

How does lithospheric strength, mantle hydration and slab flexure relate to seismicity in the southern Central Andes?

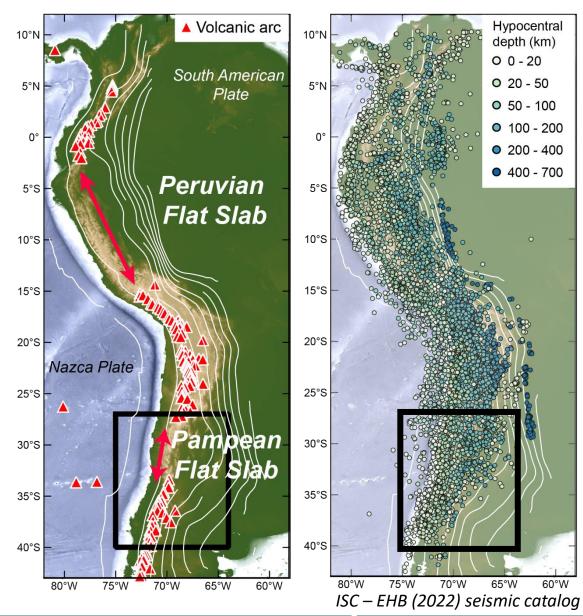
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Motivation



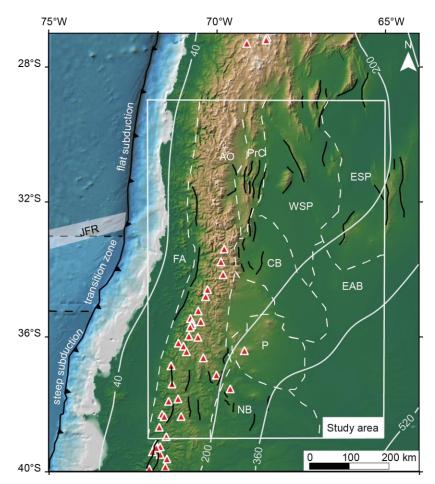
- The Southern Central Andes host one of the most actively seismic areas in South-America
- Oceanic plate changes from flat (< 5°) in the north to steep (~30°) subduction in the south

How do continental-plate structural inheritance and oceanic-plate geometry affect the localization of seismicity?

Introduction Methods Results & discussion Conclusions Conclusions

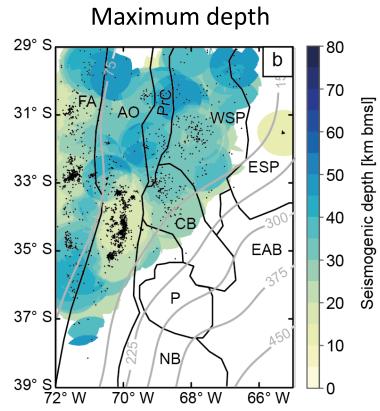
Observations: continental plate seismicity





Abbreviations of main tectonic provinces and features: AO = Andean orogen, CB = Cuyo Basin, ESP = Eastern Sierras Pampeanas, EAB = extra-Andean basins, FA= forearc, NB = Neuquén Basin, P = Payenia volcanic province, Prc = Precordillera, WSP = Western Sierras Pampeanas. JFR= Juan Fernández Ridge

Introduction



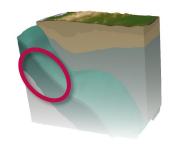
Reviewed bulletin of the International Seismological Centre (2021)

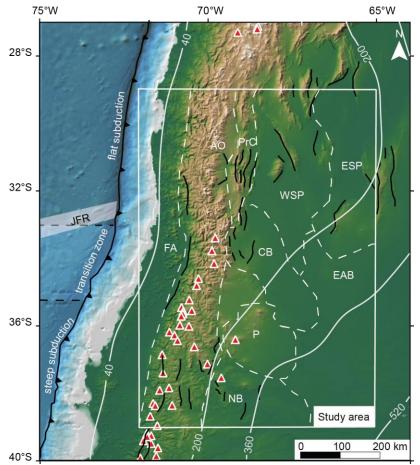
- Shallow seismicity in forearc and central orogen
- Deep seismicity in the Sierras
 Pampeanas

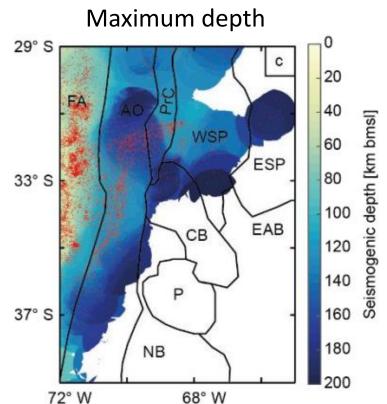
Rodriguez Piceda et al. (2022, GCubed)

Methods Results & discussion Conclusions

Observations: oceanic plate seismicity





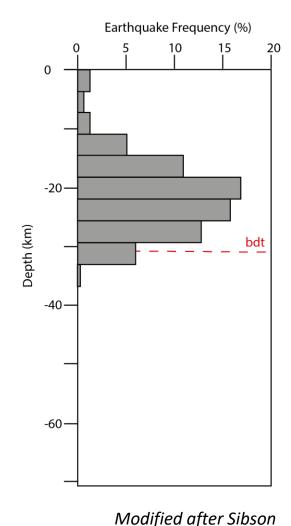


Reviewed bulletin of the International Seismological Centre (2021)

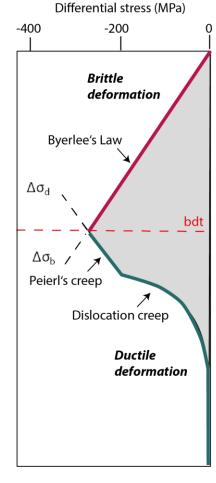
- Shallow seismicity below the forearc
- Deep seismicity within the flat slab segment

Rodriguez Piceda et al. (2022; under review, Nat. Commun. Earth Environ.)

Seismic depth distribution depends on strength



(1983), Scholz (1998)



Frictional strength

Byerlee's Law

$$\Delta \sigma_{\rm b} = f_{\rm f} (1 - f_{\rm p}) \rho_{\rm b} gz$$

Frictional parameters (incl. fluids) density, thickness

Viscous strength

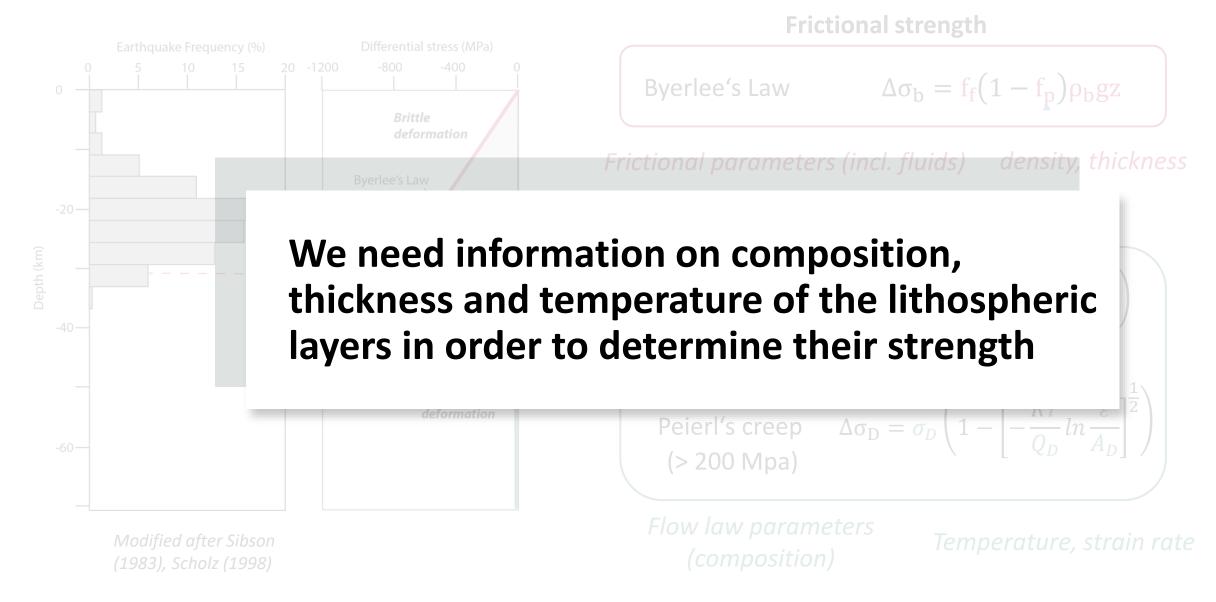
Dislocation creep
$$\Delta \sigma_{\rm d} = \left(\frac{\dot{\varepsilon}}{A_p}\right)^{\frac{1}{n}} \cdot \exp\left(\frac{Q_p}{nRT}\right)$$
 (< 200 Mpa)

Peierl's creep
$$\Delta \sigma_{\rm D} = \sigma_{\!D} \left(1 - \left[-\frac{RT}{Q_D} ln \frac{\dot{\varepsilon}}{A_D} \right]^{\frac{1}{2}} \right)$$
 (> 200 Mpa)

Flow law parameters (composition)

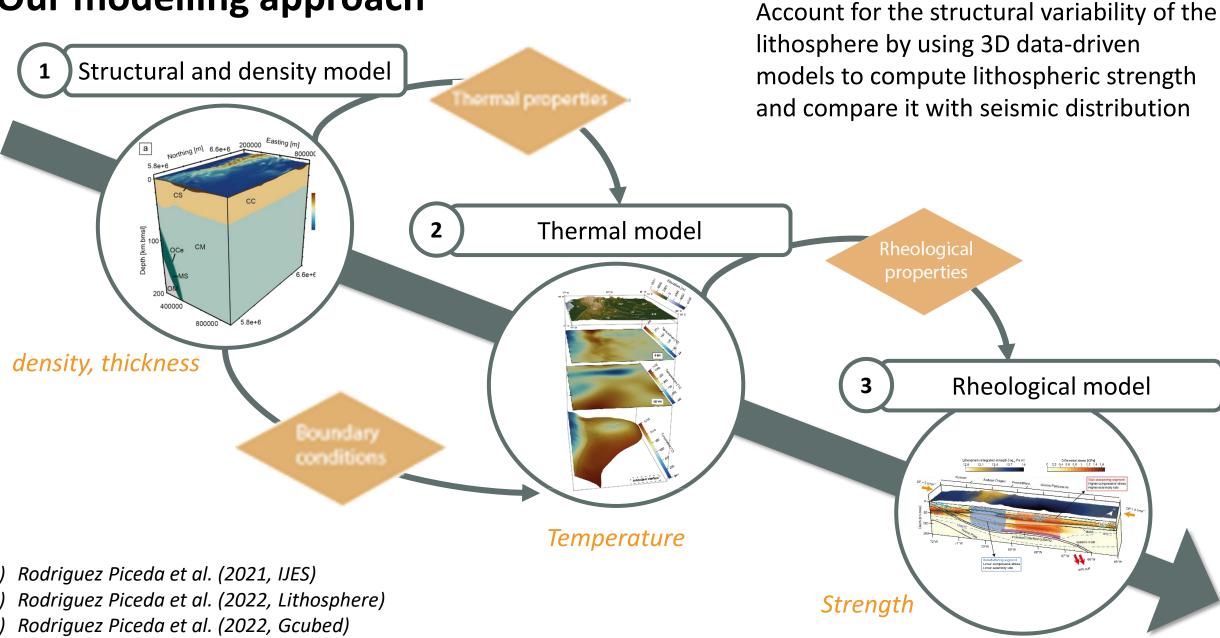
Temperature, strain rate

Seismic depth distribution depends on strength



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Our modelling approach



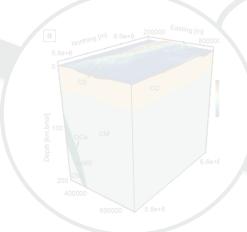
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Conclusions

Introduction Methods Results & discussion

Our modelling approach

1 Structural and density model



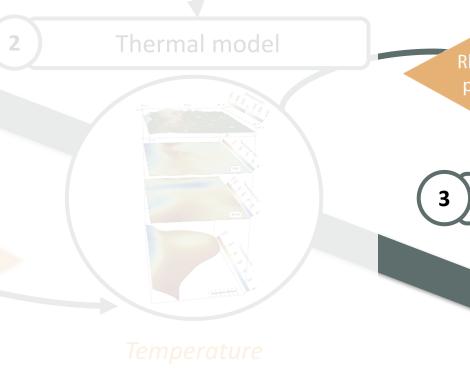
density, thickness

(1) Rodriguez Piceda et al. (2021, IJES

(2) Rodriguez Piceda et al. (2022, Lithosphere)

(3) Rodriguez Piceda et al. (2022, Gcubed)

Account for the structural variability of the lithosphere by using 3D data-driven models to compute lithospheric strength and compare it with seismic distribution

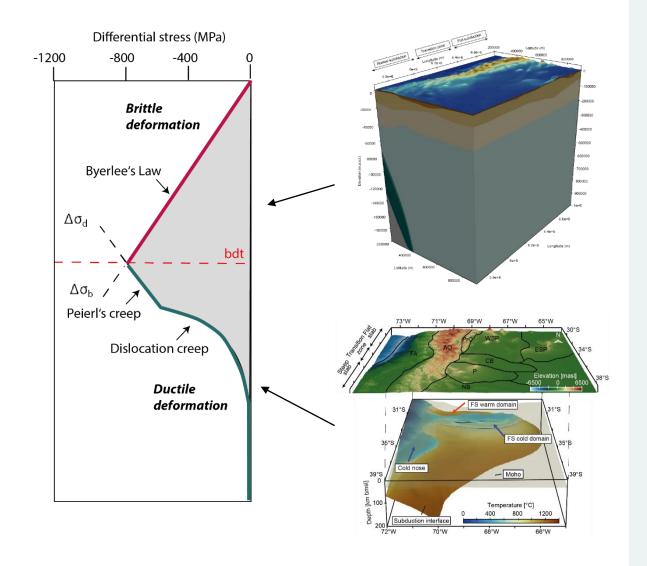


Rheological model Strength

Conclusions



Rheological model



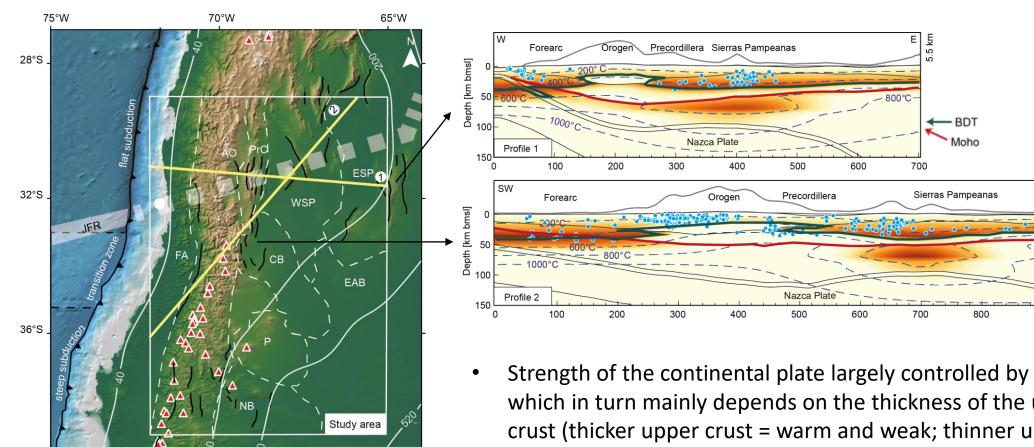
- Computation of yield strength and integrated strength
- Geometries and densities from 3D structural model (Rodriguez Piceda et al., 2021, IJES)
- Temperatures from 3D thermal model (Rodriguez Piceda et al., 2022, Lithosphere)
- Assigned frictional and creep parameters background strain rate

Rodriguez Piceda et al. (2022, Gcubed)

© <u>(1)</u>

Continental plate seismicity

Lithospheric strength & continental plate seismicity



- Strength of the continental plate largely controlled by the thermal field which in turn mainly depends on the thickness of the upper radiogenic crust (thicker upper crust = warm and weak; thinner upper crust = cold and strong)
- Brittle-ductile transition bounds the depth extent of the seismogenic zone

Conclusions

- Deeper earthquakes in cold and strong domains
- Shallower earthquakes in warm and weak domains

Rodriguez Piceda et al. (2022, Gcubed)



strong

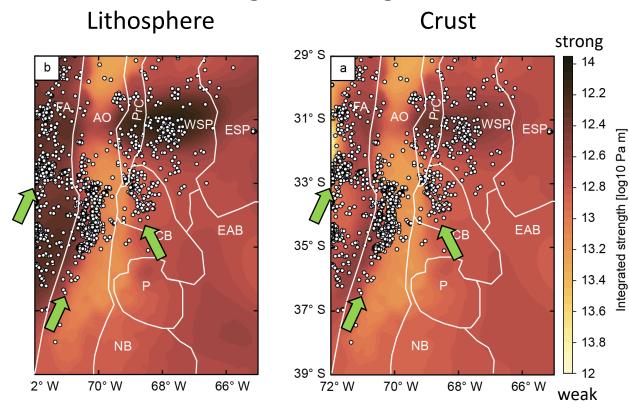
0.2

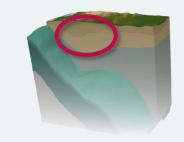
weak

NE Ē

Strength & continental plate seismicity

Integrated strength





- Lateral strength contrast needed to localize seismicity
- Seismicity at the boundaries between strong and weak crustal domains
- Sierras Pampeanas: Seismicity within strong flat slab domain (yield strength > horizontal tectonic stresses)

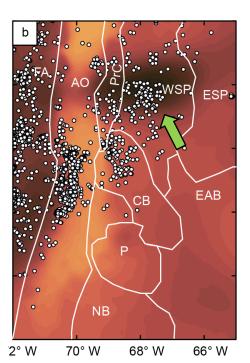
Rodriguez Piceda et al. (2022, Gcubed)

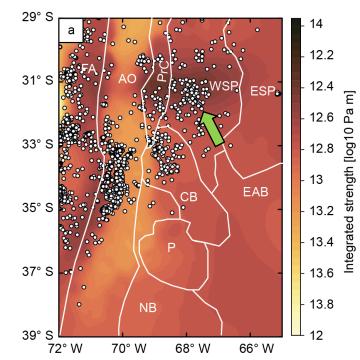
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Strength & continental plate seismicity

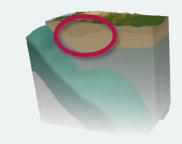
Integrated strength

Lithosphere





Crust



- Lateral strength contrast needed to localize seismicity
- Seismicity at the boundaries between strong and weak crustal domains
- Sierras Pampeanas: Seismicity within strong flat slab domain (yield strength > horizontal tectonic stresses)

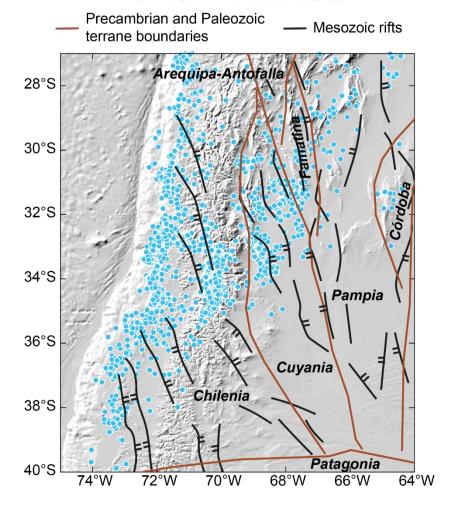
Additional mechanisms to weaken the lithosphere in the Sierras Pampeanas?

Rodriguez Piceda et al. (2022, Gcubed)

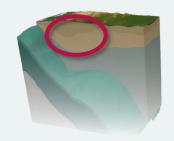


Continental plate seismicity

Pre-Andean inherited structures



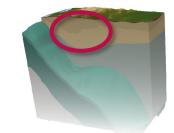
Modified Ramos (2009), Ramos et al. (2010), Wimpenny (2022)

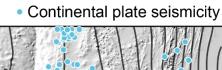


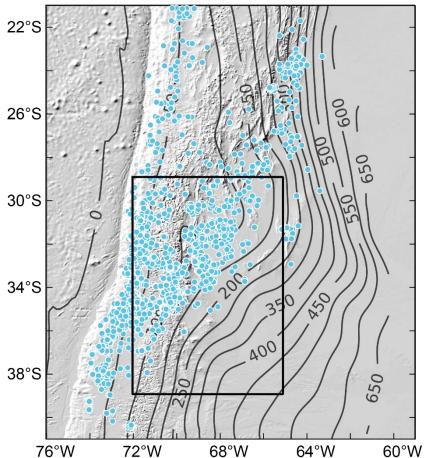
- 1st weakening mechanism: crustal-scale discontinuities (terrane sutures, Mesozoic rifts)
- However, not all inherited structures are seismically active

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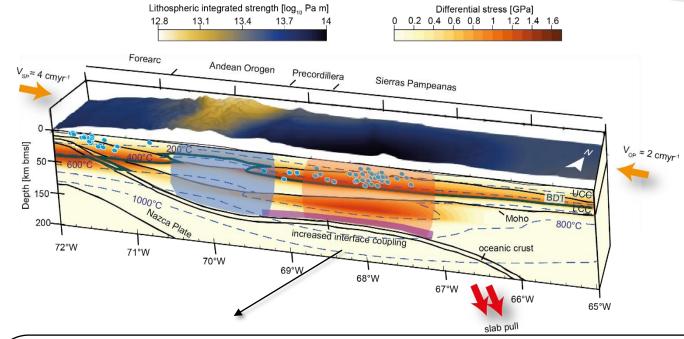
Continental plate seismicity in the Sierras Pampeanas







2nd weakening mechanism: oceanic slab



Slab steepening segment

- Increased interplate coupling allows inland transmission of stresses
- Increased negative buoyancy when slab returns to steep subduction weakens continental plate

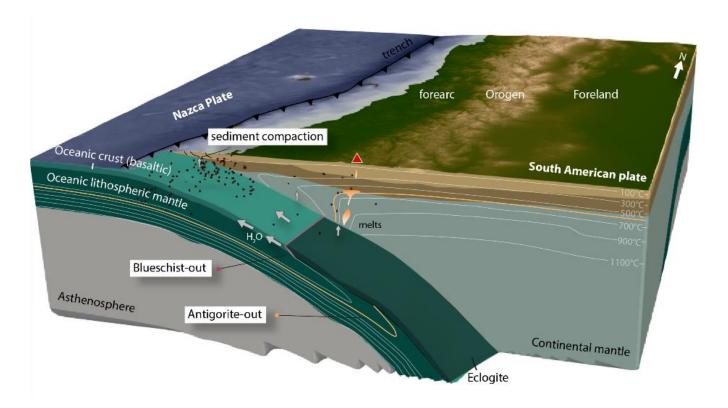
Higher seismicity rate

Rodriguez Piceda et al. (2022, Gcubed)

Introduction

Oceanic plate seismicity

Oceanic plate seismicity and dehydration



Frictional weakening from release of fluids related to:

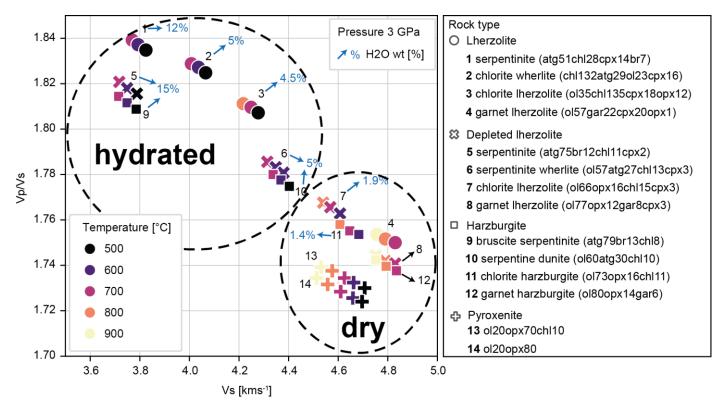
- Sediment consolidation (interface)

 Saffer & Marone (2003), Han et al. (2017), etc.
- Metamorphic reactions (oceanic crust and mantle)
 Kirby (1995), Peacock (2001), Ferrand et al. (2017), etc.

Affect state of hydration of slab and overriding plate mantle

Hydrated regions more seismically active than dry regions

Oceanic plate seismicity and dehydration

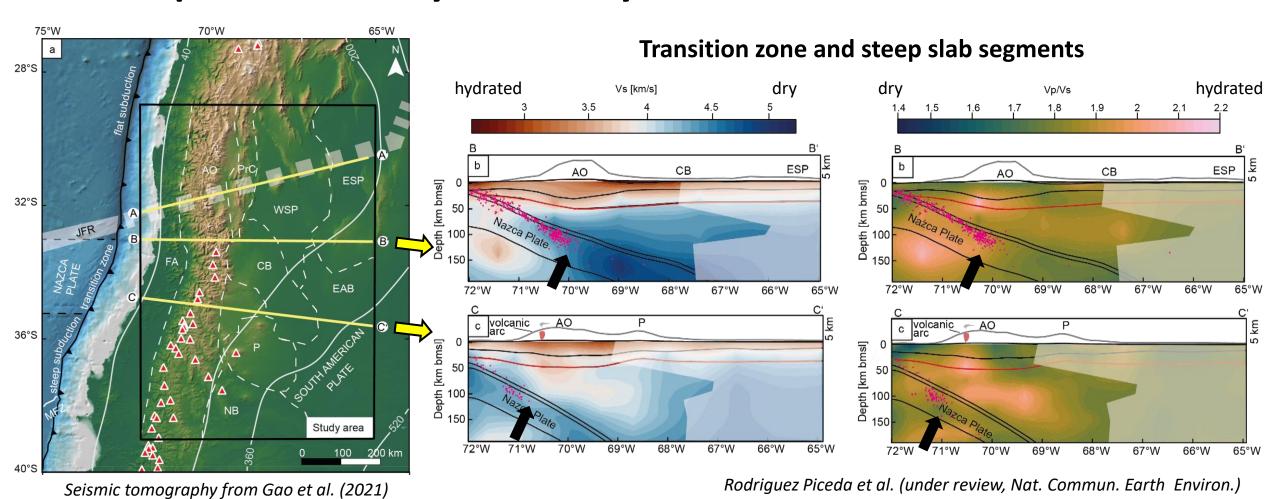


Modified after Linkimer et al. (2020)

- Vp/Vs ratio and Vs as proxy for the state of hydration of the mantle
- Dry regions characterized by high Vs and low Vp/Vs
- Hydrated regions characterized by low Vs and high Vp/Vs



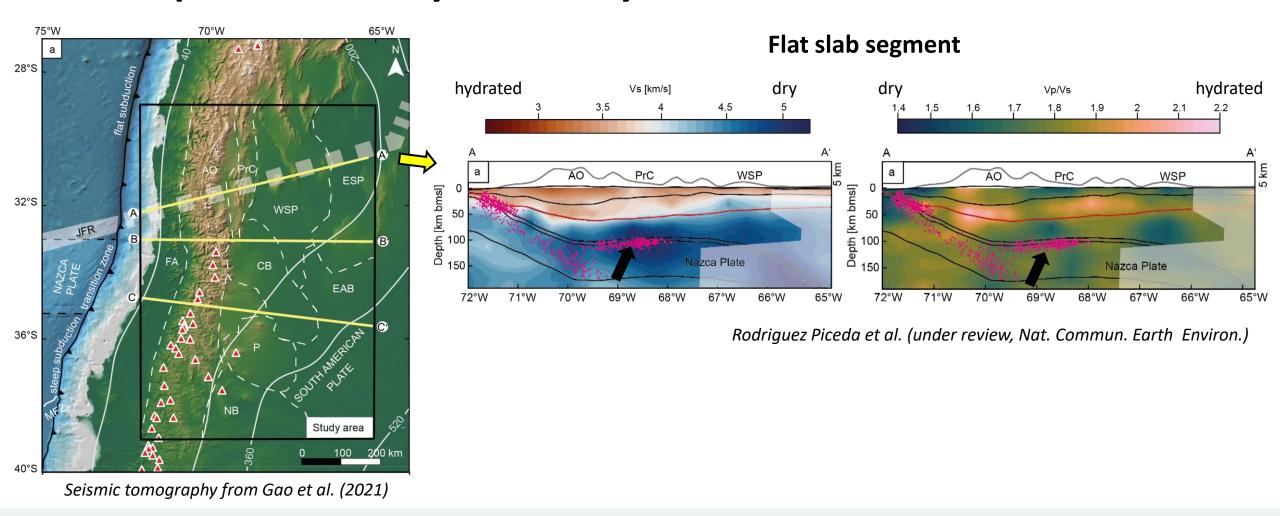
Oceanic plate seismicity and deshydration



Seismicity within hydrated mantle (low Vs; high Vp/Vs) → fluid mediated

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Oceanic plate seismicity and deshydration

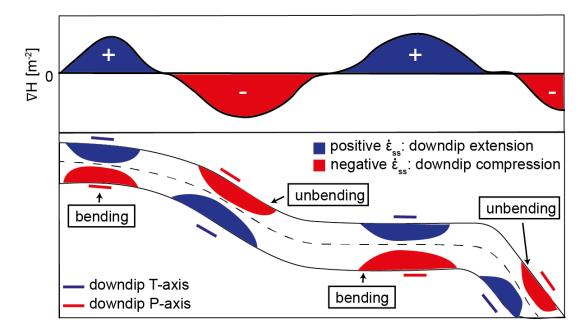


Seismicity within dry mantle (high Vs; low Vp/Vs) → not fluid mediated

Hypothesis: seismicity in the flat slab driven by flexural stresses

CC (I)

Oceanic plate seismicity and flexure



Modified after Sandiford et al. (2020)

Relationship between curvature gradient (∇H)
 and axis orientation of focal mechanisms

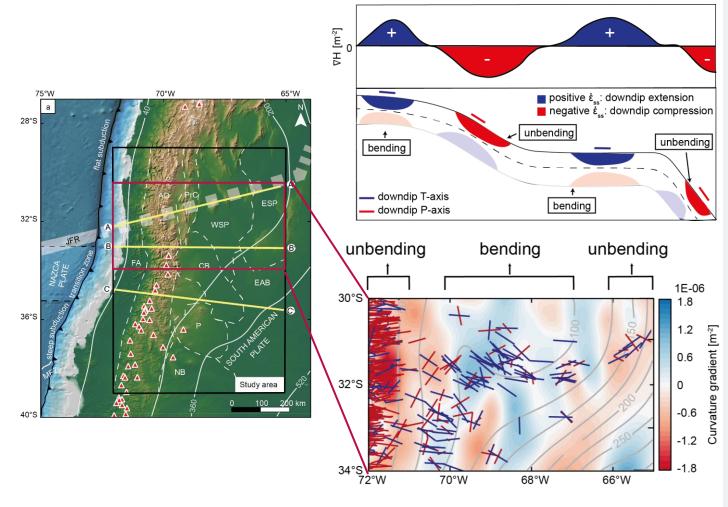
Engdahl and Scholz (1977), Isacks and Molnar (1971)

Top slab

- Positive curvature gradient: downdip T-axis
- Negative curvature gradient: downdip P-axis



Oceanic plate seismicity and flexure



GCMT catalog (Ekström et al., 2012)

- Correlation between orientation of principal axes of focal mechanisms and curvature gradient of the slab
- Slab seismicity in the flat slab segment driven by flexural stresses due to variations in dip angle

Rodriguez Piceda et al. (under review, Nat. Commun. Earth Environ.)

(C) (1)

Conclusions

How do continental-plate structural inheritance and oceanic-plate geometry affect the localization of seismicity?

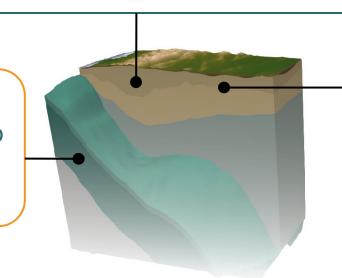
Continental plate seismicity

• By long-term strength variations related to the layer thickness (upper radiogenic crust)

Oceanic plate seismicity

- By deshydration of oceanic slab
- Flat slab: by internal flexural stresses

Introduction



Sierras Pampeanas seismicity

- By shallow inherited faulting
- By propagating the deformation to the east and exherting additional forces related to steepening

Methods Results & discussion Conclusions Conclusions

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