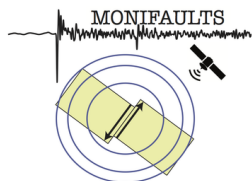


Seismological characterization of dynamics parameter of the Hunga Tonga explosion from teleseismic waves

Piero Poli Nikolai Shapiro

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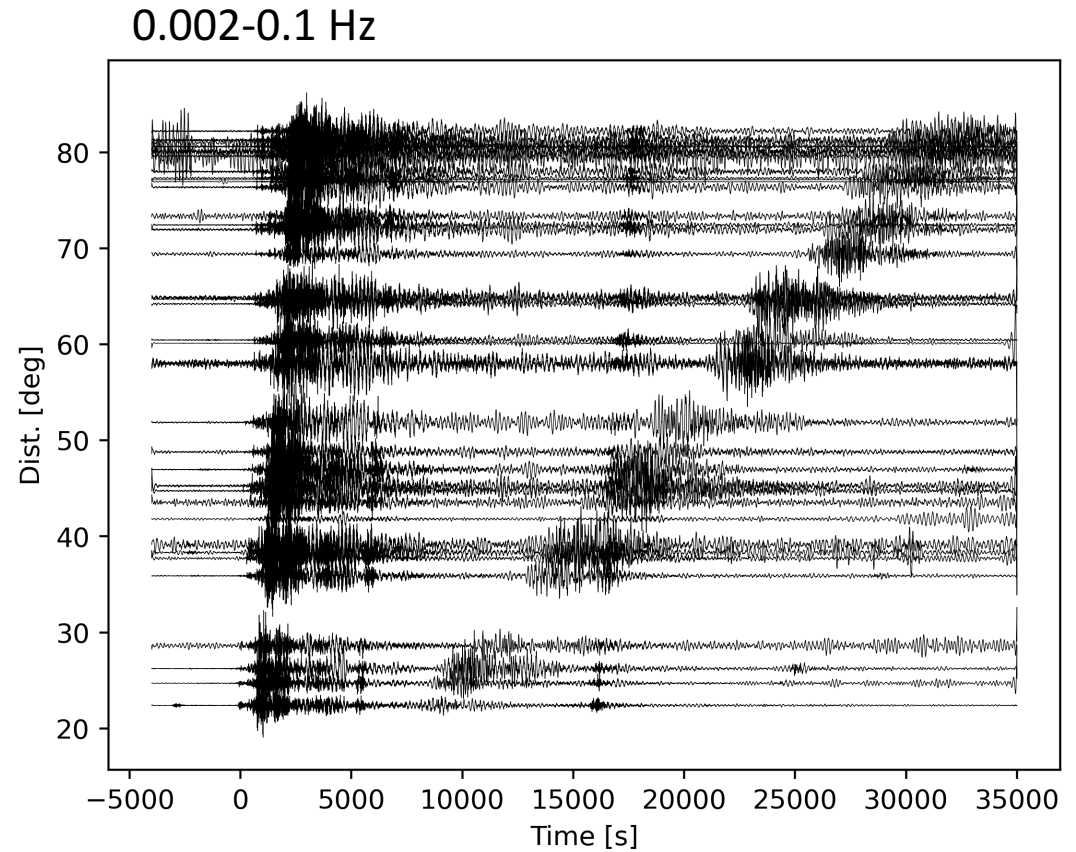
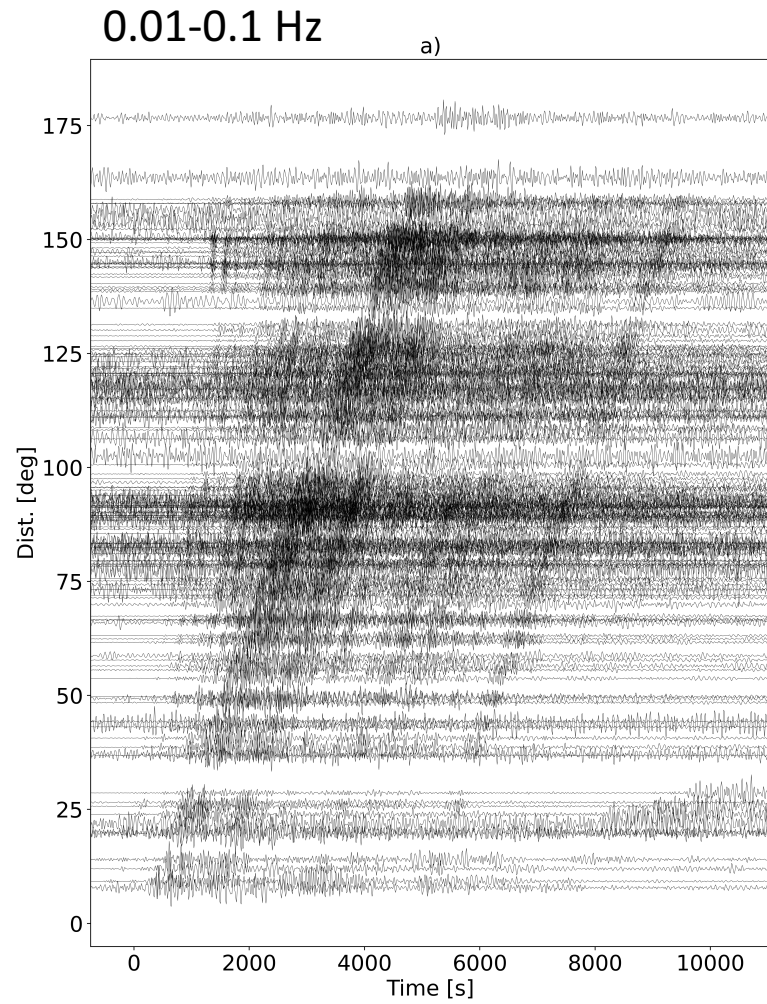


Outline

- seismological detection of volcanic explosions
- seismological characterization of explosion size
- results for the Hunga-Tonga explosion

Hunga Tonga volcanic explosion

records of global seismological network



Data

- **NETWORKS:** _GSN, GE, G, _ARCTIC, _ANTARTICA
- **DATA CENTERS:** IRIS, GEOSCOPE, GFZ
- **CHANNELS:** LHZ, VHZ

Data downloaded every day

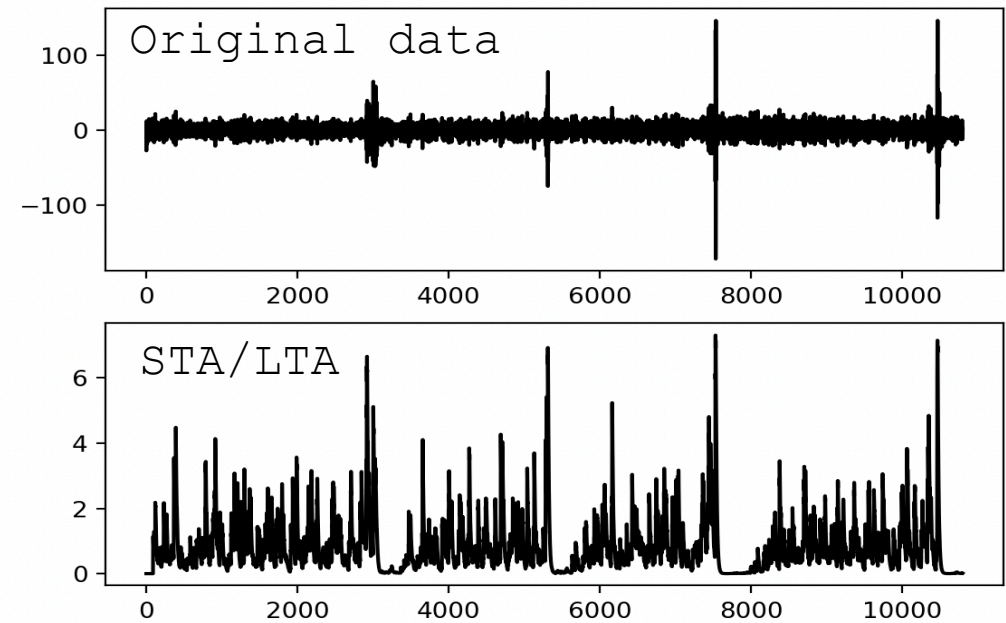
$T_0 = 00\text{h}00\text{m}00\text{s}$ (for given day)

Data span = $T_0 - 3\text{h} \mid T_0 + 3\text{h}$ (To account for detections
around midnight)

- **AVERAGE STATIONS NUMBER:** 130

Detection algorithm

- **FB:** 0.01-0.03Hz
- **Deconvolution Instrumental Response**
- **Characteristic function:** Recursive STA/LTA¹ (STA=120s LTA=900s)
- **Source grid:** One source every 2.5deg at 0Km depth (9940 sources)



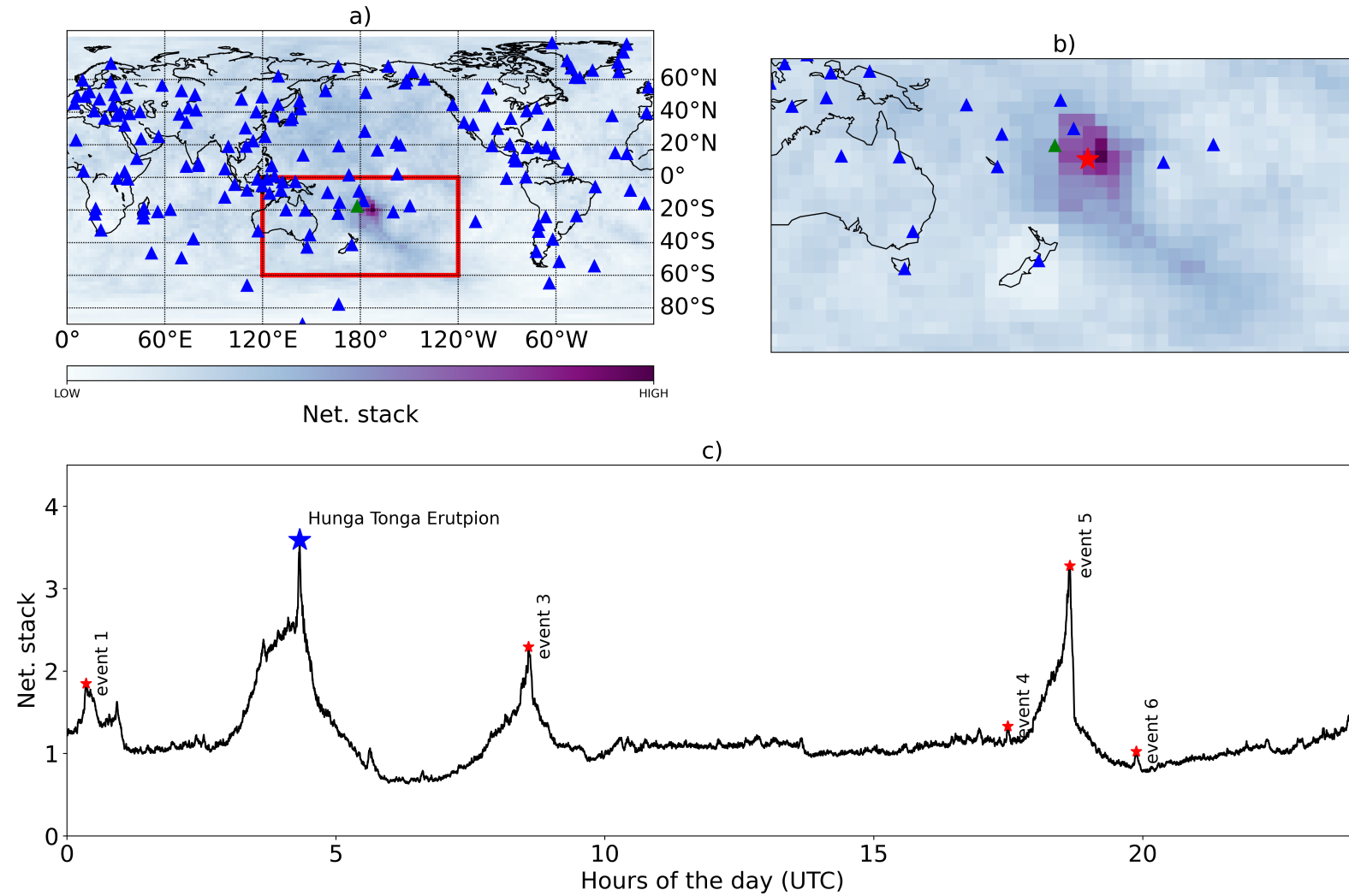
¹ -Withers, M., Aster, R., Young, C., Beiriger, J., Harris, M., Moore, S., and Trujillo, J. (1998), **A comparison of select trigger algorithms for automated global seismic phase and event detection**, *Bulletin of the Seismological Society of America*, 88 (1), 95-106. <http://www.bssaonline.org/content/88/1/95.abstract>

Detection algorithm: delay and stack

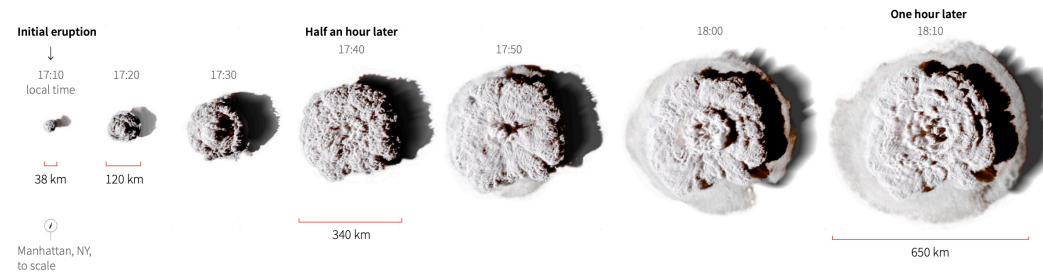
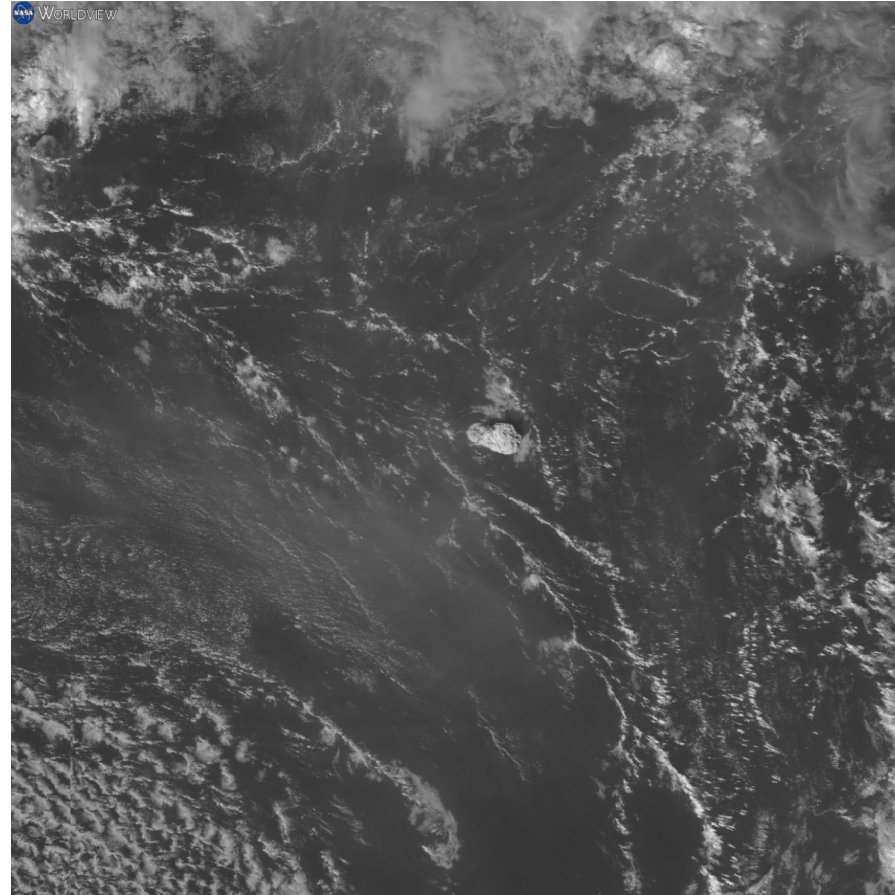
- **i = sources**
- **j = stations**
- $E(i, t) = \sum_{j=1}^N u_j(t - dt)$ **Matrix = [Sources;Time]**
- ***[max E] (t)***

dt is the propagation estimated from distance from the sources assuming **Rayleigh waves** speed from **PREM** model

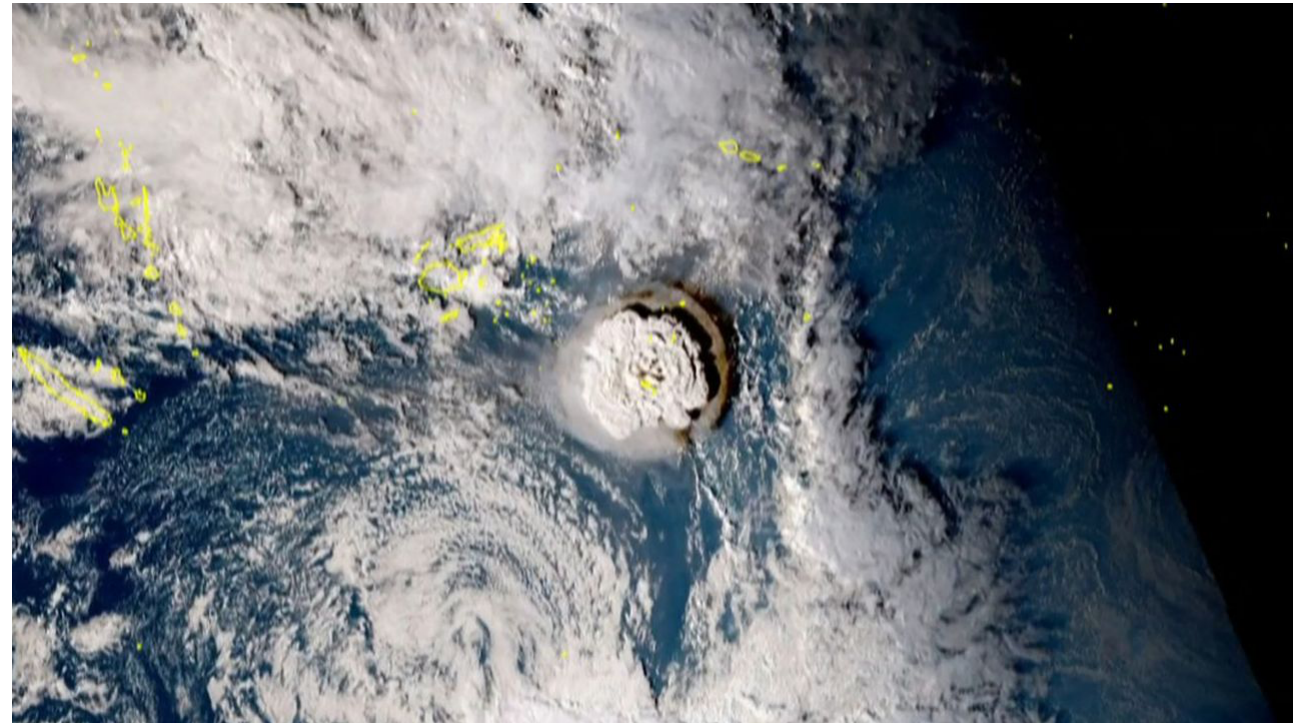
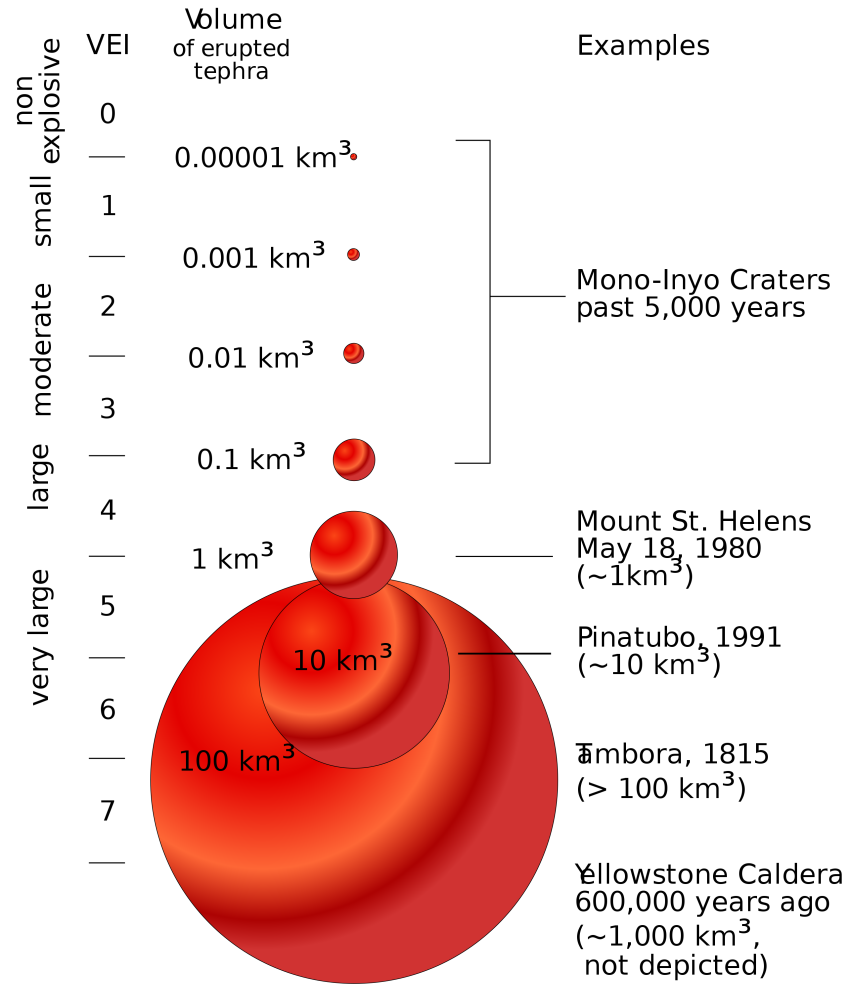
Hunga Tonga volcanic explosion



Hunga Tonga volcanic explosion

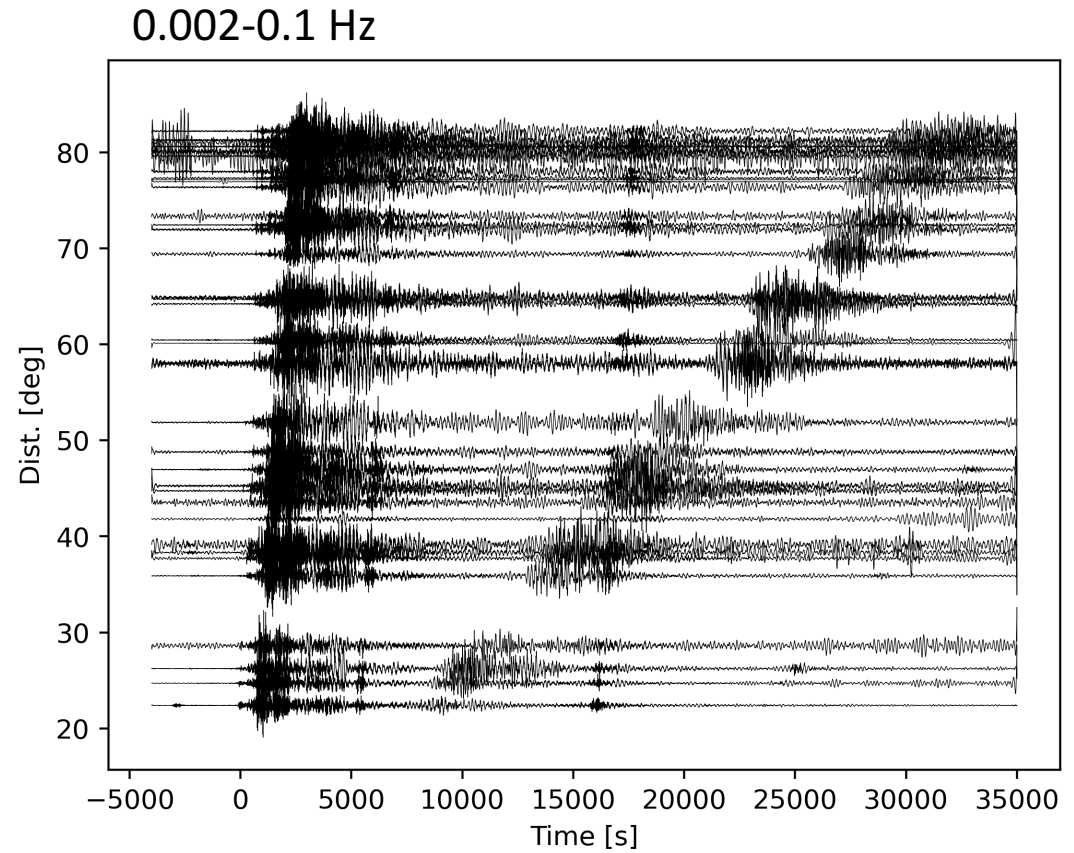
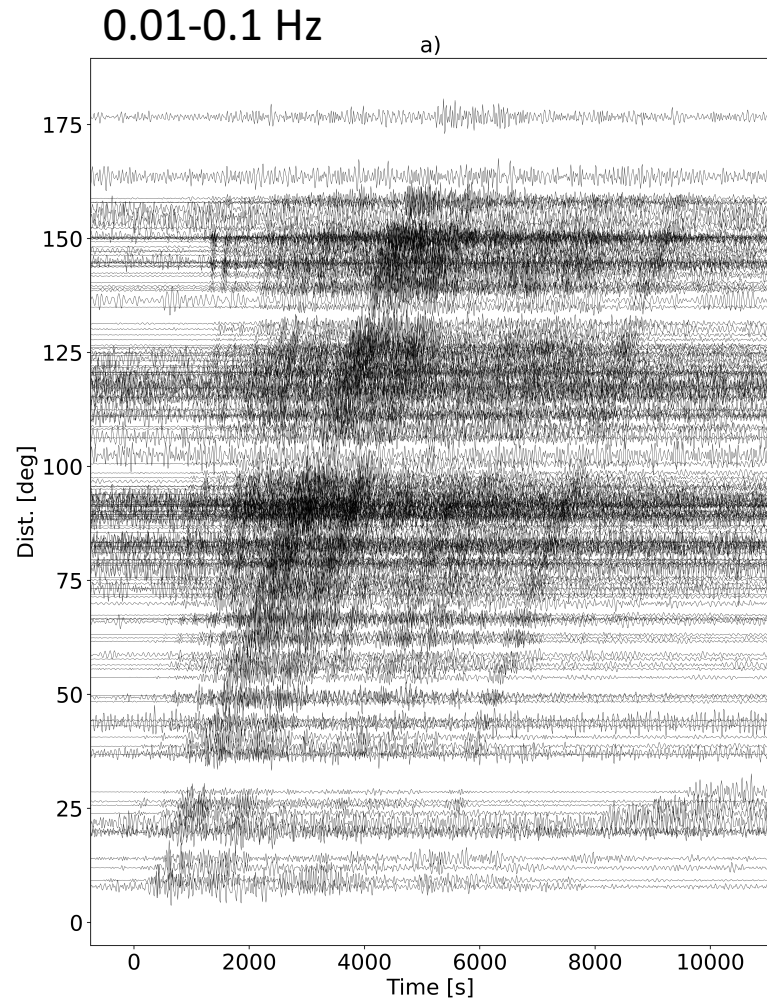


Seismological assessment of the Hunga Tonga volcanic explosion



Hunga Tonga volcanic explosion

records of global seismological network



Hunga Tonga volcanic explosion

Simplified seismic source model: vertical force

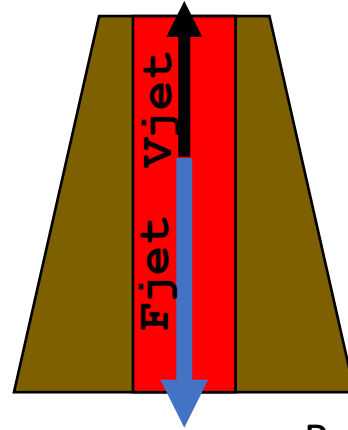
Data Source Propagation

$$u_z(t, r) = F_z(t) * G_{zz}(t - t_s, r, r_s)$$

$$|u_z(\omega)| = |F_z(\omega)| |G_{zz}(\omega)|$$

Gzz = 1N vertical force

$$\frac{|u_z(\omega)|}{|G_{zz}(\omega)|} = |F_z(\omega)|$$

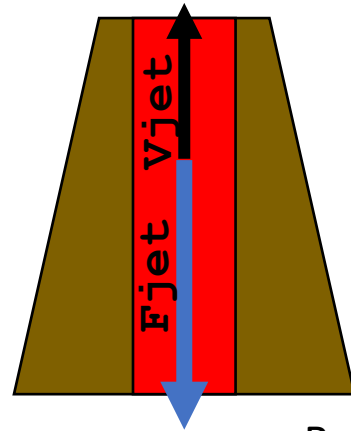


Brodsky et al., 1999

Hunga Tonga volcanic explosion

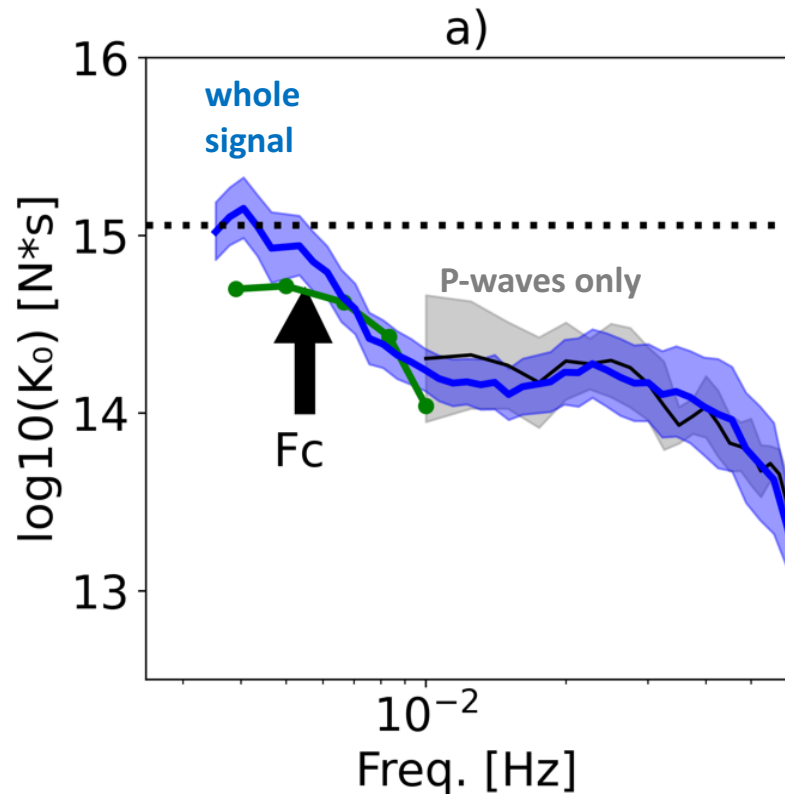
Main seismological source parameter: **Explosion impulse K_0**

$$\frac{|u_z(\omega)|}{|G_{zz}(\omega)|} = |F_z(\omega)|$$



Brodsky et al., 1999

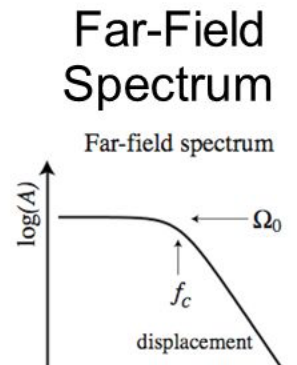
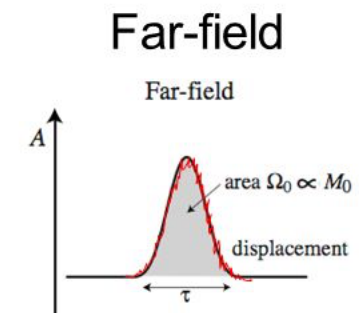
*analogy with seismic moment
for tensorial sources*



$$F_z(\omega) = \int F_z(t) e^{i\omega t} dt$$

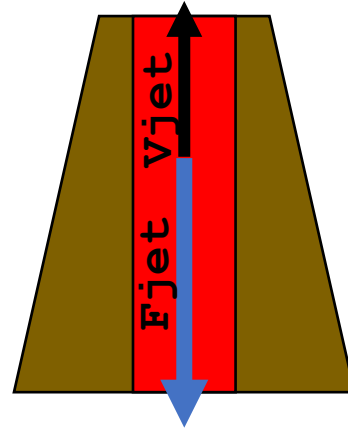
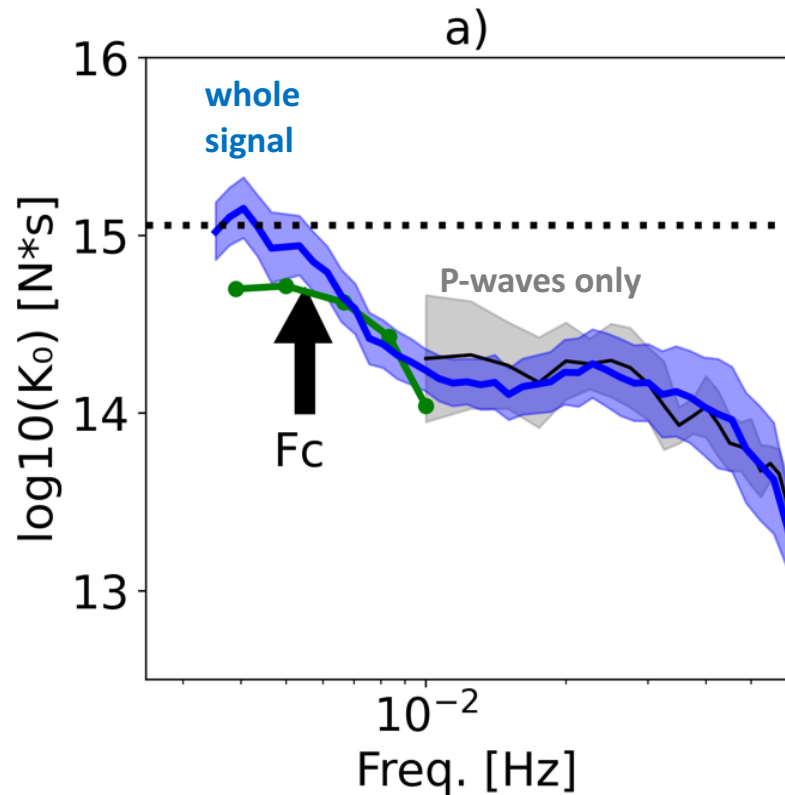
$$\lim_{\omega \rightarrow 0} F_z(\omega) = \int F_z(t) dt = K_0$$

$$f_c \approx 1/\tau_{exp}$$



Hunga Tonga volcanic explosion

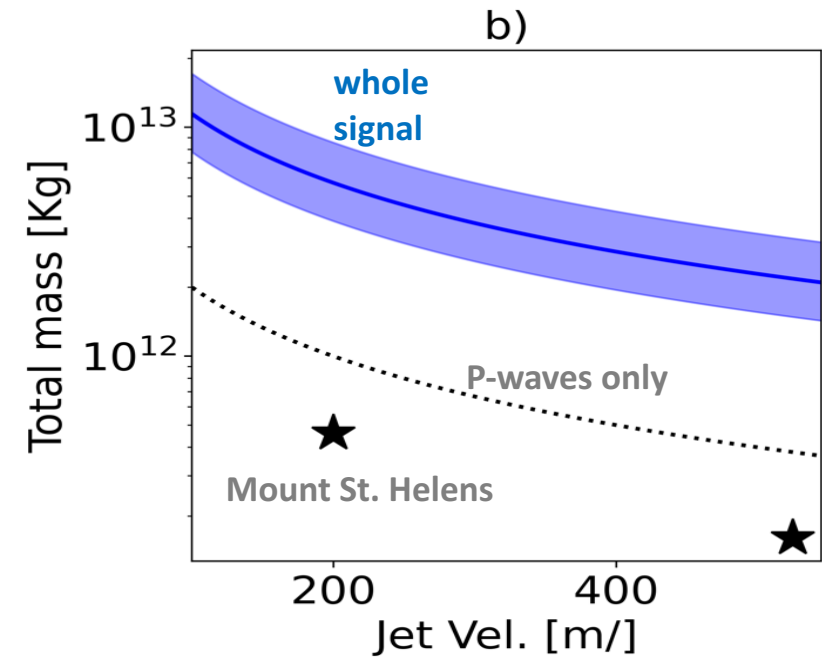
Simplified model to relate explosion impulse with ejected mass



$$F_{jet}(t) = V_{jet} \dot{m}(t)$$

$$m_{total} = \int \dot{m} dt$$

$$m_{total} = K_0 / V_{jet}$$



Brodsky et al., 1999

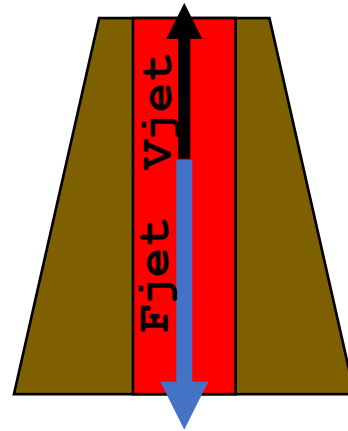
Hunga Tonga volcanic explosion

Source parameters

$$VEI = \log(V_{total}/10^9) + 5$$

$$VEI = \log(K_0/V_{jet}/\rho_{tephra}) - 4$$

$$\rho_{tephra} = 1000 \text{ kg/m}^3$$



$$F_{jet}(t) = V_{jet}\dot{m}(t)$$

$$m_{total} = \int \dot{m} dt$$

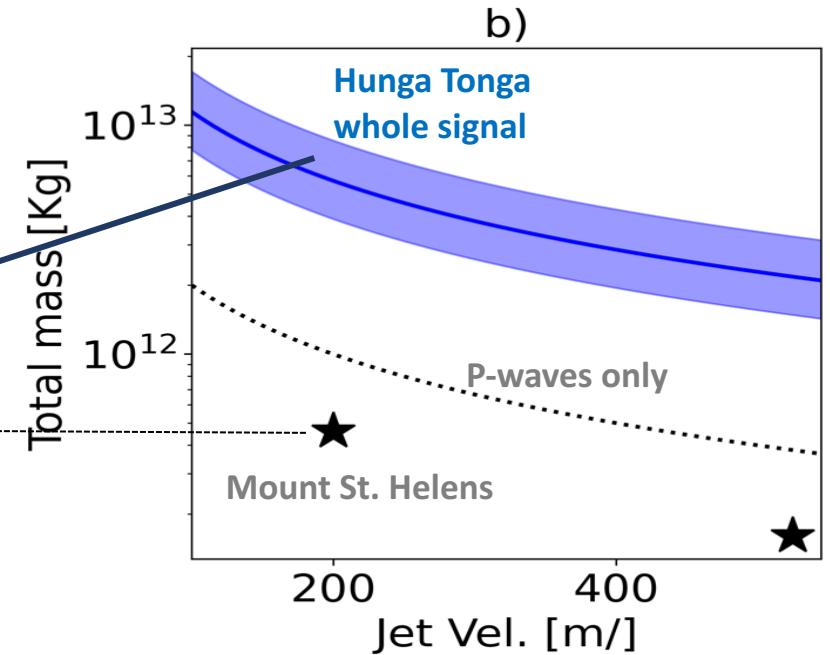
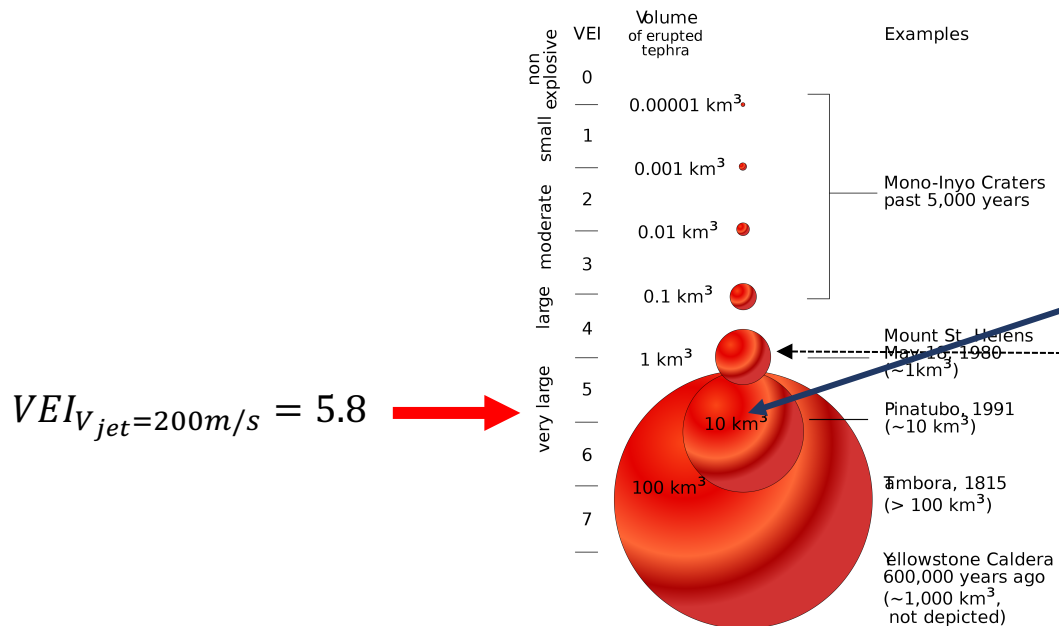
$$m_{total} = K_0/V_{jet}$$

Some estimates:

$$K_0 \approx 1.3 \cdot 10^{15} \text{ Ns}$$

$$m_{total} \approx 1.3 \cdot 10^{13} \text{ kg}$$

$$VEI_{V_{jet}=200\text{m/s}} = 5.8$$



Some conclusions

- Continuous analysis of seismic data can help to monitor (in near real time) volcanic explosions especially in remote places
- Rapid characterization of newly detected signals is easy to implement and provide rapid estimates of explosion dynamics parameters
- For the Hunga-Tonga eruption several parameters are derived and helped to constraint a first estimate of a VEI~6