

Recent deformation in the frontal Jura fold-and-thrust belt from a deep-seated thrust fault: evidence from Late Quaternary fluvial terraces

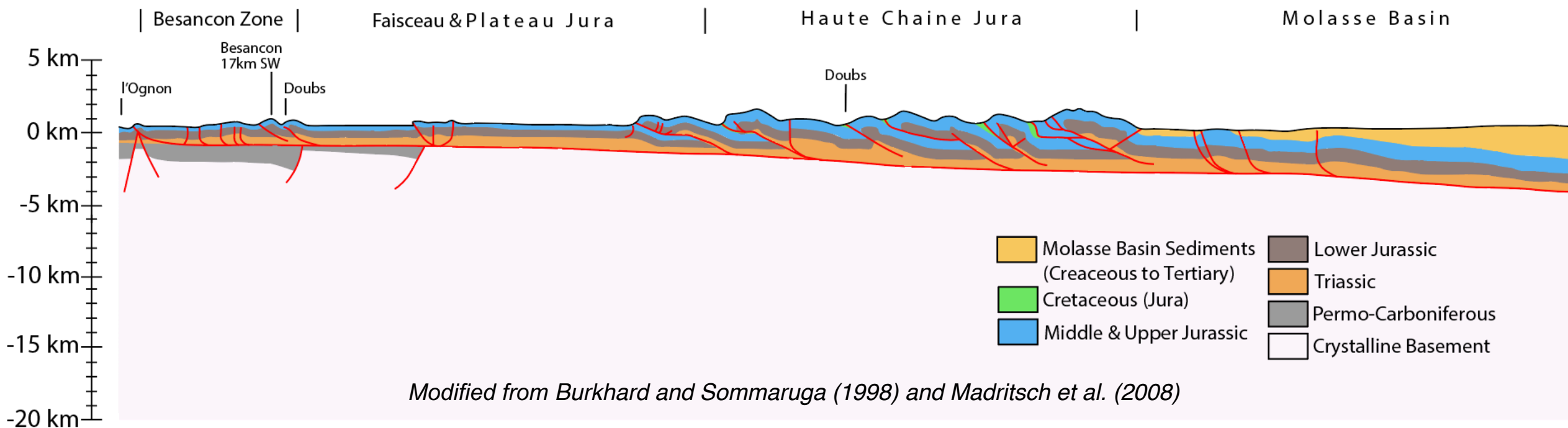
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An Overview of the Jura Mountains

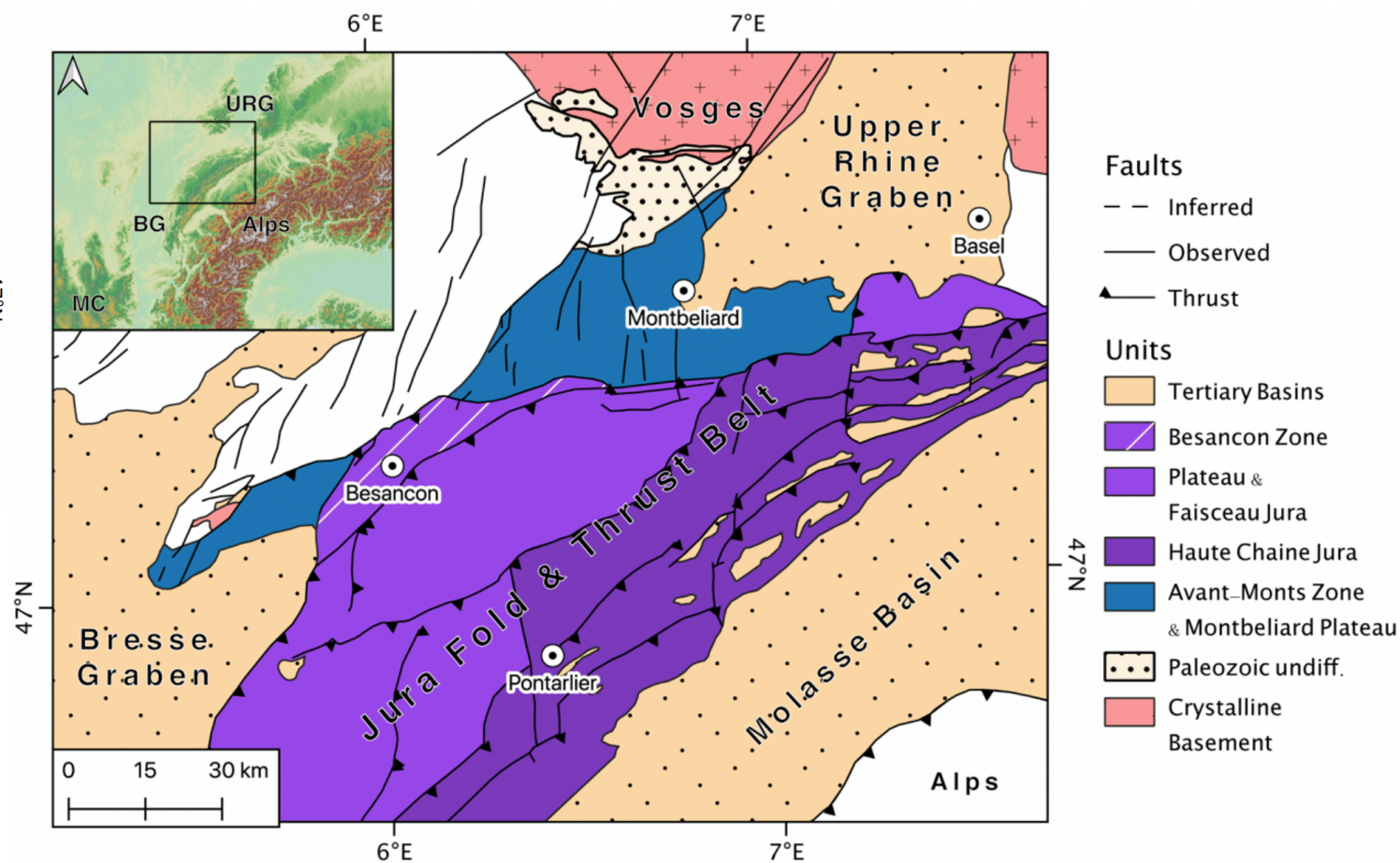
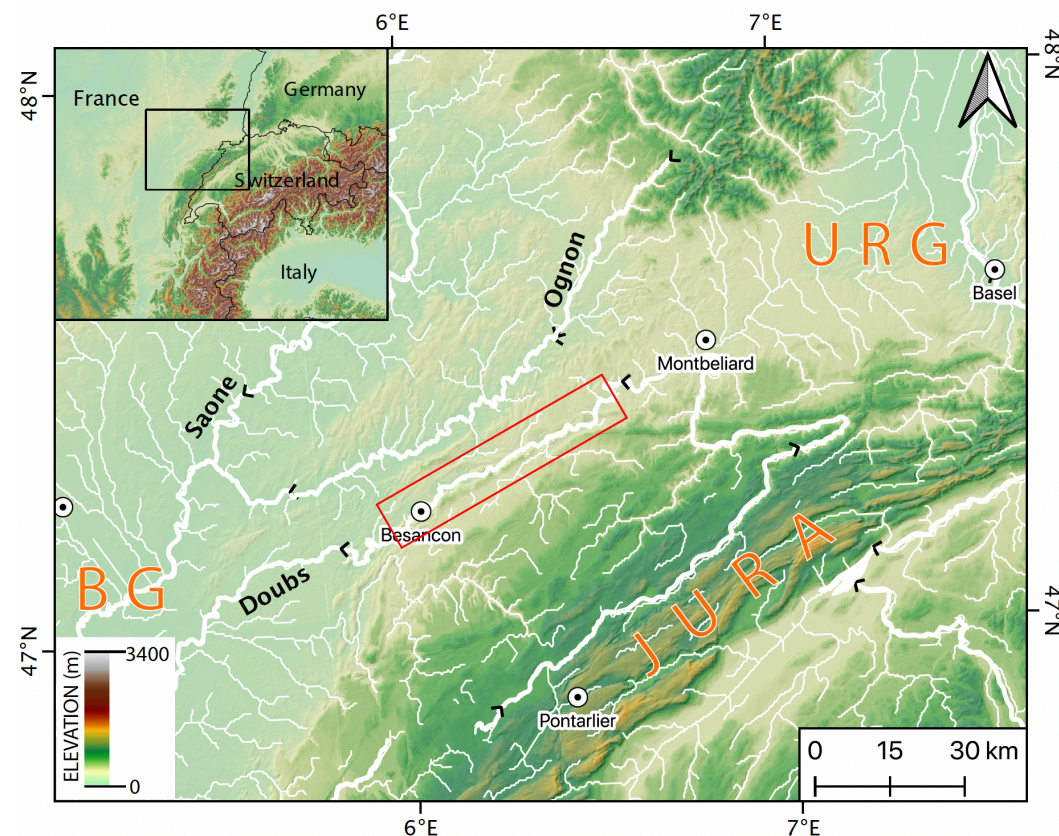


How is the northwestern edge of the Jura fold-and-thrust belt currently deforming?

- Inactive?
- Continuing thin-skinned deformation?
 - Nivière and Marquis (2000): abandonment of terrace risers (Upper Rhine Graben)
- Change to a thick-skinned regime?
 - Becker (2000): borehole in situ stress measurements (Haute Chaine Jura)
 - Ustaszewski and Schmid (2006): Paleostress reconstruction (Upper Rhine Graben)



The Doubs River and the Besançon Zone



Previous Literature

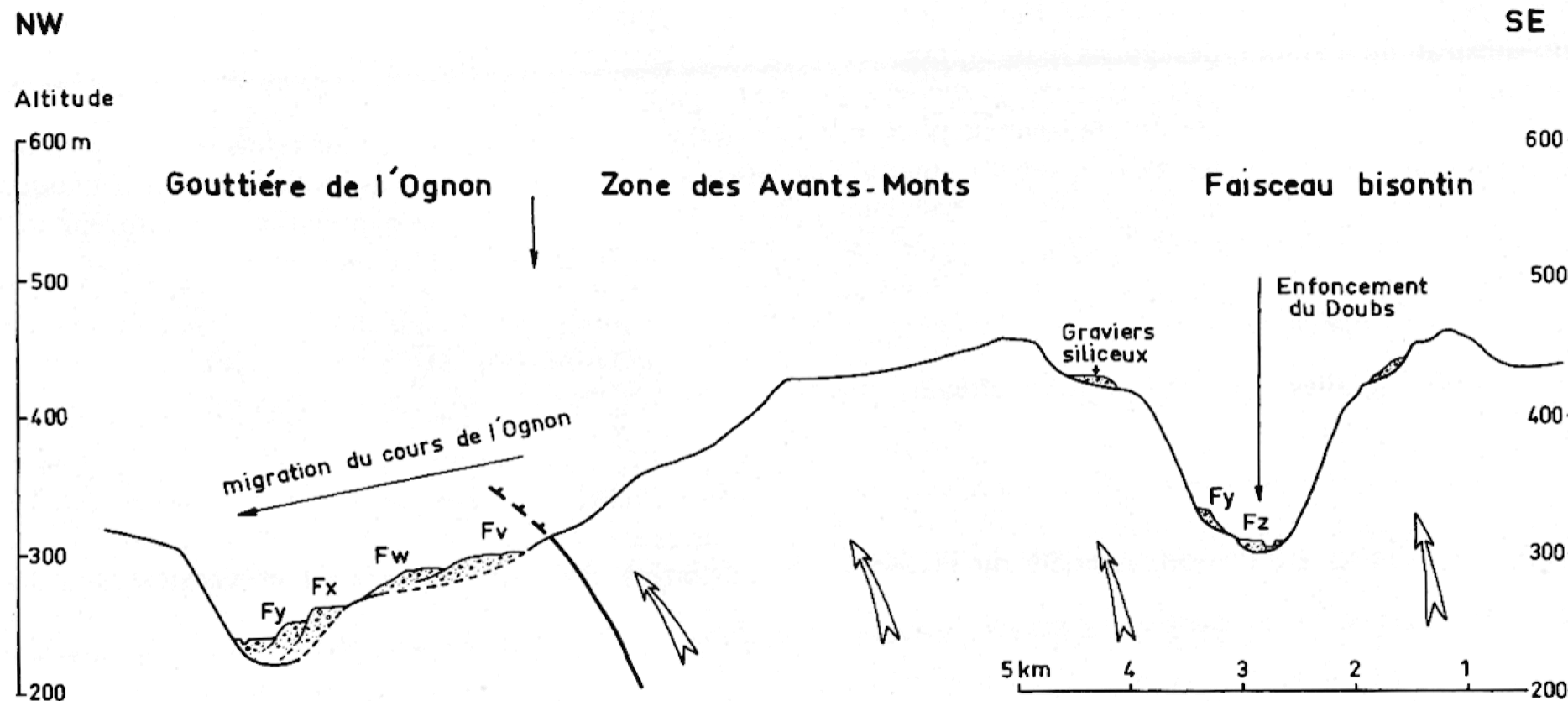
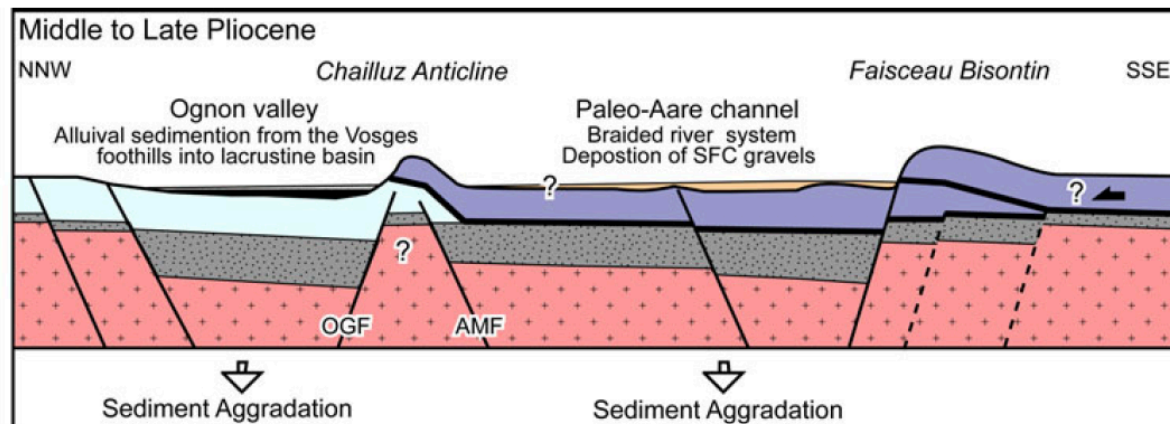


Fig. 3. — Coupe transversale comparée des vallées du Doubs et de l'Ognon et position respective des dépôts alluviaux.
Fig. 3. — Comparaison in transversal outcrop of Doubs valley and Ognon valley. Respective position of alluvial deposits.

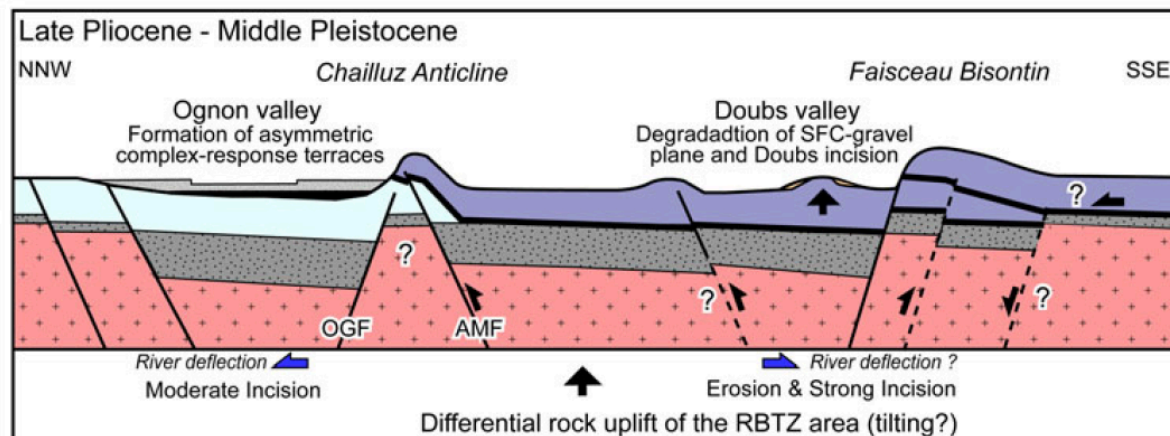
Campy, 1984



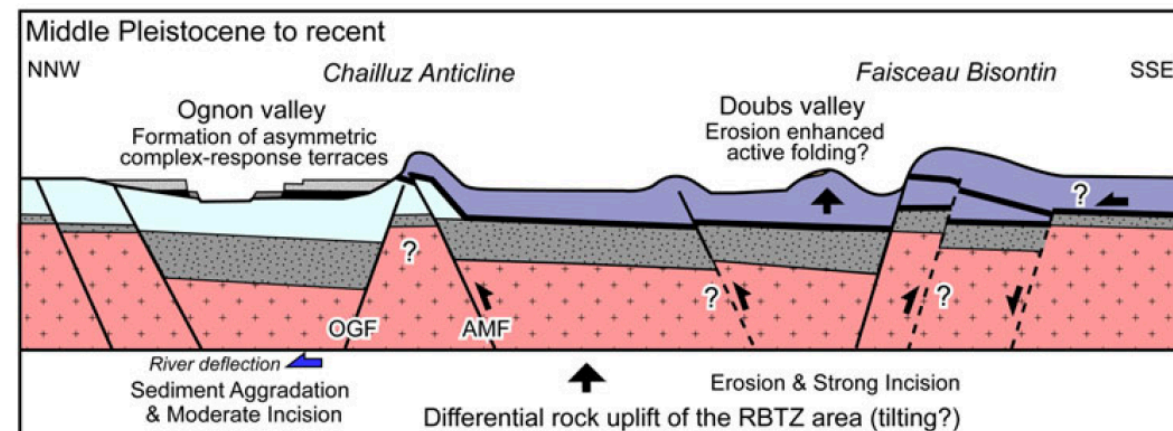
Previous Literature



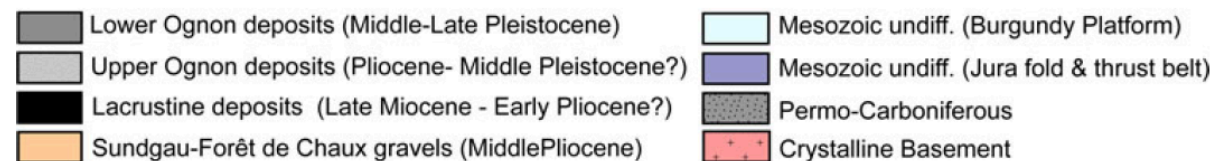
(a)



(b)



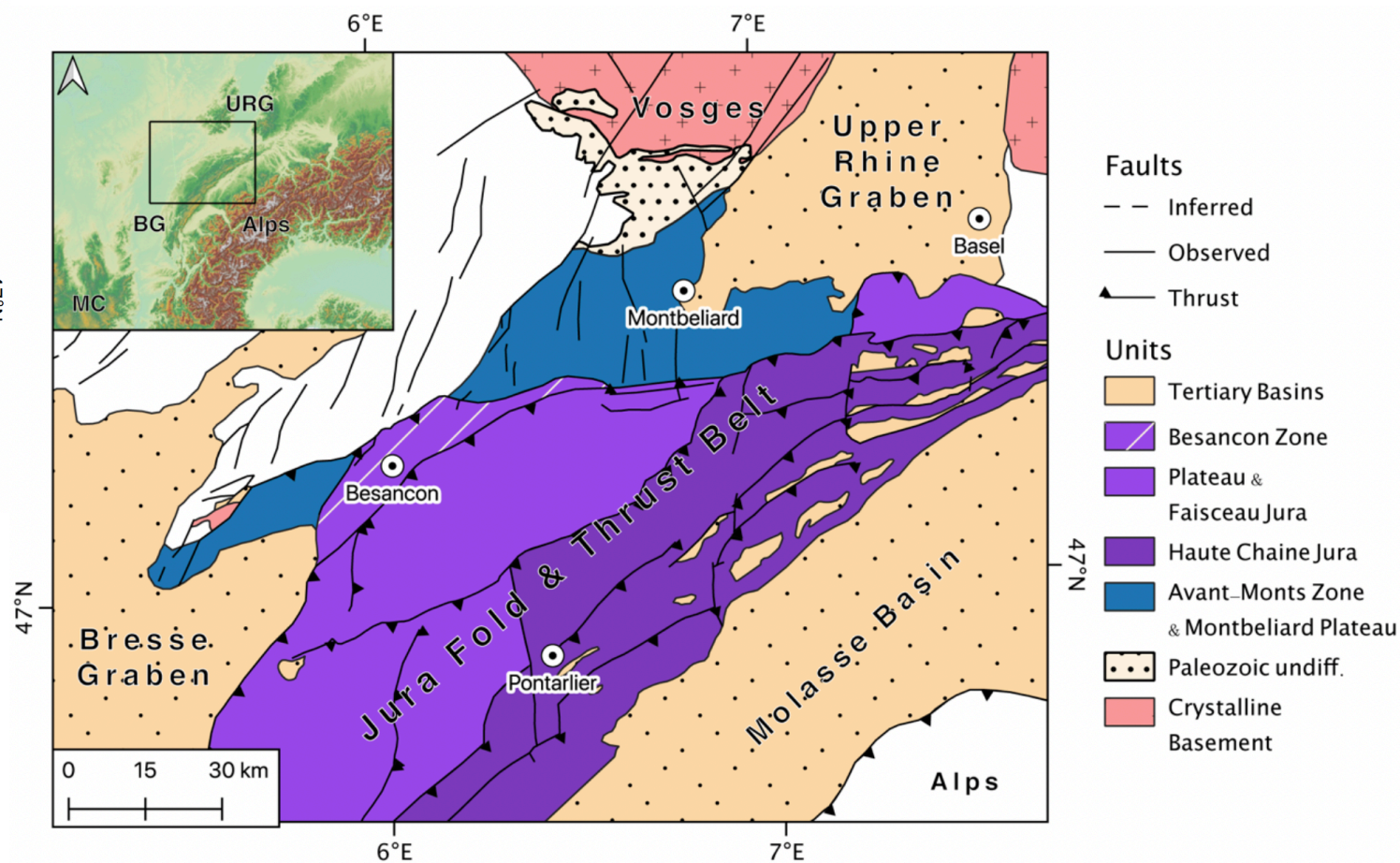
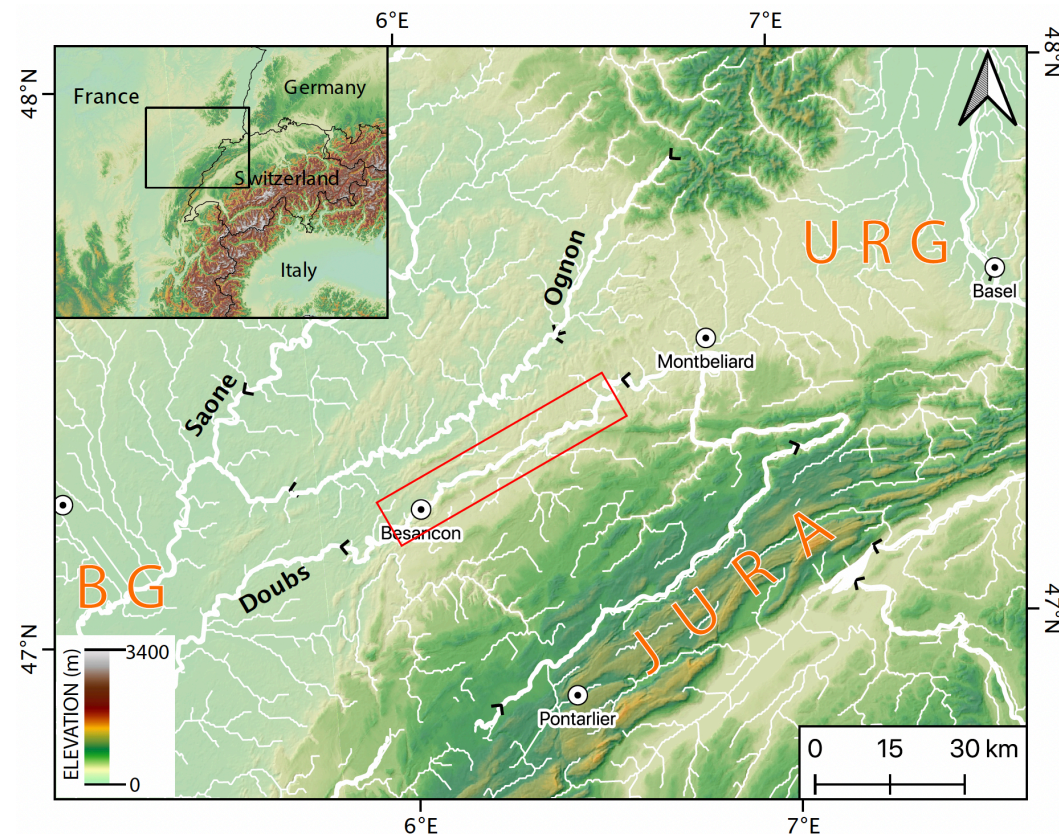
(c)



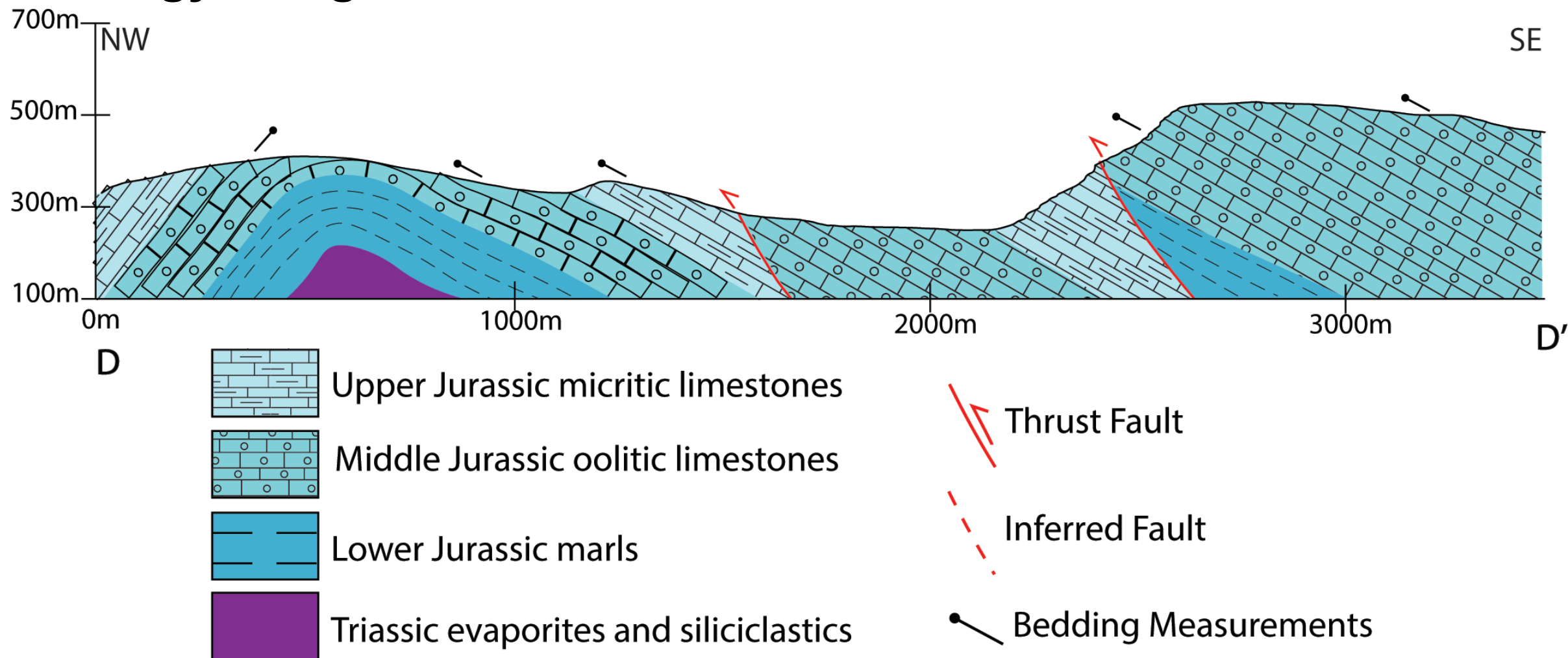
Madritsch et al., 2010b



The Doubs River and the Besançon Zone



Geology along the Doubs River



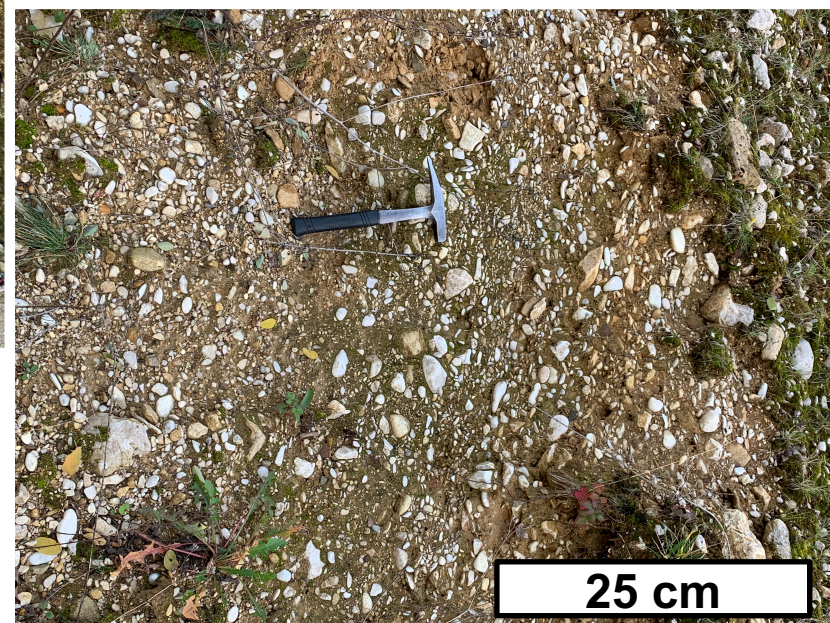
Fluvial Terraces along the Doubs River: T0



Tread Elevation = 1.5m
Strath Elevation = 1m



Fluvial Terraces along the Doubs River: T1



Tread Elevation = $\leq 25\text{m}$
Strath Elevation = $\sim 5\text{m}$

Fluvial Terraces along the Doubs River: T2 (Clerval)



Tread Elevation = ?
Strath Elevation = ~20m



OSL Dating Results (T1)

Sample	n	Water (%)	D (Gy/ka)	De (Gy)	Age (ka)
DEL-01-Q	24	10 ± 5	1.223 ± 0.06	79.92 ± 10.19	65.4 ± 8.9
		15 ± 5	1.168 ± 0.056		68.4 ± 9.3
		20 ± 5	1.119 ± 0.052		71.4 ± 9.7
DEL-02-Q	16	10 ± 5	1.378 ± 0.063	58.87 ± 12.08	42.7 ± 9.0
		15 ± 5	1.315 ± 0.058		44.8 ± 9.4
		20 ± 5	1.259 ± 0.054		46.8 ± 9.6

Sample	n	Water (%)	D (Gy/ka)	De (Gy)	Age (ka)	g (%)	Unfaded Age (ka)
DEL-01-F	15	10 ± 5	2.067 ± 0.152	277.79 ± 23.9	134.3 ± 15.2	0.62 ± 0.58	142.1 ± 18.1
		15 ± 5	2.008 ± 0.15		138.3 ± 15.8		146.3 ± 18.6
		20 ± 5	1.955 ± 0.148		142.1 ± 16.3		150.3 ± 19.1
DEL-02-F	15	10 ± 5	2.166 ± 0.095	95.16 ± 17.59	43.9 ± 8.4	0.62 ± 0.61	46.3 ± 9.2
		15 ± 5	2.098 ± 0.092		45.4 ± 8.6		47.8 ± 9.6
		20 ± 5	2.037 ± 0.089		46.7 ± 8.9		49.3 ± 9.8
DEL-02-F-SG	55	10 ± 5	2.166 ± 0.095	69.42 ± 8.87	32.1 ± 4.3	0.62 ± 0.61	33.8 ± 5.0
		15 ± 5	2.098 ± 0.092		33.1 ± 4.5		34.9 ± 5.1
		20 ± 5	2.037 ± 0.089		34.1 ± 4.6		35.9 ± 5.3

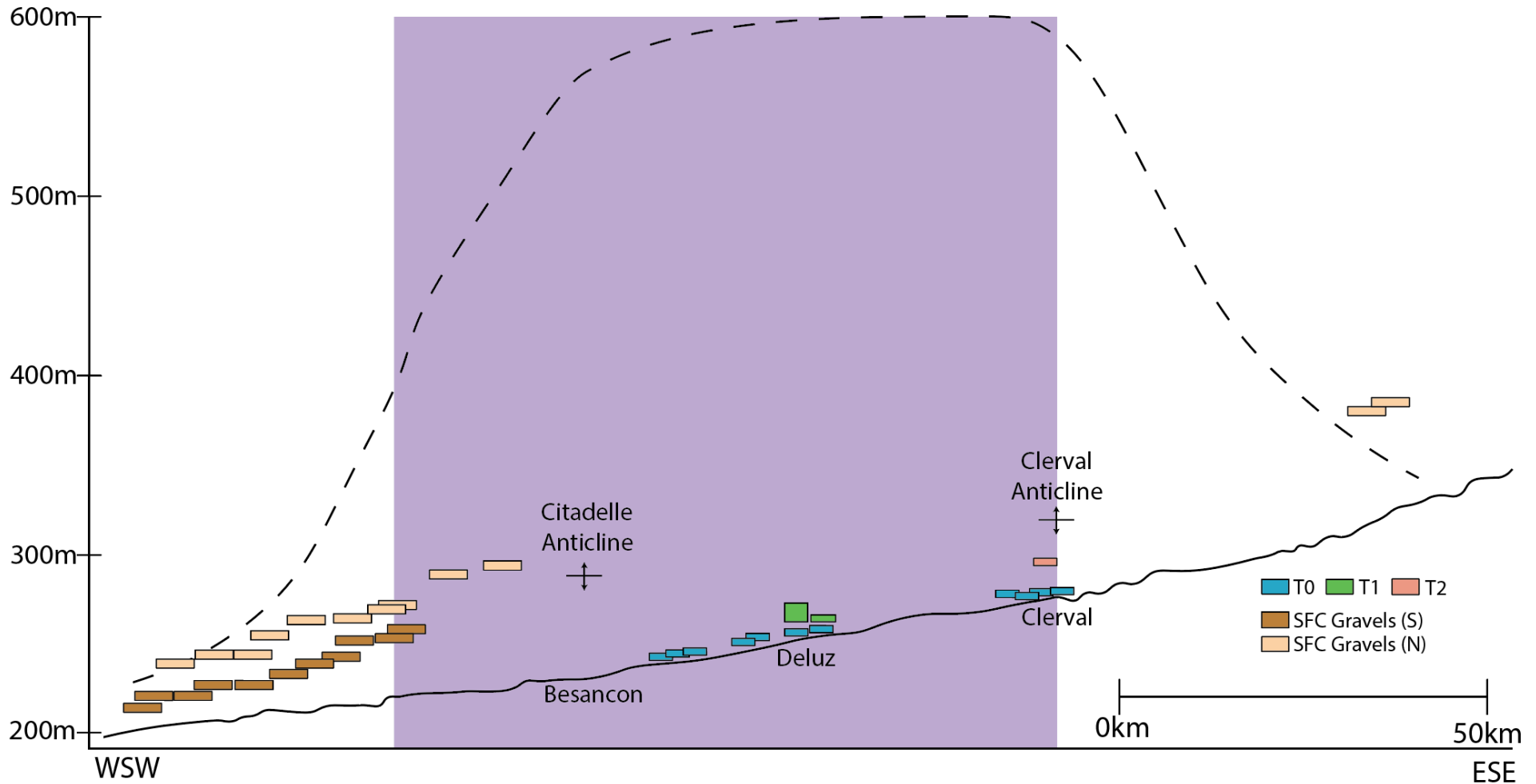


OSL Dating Results

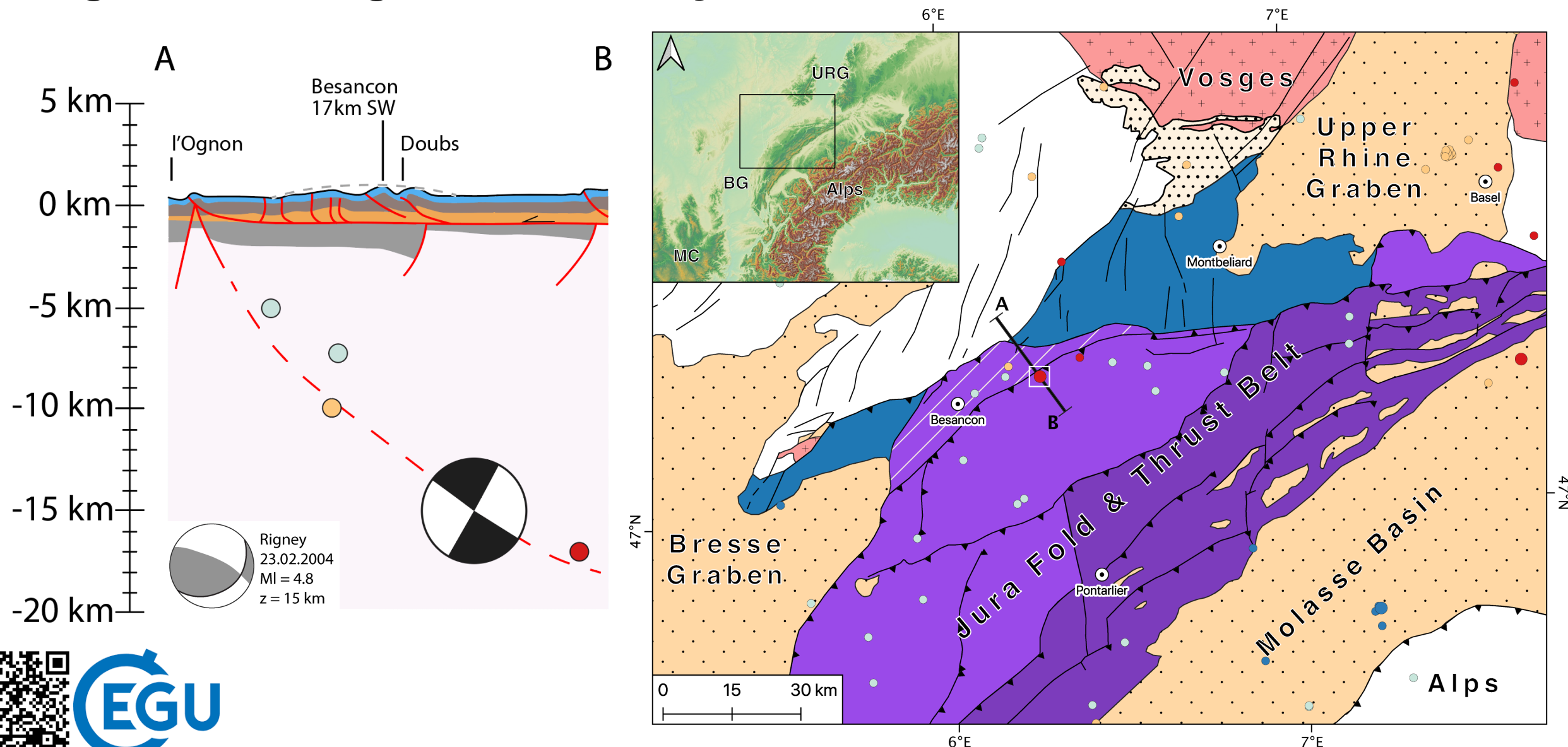
- T1 (5m) dated at 35 ± 5 ka gives an uplift rate of **0.14mm/yr**
 - *0.17 mm/yr from Madritsch et al. 2010a
 - *onset of Last Glacial Maximum = $\sim 33 - 26.5$ ka
- Assuming that the uplift rate is stable through time:
 - T0 (1m) = ~ 7 ka (same age as in Madritsch et al. 2010a)
 - T2 (20m) = ~ 140 ka
- T2 gives an interesting computed age, as the “oldest” ages from T1 are ~ 145 ka
 - *Penultimate Glacial Maximum = ~ 140 ka



Fluvial Terraces along the Doubs River



Regional Setting of the Besançon Zone



References

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