An aerial photograph of the ocean surface, showing a dense field of surface drifters. The drifters are arranged in complex, swirling patterns, indicating the presence of oceanic eddies and currents. The water is a deep blue, and the drifters appear as lighter, textured patches against the darker water.

Direct evidence of an oceanic dual kinetic energy cascade and its seasonality from surface drifters

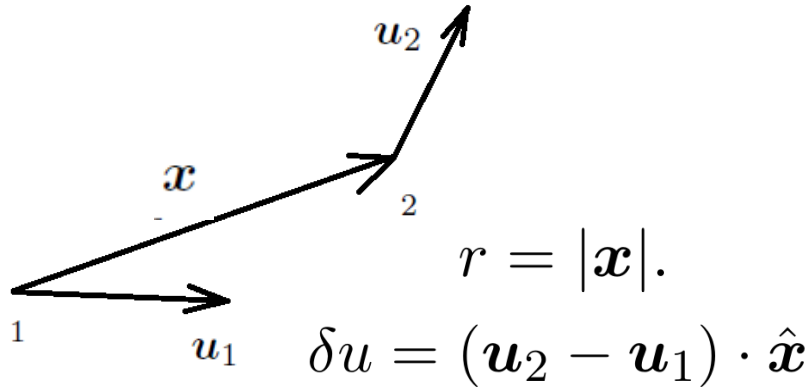
Jin-Han Xie (Peking University)

Dhruv Balwada (Columbia)

Raffaele Marino and Fabio Feraco (ENS de Lyon)

Inertial range theory

- 3D: Richardson (1922) → Kolmogorov (1941)'s *“4/5-law” for third-order structure function*.
(Two-point measurements)



$$l_f \gg r \gg l_\nu$$

$$\overline{\delta u^3} = -\frac{4}{5}\epsilon r$$

- Dissipation scale:

$$l_\nu = \left(\frac{\nu^3}{\epsilon}\right)^{1/4}$$

The “4/5-law” is a result calculated from NS (KHM) equation.

- 2D: Lindborg 1999, Bernard 1999, Yakhot 1999:

Energy inertial range:

$$\overline{\delta u^3} = \frac{3}{2}\epsilon r$$

Enstrophy inertial range:

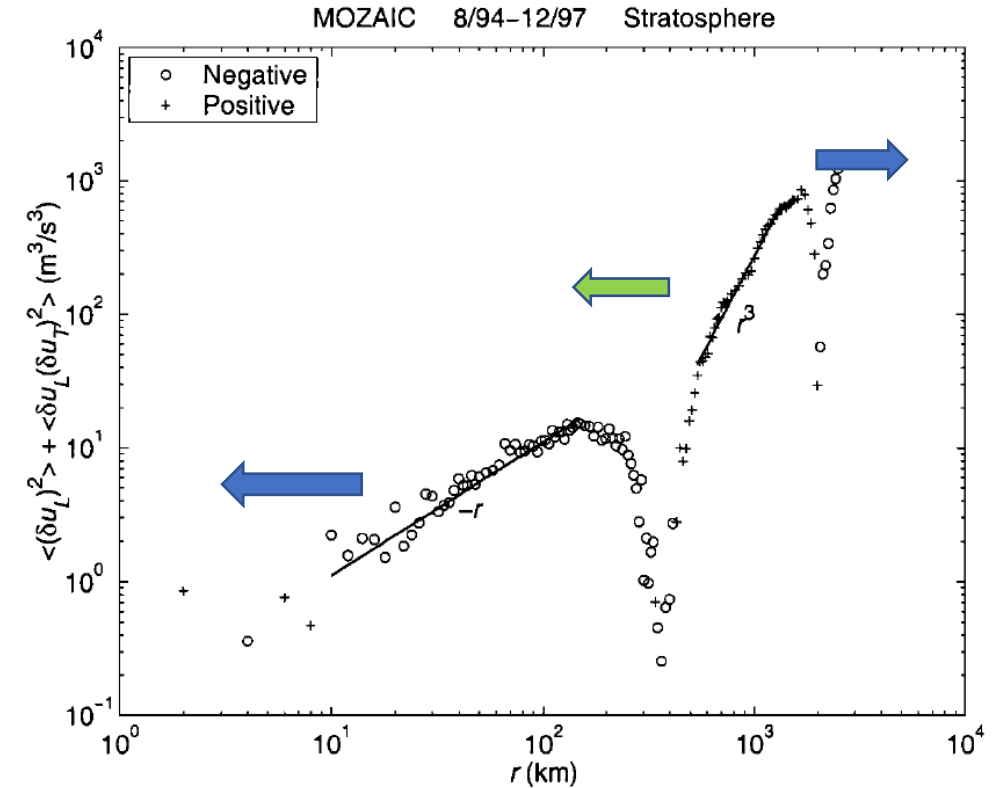
$$\overline{\delta u^3} = \frac{1}{8}\eta r^3$$

Defects of previous inertial-range theories

- † Previous theories apply to *local* inertial ranges, how to obtain a *global* description?
 - How to determine inertial ranges? (power-law regimes are asymptotic)
- † Previous theories apply to unidirectional energy transfer scenarios. How about the bidirectional energy transfer?
- † How to obtain the forcing scale?
 - The forcing information is always unknown but important to geophysics.

❖ Our solution: from Navier-Stokes equation derive

A global forcing-scale-resolving theory that captures the bidirectional energy transfer.



Third-order structure function obtained from MOZAIC data (Cho & Lindborg 2001)

Derivation of structure-function expression

(with X & Bühler 2019)

- In the KHM equation, under the assumption of isotropy, we obtain a new expression for the third-order structure function:

$$F = -\epsilon_u + (\epsilon_u + \epsilon_d) H(K - 1/l_f),$$

$$V(r) = -4r \int_0^\infty \frac{1}{K} F(K) J_2(Kr) dK.$$



$$V = 2\epsilon_u r - 4(\epsilon_u + \epsilon_d) l_f J_1(r/l_f).$$

$$\mathbf{V} = \overline{\delta \mathbf{u} |\delta \mathbf{u}|^2} = V(r) \mathbf{e}_r$$

$$V = \overline{\delta u_L^3} + \overline{\delta u_L \delta u_T^2}$$

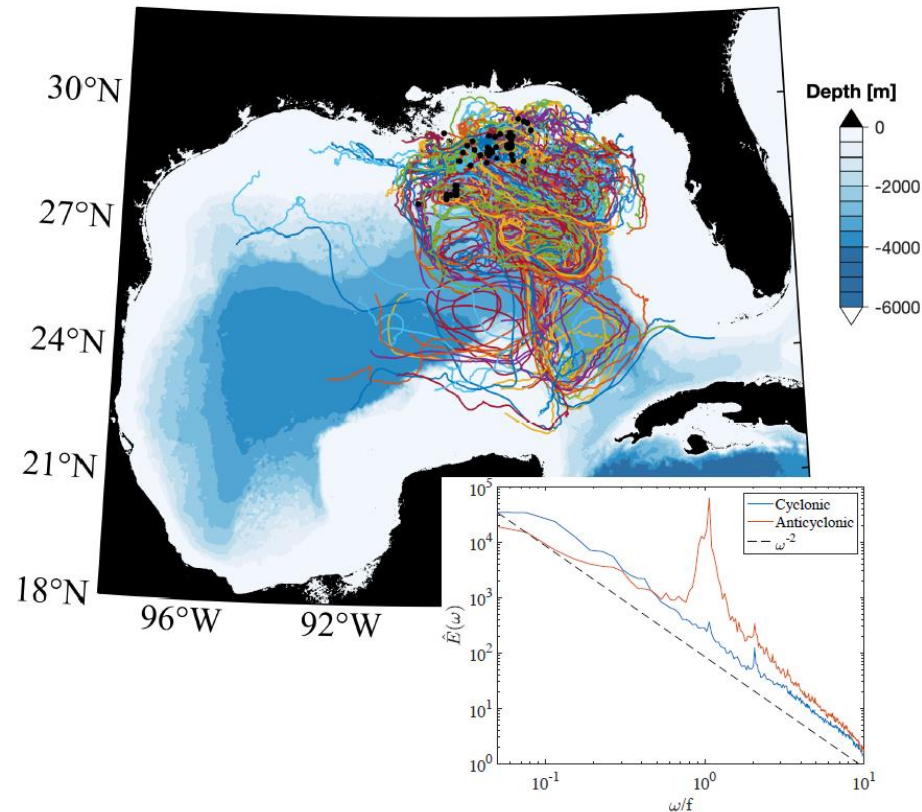
- (1) *bidirectional energy transfer* (2) *forcing scale l_f* (3) *Asymptotically recovers the inertial-range theories:*

$$V = \begin{cases} \underbrace{-2\epsilon_d r}_{\text{downscale energy}} + \underbrace{\frac{1}{4} \frac{\epsilon}{l_f^2} r^3}_{\text{"enstrophy"}} + O((r/l_f)^5), & \text{when } r/l_f \ll 1, \\ \underbrace{2\epsilon_u r}_{\text{upscale energy}} + O((r/l_f)^{-1/2}), & \text{when } r/l_f \gg 1. \end{cases}$$

Oceanic turbulence

- The energy injection scale and energy flux direction of oceanic flow is unknown.
- Can we obtain these information from measured data?

Grand Lagrangian Deployment (GLAD)
Summer 2012



Lagrangian Submesoscale Experiment (LASER)
Winter 2016

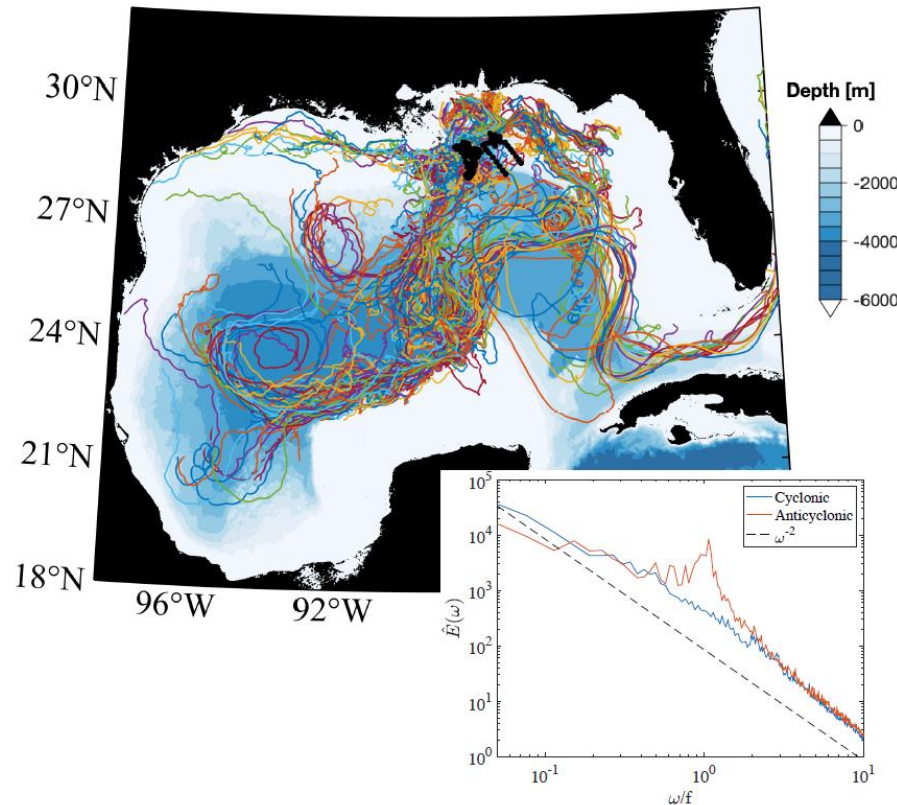
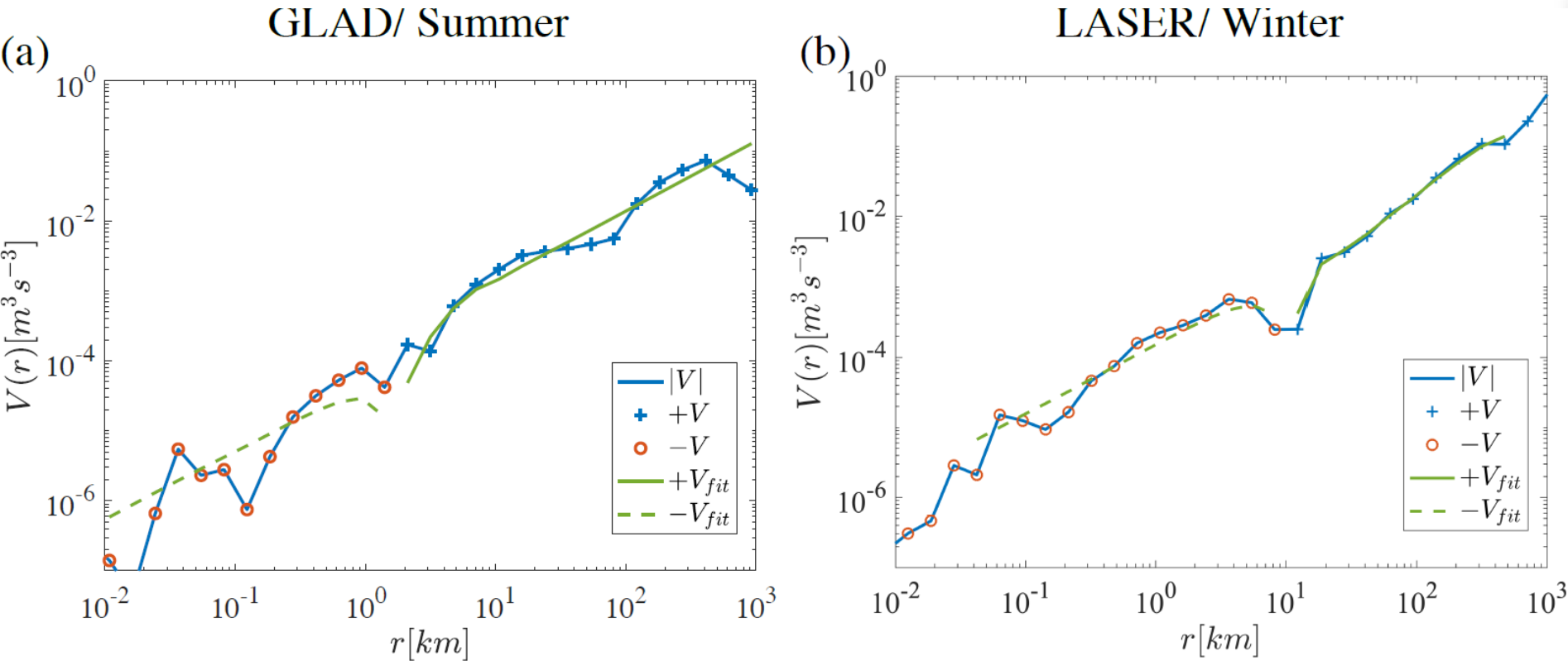


Figure: ocean drifter experiments. (with Balwada, Marino and Feraco)

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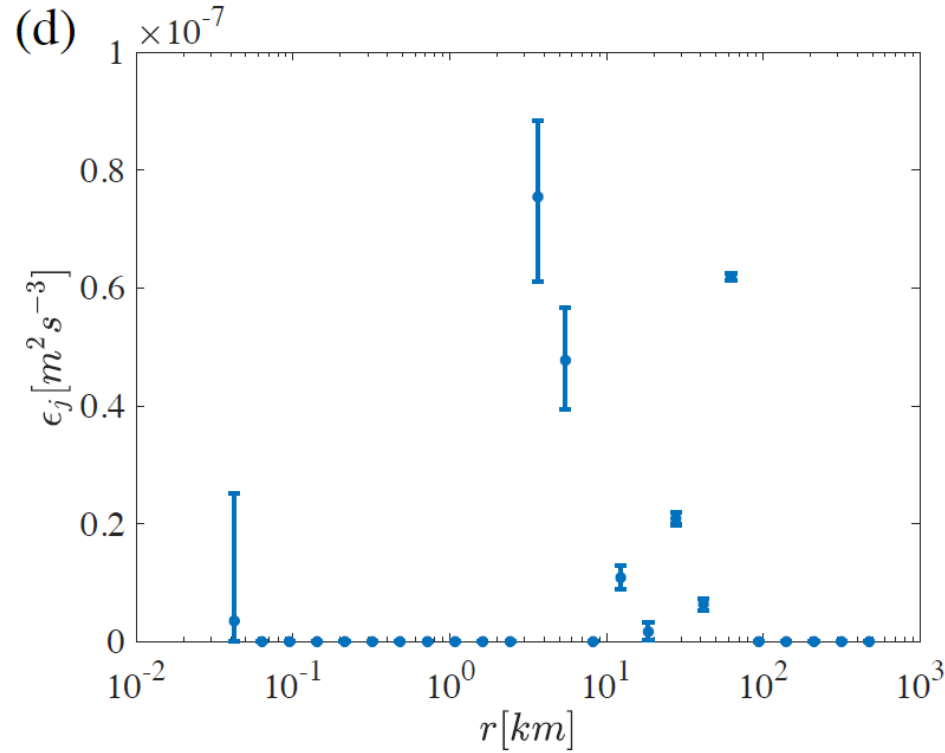
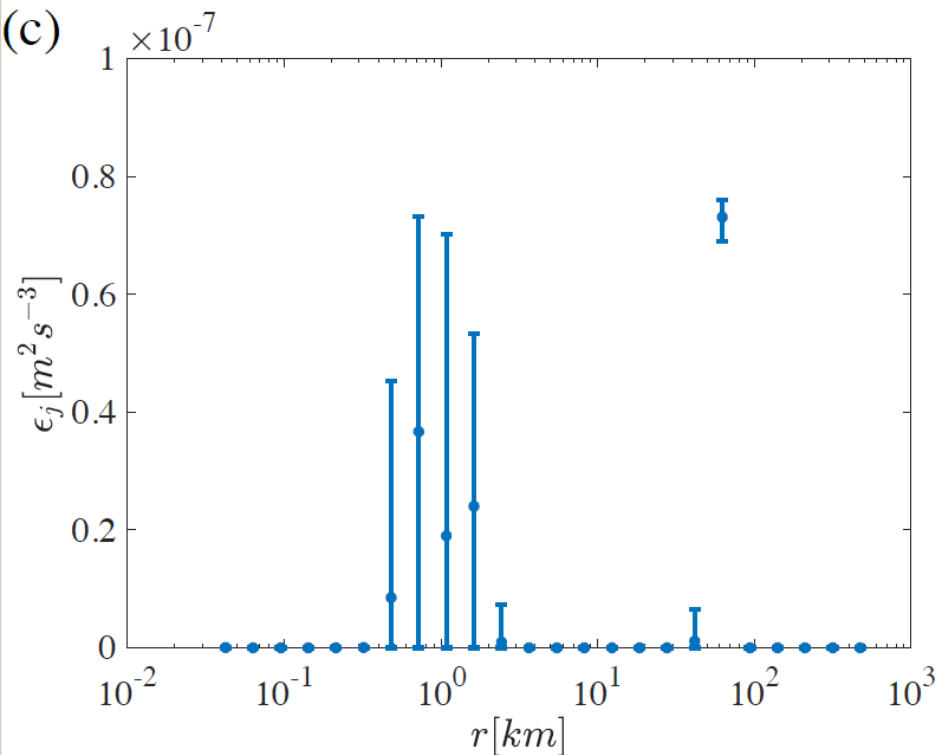


$$V = 2\epsilon_u r - 4 \sum_{j=1}^N \frac{\epsilon_j}{k_{fj}} J_1(k_{fj} r)$$

Figure: third-order structure functions from measurements.

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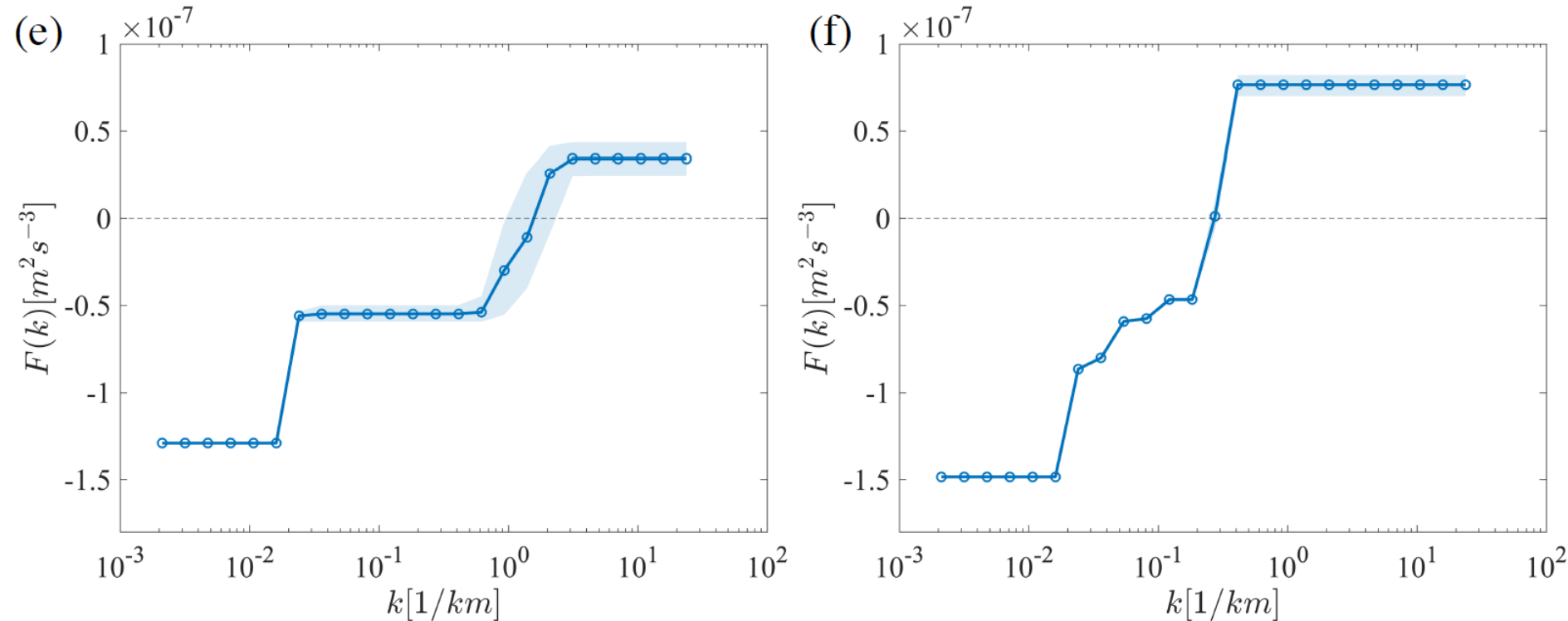


$$V = 2\epsilon_u r - 4 \sum_{j=1}^N \frac{\epsilon_j}{k_{fj}} J_1(k_{fj} r)$$

Figure: detected energy injection at different scales.

Oceanic turbulence

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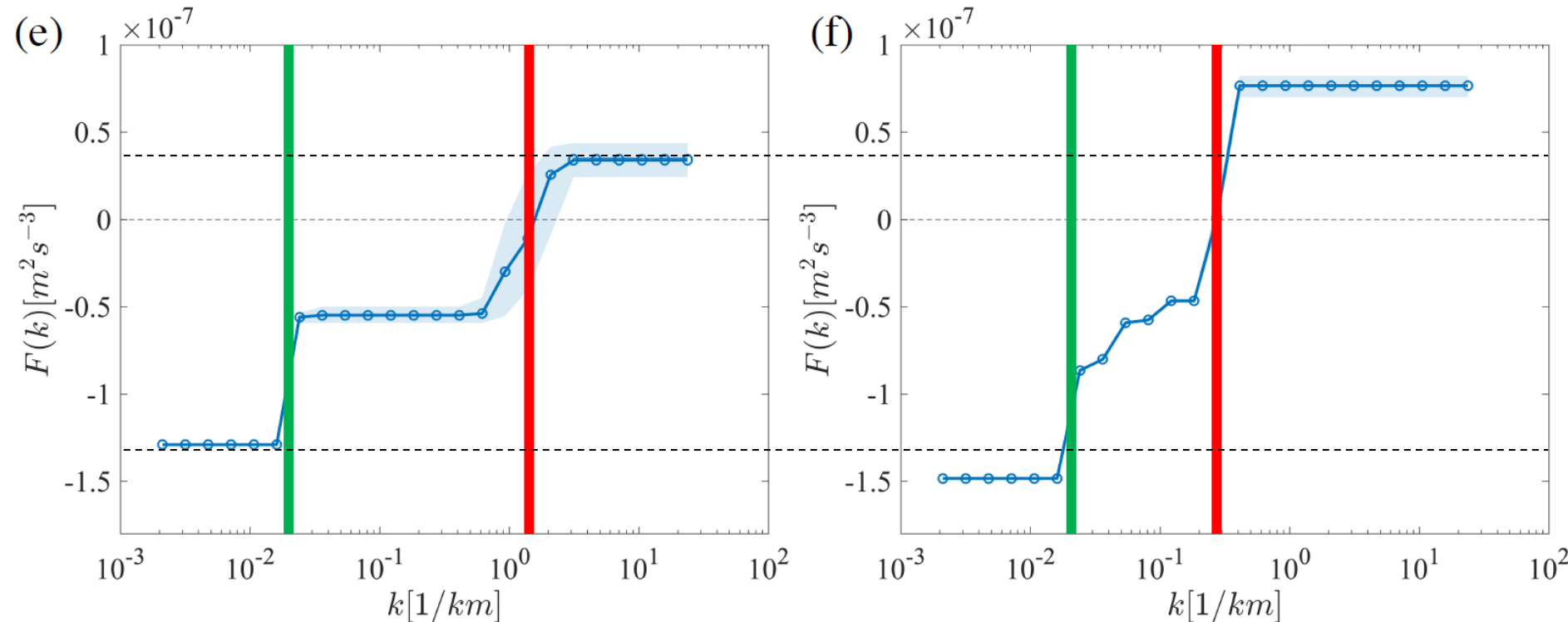


$$V = 2\epsilon_u r - 4 \sum_{j=1}^N \frac{\epsilon_j}{k_{fj}} J_1(k_{fj} r)$$

Figure: energy flux across scales.

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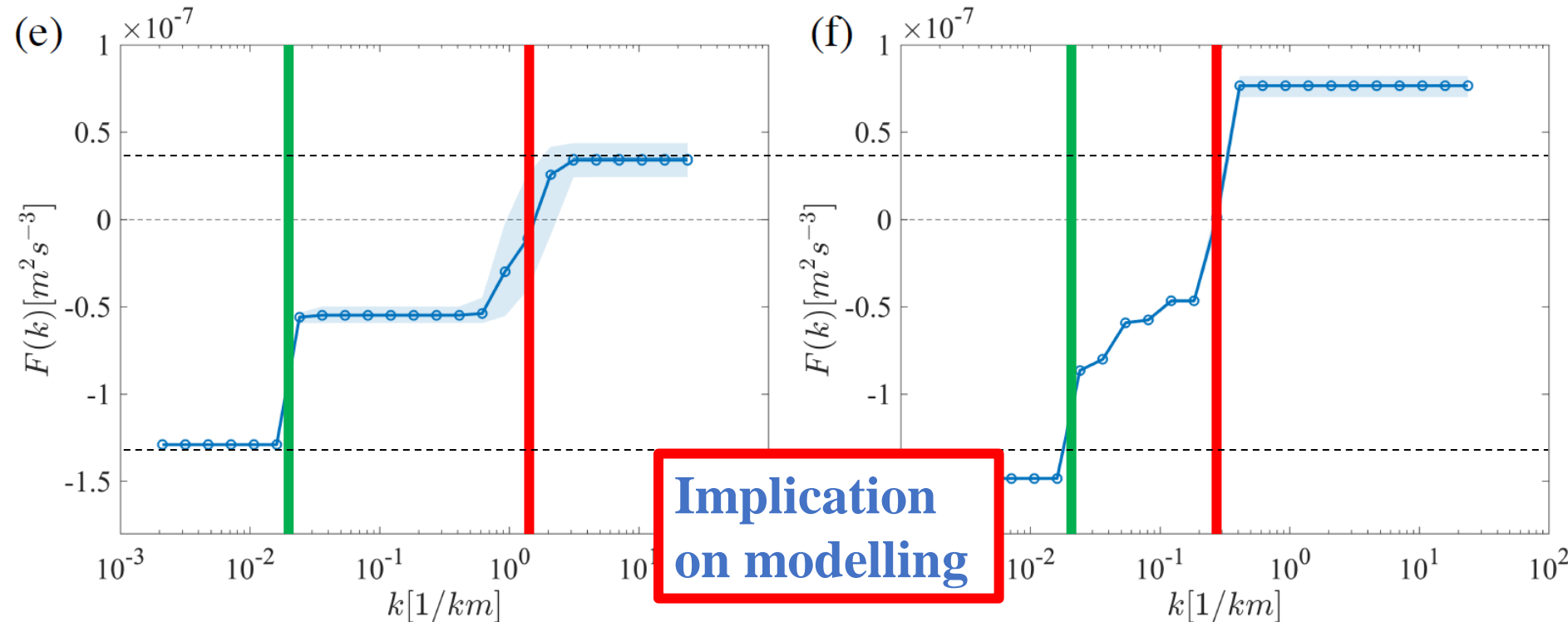
$$V = 2\epsilon_u r - 4 \sum_{j=1}^N \frac{\epsilon_j}{k_{fj}} J_1(k_{fj} r)$$

Figure: energy flux across scales.
baroclinic instability?:
Ocean depth: —
Mixed layer: —



- Large and small-scale energy injection corresponds to baroclinic instability associated with ocean depth and mixed-layer depth, respectively. (winter has deeper mixed layer)

Oceanic turbulence

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Figure: energy flux across scales.
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Summary

- For third-order structure function in *isotropic* turbulence we obtain a *global forcing-scale-resolving* theory that captures the *bidirectional energy transfer*.
- Application to drifter measurements in GoM, we find
 - Bidirectional energy flux;
 - Stronger downscale flux and larger energy-injection scale in the winter;
 - Forcing scales relate to deformation radius.
- Flaws: inhomogeneity, anisotropy, rotation, stratification, mean flow...

Thank you very much!