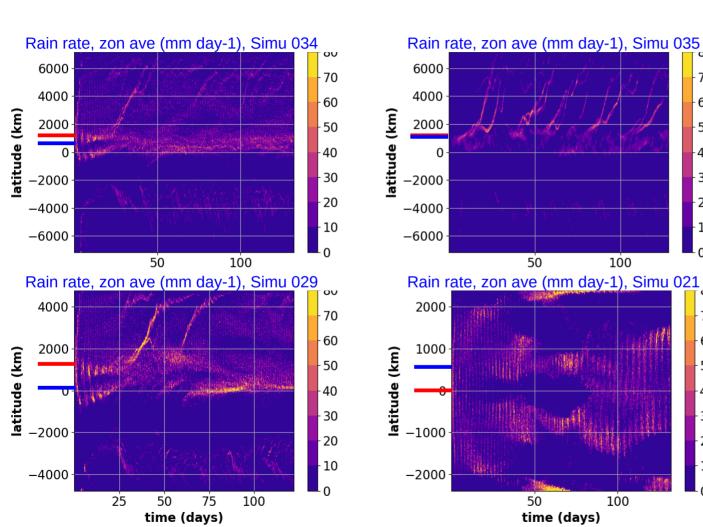
An ITCZ shifted to the North, but not as far as the SST peak





- 034: Big domain + explicit convection
- 035: Big domain + KF convection
- 029: Smaller domain, higher resolution
- 021: SST peak at the Equator, even smaller domain

An ITCZ shifted to the North, but not as far as the SST peak



034: Big domain + explicit convection

-70

60

-50

40

- 30

- 20

- 10

- 70

- 60

- 50

40

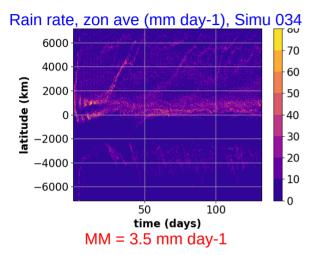
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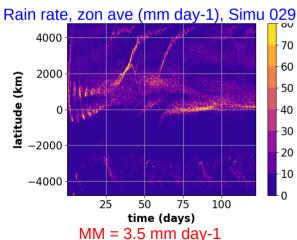
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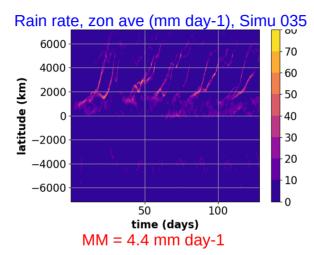
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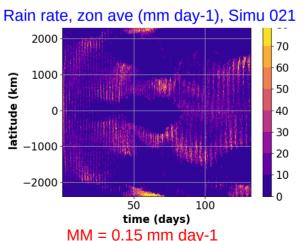
- 035: Big domain + KF convection
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A monsoon metrics









Monsoon metrics based on difference between tropical precipitation P in each hemisphere:

- 034: Big domain + explicit convection
- 035: Big domain + KF convection
- 029: Smaller domain, higher resolution
- 021: SST peak at the Equator, even smaller domain

Conclusion and future work

- These control simulations manage to capture some general features associated with real monsoons: tropical cyclones, large-scale circulation (subtropical jets, Trade winds), precipitation distribution with latitude.
- The simulation with a convection scheme is the most different.
- We have defined a monsoon metrics suitable for our simulations

- Series of simulations with various SST gradients: how do they influence monsoon intensity? And with asymmetric land?
- Control of monsoon intensity by T or MSE or other thermodynamic variables ?
- Linear or non-linear relationships?

Thank you

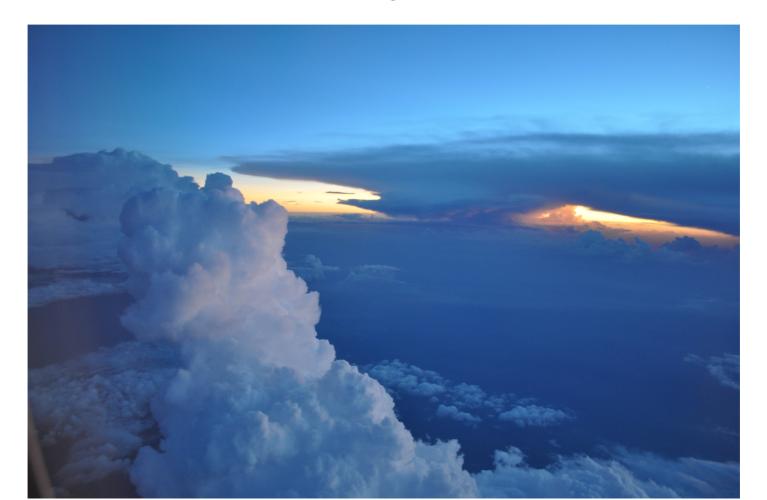


Photo by Maxime Colin