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Identifying efficient ensemble perturbations for initializing subseasonal-to-seasonal prediction

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Initialization of S2S predictions

- S2S predictions is beyond atmosphere predictability limit
 - Coupled Earth system models must be used
- Usually done with ensemble:
 - How to initialize them consistently to obtain reliable results?
- Already tested, use of local properties:
 - Bred vectors (Peña & Kalnay, 2004; Yang et al., 2008; O'Kane et al., 2019)
 - Backward Lyapunov Vectors (BLVs, related to singular vectors) (Vannitsem & Duan, 2020) to tune the initialization of the models.

In the present work, we study the projections of the initial conditions on mainly:

- Covariant Lyapunov vectors (CLVs) and their adjoint
- Dynamical Mode Decomposition adjoint modes
- Perron-Frobenius Mode Decomposition adjoint modes for predictions experiments with a low-order coupled ocean-atmosphere model (MAOOAM).

Dynamical Modes Decomposition (DMD)

• Considering 2 collections of states of a dynamical system $X = [\mathbf{x}_0 \dots \mathbf{x}_{K-1}]$ and $Y = [\mathbf{x}_1 \dots \mathbf{x}_K]$, then one define

$$M^{\mathrm{DMD}} = Y X^{+}$$
 where X^{+} is the pseudoinvere

as the DMD decomposition of the observable g(x)=x the system. Related to Linear Inverse Modeling

Penland (1989)

- The left eigenvectors \mathbf{w}_i of M^{DMD} provides approximation of the system's Koopman operator eigenfunctions and are called *adjoint* DMD modes. (Tu et al., 2014)
- The Koopman K operator is an ∞ -dimensional linear operator propagating the observable of the system. For an observable g:

$$\mathcal{K}^{ au}\,g(m{x}) = g\left(m{\Phi}^{ au}(m{x})
ight)$$
 where ϕ is the flow of the system $\dot{m{x}} = m{f}(m{x})$

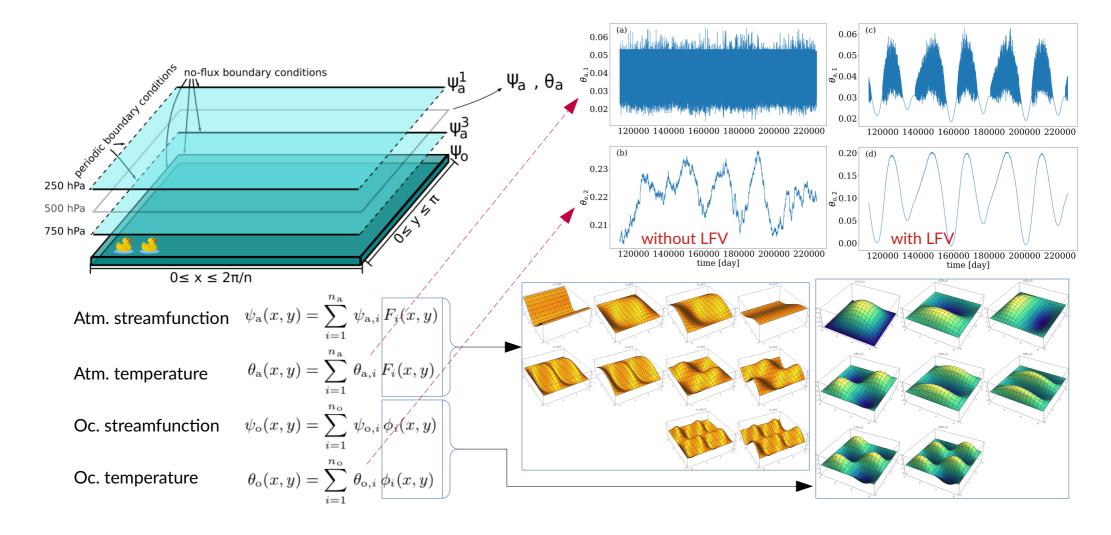
The action of the Koopman operator can then be approximated with the DMD as

$$\sum_{i=1}^{P} \boldsymbol{c}_{i}^{\text{DMD}} \, \lambda_{i}^{\text{DMD}} \, \boldsymbol{w}_{i}^{*} \quad \boldsymbol{x}$$
 where the $\boldsymbol{c}_{i}^{\text{DMD}}$ are modes depending on the observable g

The \boldsymbol{w}_{i} define approximate invariant manifolds for the Koopman operator

MAOOAM: an ocean-atmosphere coupled model

QG atmosphere coupled to a shallow-water ocean



Experiments design

N ensemble forecasts along a reference trajectory:

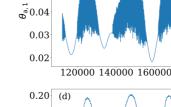
Control run IC:
$$m{x}_n^{ ext{ctrl}}(0) = m{x}_n(0) + m{\delta} m{x}_0^{ ext{ctrl}}$$
 n = 1,...,N

$$\boldsymbol{y}_{m,n}(0) = \boldsymbol{x}_n^{\mathrm{ctrl}}(0)$$

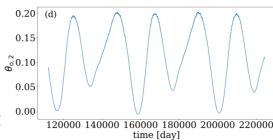
Perfect ensemble IC:
$$\boldsymbol{y}_{m,n}(0) = \boldsymbol{x}_n^{\text{ctrl}}(0) + \boldsymbol{\delta x}_0^m$$
 , $m = 1, \dots, M-1$

$$\delta x_0^{M} = 0$$

Same distribution U[$-\epsilon/2,\epsilon/2$]



0.05



Experiments

Projection onto subspaces spanned by selected vectors:

Projector:

 $\Pi = B(B^*B)^{-1}B^*$ where B is the column matrix of the selected vectors

Projected perturbations: $\delta x_0^{\prime m} = \Pi \delta x_0^m$

$$oldsymbol{\delta x}_0^{\prime m} = \Pi \, oldsymbol{\delta x}_0^m$$

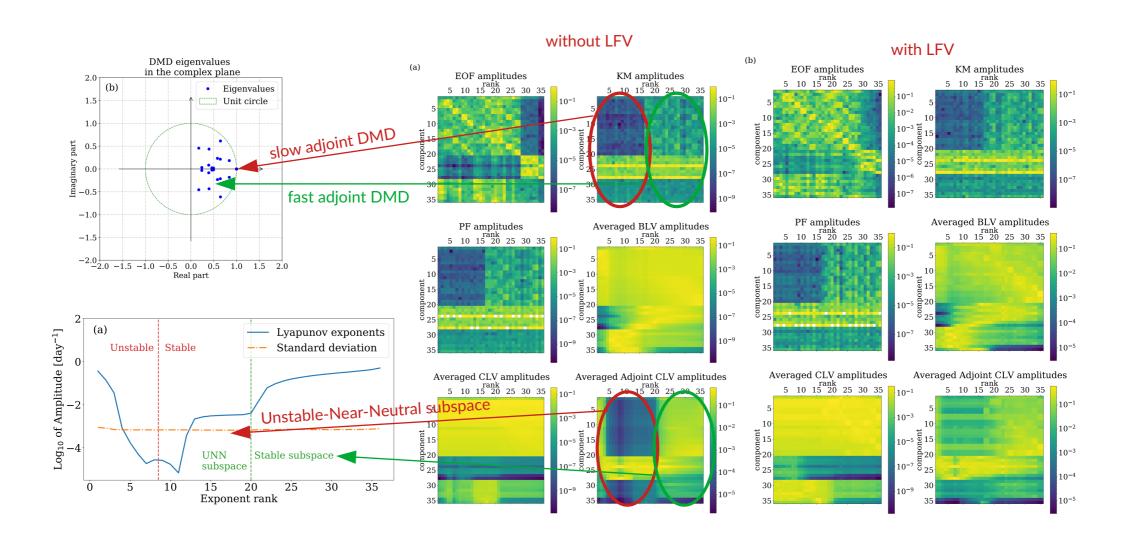
New ensembles:
$$m{y}_{m,n}'(0) = m{x}_n^{ ext{ctrl}}(0) + m{\delta}m{x}_0'^m$$
 , $m=1,\ldots,M$

Goal:

→ Obtain forecasts as reliable as the ones provided by the perfect ensembles.

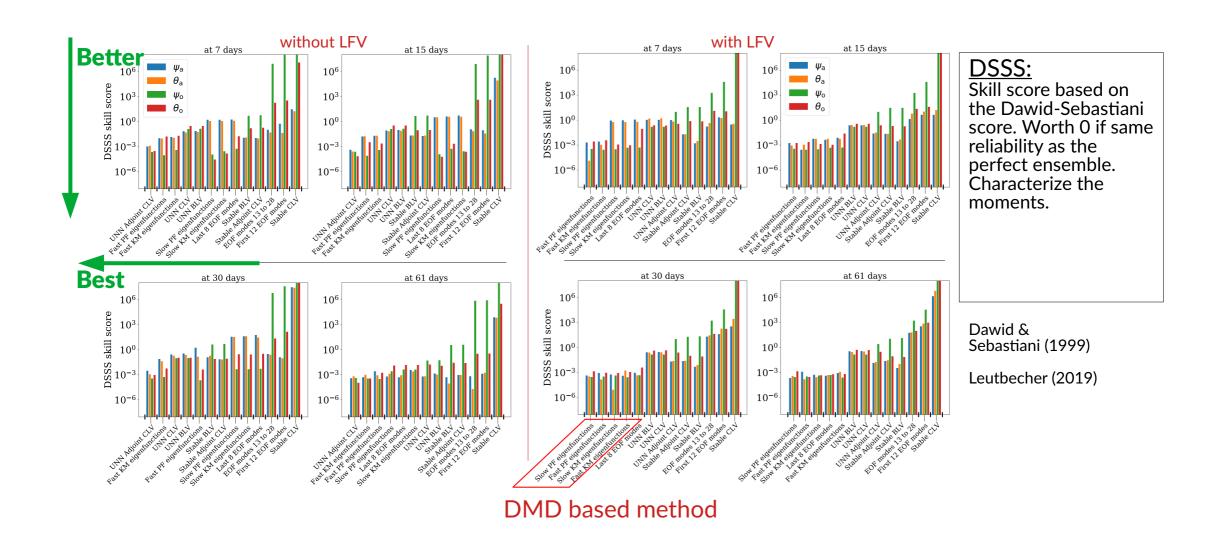


Selected bases: analysis of the reference trajectory





Results: DSSS skill scores at different lead times





Key results:

- Approximated KM and PF eigenfunctions obtained using DMD provide reliable ensemble forecasts, and are "easy" to compute.
- Adjoint CLVs also provide reliable forecasts (sometimes the most reliable ones), but are notably hard to compute.
- A consistent link exists between the two frameworks, which explains the results.
- Results seem to not depend on the regime (with or withour LFV)
- Results of Vannitsem & Duan (2020) with the BLVs can also be explained with the DMDs.

Forthcoming developments:

- Many, but most notably, replication of the study with a higherdimensional system, with a non-trivial dimensionality reduction to make DMD tractable.
- Ultimately, development of the approach in a realistic S2S framework.

Notice finally that the first forward Liapunov vector is not the direction of the fastest growing perturbation for large future time, since all perturbations contained within $F_1^+(t_1) \setminus F_2^+(t_1)$ eventually grow at the same average rate λ_1 . In other words, perturbations starting from, say, the subspace spanned by $f_1^+(t_1)$ and $f_2^+(t_1)$ are, in practice, not growing faster than almost any random perturbation. This indicates that, for long-range forecasts, the use of singular vectors for initial perturbations looses its relevance.

Legras & Vautard (1996) A guide to Liapunov vectors. ECMWF seminar on predictability.

Takeaway message

To initialize ensemble forecast for S2S:

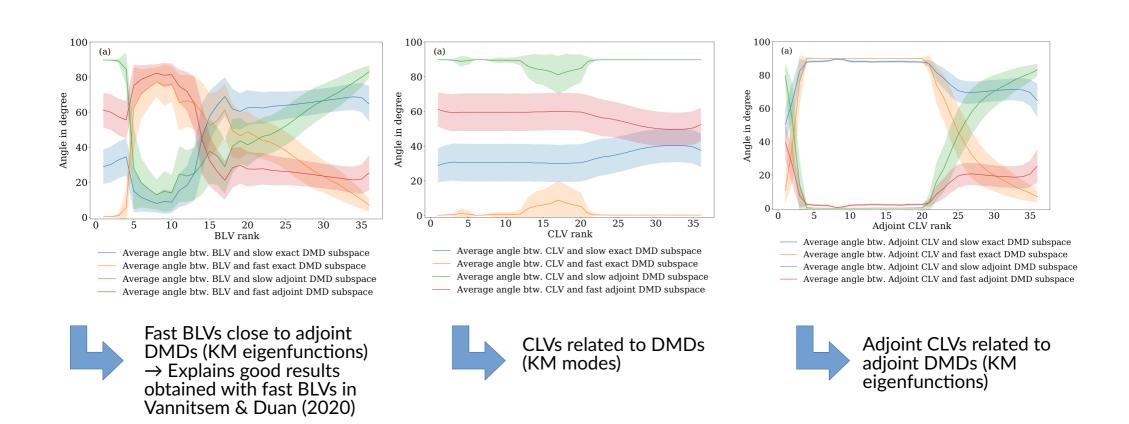
- Local stability properties are less important
- However, ensembles initial conditions can be constructed in more relevant subspaces
 - → constructed with *measures* of the system valid at these timescales (through DMD)

<u>Reference</u>

Jonathan Demaeyer, Stephen G. Penny, Stéphane Vannitsem. Identifying efficient ensemble perturbations for initializing subseasonal-to-seasonal prediction. JAMES, in press, doi: 10.1029/2021MS002828



Backup: Correspondence between bases



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